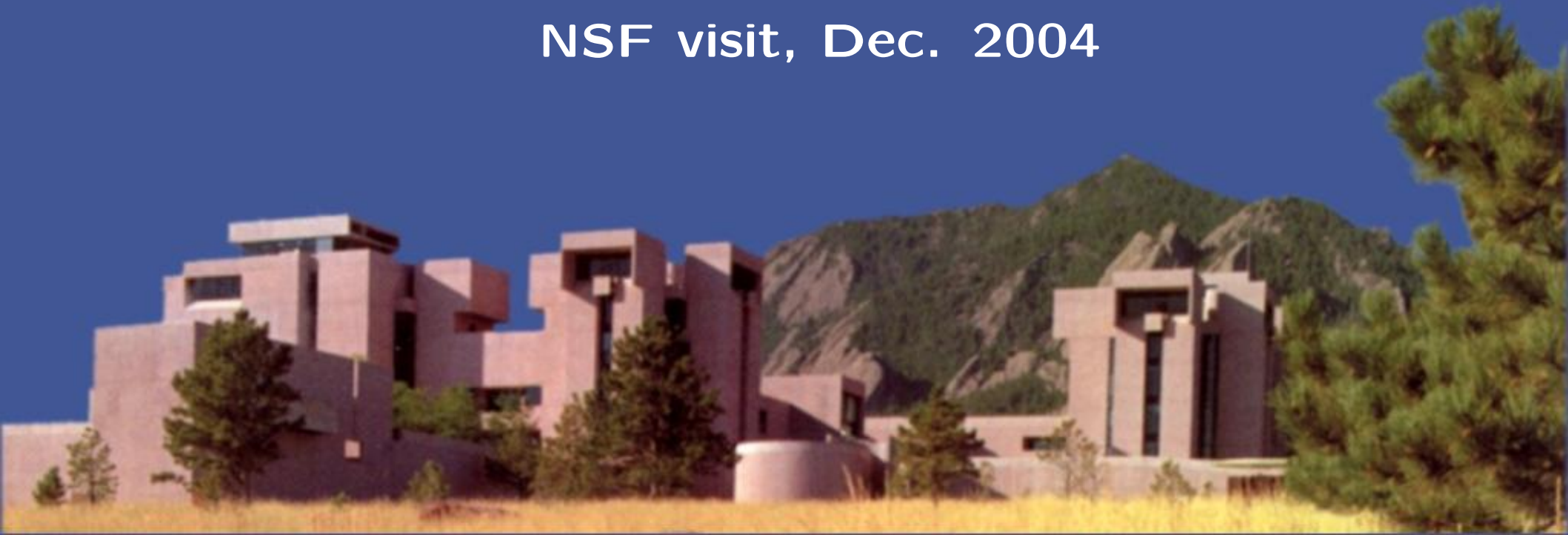


Overview of three Projects During my PostDoc Life

NSF visit, Dec. 2004



Introduction

- Tapering in ensemble Kalman filter
Thomas Bengtsson (Berkeley), Jeff Anderson (NCAR)
- Tapering in spatial prediction
Doug Nychka (NCAR), Marc Genton (Texas A&M)
- Bayes model for climate projections
Tom Wigley (NCAR), Steve Sain (CU Denver)

Tapering in Ensemble KF

Let the state-space model be

$$\begin{aligned}\mathbf{y}_t &= \mathbf{H}_t \mathbf{x}_t + \mathbf{w}_t & \mathbf{w}_t &\sim \mathcal{N}_r(\mathbf{0}, \mathbf{R}_t) \\ \mathbf{x}_t &= \mathbf{G}_t \mathbf{x}_{t-1} + \mathbf{v}_t & \mathbf{v}_t &\sim \mathcal{N}_q(\mathbf{0}, \mathbf{Q}_t)\end{aligned}$$

and assume $[\mathbf{x}_t | \mathbf{y}_{t-1}, \dots, \mathbf{y}_0] \sim \mathcal{N}_q(\mathbf{x}_t^f, \mathbf{P}_t^f)$.

“Filter” the state \mathbf{x}_t given the new observation \mathbf{y}_t :

$$\begin{aligned}[\mathbf{x}_t | \mathbf{y}_t, \dots, \mathbf{y}_0] &\sim \mathcal{N}_q(\mathbf{x}_t^a, \mathbf{P}_t^a) \\ \mathbf{x}_t^a &= \mathbf{x}_t^f + \mathbf{K}_t (\mathbf{y}_t - \mathbf{H}_t \mathbf{x}_t^f) \\ \mathbf{P}_t^a &= (\mathbf{I} - \mathbf{K}_t \mathbf{H}_t^\top) \mathbf{P}_t^f \\ \mathbf{K}_t &= \mathbf{P}_t^f \mathbf{H}_t^\top (\mathbf{H}_t \mathbf{P}_t^f \mathbf{H}_t^\top + \mathbf{R}_t)^{-1}\end{aligned}$$

Instead, update a sample from the forecast distribution.

Tapering in Ensemble KF

Let $\{\mathbf{x}_i\}$, $i = 1, \dots, n$, iid $\mathcal{N}_q(\mathbf{0}, \Sigma)$ with n small and q huge. Denote \mathbf{S} the sample covariance matrix.

For a positive definite matrix \mathbf{C} , consider $\mathbf{S} \circ \mathbf{C}$.

With an appropriate matrix norm describe the dependence of

$$\left\| \Sigma - \tilde{\Sigma} \right\| \quad \left\| (\Sigma + \mathbf{I})^{-1} - (\tilde{\Sigma} + \mathbf{I})^{-1} \right\|$$

with

$$\tilde{\Sigma} = \mathbf{S} \quad \text{or} \quad \tilde{\Sigma} = \mathbf{S} \circ \mathbf{C}$$

on ensemble size n , state dimension q , covariance matrix Σ and taper \mathbf{C} .

Tapering in Prediction

Suppose a spatial process Z with $E(Z) = \mathbf{0}$ and $\text{Cov}(Z) = \mathbf{C}$.

Objective: predict Z at many locations
given the observations $\mathbf{Z} = (Z(\mathbf{x}_1), \dots, Z(\mathbf{x}_n))^T$.

The kriging predictor (BLUP) is

$$\hat{Z}(\mathbf{x}_0) = \mathbf{c}^T \mathbf{C}^{-1} \mathbf{Z}$$

with $\mathbf{c}_i = \text{Cov}(Z(\mathbf{x}_0), Z(\mathbf{x}_i))$.

Approach:

Introduce sparseness in \mathbf{C} to gain computational advantages

Tapering in Prediction

Taper Condition: Let f_θ be the spectral density of the taper covariance, C_θ , and for some $\epsilon > 0$ and $M < \infty$

$$f_\theta(\rho) < \frac{M}{(1 + \rho^2)^{\nu+d/2+\epsilon}}$$

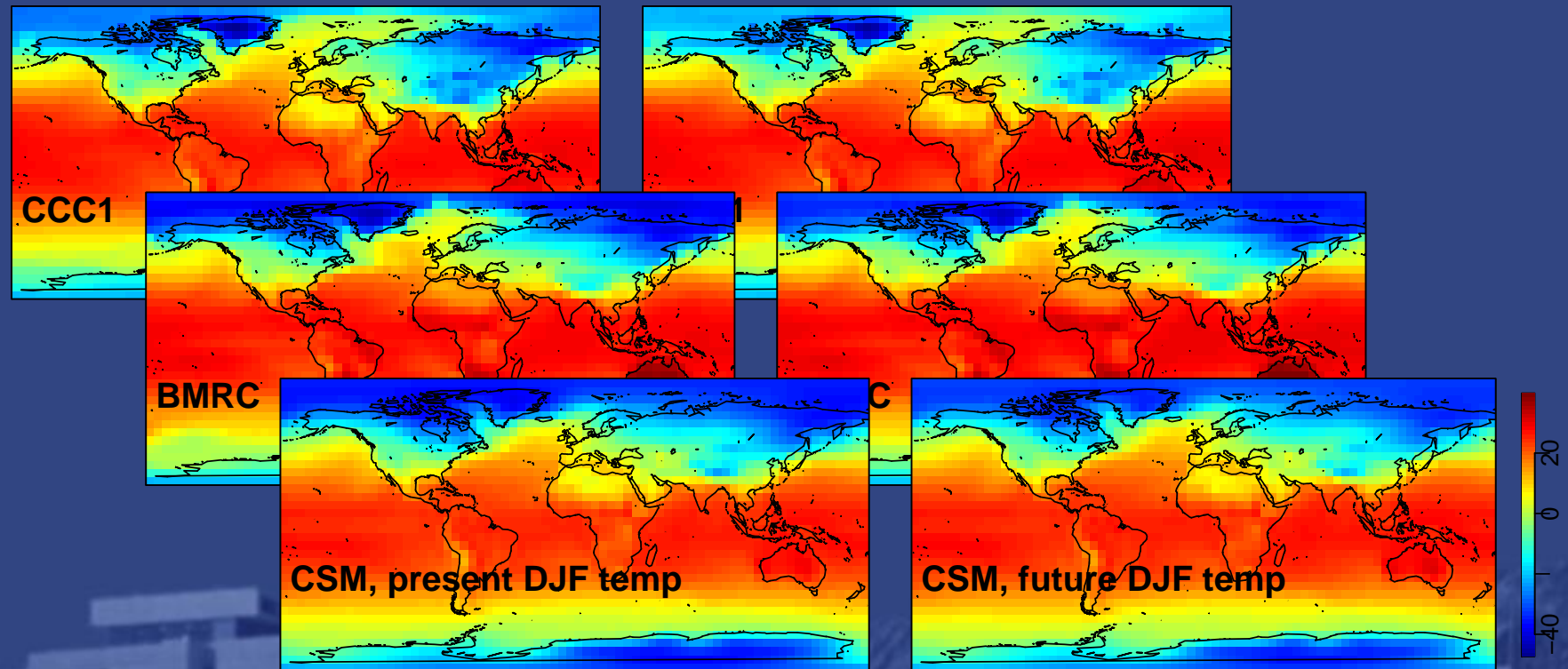
Taper Theorem: Assume that $C_{\alpha,\nu}$ is a Matérn covariance with smoothness parameter ν and that the Taper Conditions and ... hold. Then

$$\lim_{n \rightarrow \infty} \frac{\text{MSE}(\mathbf{x}^*, C_{\alpha,\nu} C_\theta)}{\text{MSE}(\mathbf{x}^*, C_{\alpha,\nu})} = 1$$

$$\lim_{n \rightarrow \infty} \frac{\rho(\mathbf{x}^*, C_{\alpha,\nu} C_\theta)}{\text{MSE}(\mathbf{x}^*, C_{\alpha,\nu})} = 1$$

Models for Climate Projections

AOGCMs: Atmospheric-Ocean General Circulation Models



Goal: Probabilistic description of modeled climate change

Models for Climate Projections

“Large” linear fixed effects model

$$\mathbf{X}_0 = \boldsymbol{\mu}_x + \boldsymbol{\varepsilon} \quad (\text{observed present climate})$$

$$\mathbf{X}_i = \boldsymbol{\mu}_x + \mathbf{u}_i + \boldsymbol{\sigma}_i \quad (\text{simulated present climate})$$

$$\mathbf{Y}_i = \boldsymbol{\mu}_y + \mathbf{v}_i + \boldsymbol{\nu}_i \quad (\text{simulated future climate})$$

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$$\boldsymbol{\mu}_x = \mathbf{M}_c \boldsymbol{\theta} \quad \boldsymbol{\varepsilon} \sim \mathcal{N}_n(\mathbf{0}, \mathbf{S})$$

$$\boldsymbol{\mu}_x = \mathbf{M}_c \boldsymbol{\theta} \quad \mathbf{u}_i = \mathbf{M}_b \boldsymbol{\beta}_i \quad \boldsymbol{\sigma}_i \sim \mathcal{N}_n(\mathbf{0}, \boldsymbol{\Sigma}_i)$$

$$\boldsymbol{\mu}_y = \mathbf{M}_c \boldsymbol{\eta} \quad \mathbf{v}_i = \mathbf{M}_b \boldsymbol{\gamma}_i \quad \boldsymbol{\nu}_i = \boldsymbol{\omega}_i + \rho_i \boldsymbol{\sigma}_i$$

$$\boldsymbol{\omega}_i \sim \mathcal{N}_n(\mathbf{0}, \boldsymbol{\Omega}_i) \quad \boldsymbol{\omega}_i \perp \boldsymbol{\sigma}_i$$

Use a Gibbs sampler to obtain posterior climate changes.

Opportunity

- Collaboration with scientists
Ammann, Anderson, Baker, Berner, Cassou, Collins, Fournier, Mahowald, Mearns, Meehl, Saravanan, Schimel, Snyder, Tribbia, Wigley, . . .
- Pursue cutting edge research
KriSp, IPCC, . . .
- Establishing a scientific network
“nobody ever leaves NCAR” . . .