

An Introduction to Climate Modeling

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Outline

- What is Climate & why do we care
- Hierarchy of atmospheric modeling strategies
 - 1D Radiative Convective models
 - 3D General Circulation models (GCMs)
- Conceptual Framework for General Circulation Models
- Scale interaction problem
 - concept of resolvable and unresolvable scales of motion
- Parameterization of physical processes
 - approaches rooted in budgets of conserved variables
- Model Validation and Model Solutions

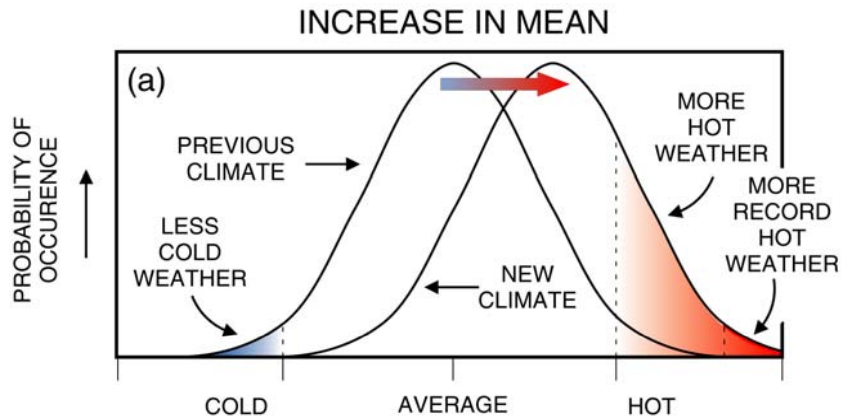


Question 1: How can we predict Climate (50 yrs)
if we can't predict Weather (10 days)?

Question 2: What is Climate?

- Average Weather
- Record high and low temperatures
- The temperature range
- Distribution of possible weather
- Extreme events

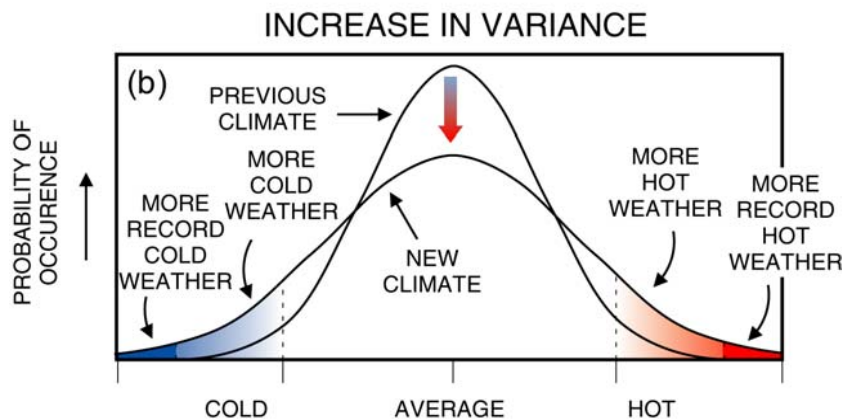
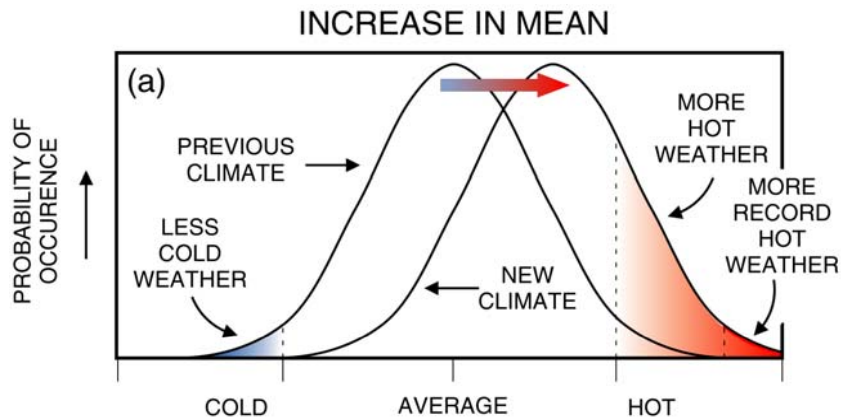




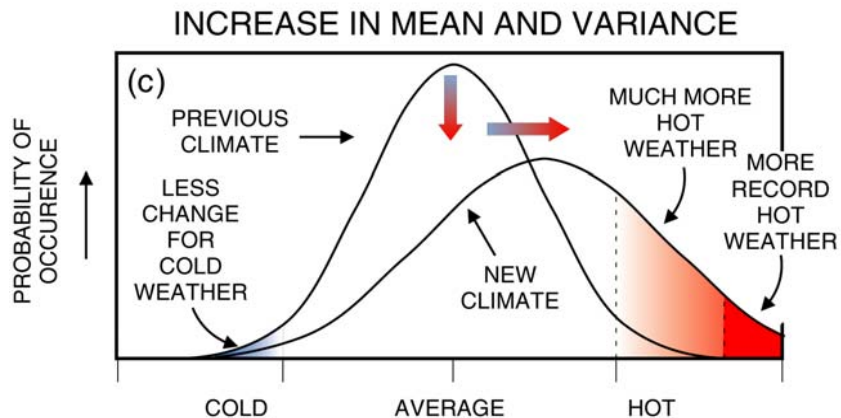
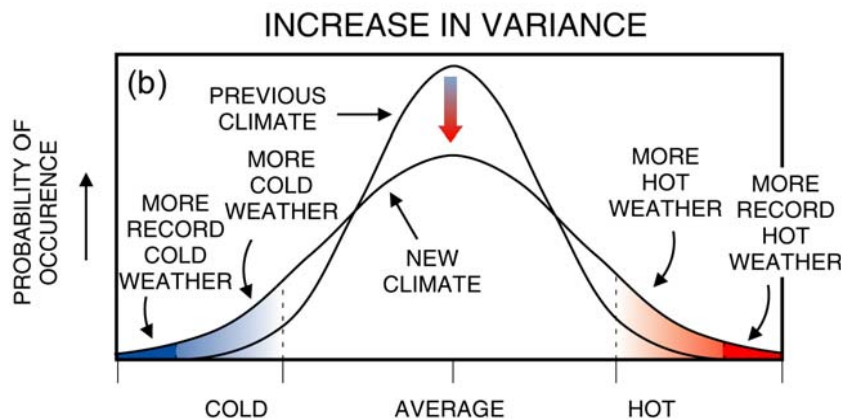
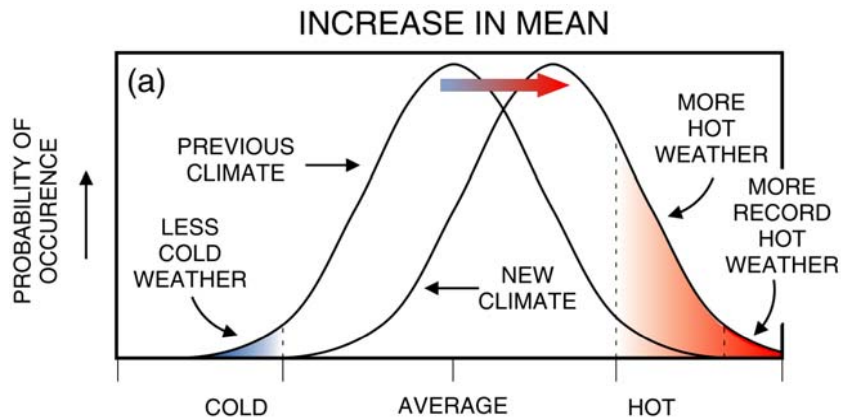
(1) What is Climate?

***Climate change
and its manifestation
in terms of weather
(climate extremes)***

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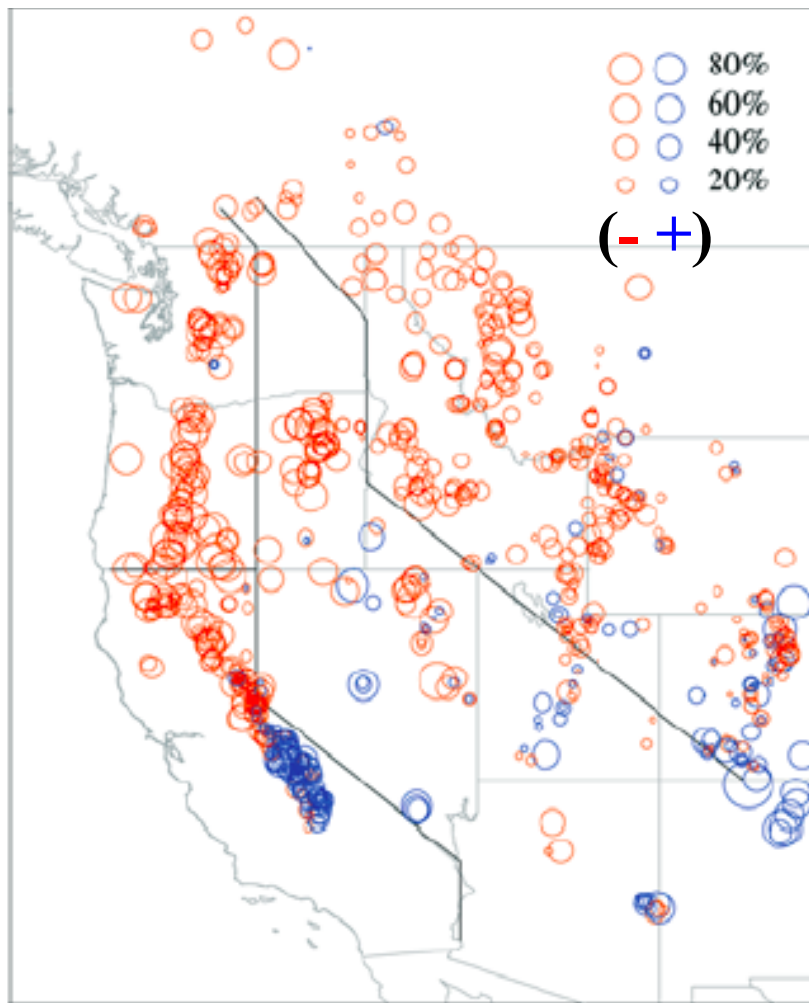
***Climate change
and its manifestation
in terms of weather
(climate extremes)***



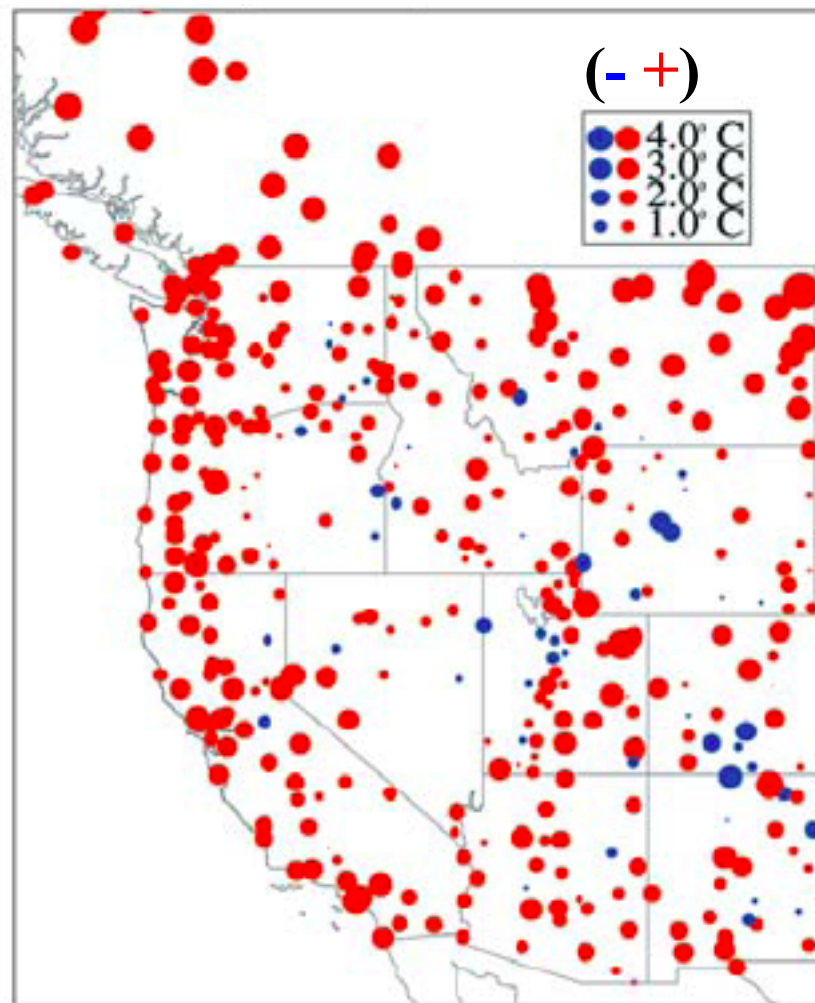
Impacts of Climate Change

Observed Change 1950-1997

Snowpack

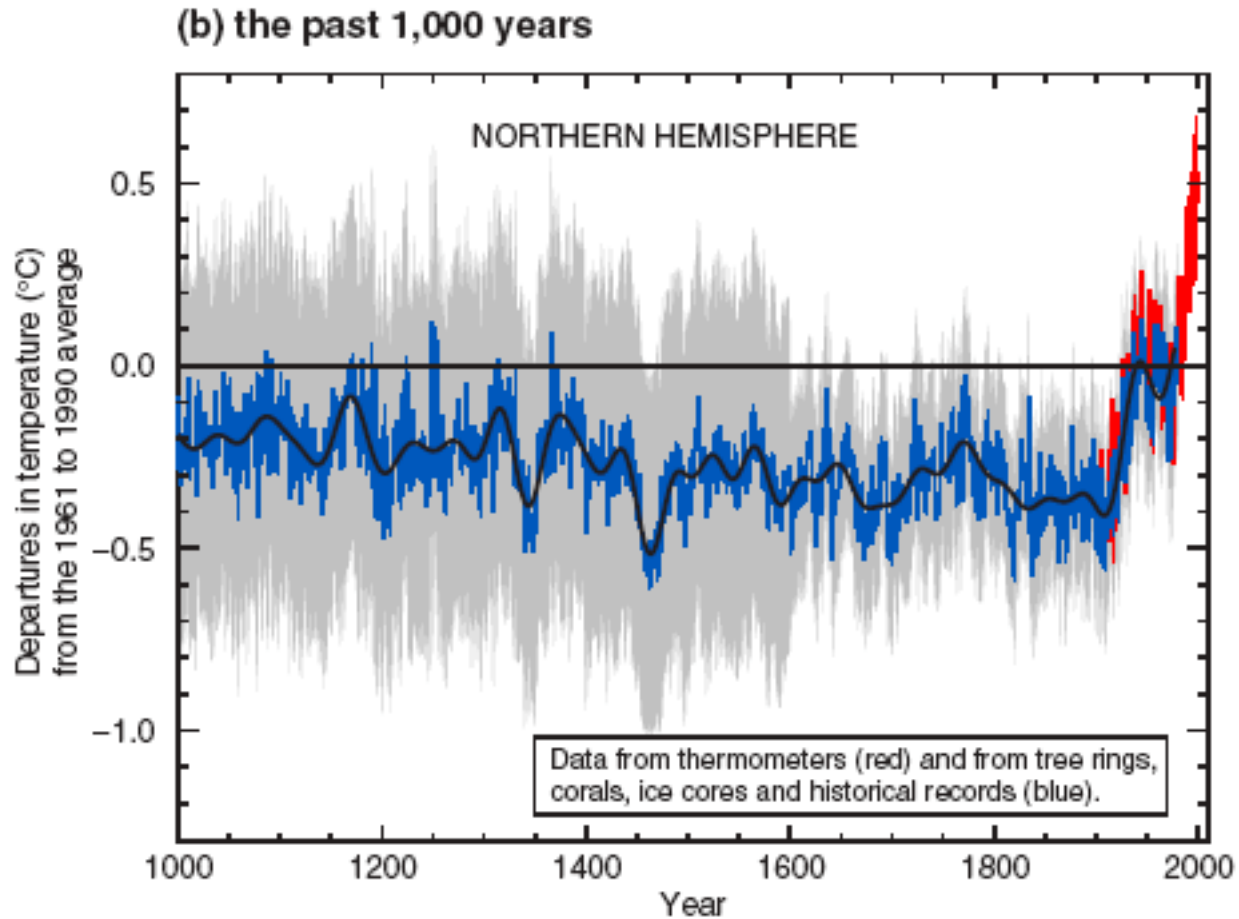


Temperature



Mote et al 2005

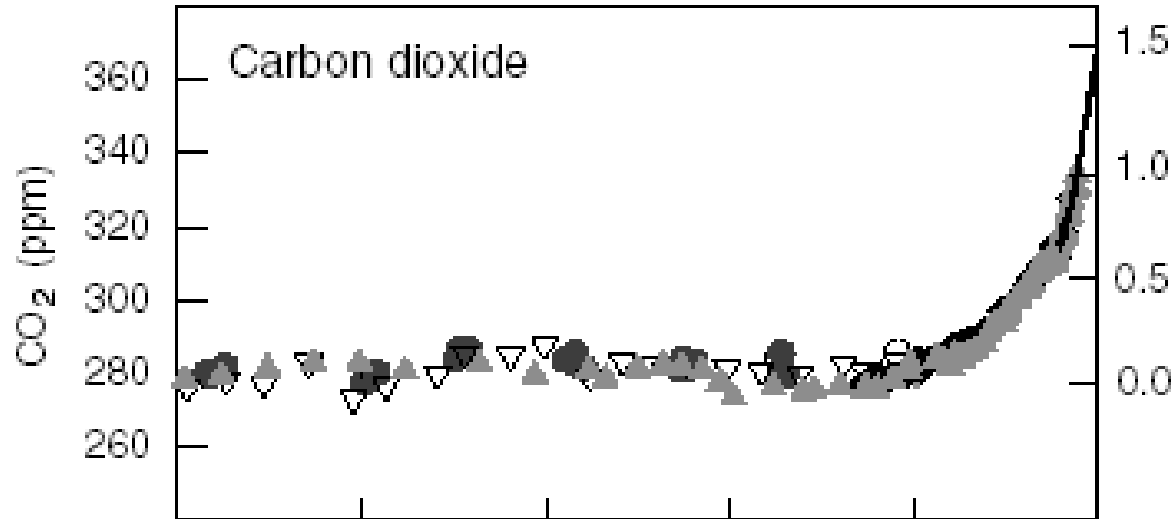
Observed Temperature Records



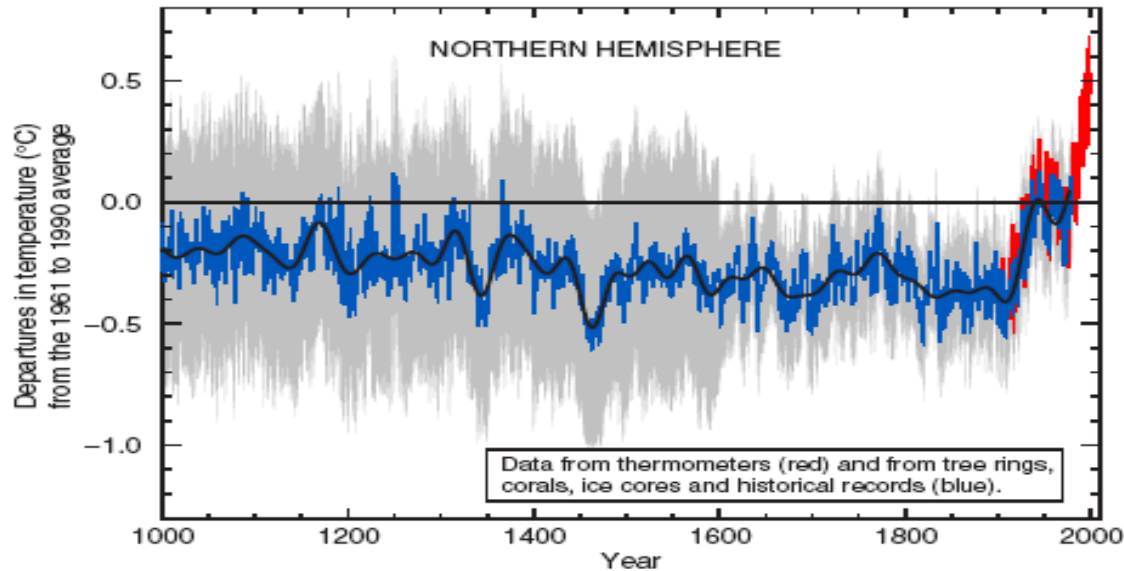
IPCC, 3rd Assessment, Summary For Policymakers

'Anthropogenic' Changes

Atmospheric concentration



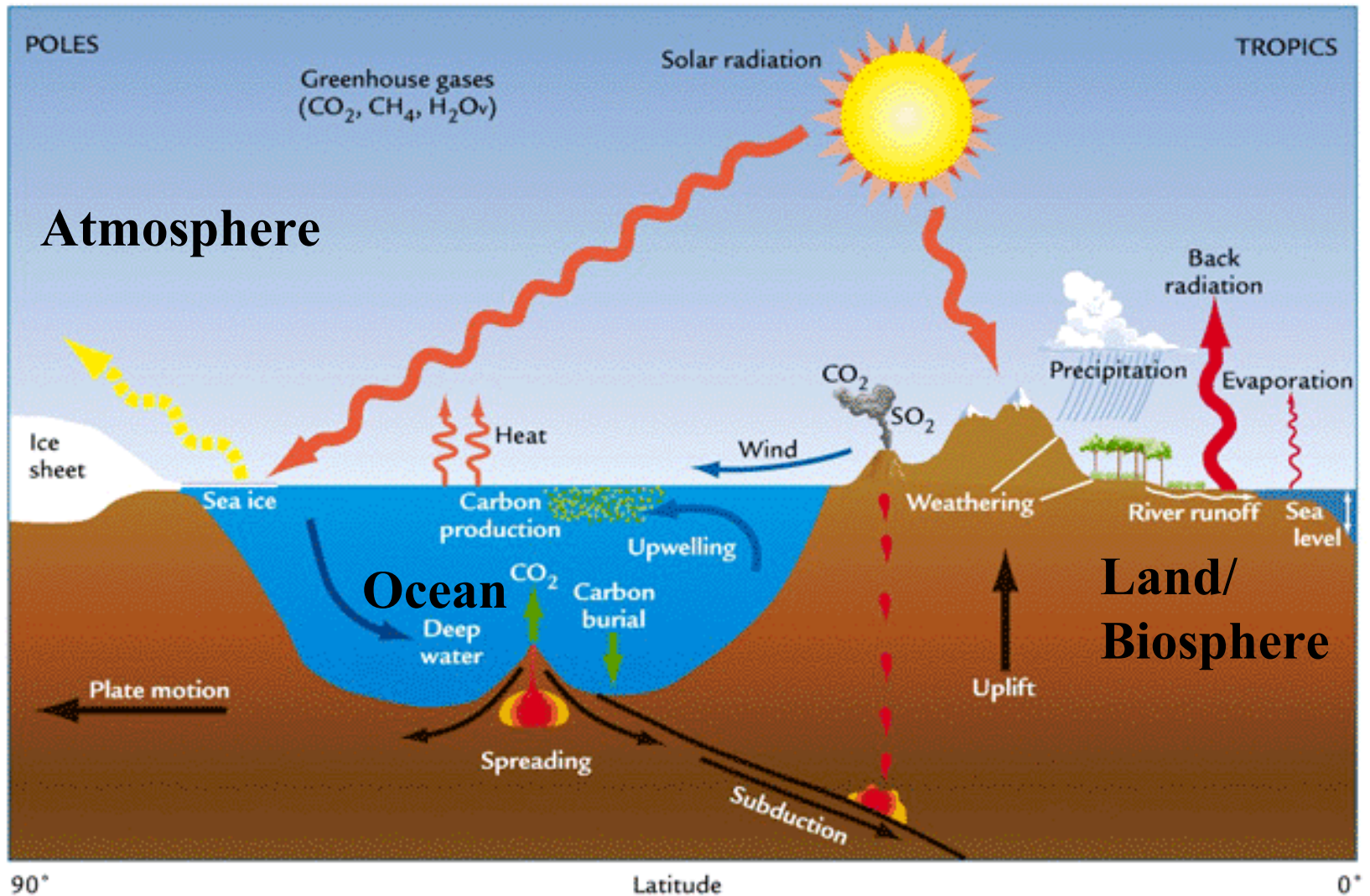
(b) the past 1,000 years



Radiative Forcing (Wm⁻²)



The Earth's climate system



90°

Latitude

0°

Pole

Equator

Principles of Atmospheric Modeling

- Scientific basis for atmospheric simulation
 - rooted in laws of classical mechanics/thermodynamics
 - developed during 18th and 19th centuries (see Thompson, 1978)
 - early mathematical model described by Arrhenius (1896)
 - surface energy balance model
- Two modeling approaches developed over last century
 - based on energy balance requirements
 - dynamical models (e.g., explicit transports)



Conceptual Framework for Modeling

- Can't resolve all scales, so have to represent them
- Energy Balance / Reduced Models
 - Mean State of the System
 - Energy Budget, conservation, Radiative transfer
- Dynamical Models
 - Finite element representation of system
 - Fluid Dynamics on a rotating sphere
 - Basic equations of motion
 - Physical Parameterizations for moving energy

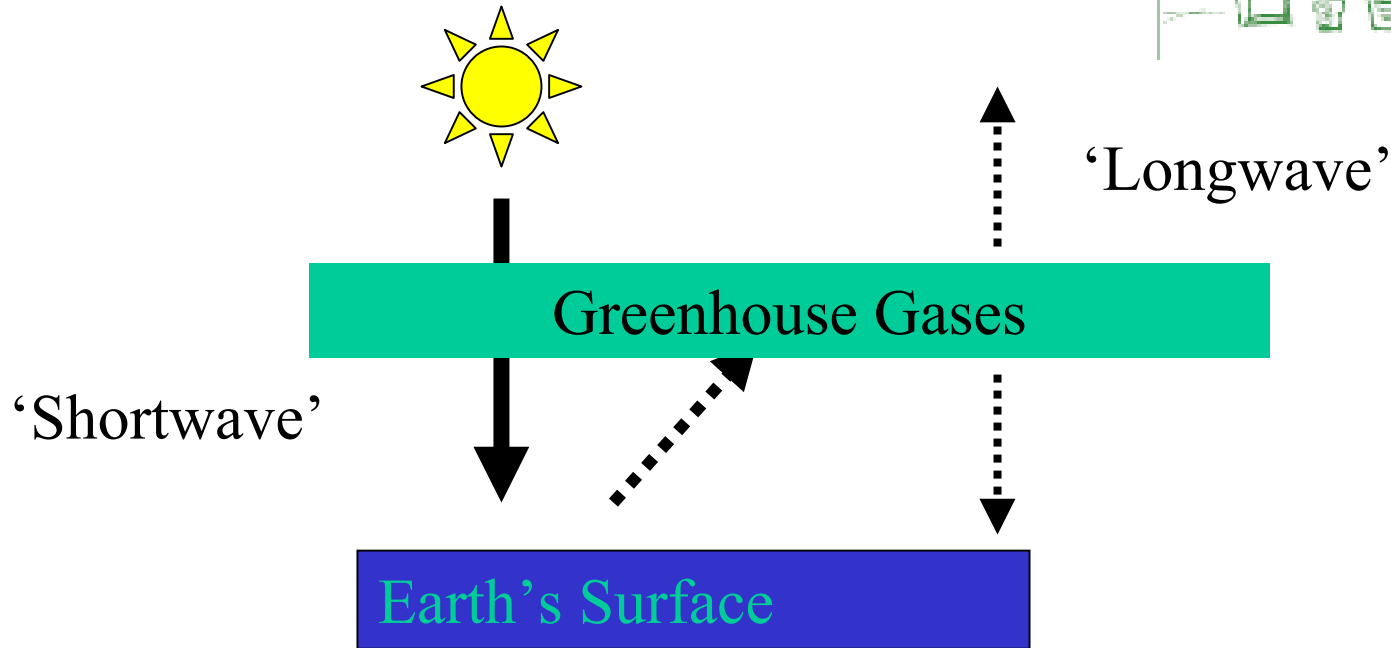


What is the greenhouse Effect?

Greenhouses Trap Heat

Analogous to Atmosphere:

- Glass is transparent in Visible(shortwave)
- Glass absorbs in Infrared (longwave)



Major Greenhouse Gases: Water Vapor, Carbon Dioxide, Methane

Atmospheric modeling hierarchy

Understanding has been aided by a hierarchy of approaches

Consider the flux form of thermodynamic energy equation

$$c_p \frac{\partial T}{\partial t} = -c_p \nabla \cdot (\mathbf{V}T) - c_p \frac{\partial(\omega T)}{\partial p} + c_p \frac{\kappa \omega T}{p} + Q_{\text{rad}} + Q_{\text{conv}} \quad (1)$$

where T - temperature; \mathbf{V} - horizontal wind vector; p - pressure; ω - vertical pressure velocity; Q_{rad} and Q_{conv} - net radiative and convective heating

- Simple Zero-Dimensional (Energy Balance) Climate Model

- Averaging (1) over horizontal and vertical space dimensions yields

$$c_p \frac{\partial \langle \hat{T} \rangle}{\partial t} = \langle S \rangle - \langle F \rangle$$

where S is net absorbed solar radiation and F is longwave radiation emitted to space

For a long-term stable climate, $\langle S \rangle - \langle F \rangle = 0$



Atmospheric modeling hierarchy

- Simple One-Dimensional (Radiative-Convective) Climate Model
 - Averaging (1) over horizontal space dimensions yields

$$c_p \frac{\partial \langle T \rangle}{\partial t} = \langle Q_{\text{rad}} \rangle + \langle Q_{\text{conv}} \rangle$$

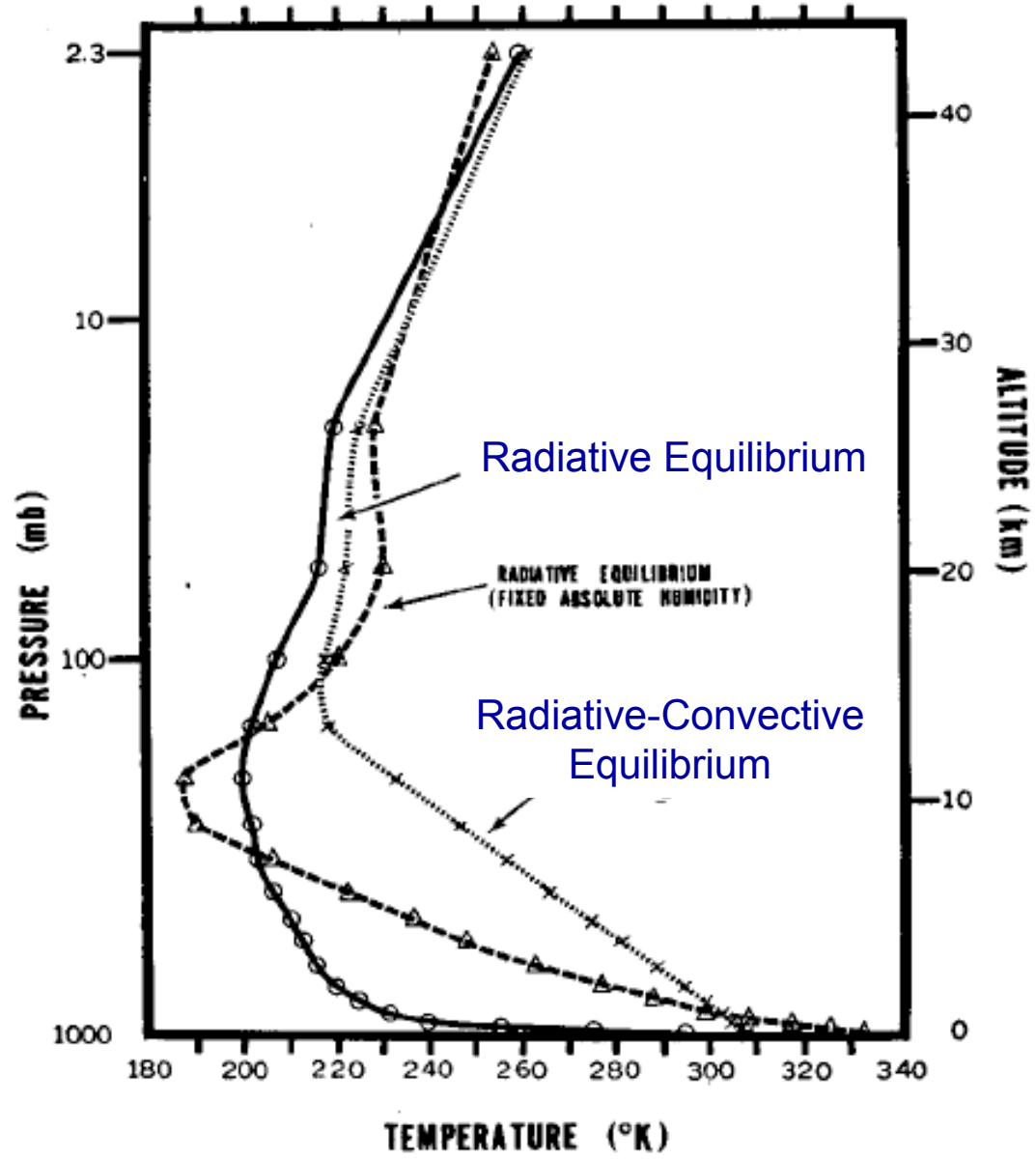
where a globally averaged vertical profile of T can be determined from expressions for $\langle Q_{\text{rad}} \rangle$ and $\langle Q_{\text{conv}} \rangle$

- Higher-order models determined by form of averaging operators



1D Radiative Convective Model

Manabe
& Wetherald 1967



1D models: Doubling CO2

TABLE 5. Change of equilibrium temperature of the earth's surface corresponding to various changes of CO₂ content of the atmosphere.

Change of CO ₂ content (ppm)	Fixed absolute humidity		Fixed relative humidity	
	Average cloudiness	Clear	Average cloudiness	Clear
300 → 150	-1.25	-1.30	-2.28	-2.80
300 → 600	+1.33	+1.36	+2.36	2.92

Manabe & Wetherald 1967

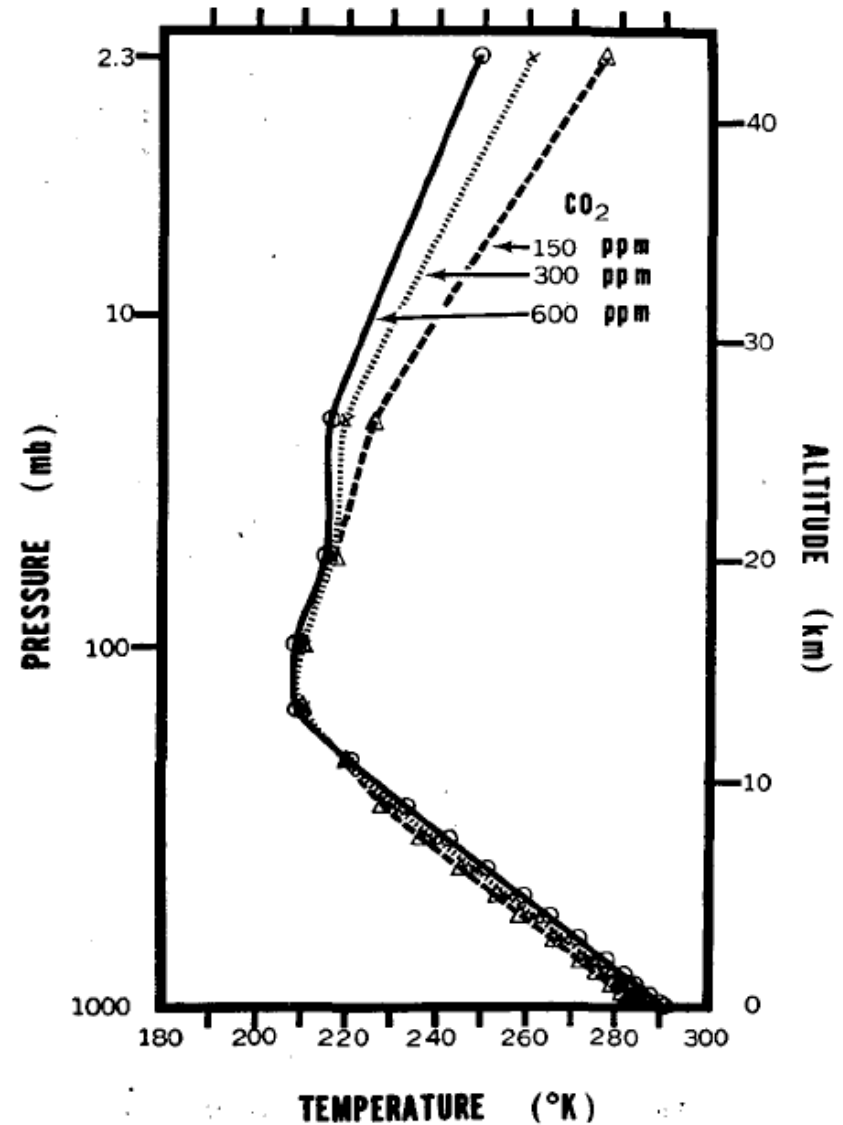
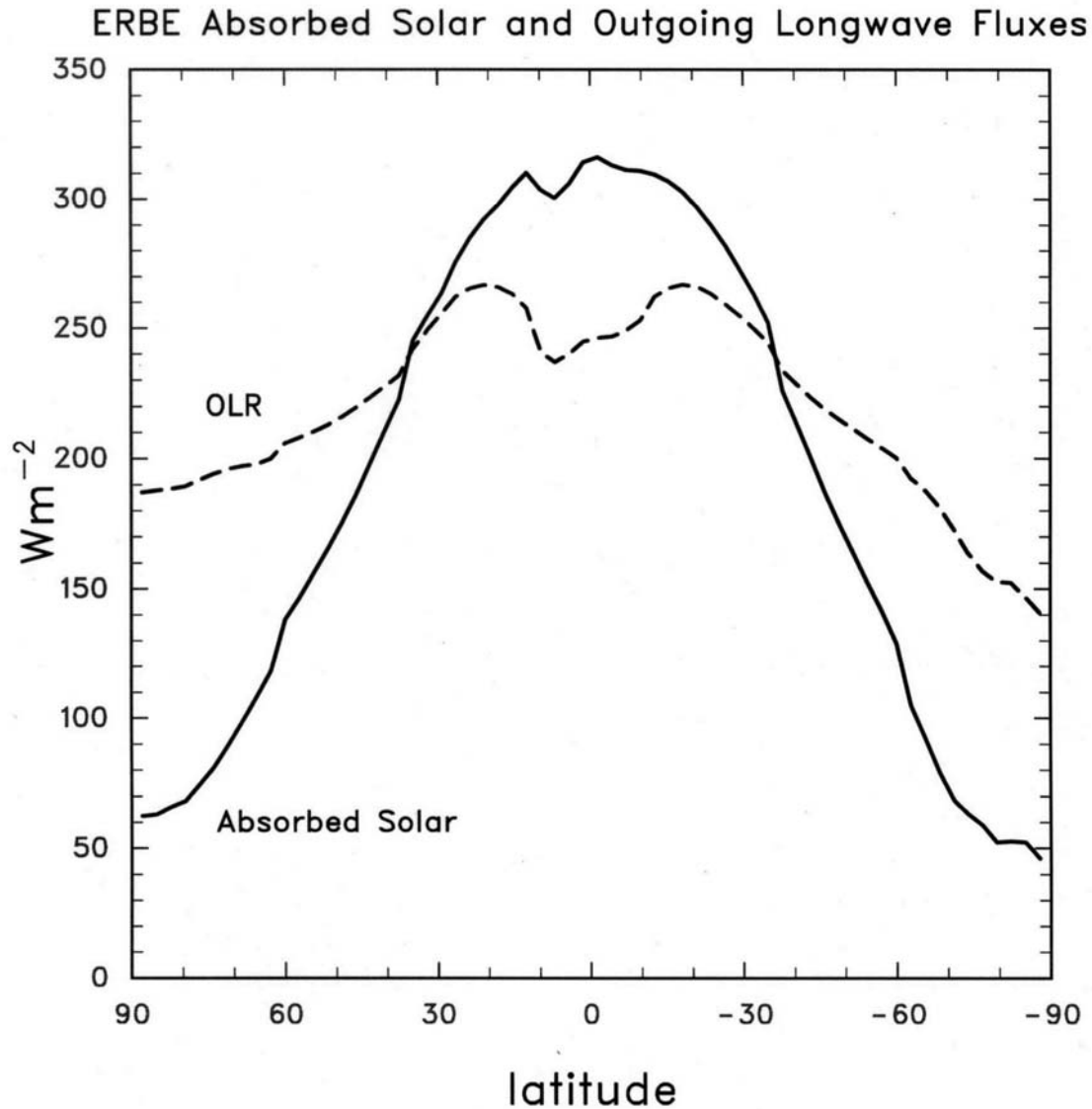


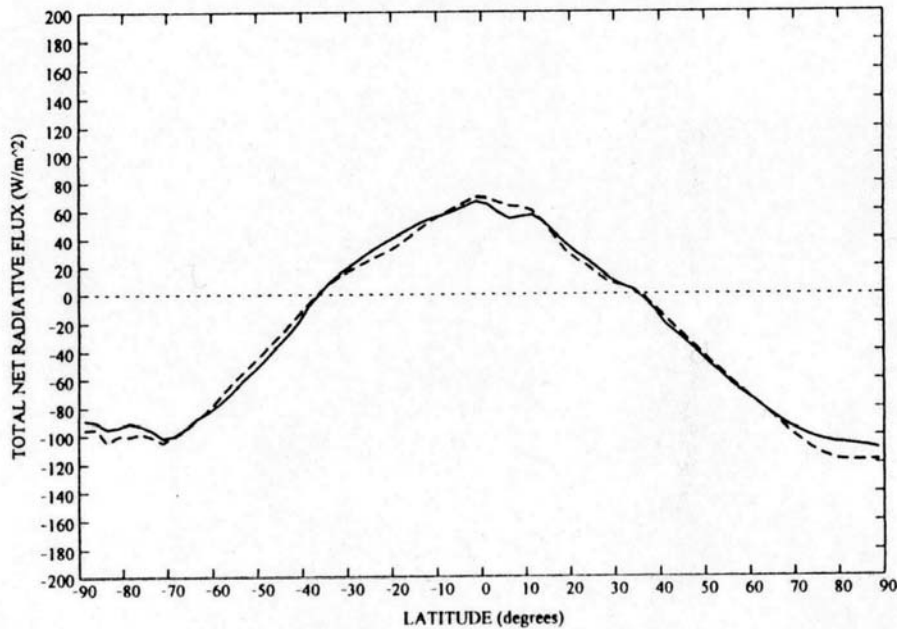
FIG. 16. Vertical distributions of temperature in radiative convective equilibrium for various values of CO₂ content.

Top of Atmosphere Radiation Component Fluxes

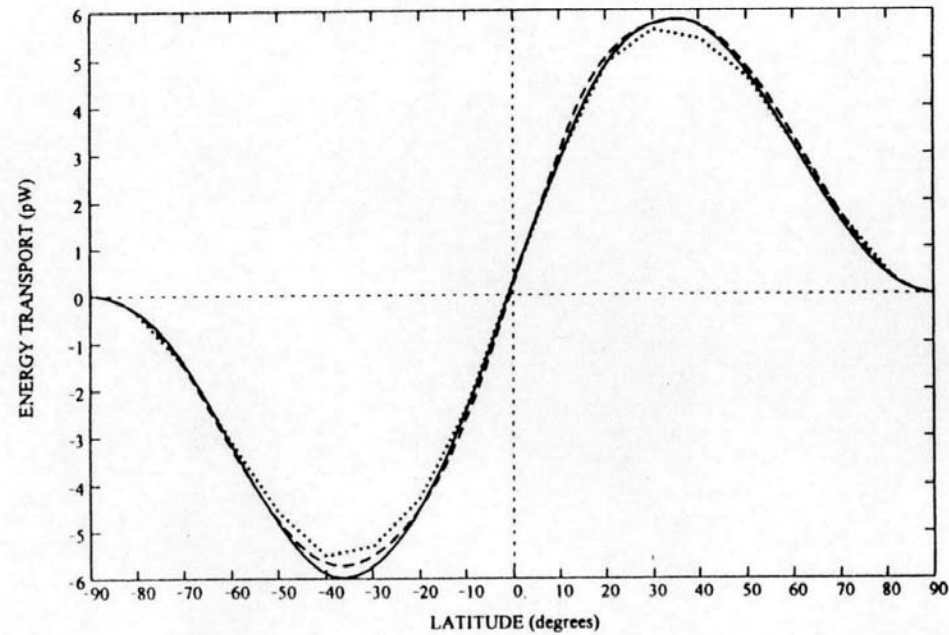


Top of Atmosphere Net Radiation Budget and Implied Meridional Energy Transport

TOTAL ANNUAL MEAN NET RADIATION AT TOA FOR ISCCP-FC & ERBE (ADJUSTED)



TOTAL NORTHWARD ENERGY TRANSPORT FROM ISCCP-FC, ERBE AND PEIXOTO & OORT



Zhang and Rossow (1997)

Atmospheric General Circulation Models and Climate Simulation

- Reduced models of the climate system
 - apply “averaging operator” to governing equations
- Atmospheric General Circulation Models (AGCMs)
 - simulate detailed “weather” fluctuations in the fluid system
 - day-to-day solution details are non-deterministic (Lorenz, 1962)
 - apply “averaging operator” to detailed solution sequence
 - utility lies in prediction of statistical properties of the fluid system
 - *chronological sequence of intermediate states unimportant*



Physical processes regulating climate

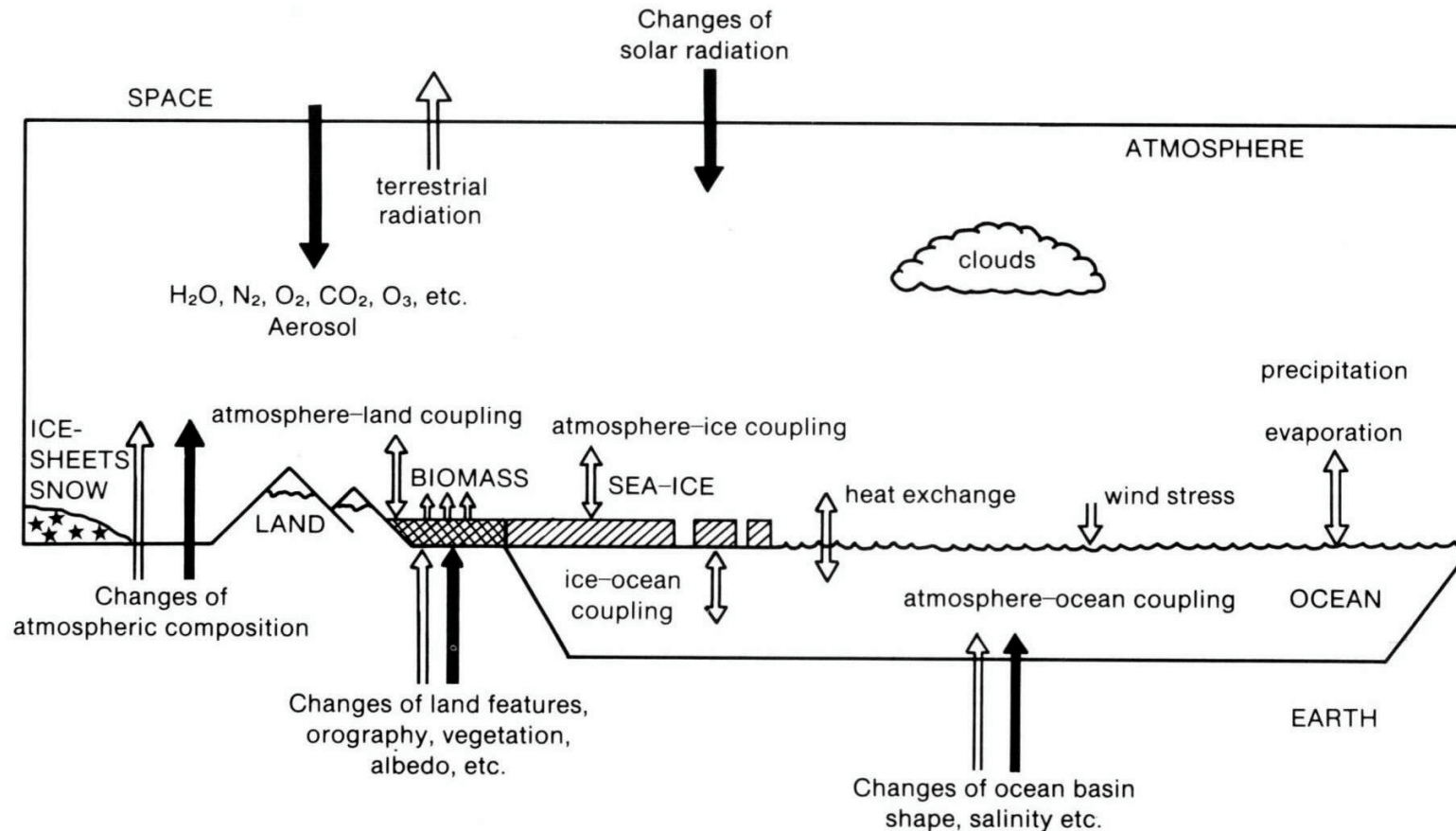


Figure 3.1: Schematic illustration of the components of the coupled atmosphere-ocean-ice-land climatic system. The full arrows are examples of external processes, and the open arrows are examples of internal processes in climatic change (from Houghton, 1984).

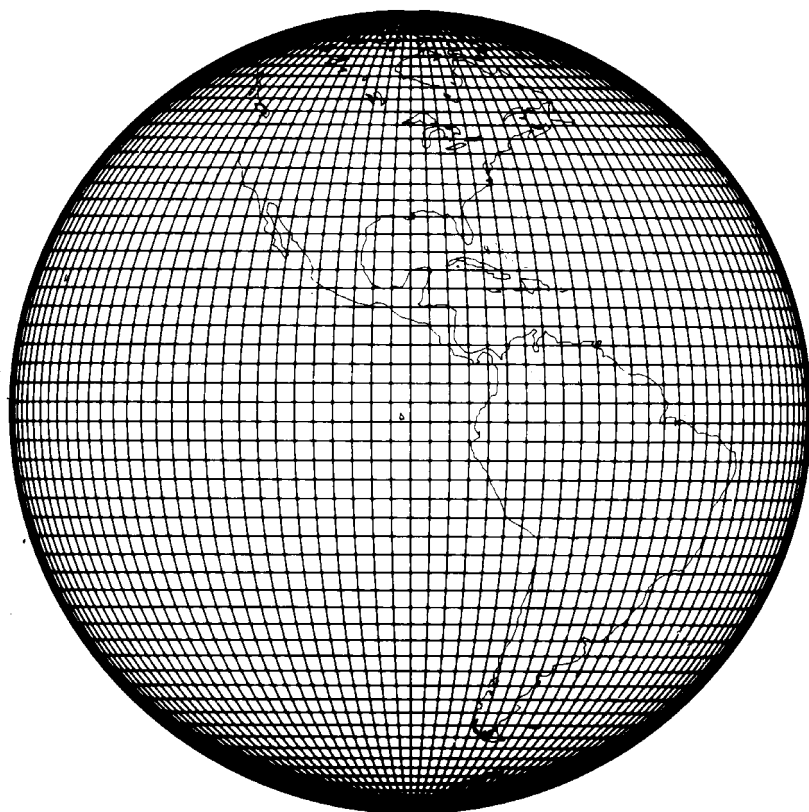
Modeling the Atmospheric General Circulation

Understanding of climate & global scale dynamics

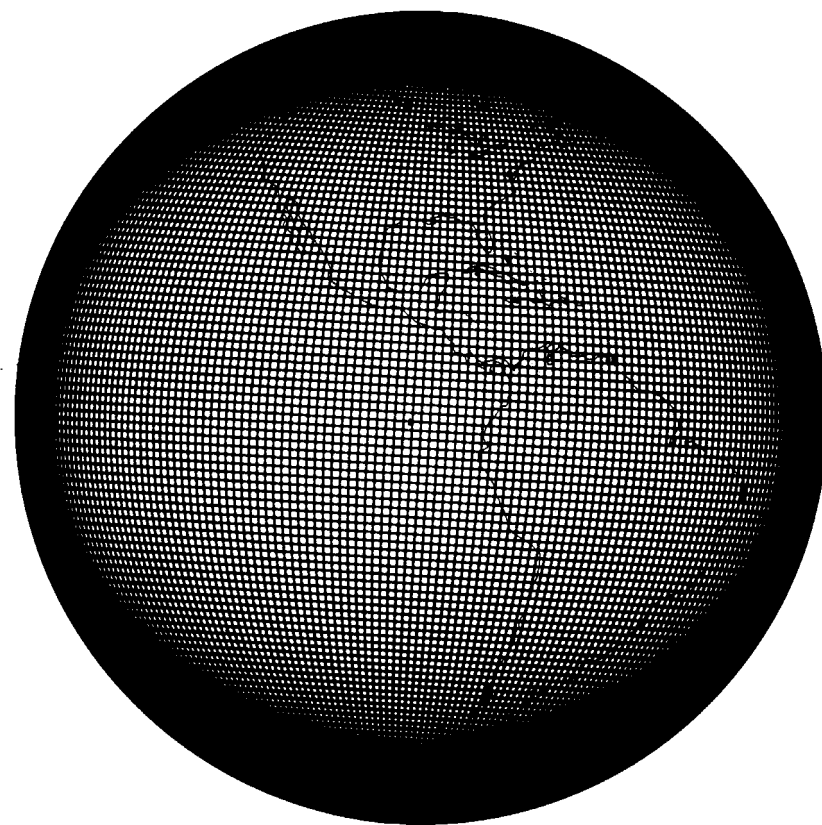
- atmospheric predictability/basic fluid dynamics
- physics/dynamics of phase change
- radiative transfer (aerosols, chemical constituents, etc.)
- atmospheric chemistry (trace gas sources/sinks, acid rain, etc.)
- interactions between the atmosphere and ocean (e.g., El Nino, etc.)
- solar physics (solar-terrestrial interactions, solar dynamics, etc.)
- impacts of anthropogenic and other biological activity



Examples of Global Model Resolution



Typical Climate Application



Next Generation Climate Applications



Meteorological Primitive Equations

- Applicable to wide scale of motions; $> 1\text{hour}$, $> 100\text{km}$

$$d\bar{\mathbf{V}}/dt + f\mathbf{k} \times \bar{\mathbf{V}} + \nabla\bar{\phi} = \mathbf{F}, \quad (\text{horizontal momentum})$$

$$d\bar{T}/dt - \kappa\bar{T}\omega/p = Q/c_p, \quad (\text{thermodynamic energy})$$

$$\nabla \cdot \bar{\mathbf{V}} + \partial\bar{\omega}/\partial p = 0, \quad (\text{mass continuity})$$

$$\partial\bar{\phi}/\partial p + R\bar{T}/p = 0, \quad (\text{hydrostatic equilibrium})$$

$$d\bar{q}/dt = S_q. \quad (\text{water vapor mass continuity})$$

Harmless looking terms \mathbf{F} , Q , and $S_q \implies$ “physics”



Global Climate Model Physics

Terms F , Q , and S_q represent physical processes

- Equations of motion, F
 - turbulent transport, generation, and dissipation of momentum
- Thermodynamic energy equation, Q
 - convective-scale transport of heat
 - convective-scale sources/sinks of heat (phase change)
 - radiative sources/sinks of heat
- Water vapor mass continuity equation
 - convective-scale transport of water substance
 - convective-scale water sources/sinks (phase change)



Model Physical Parameterizations

Physical processes breakdown:

- Moist Processes
 - Moist convection, shallow convection, large scale condensation
- Radiation and Clouds
 - Cloud parameterization, radiation
- Surface Fluxes
 - Fluxes from land, ocean and sea ice (from data or models)
- Turbulent mixing
 - Planetary boundary layer parameterization, vertical diffusion, gravity wave drag



Basic Logic in a GCM (Time-step Loop)

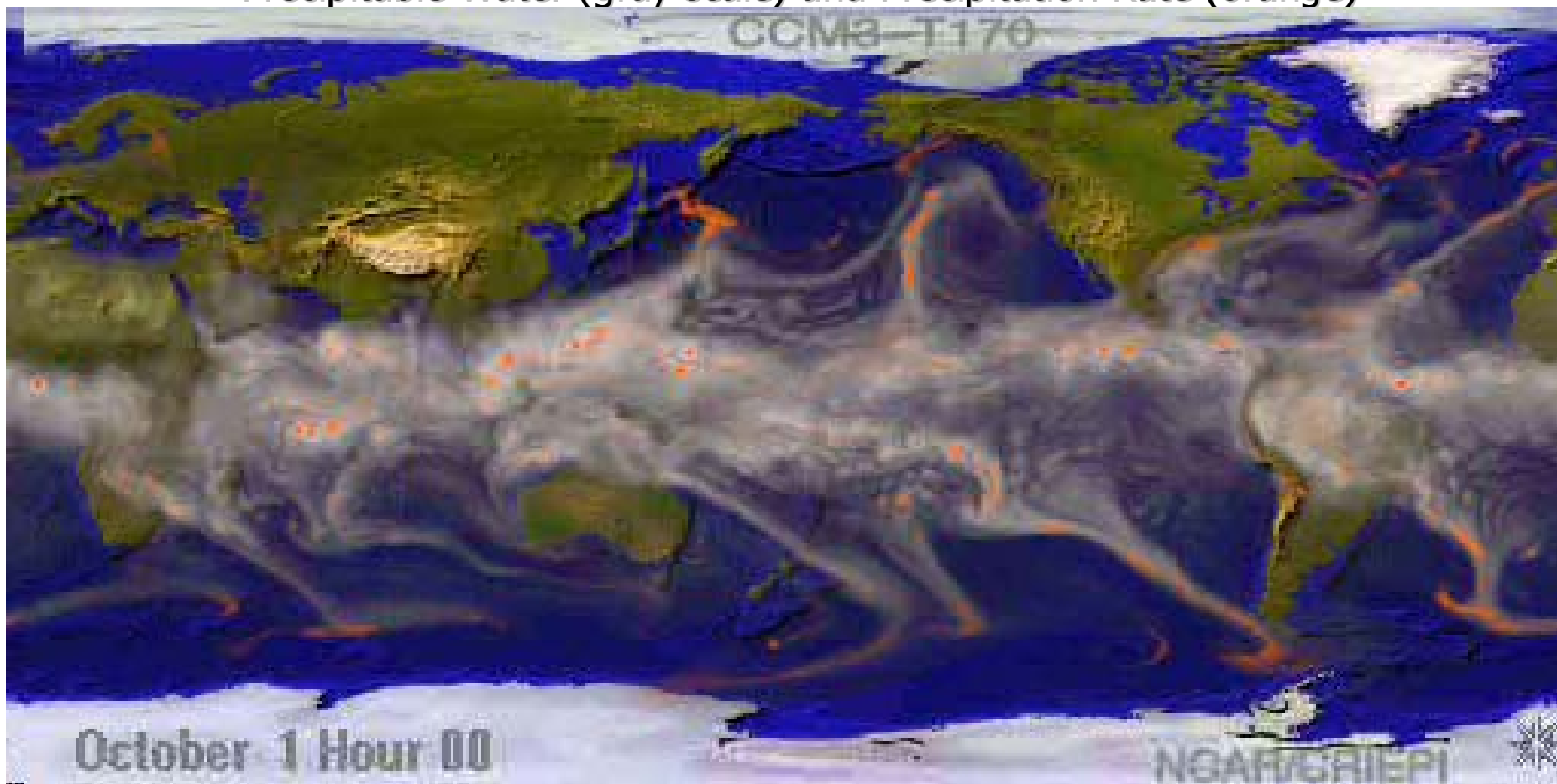
For a grid of atmospheric columns:

- 'Dynamics': Iterate Basic Equations
Horizontal momentum, Thermodynamic energy,
Mass conservation, Hydrostatic equilibrium,
Water vapor mass conservation
- Transport 'constituents' (water vapor, aerosol, etc)
- Calculate forcing terms ("Physics") for each column
Clouds & Precipitation, Radiation, etc
- Update dynamics fields with physics forcings
- Next time step (repeat)



Example of State of the Art Global Model Simulation

Precipitable Water (gray scale) and Precipitation Rate (orange)



Animation courtesy of NCAR SCD Visualization and Enabling Technologies Section

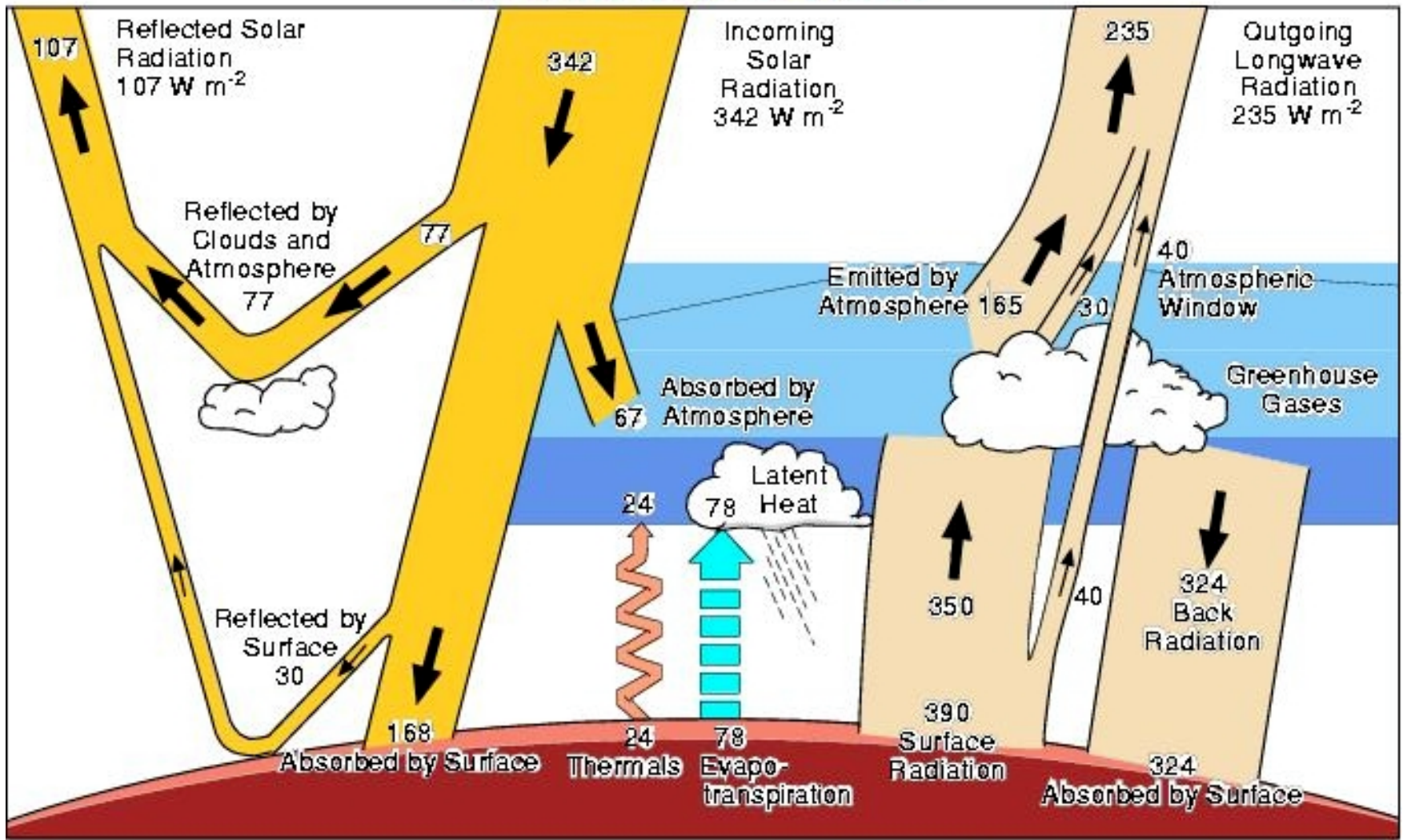
Physical Parameterization

To close the governing equations, it is necessary to incorporate the effects of physical processes that occur on scales below the numerical truncation limit

- Physical parameterization
 - express unresolved physical processes in terms of resolved processes
 - generally empirical techniques
- Examples of parameterized physics
 - dry and moist convection
 - cloud amount/cloud optical properties
 - radiative transfer
 - planetary boundary layer transports
 - surface energy exchanges
 - horizontal and vertical dissipation processes
 - ...



Radiation

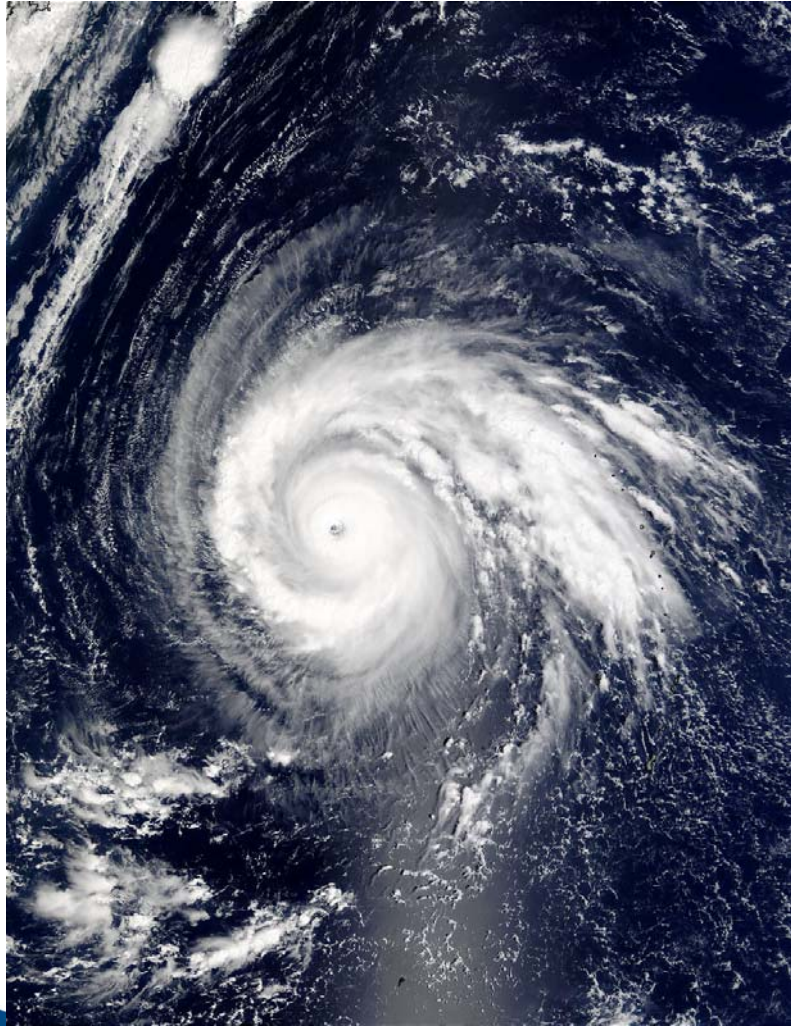


Kiehl and Trenberth 1997

Atmospheric Energy Transport

Synoptic-scale mechanisms

- hurricanes



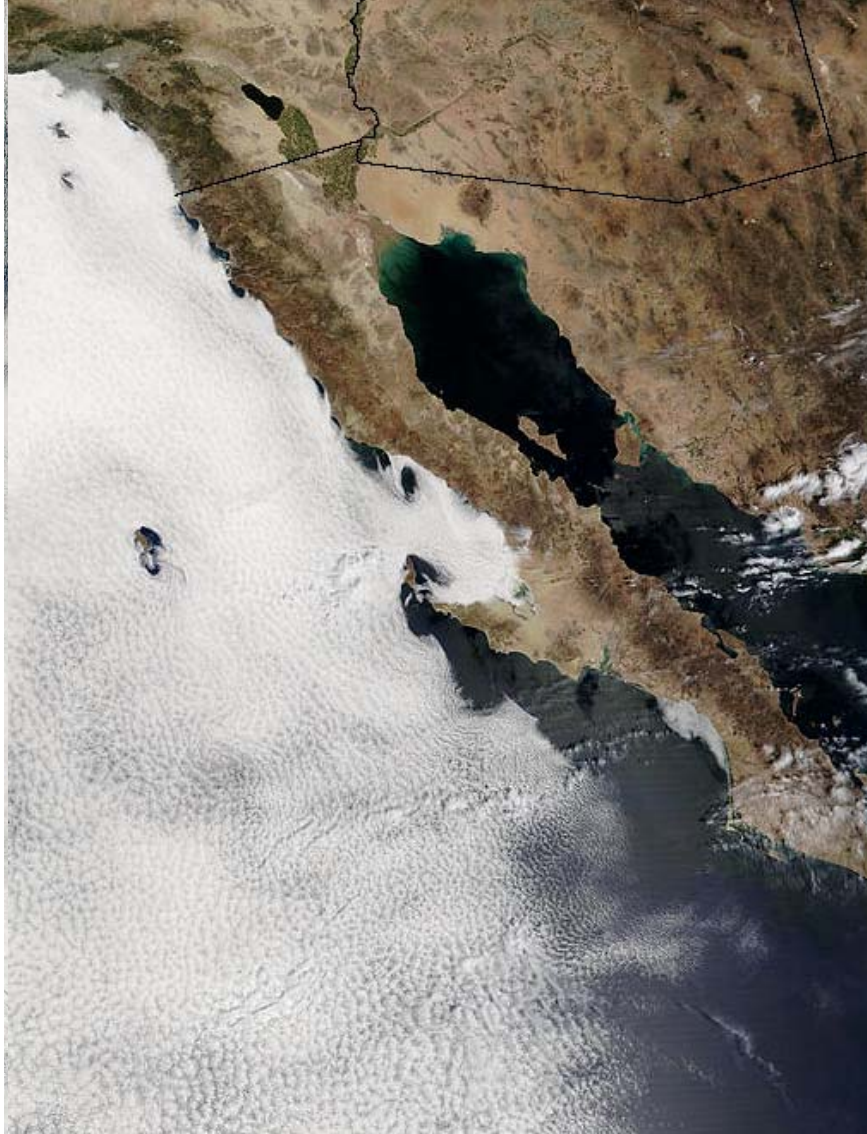
- extratropical storms



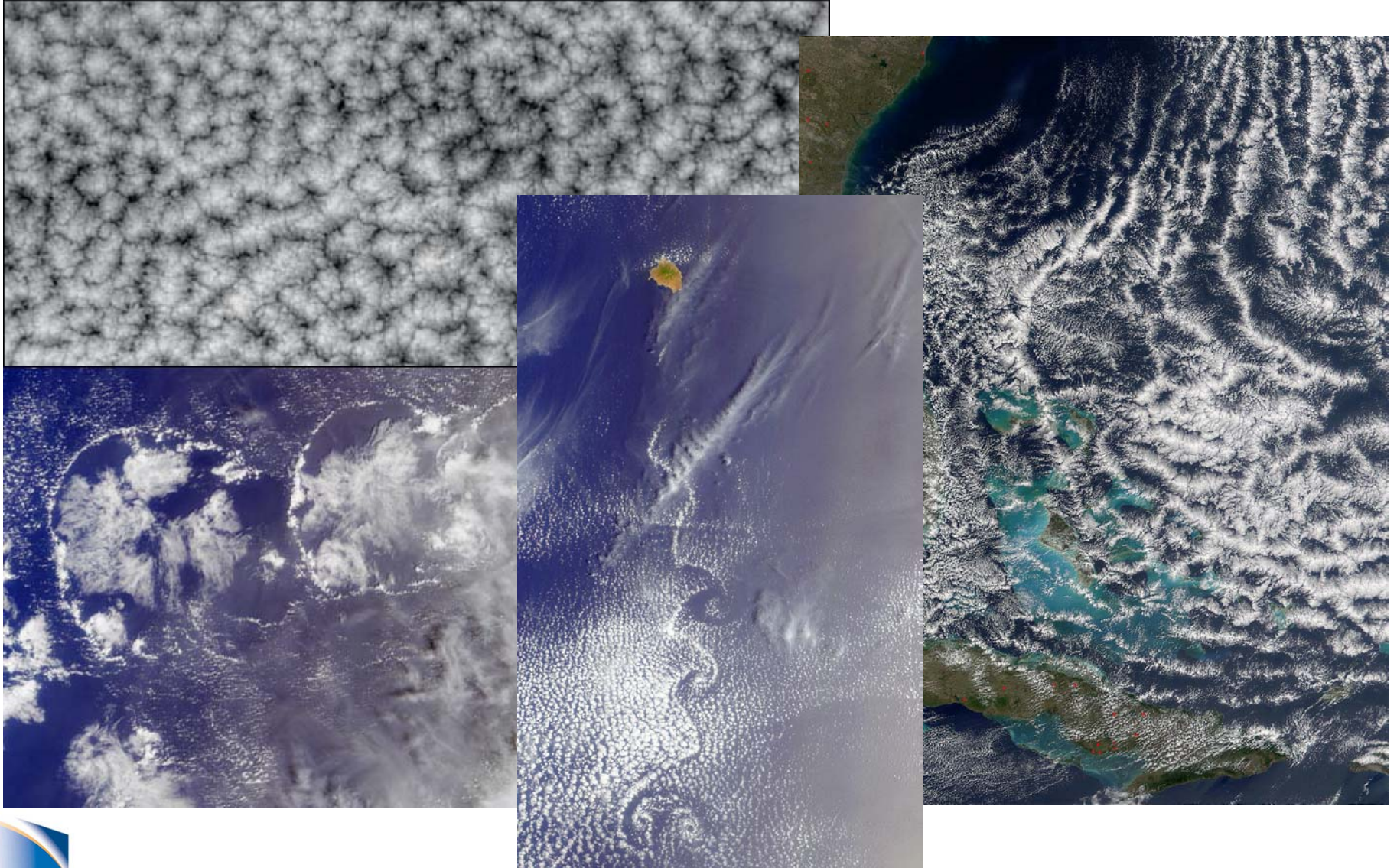
Clouds are a fundamental component of larger-scale organized energy transport mechanisms



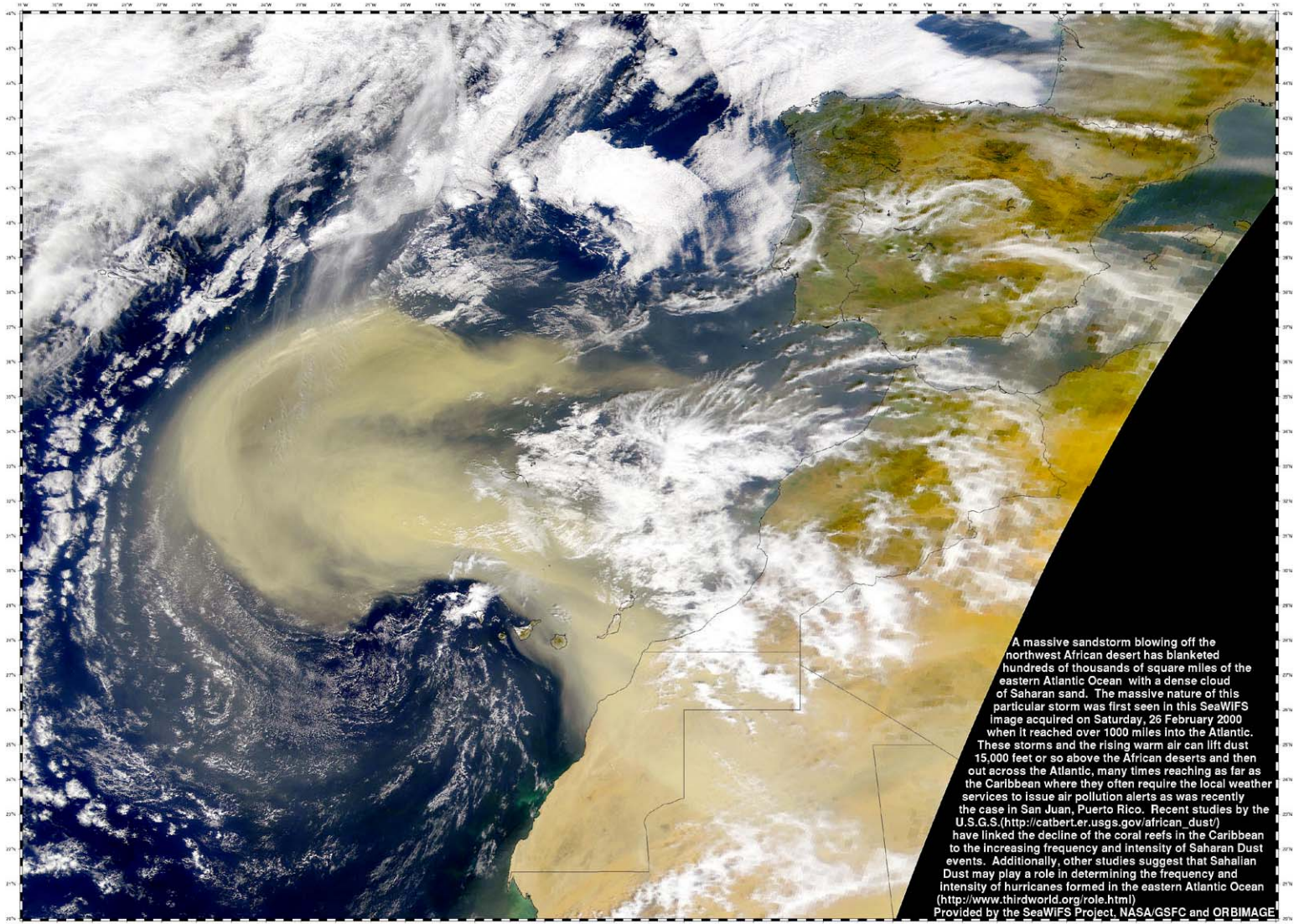
Other Energy Budget Impacts From Clouds



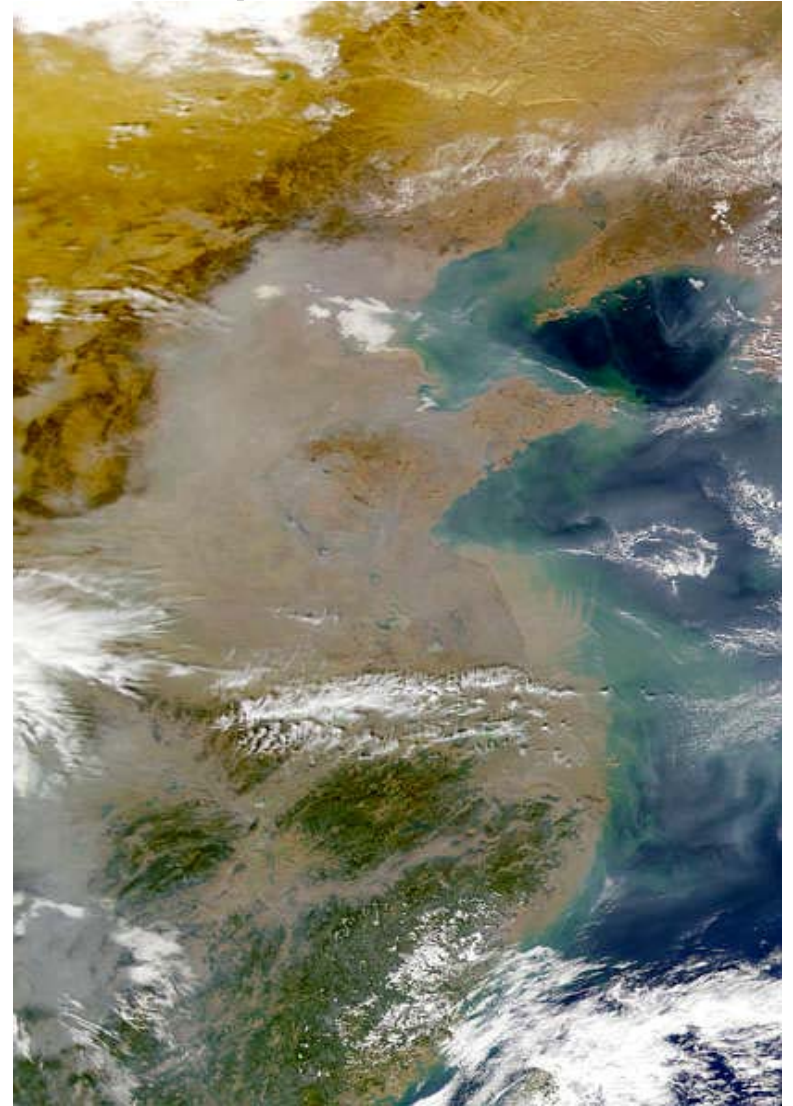
Other Energy Budget Impacts From Clouds



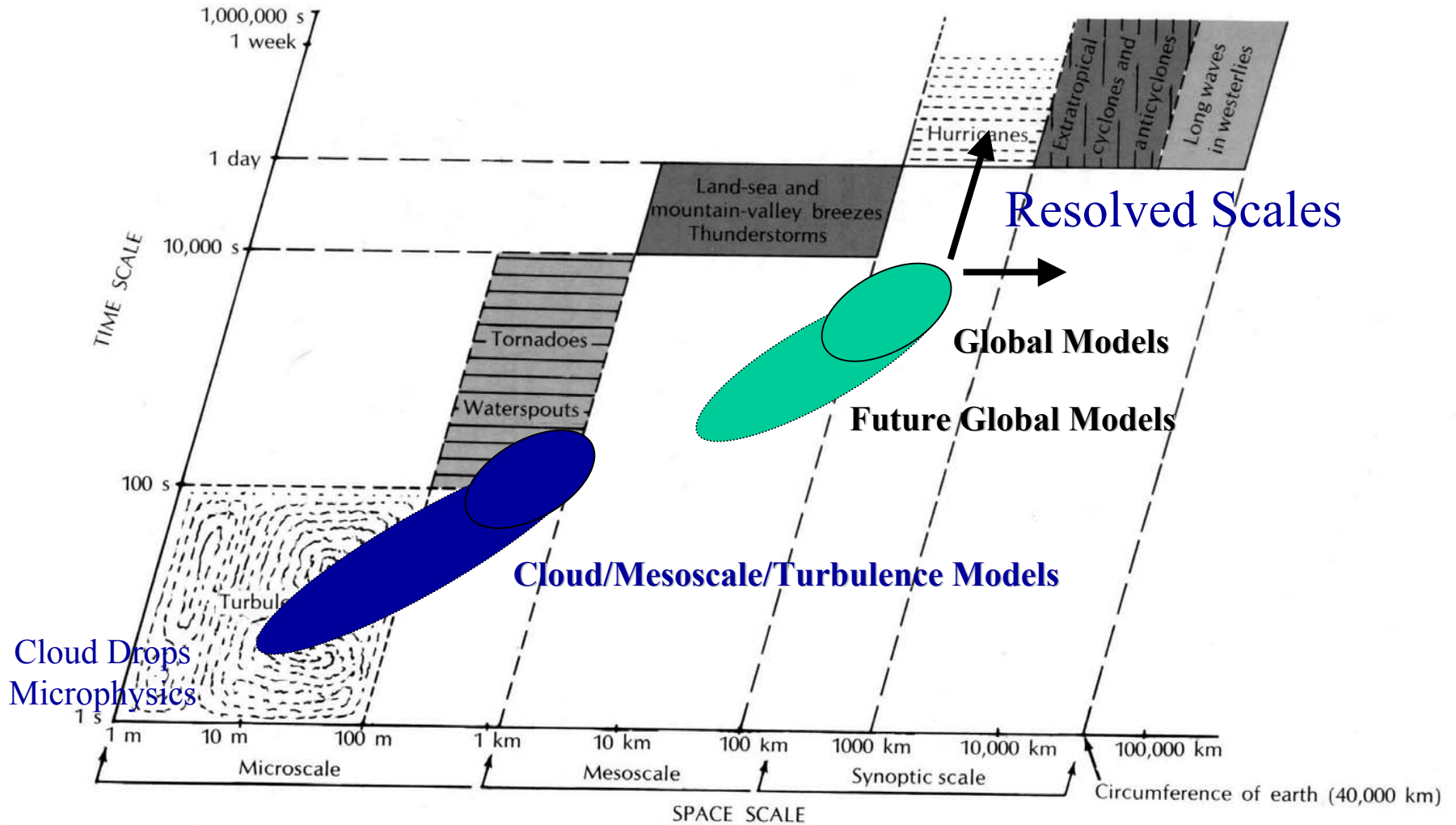
Energy Budget Impacts of Atmospheric Aerosol



Energy Budget Impacts of Atmospheric Aerosol

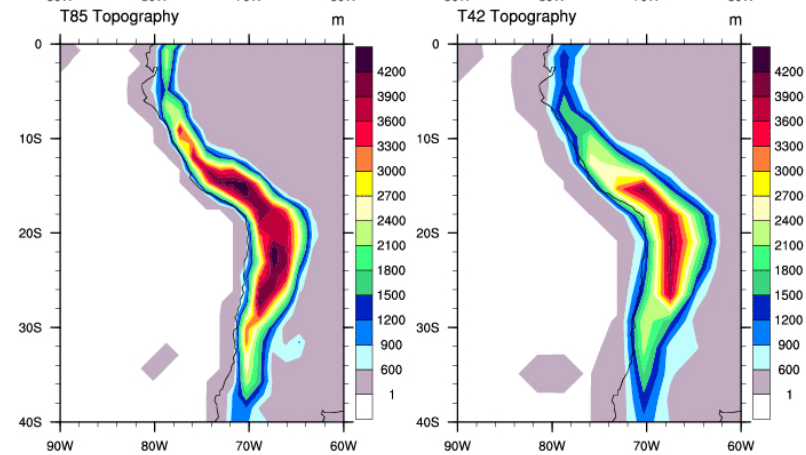
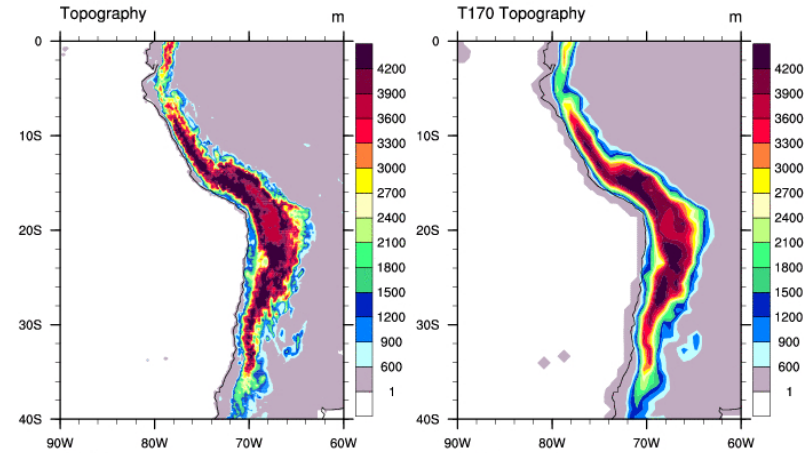
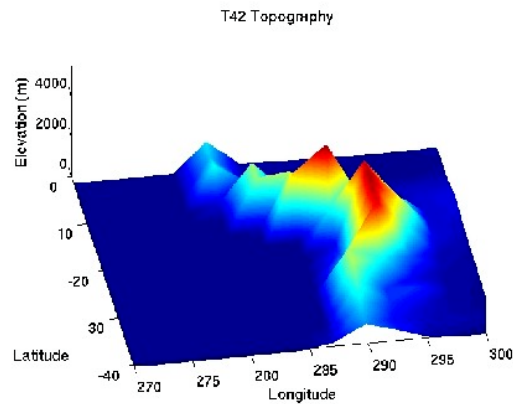
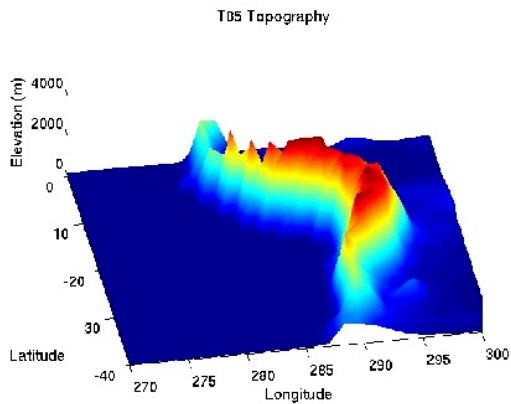
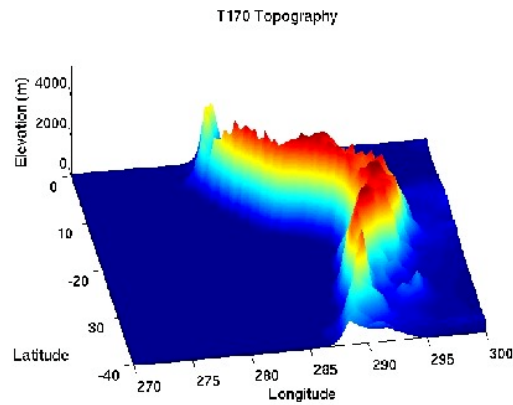
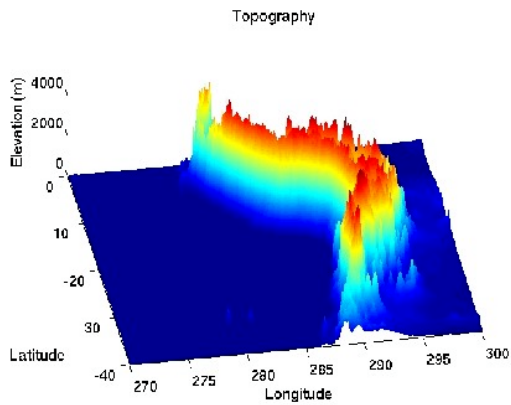


Scales of Atmospheric Motions



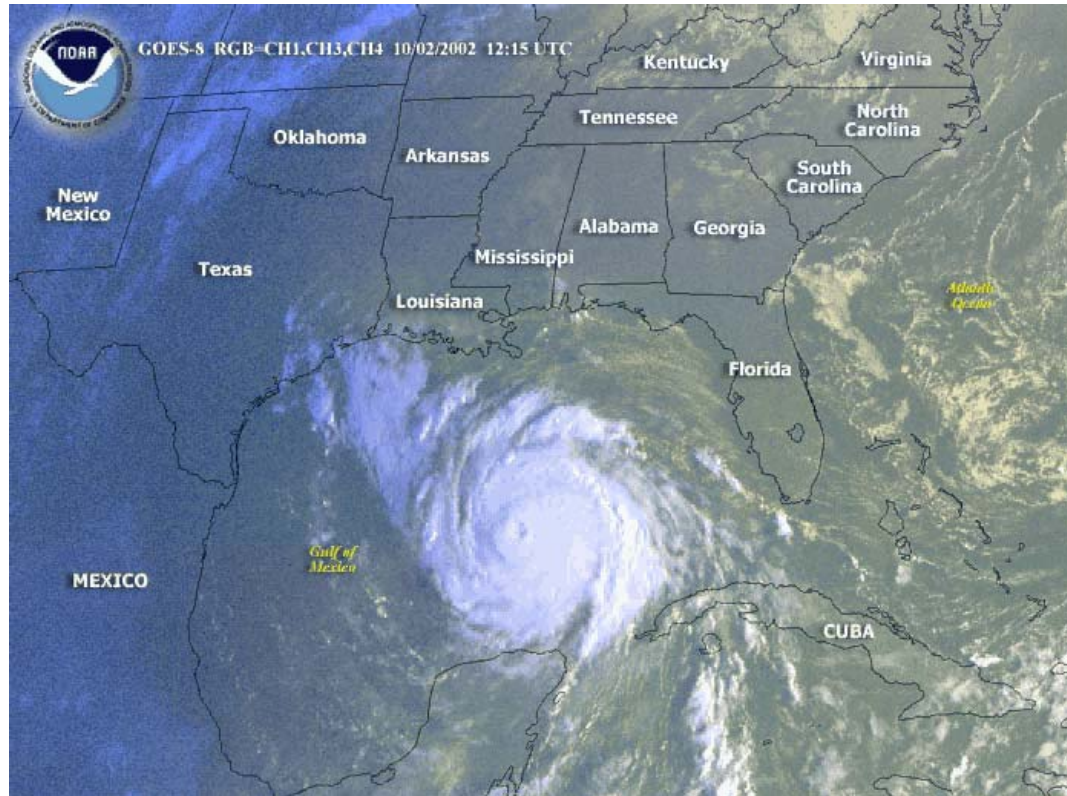
Anthes et al. (1975)

Global Modeling and Horizontal Resolution



Capturing Principle Phenomenological Scales of Motion in Global Models

Simulation of Tropical Cyclone Impacts on Climate

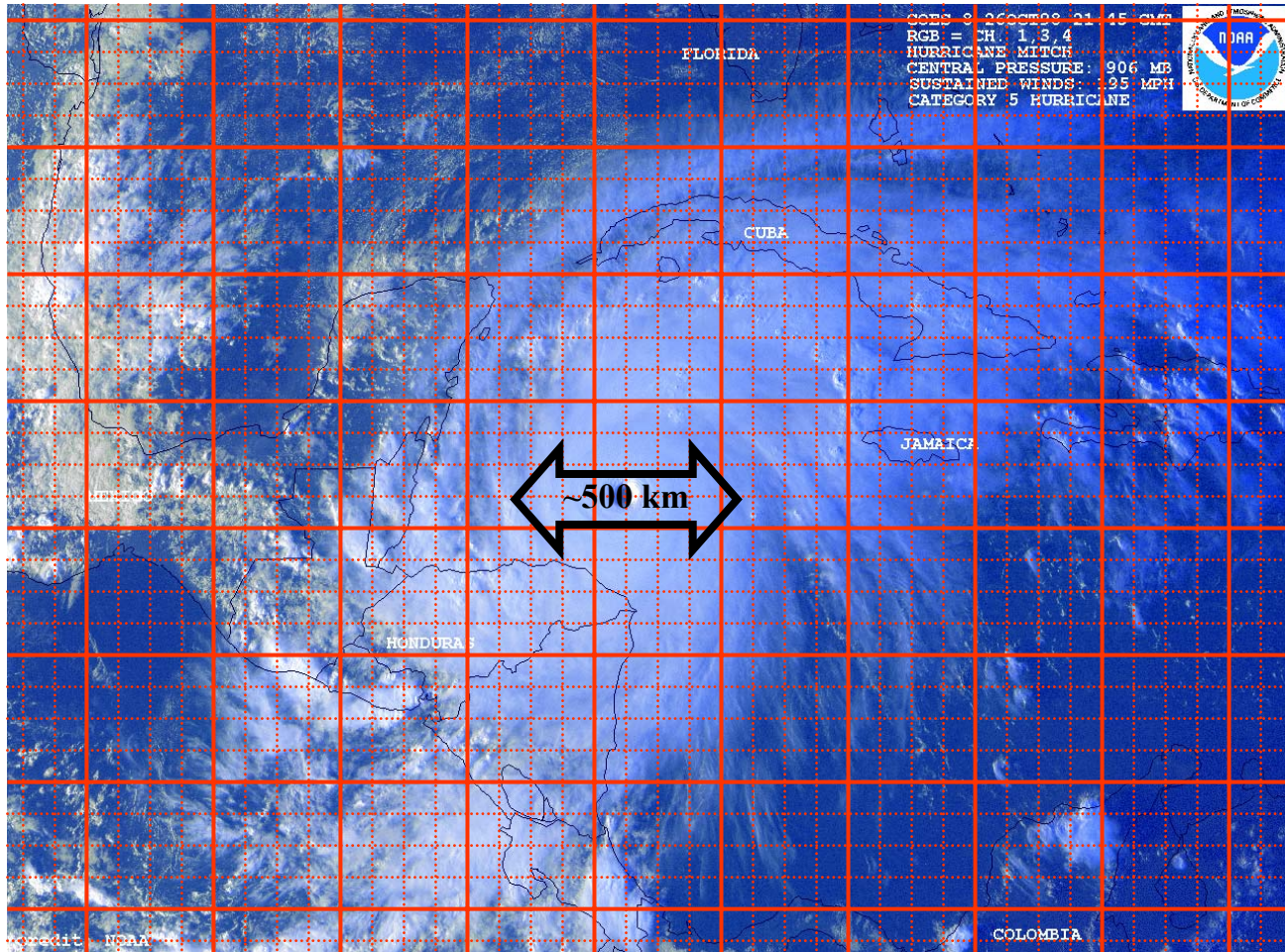


Courtesy, Raymond Zehr, NOAA CIRA

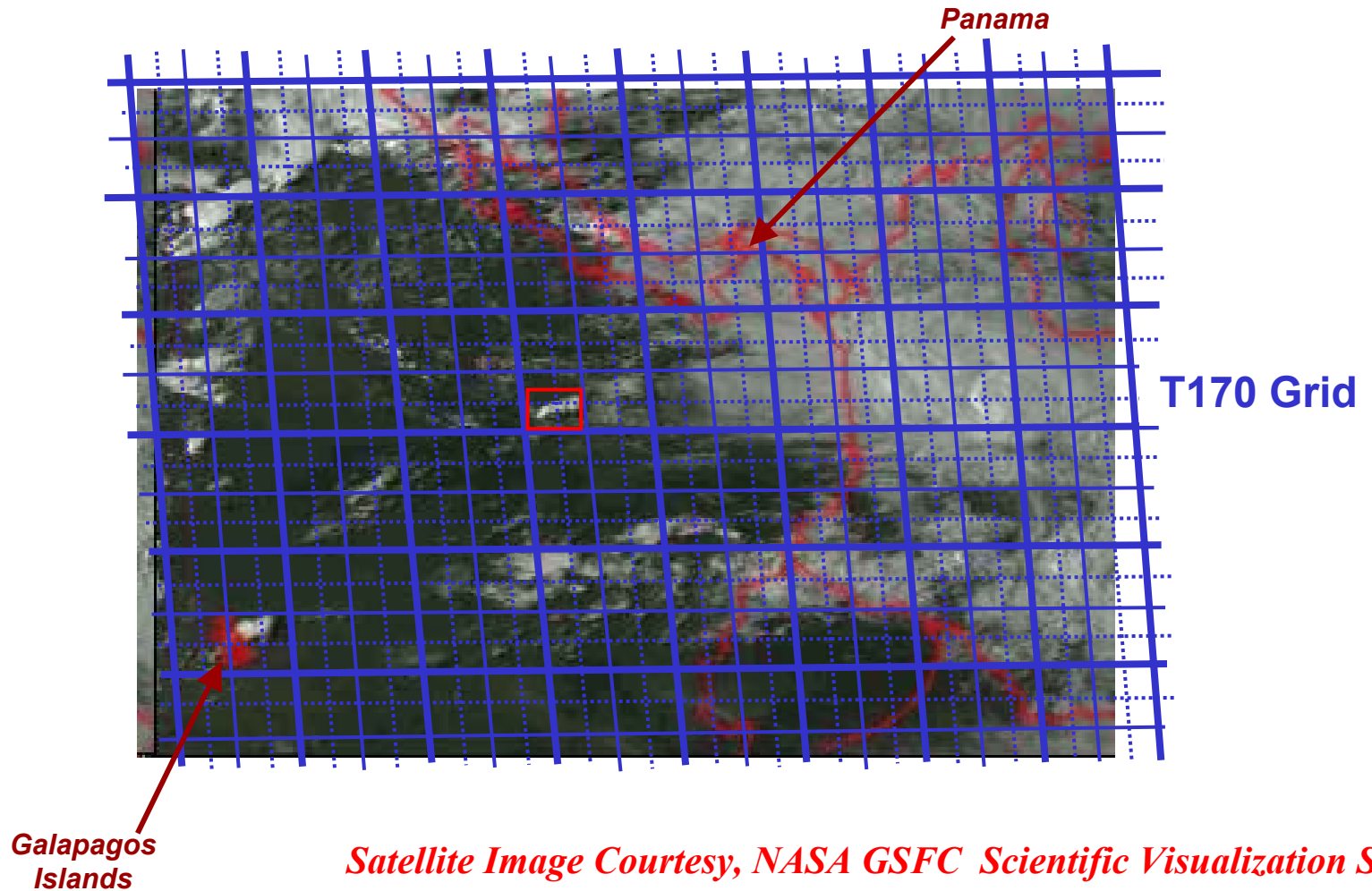
J. J. Hack/A. Gettelman: June 2005

High-Resolution Global Modeling

Simulation of Tropical Cyclone Impacts on Climate



High-Resolution Global Modeling



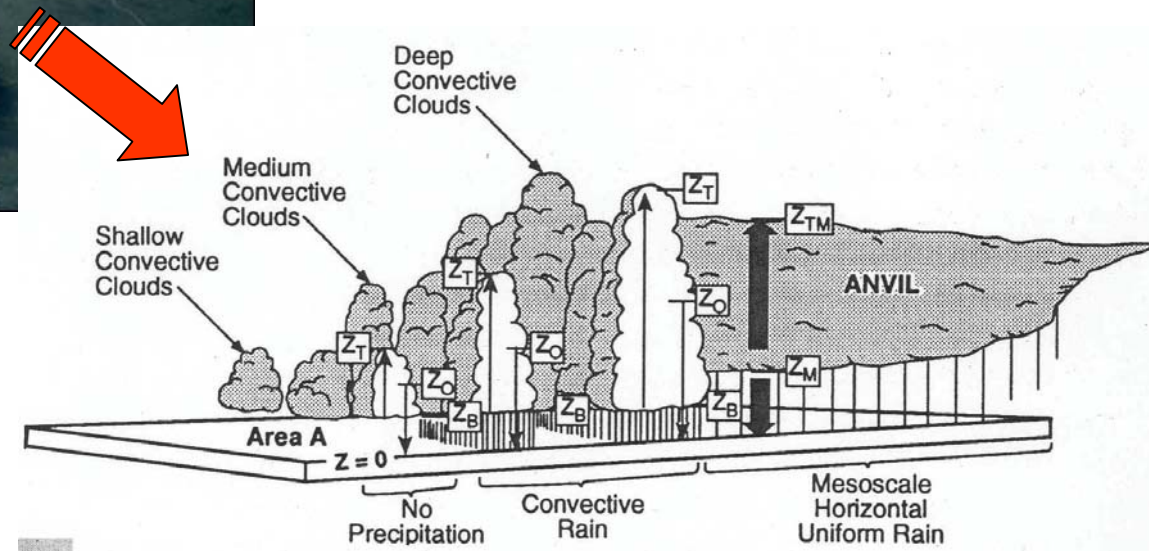
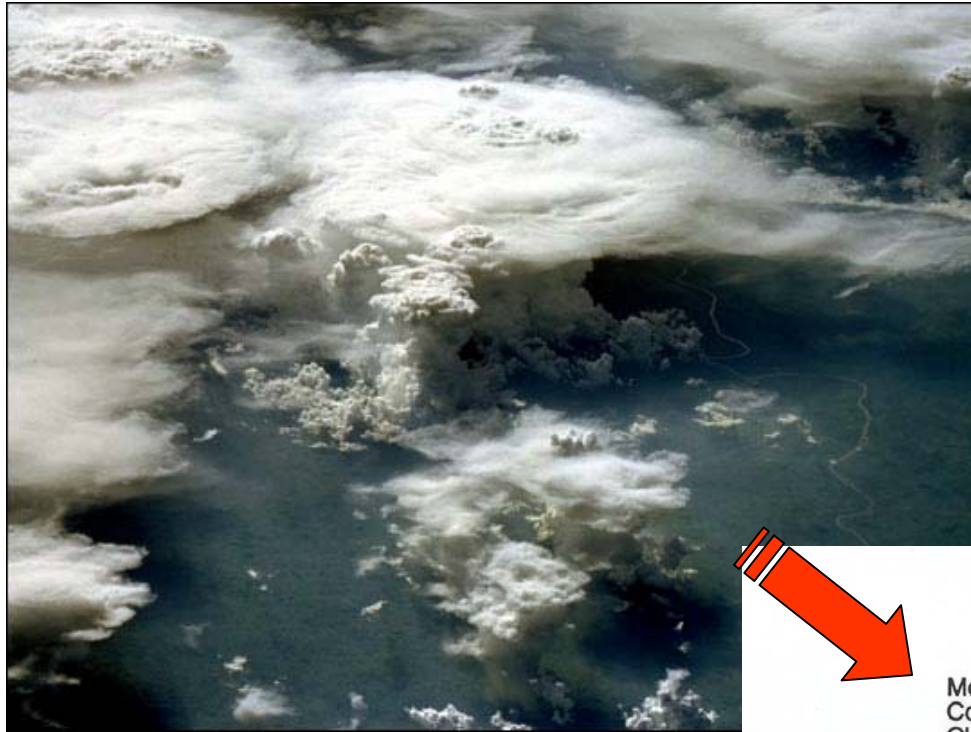
Cumulus Convection

If the atmosphere is buoyantly unstable to small vertical displacements, it can be said to be convectively unstable

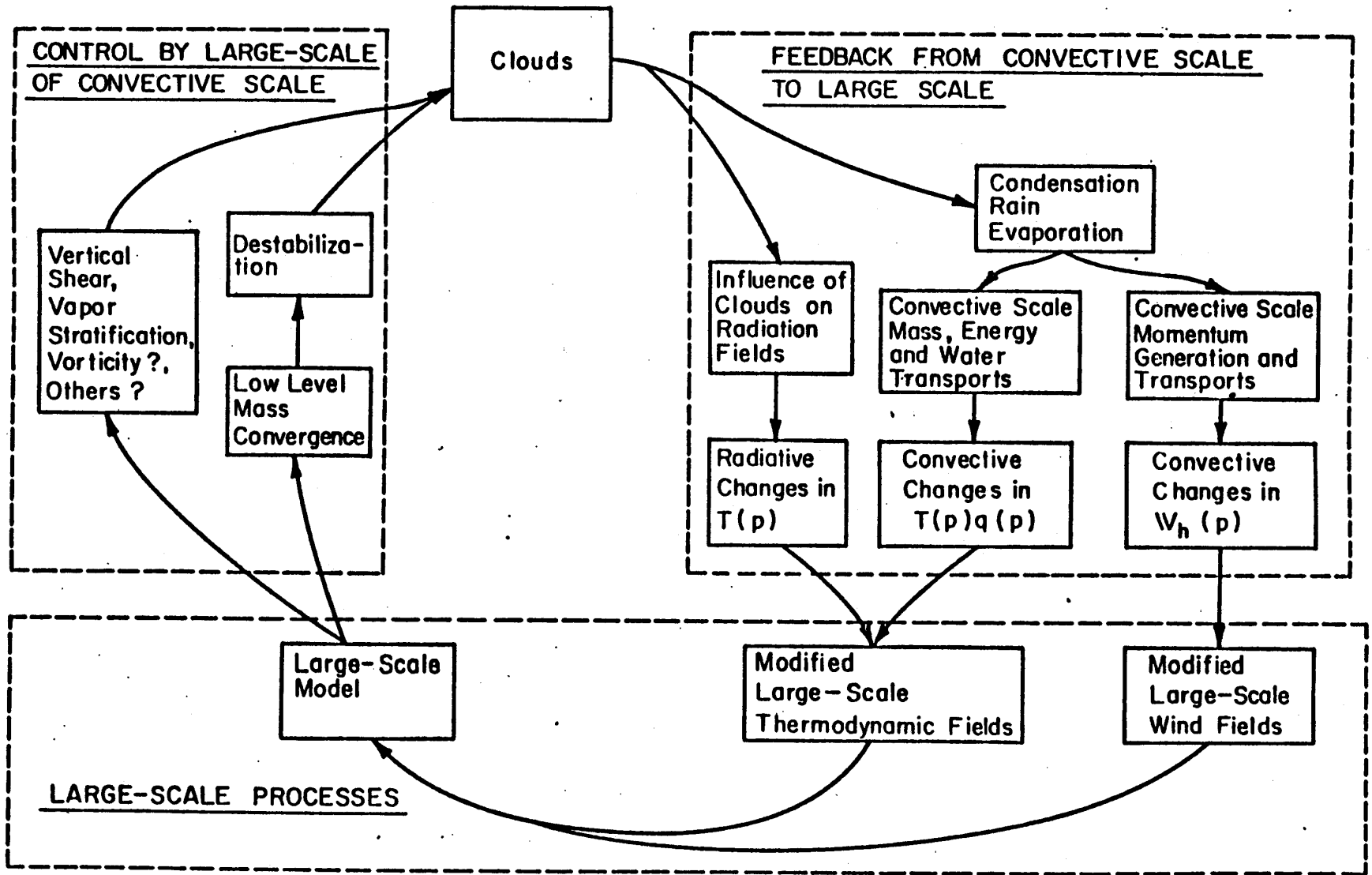
- Convective overturning
 - with or without phase change
 - space scale \sim 1-10km; time scale \sim 1 hour
- Moist convection
 - most common and energetically important
 - affects the general circulation on wide range of time scales
 - provides fundamental coupling of dynamics and hydrological cycle



Process Models and Parameterization



SIMPLIFIED LARGE-SCALE: CONVECTIVE INTERACTION



Parameterization of Cumulus Convection

To extract the details of how the observed profile is maintained by moist convection, it is necessary to use an abstraction for the collective behavior of convective motions

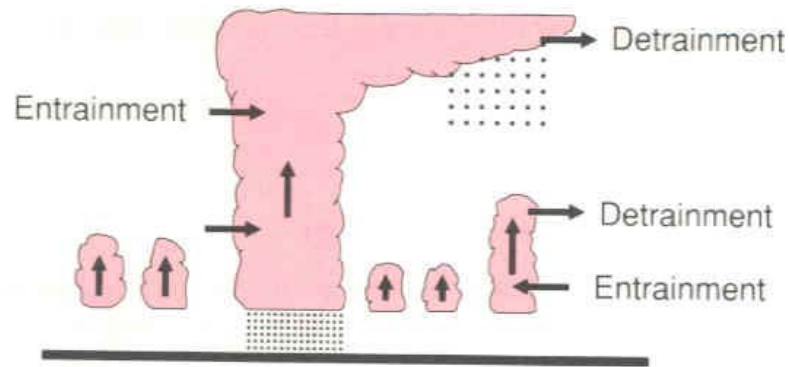
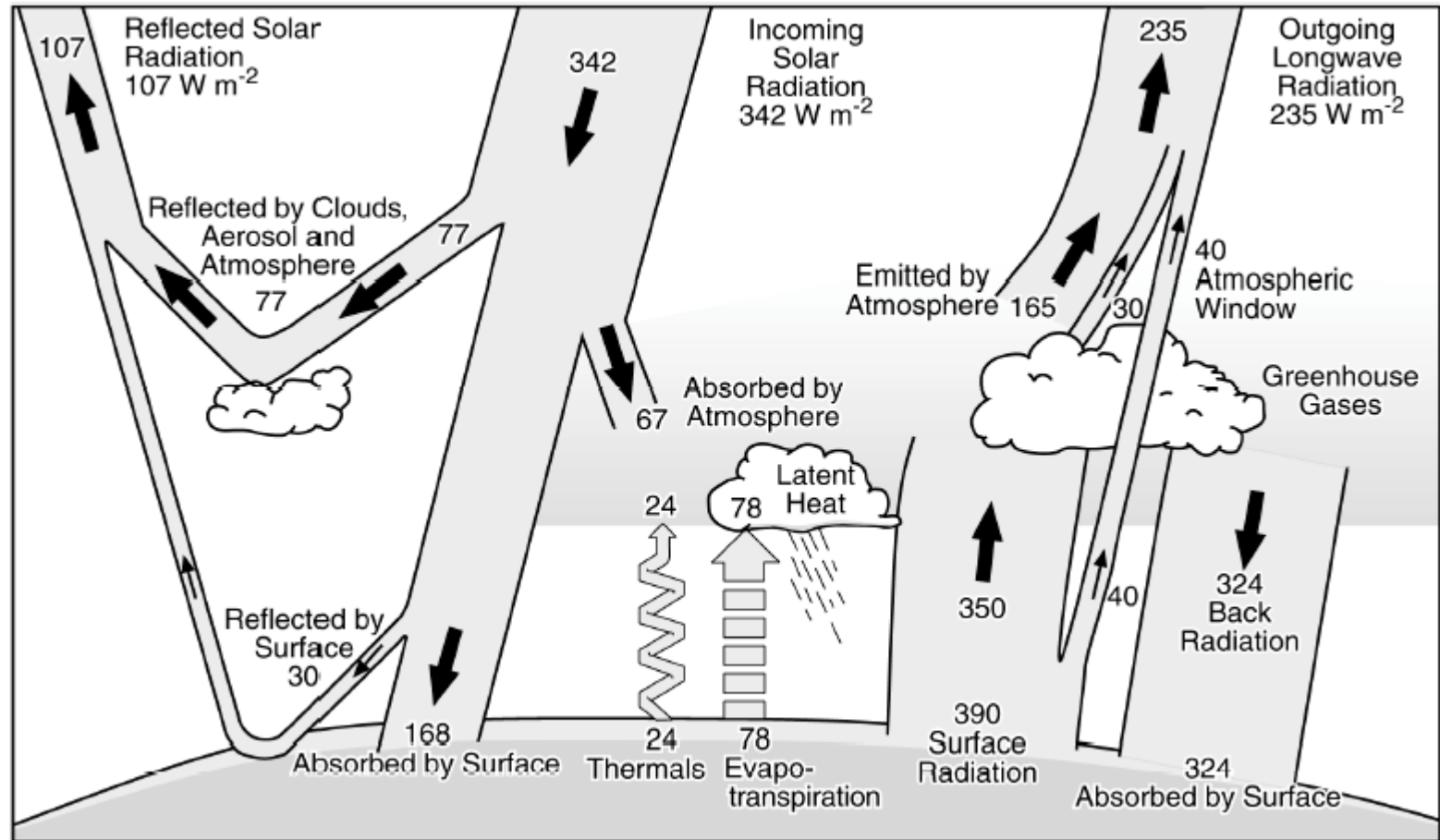


Fig. 10.16 Schematic of an ensemble of cumulus clouds. From Yanai et al. (1973).

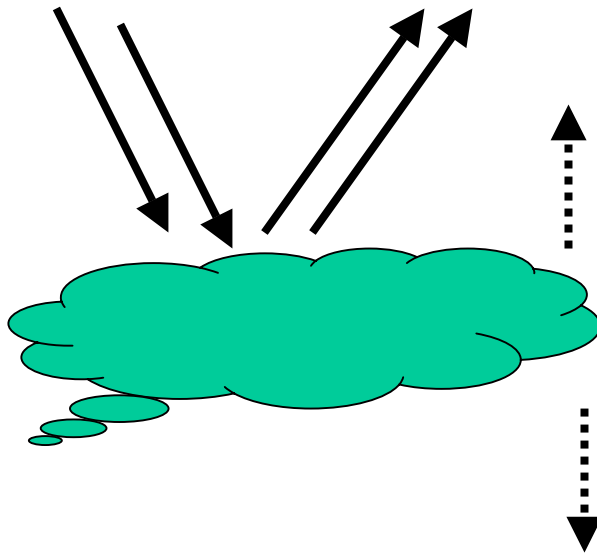
- Convective mass flux
 - how much overturning is associated with convective activity
- Breakdown of total diabatic forcing
 - where is the water condensing and/or raining out
 - what role do the convective eddy transports play

What are the key uncertainties?

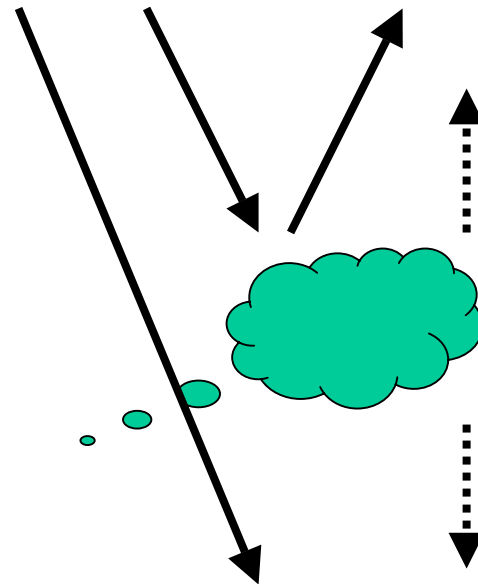


Uncertainties (1):

- Low Clouds over the ocean:
Reflect Sunlight (cool) : Dominant Effect
Trap heat (warm)

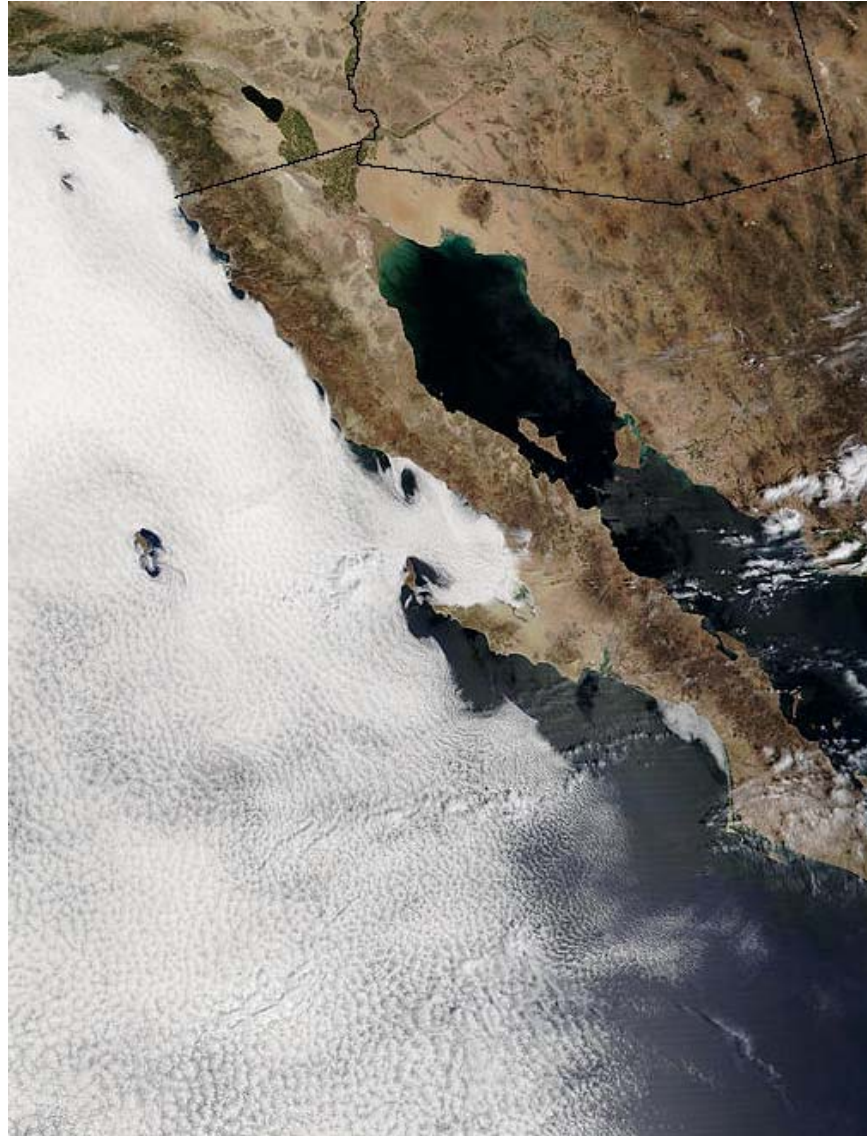


More Clouds=Cooling



Fewer Clouds=Warming

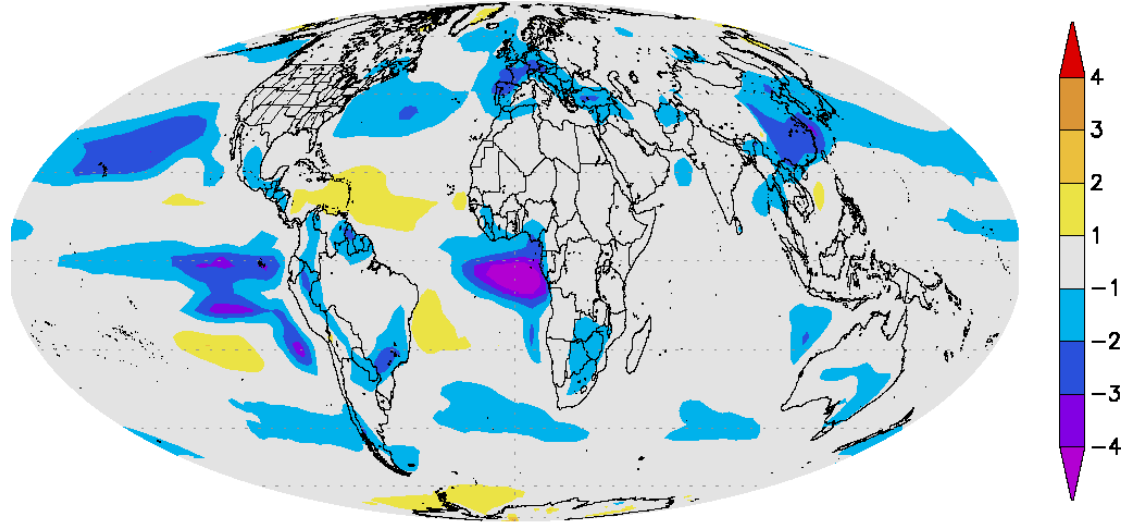
Marine Stratus: Low Clouds over the Ocean



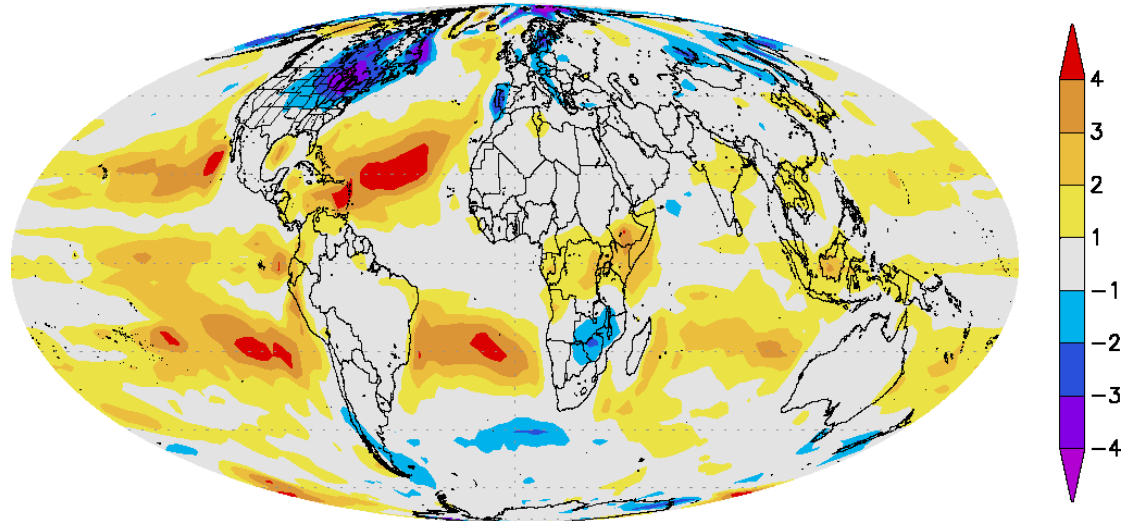
Low Clouds Over the Ocean

2 Models: Changes
are OPPOSITE!

GFDL AM2-ML ($2\times\text{CO}_2$ - CTRL)

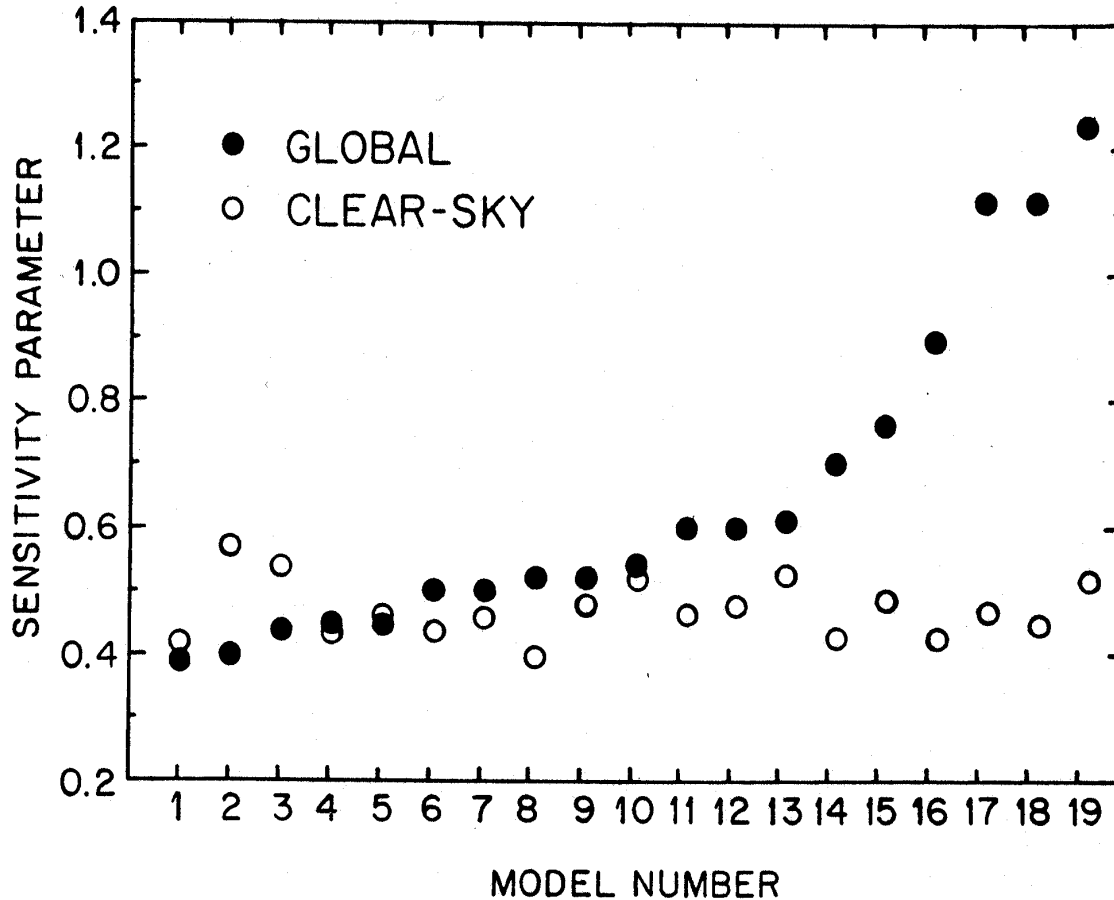


NCAR CAM2 (Year70 @1%CO₂/yr - CTRL)



Change in Low Cloud Amount (%/K)

Parameterization of Clouds



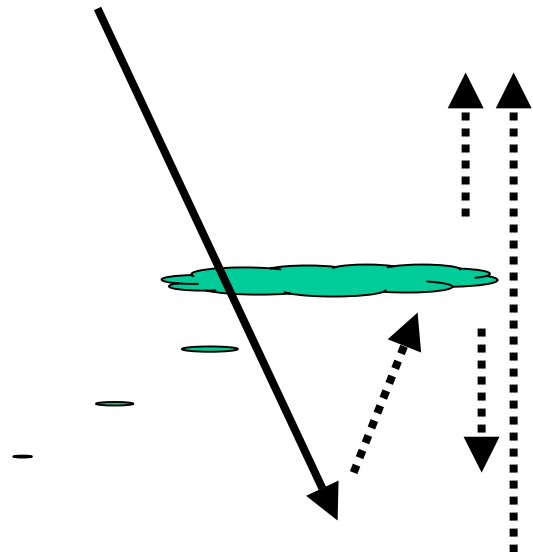
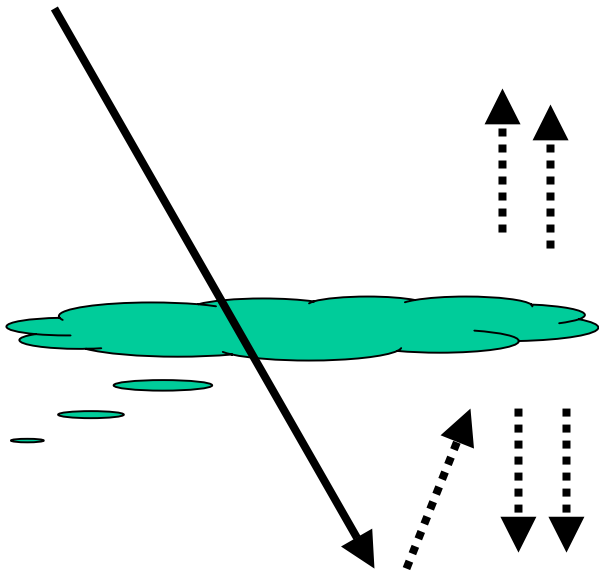
Cess et al. (1990)

Fig. 1. Clear-sky and global sensitivity parameters ($\text{K m}^2 \text{W}^{-1}$) for the 19 GCMs. The model numbers correspond to the ordering in Table 9.

Uncertainties (2):

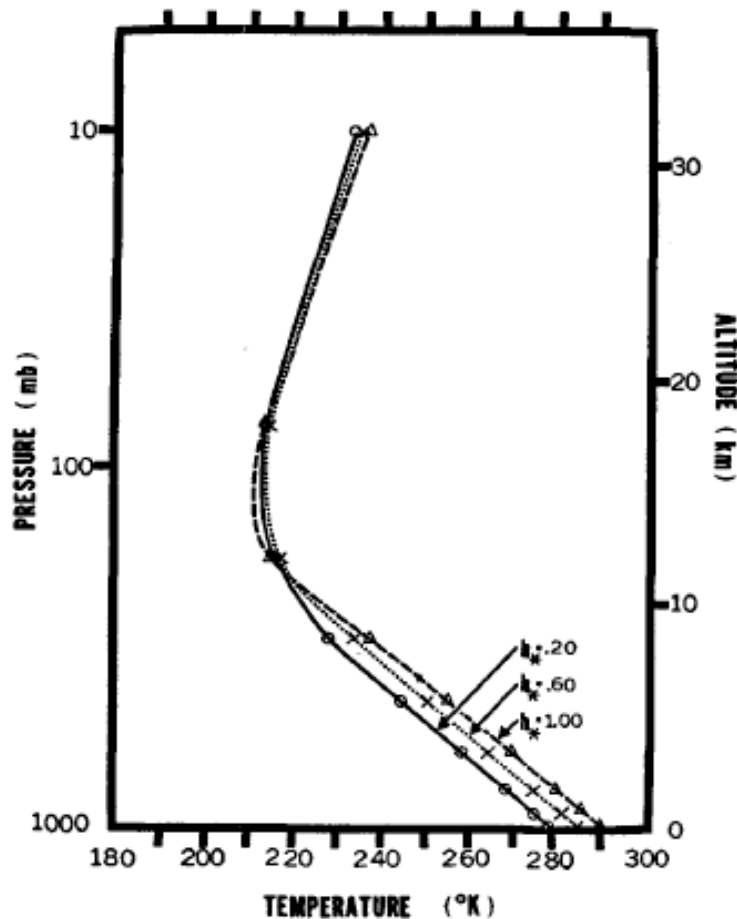
2. High Clouds:

Dominant effect is that they Trap heat (warm)



Uncertainties (3):

- Water Vapor: largest greenhouse gas
Increasing Temp=Increasing water Vapor (more greenhouse)
Effect is expected to ‘amplify’ warming through a ‘feedback’



1D Radiative-Convective Model:
Higher humidity=>warmer surface

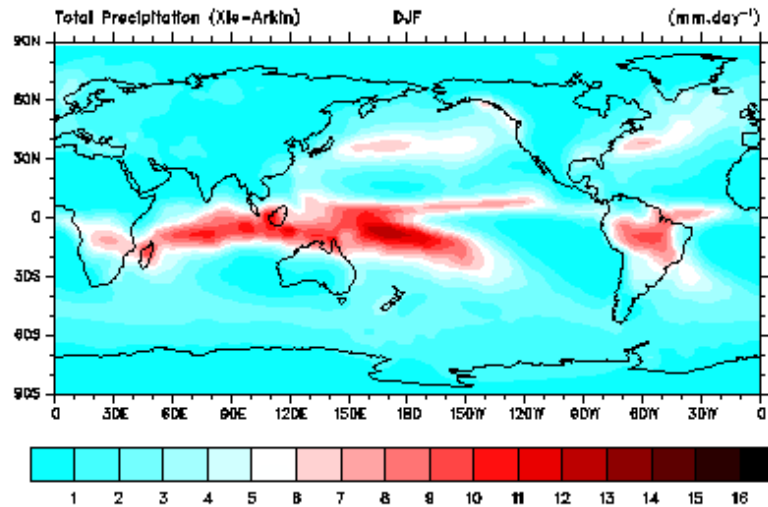
How can we evaluate simulation quality?

- Continue to compare long term mean climatology
 - average mass, energy, and momentum balances
 - tells you where the physical approximations take you
 - ***but you don't necessarily know how you get there!***
- Must also consider dominant modes of variability
 - provides the opportunity to evaluate *climate sensitivity*
 - response of the climate system to a specific forcing factor
 - evaluate modeled response on a hierarchy of time scales
 - exploit natural forcing factors to test model response
 - diurnal and seasonal cycles
 - El Niño Southern Oscillation (ENSO)
 - intraseasonal variability; e.g., MJO
 - solar variability
 - volcanic aerosol loading

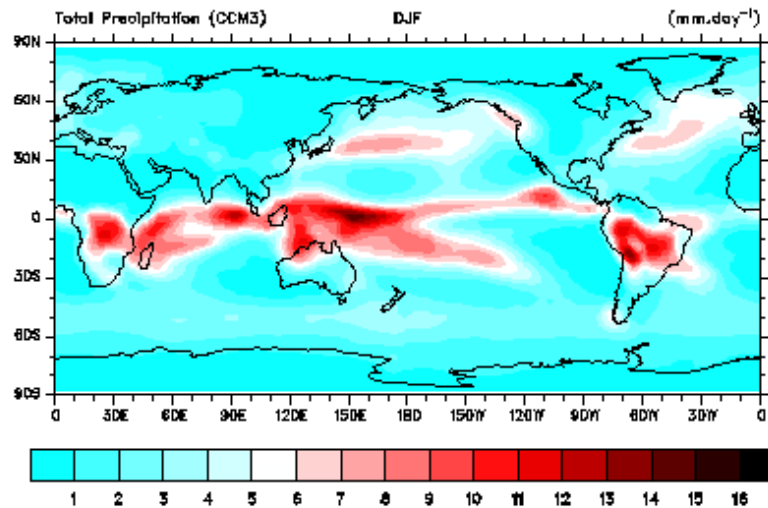


Comparison of Mean Simulation Properties

Observed
Precipitation



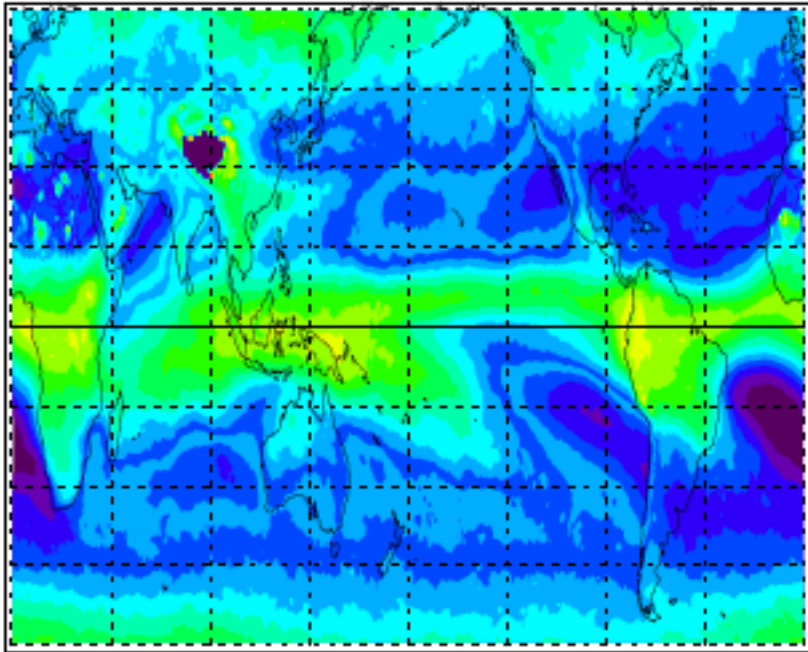
Simulated
Precipitation



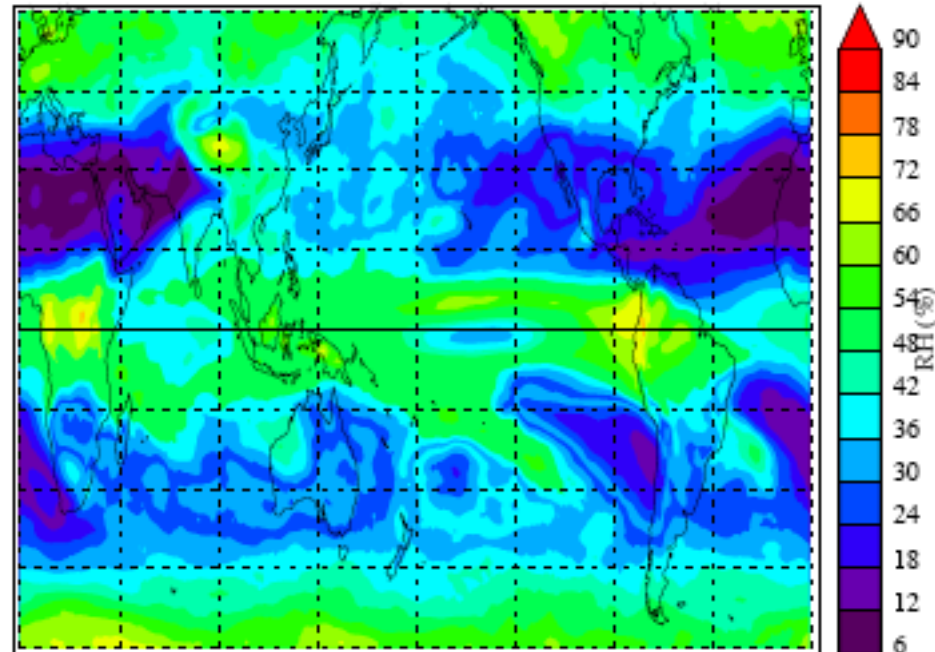
Mean Biases

Relative humidity, March-May 3km (9,000ft)

Observed



Simulated



Variability: El Niño Composite

Observed

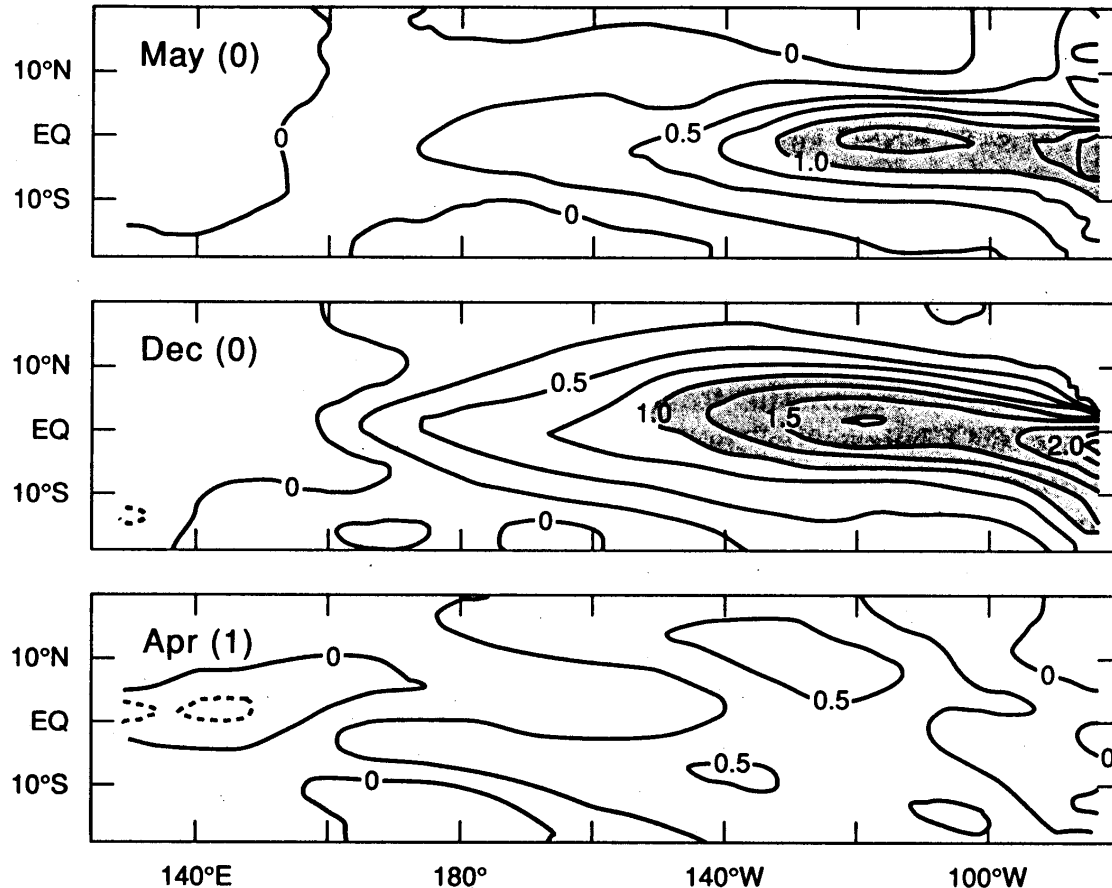


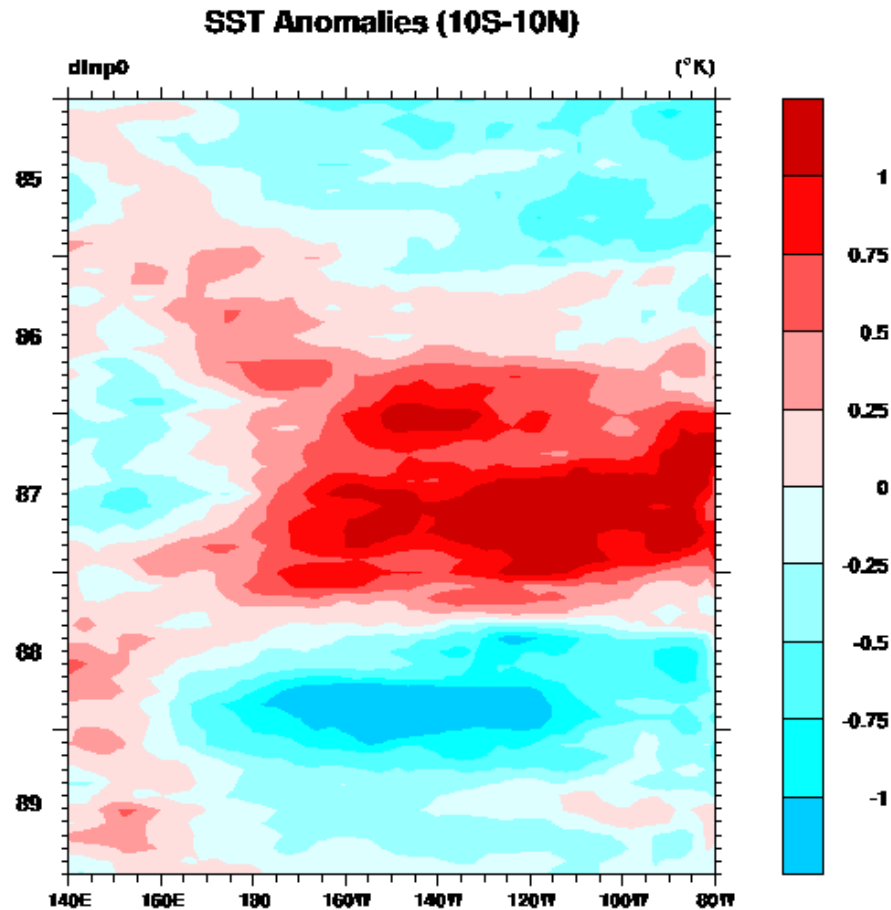
Fig. 18.2 Sea surface temperature anomalies ($^{\circ}\text{C}$) for a composite El Niño (Rasmusson and Carpenter, 1982), constructed by averaging over 6 events (1951, 1953, 1957, 1965, 1969, 1972; cf. Fig. 18.1). Shown are maps for May and December of the El Niño year and April of the following year.



Testing AGCM Sensitivity

Simulated

Pacific SST Anomalies and ENSO

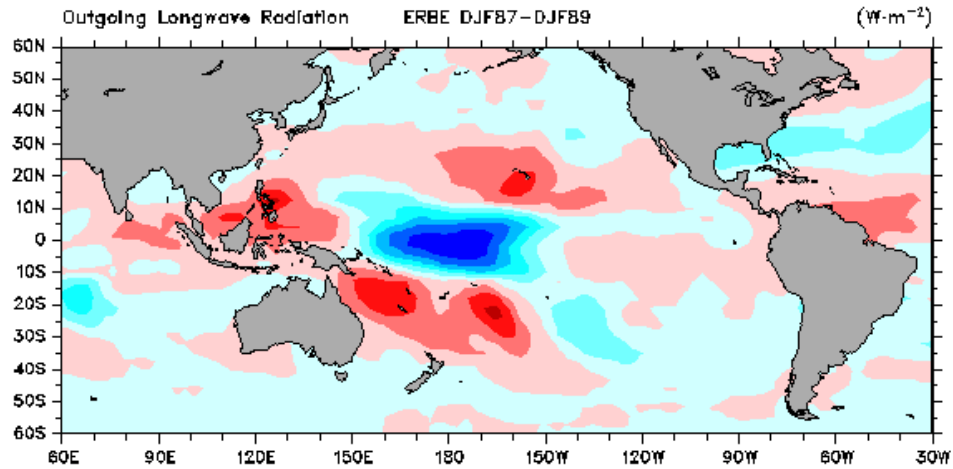


Hack (1998)

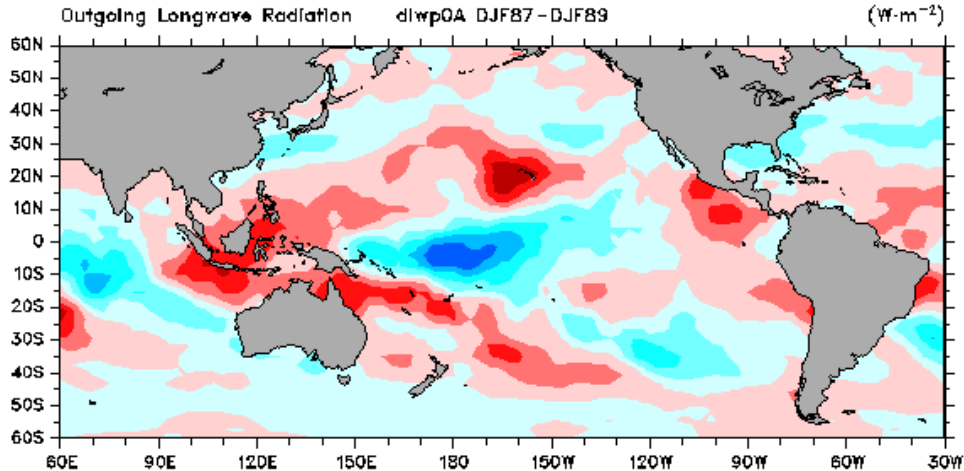
Testing AGCM Sensitivity

Cloud (OLR) Anomalies and ENSO

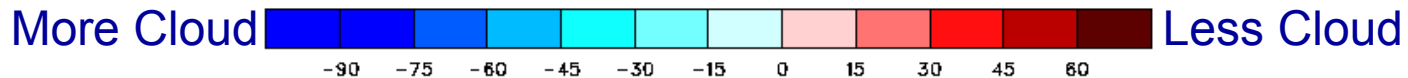
Observed



Simulated

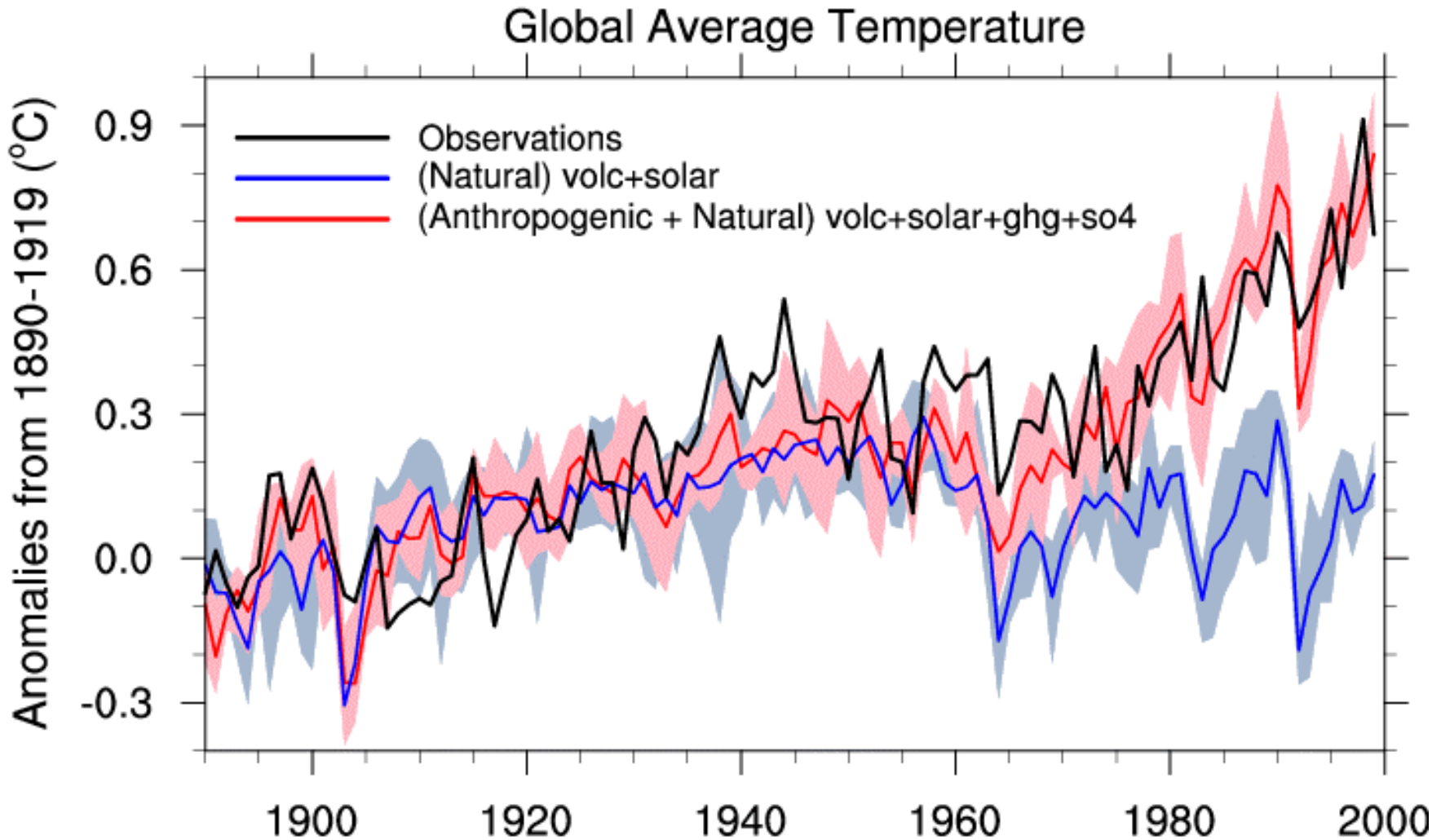


Hack (1998)



Observations: 20th Century Warming

Model Solutions with Human Forcing



Improving simulation quality

- Examine role of parameterization techniques on transient behavior
 - oversimplifications playing a role in inadequate variability?
- Understand role of scale interaction on transient and mean state

CCM2 ITCZ behavior as function of horizontal resolution

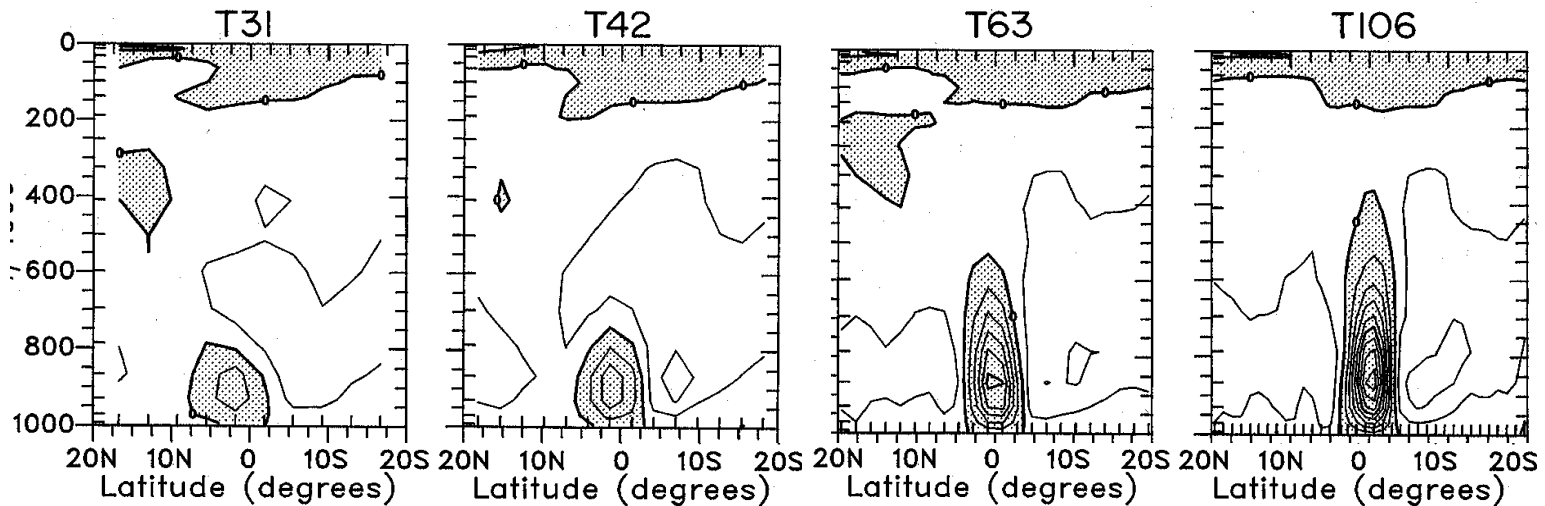
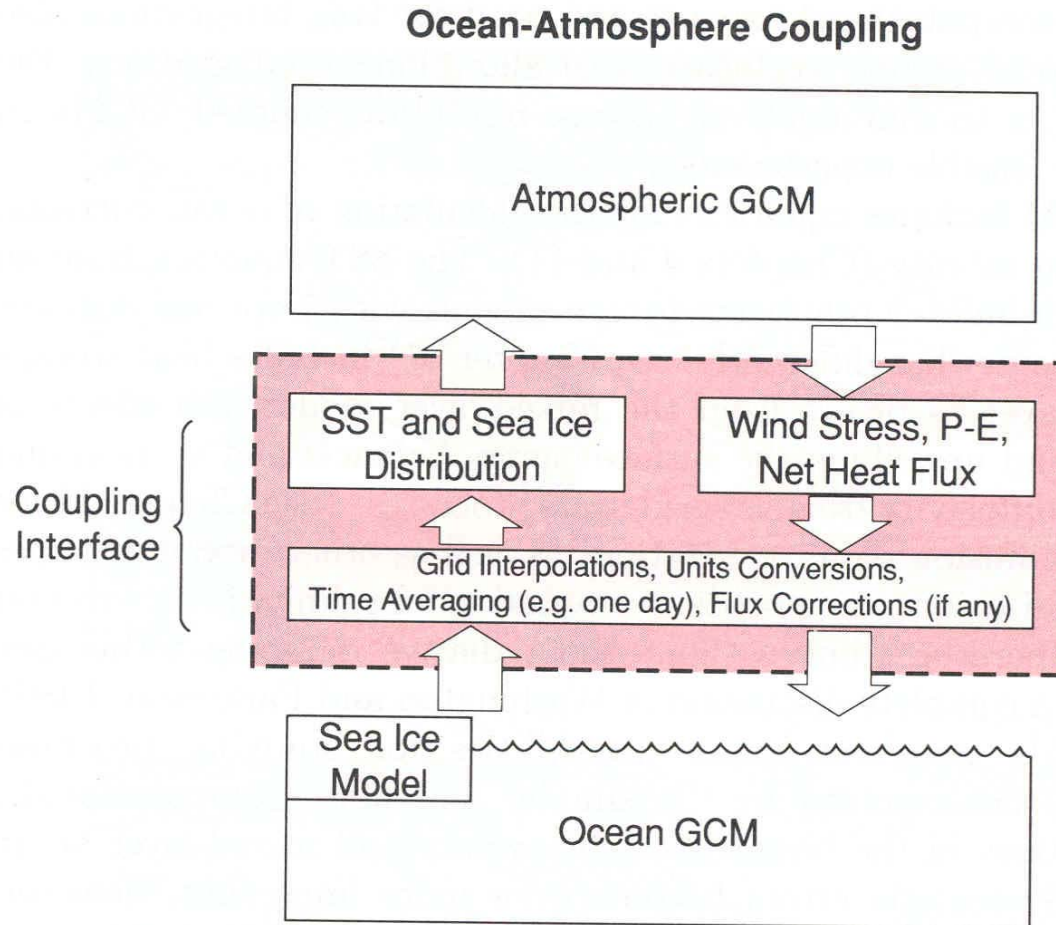


Fig. 6. January average, zonal average over Atlantic (30°W to 7.5°E) pressure vertical velocity (ω) for R15, T21, T31, T42, T63, and T106 simulations. Contour interval is 20 mb day^{-1} , negative (upward) regions *stippled*

Williamson, Kiehl, and Hack (1995)

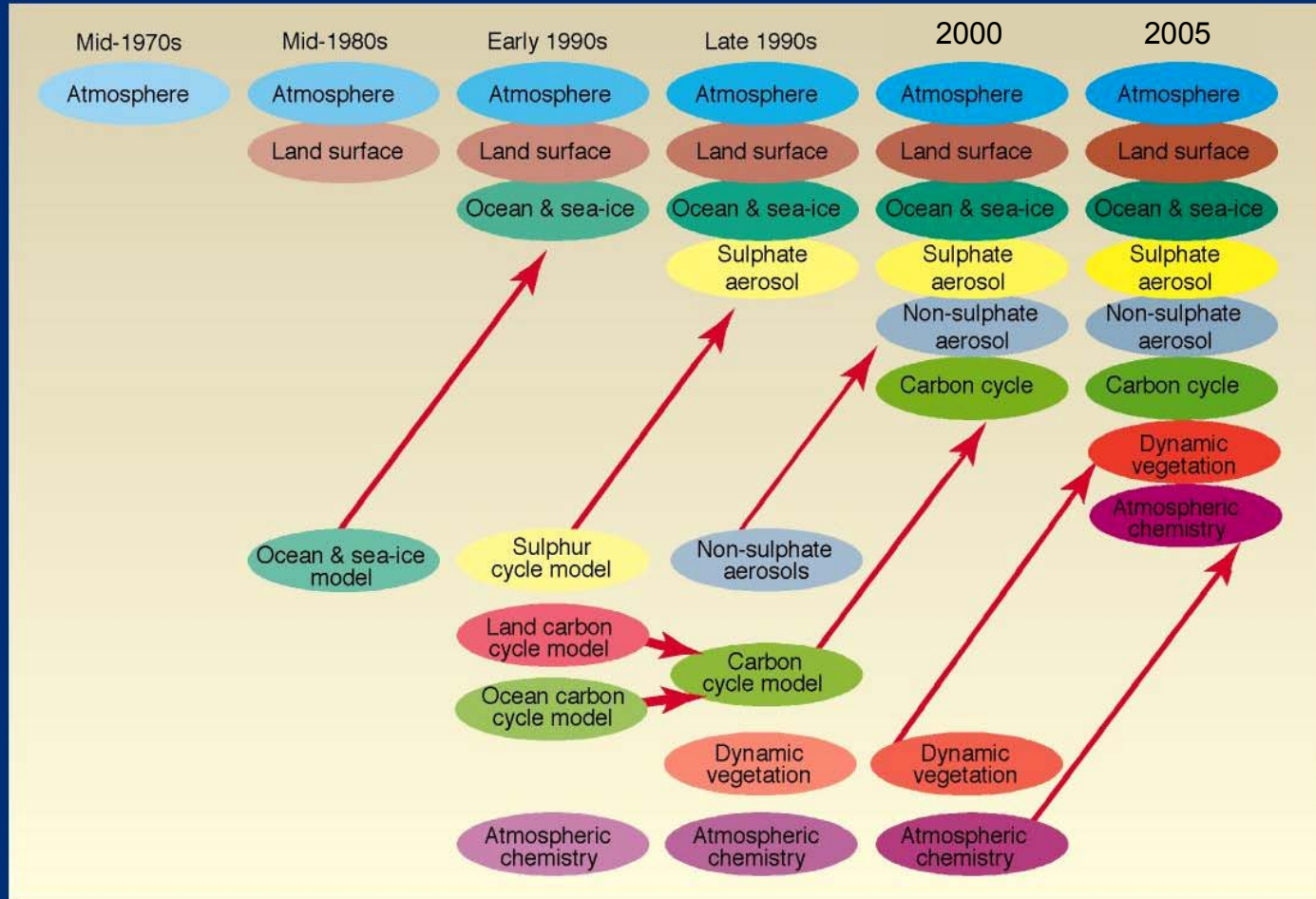
Coupled Models = Increased Technical Complexity



Note: Ocean GCM's are as complex as Atmosphere GCM's!

Climate Model 'Evolution'

The development of climate models, past, present and future



WG1 - TS BOX 3
FIGURE 1

Summary

- Global Climate Modeling
 - complex and evolving scientific problem
 - parameterization of physical processes pacing progress
 - observational limitations pacing process understanding
- Parameterization of physical processes
 - opportunities to explore alternative formulations
 - exploit higher-order statistical relationships?
 - exploration of scale interactions using modeling and observation
 - high-resolution process modeling to supplement observations
 - e.g., identify optimal truncation strategies for capturing major scale interactions
 - better characterize statistical relationships between resolved and unresolved scales



The End

