

# Anders Malmberg

- ▶ PhD in Mathematical Statistics, Lund University, Sweden, 2005. *Space-Time Prediction of Ocean Winds.*
- ▶ Starting project *Mesoscale/Tropical Balance Constraints and Data Assimilation.* Joint work with the Mesoscale and Microscale Meteorology Division, NCAR.
- ▶ Presenting *A Stochastic Transport Model for Atmospheric Carbon Monoxide* a collaboration with David Edwards, Ave Arellano, Atmospheric Chemistry Division, NCAR, Doug Nychka, IMAGE, NCAR, and Chris Wikle, Department of Statistics, University of Missouri.

# A Stochastic Transport Model for Atmospheric Carbon Monoxide

Anders Malmberg

April 27th, 2006

# Outline

## Background

Satellite Data

Statistical Challenges

## Statistical Model

Bayes' Theorem and Hierarchical Modeling

Hierarchical Stages

## Observing Systems Simulation Experiment

Experiment Setup

Gibbs Estimations

## Future Work

# Outline

## Background

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# Carbon Monoxide, CO

## ▶ **Facts**

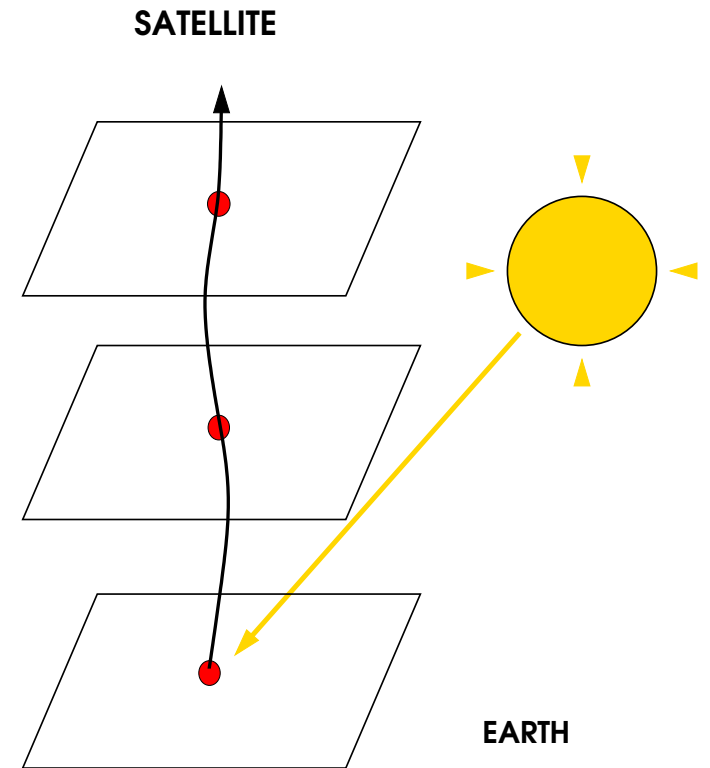
- ▶ Anthropogenic sources accounts for a major part of the atmospheric CO. Examples are incomplete combustion in automobile engines, industrial processes, and the burning of forests to expand agriculture.
- ▶ In the atmosphere, CO contributes to the abundance of ozone, methane, and other greenhouse gases.
- ▶ One of six air pollutants that the US Congress and the EPA mandates regular monitoring of since it endanger public health and the environment.

## ▶ **Atmospheric Research Interests**

- ▶ CO is used as a fingerprint for anthropogenic activities.
- ▶ CO is important for making statements about air quality and studying climate change.

# How to Remotely Sense Atmospheric CO

CO in different vertical layers in the lower portion of the atmosphere can be remotely sensed from space by measuring the surface radiance.



# Satellite Retrievals

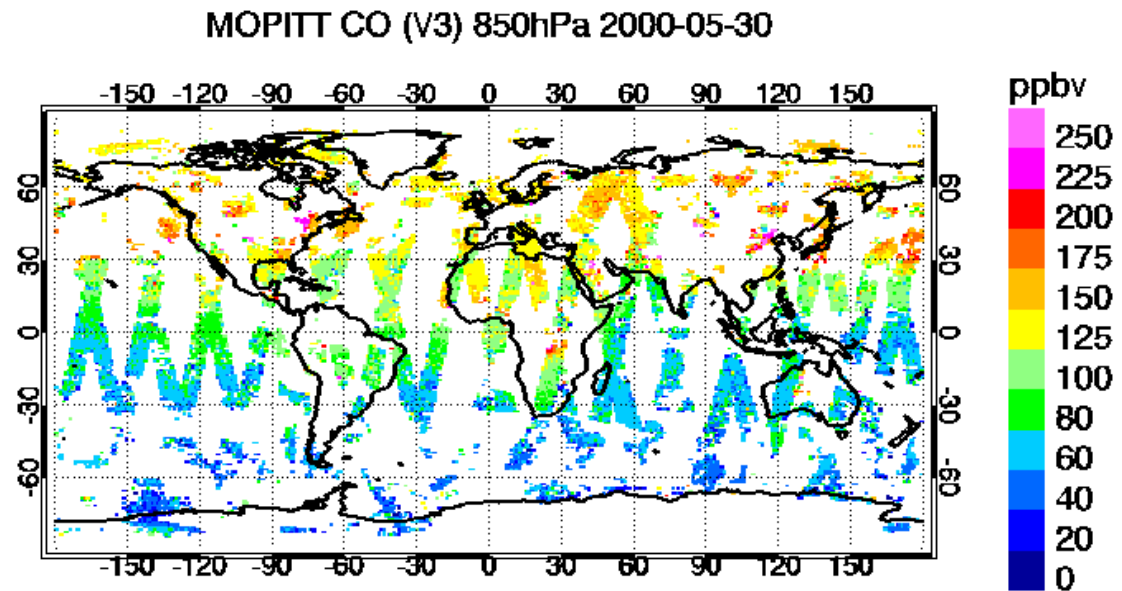
Solving the inverse problem, the column of retrieved concentrations at a longitude/latitude location  $s$  and time  $t$ , can be approximated with

$$\begin{pmatrix} Z^1(s, t) \\ Z^2(s, t) \\ \vdots \\ Z^7(s, t) \end{pmatrix} = A \begin{pmatrix} \alpha^1(s, t) \\ \alpha^2(s, t) \\ \vdots \\ \alpha^7(s, t) \end{pmatrix} + (\mathbf{I} - A)\alpha_{prior} + \begin{pmatrix} \epsilon^1(s, t) \\ \epsilon^2(s, t) \\ \vdots \\ \epsilon^7(s, t) \end{pmatrix}$$

where  $A$  is a 7 by 7 averaging matrix,  $\alpha(s, t) = (\alpha^1(s, t), \dots, \alpha^7(s, t))^T$  is the unknown true concentration that we want to estimate, and  $\epsilon$  is assumed to be additive white noise. The inverse problem is constrained by  $\alpha_{prior}$ .

# Satellite Retrievals

The Terra satellite circles the earth in 99 minutes and a near global coverage in 5 days. The plot shows daily satellite retrievals, at 850 hPa, gridded to 1 degree horizontal resolution.



Gridded at 1x1deg from MOP02-20000530-L2V5.7.1 prov.hdf (apriori fraction < 50%)



# Statistical Challenges

- ▶ **We have:**

- ▶ sparse data in space and time,
- ▶ data contaminated with observation noise.

- ▶ **We want to:**

- ▶ validate satellite data,
- ▶ have a complete space-time data set and a quantified measure of uncertainty,
- ▶ compare satellite data to chemical transport model data.

- ▶ **Statistical challenges:**

- ▶ estimate a dynamical model in four dimensions that allows us to fill in spatial and temporal gaps in the data,
- ▶ merge different data sets in an objective fashion while accounting for their different uncertainties.

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# Bayes' Theorem and Hierarchical Modeling

Assuming we want to estimate a process and some parameters given data,  $[process, parameters|data]$ , we use Bayes' Theorem to rewrite this in densities that we can specify:

$$\frac{[data|process, parameters][process|parameters][parameters]}{[process, parameters]}$$

- ▶ data stage, likelihood of data given process,
- ▶ process stage, specifying spatial and temporal dynamics,
- ▶ parameter stage, with hyper parameters.

# Data Stage, [*data*|*process*, *parameters*]

Using the approximate solution of the inverse model, we assume that the likelihood of a column of satellite data,  $Z$ , given the true process,  $\alpha$ , is,

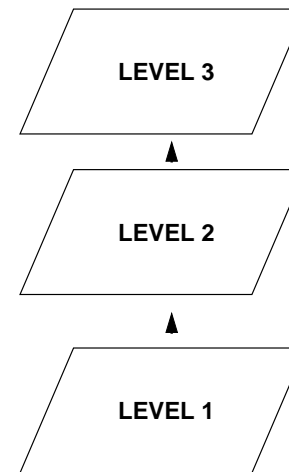
$$\begin{pmatrix} Z^1(\mathbf{s}, t) \\ Z^2(\mathbf{s}, t) \\ \vdots \\ Z^7(\mathbf{s}, t) \end{pmatrix} \mid \begin{pmatrix} \alpha^1(\mathbf{s}, t) \\ \alpha^2(\mathbf{s}, t) \\ \vdots \\ \alpha^7(\mathbf{s}, t) \end{pmatrix} \sim MVN\left(A \begin{pmatrix} \alpha^1(\mathbf{s}, t) \\ \alpha^2(\mathbf{s}, t) \\ \vdots \\ \alpha^7(\mathbf{s}, t) \end{pmatrix} + (\mathbf{I} - A)\alpha_{prior}, \Sigma_{\epsilon}\right),$$

where  $\Sigma_{\epsilon}$  is the variance-covariance matrix of the observations.

# Process Stage, [*process*|*parameters*] I of II

In the vertical, we assume a Markov property. At each time we assume the vertical levels are conditionally independent,

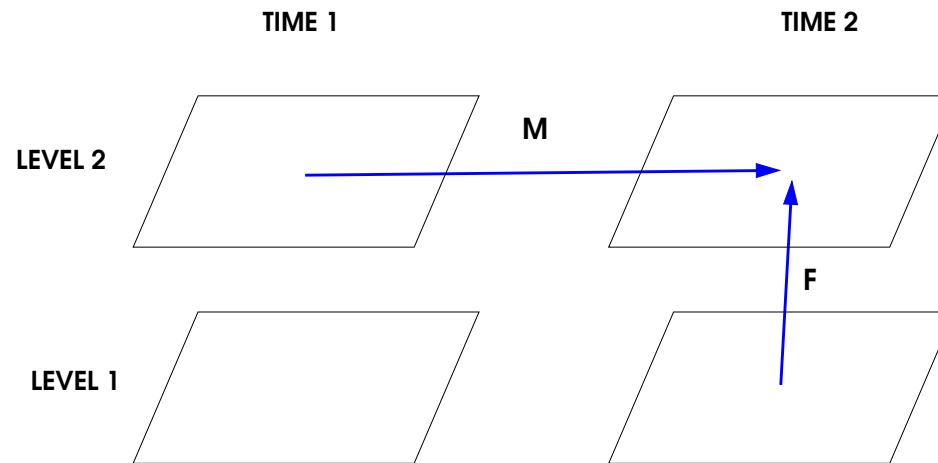
$$[\alpha^3(t), \alpha^2(t), \alpha^1(t)] =$$
$$[\alpha^3(t) | \alpha^2(t)] [\alpha^2(t) | \alpha^1(t)] [\alpha^1(t)]$$



The temporal dynamics are assumed to follow an advection diffusion equation,

$$\alpha^2(t_2) | \alpha^1(t_2), \alpha^2(t_1) \sim MVN(M(t_1)\alpha^2(t_1) + F\alpha^1(t_2), \Sigma_\eta),$$

where M is the advection diffusion matrix, and the forcing matrix F, origins in the vertical Markov assumption.



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# Observing Systems Simulation Experiment

## General Setup:

- ▶  $15 \times 15$  longitude latitude domain in the North Pacific.
- ▶ For the **vertical** model we assume that,

$$\alpha^j(s_i, t) = f^j \alpha^{j-1}(s_i, t).$$

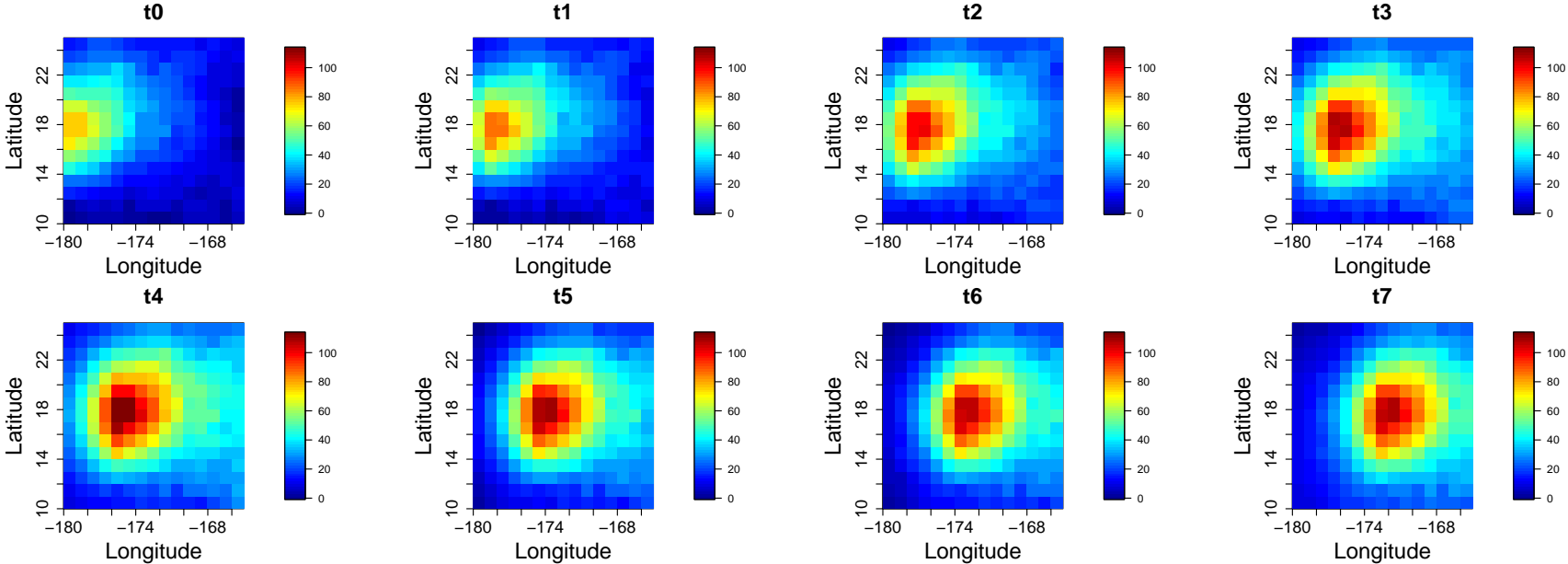
- ▶ For the **horizontal** model we assume a simple translation,

$$\alpha^j(s_i, t) = m^j \alpha^j(s_i - \Delta_x, t - 1).$$

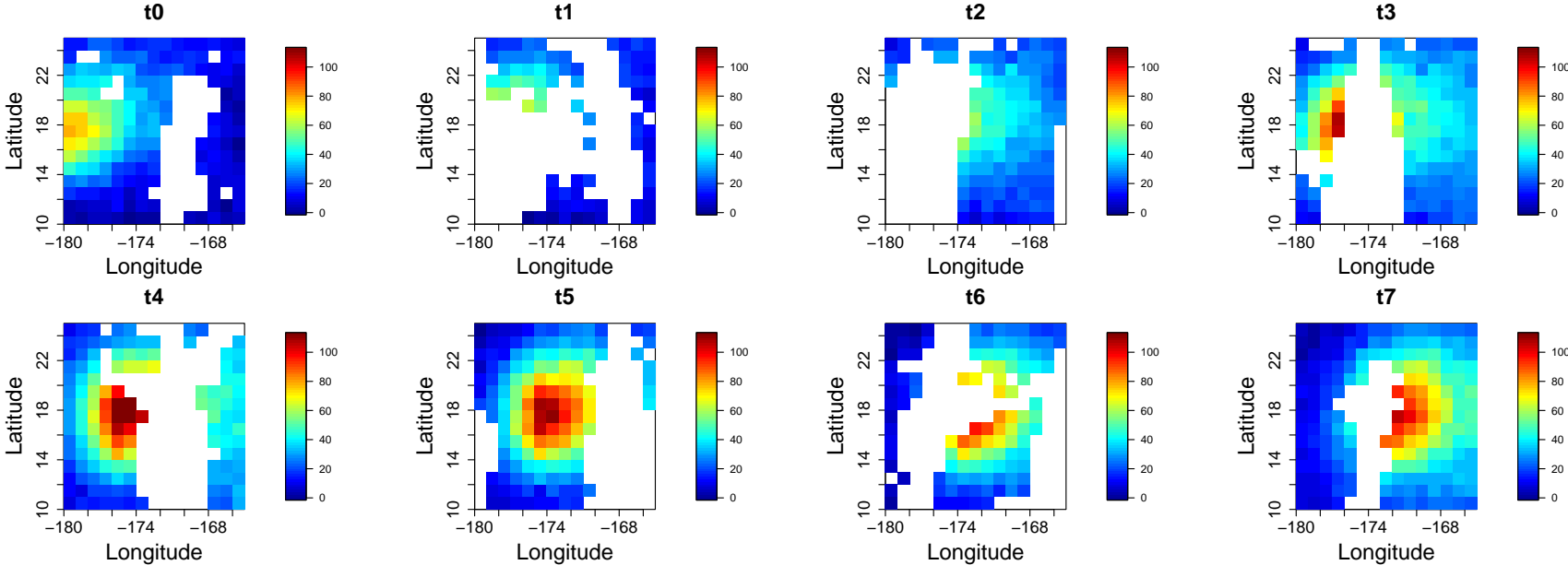
- ▶ Using these parameters and dynamical noise a “true” atmospheric plume is simulated.
- ▶ Apply Gibbs sampler to estimate  $f^j$ ,  $m^j$ , and  $\alpha^j$ .



# Simulated Truth 700 hPa

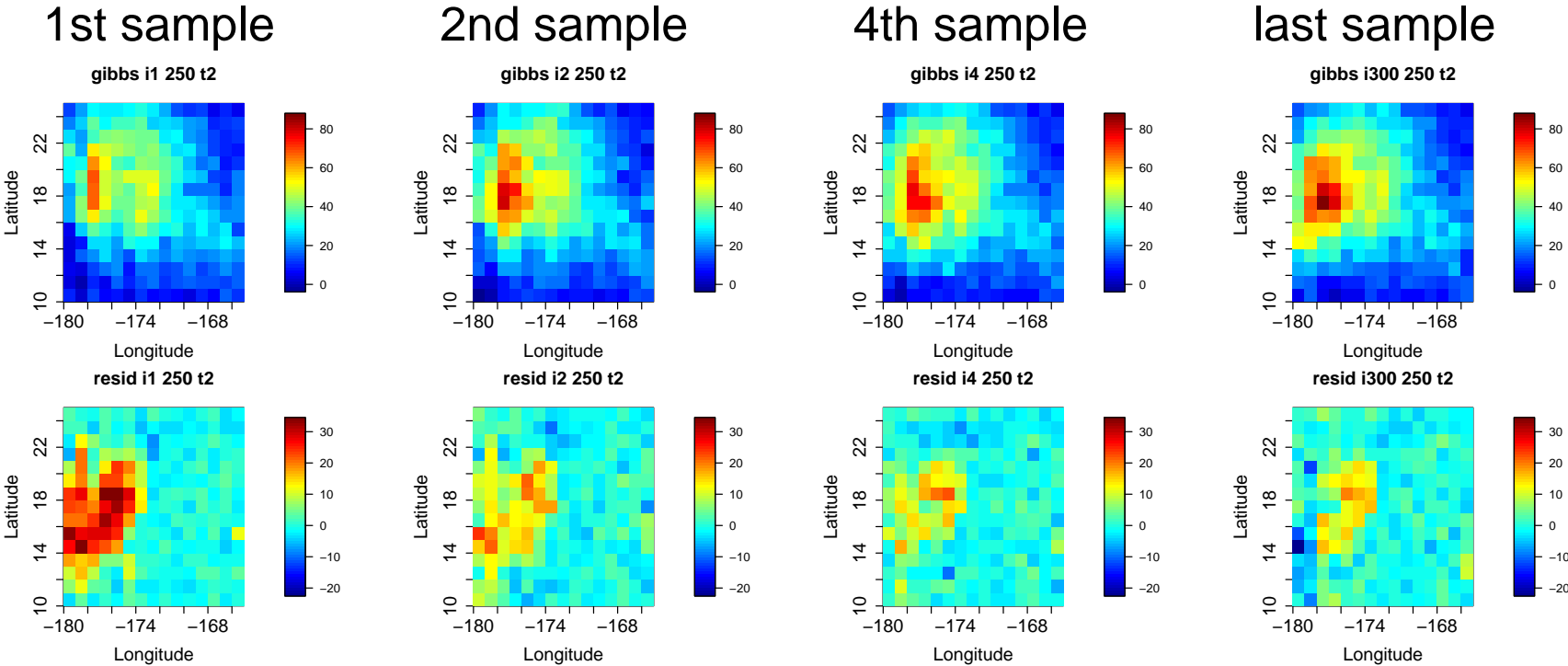


# Simulated Satellite Data 700 hPa



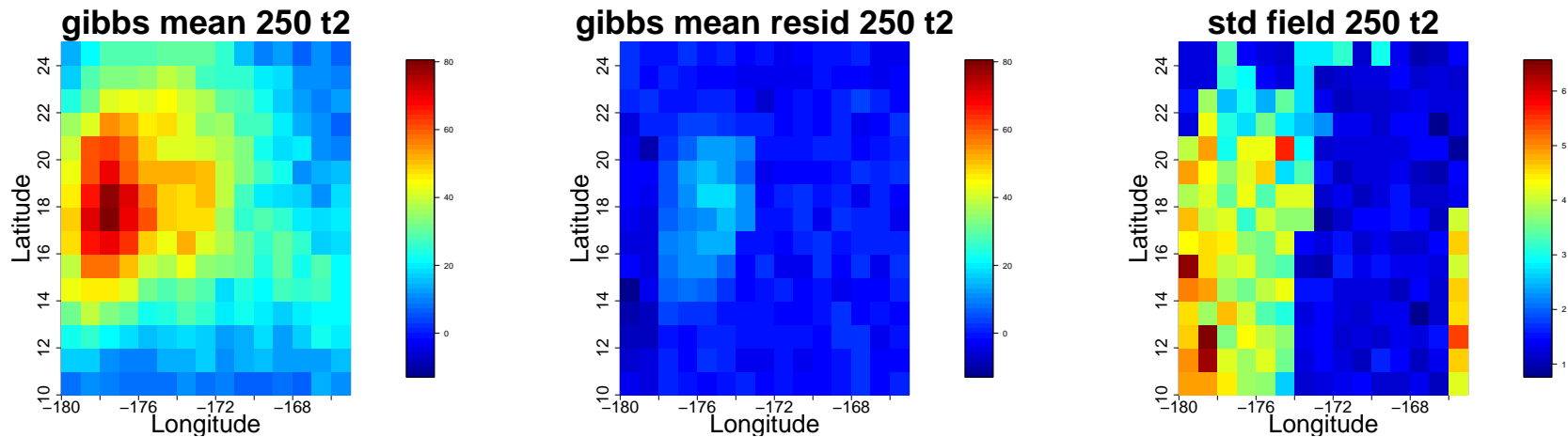
# Gibbs Samples 250 hPa, time $t_2$ , with strong prior

Upper levels has proved to be hard to estimate in this model.



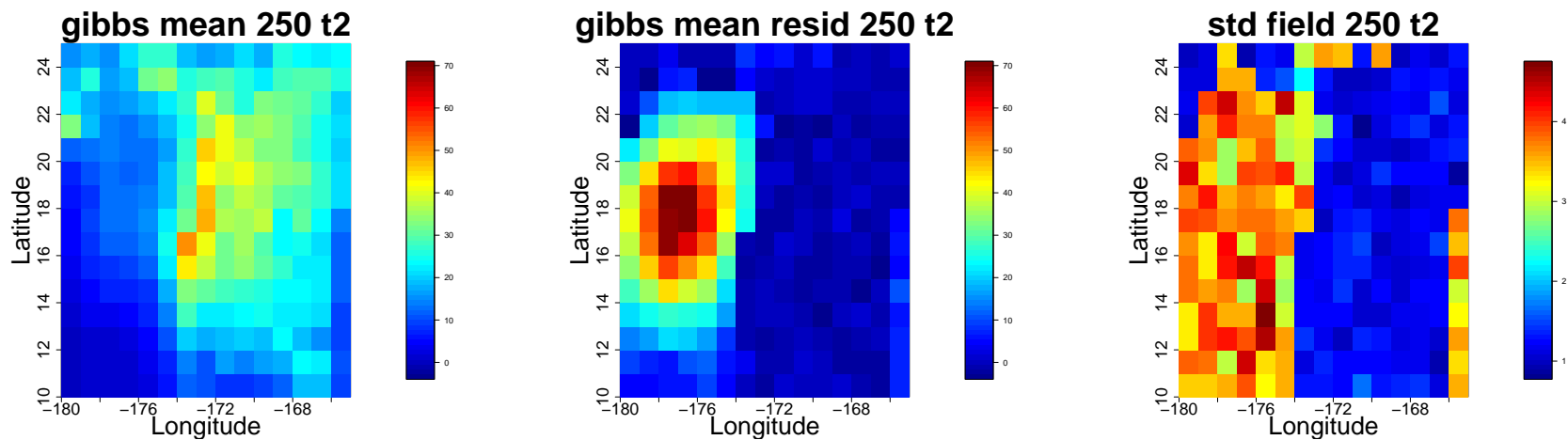
# Gibbs Statistics 250 hPa, time $t_2$ , with strong prior

Using close to true parameters for  $f$  and  $m$  the plume is traceable.  
Locally, the data from the satellite decreases the standard deviation.

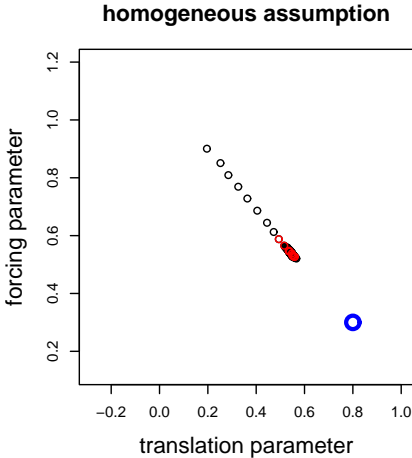
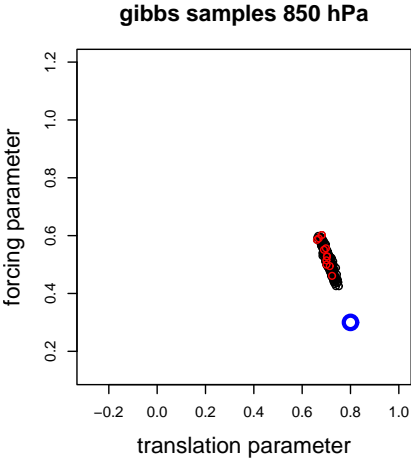
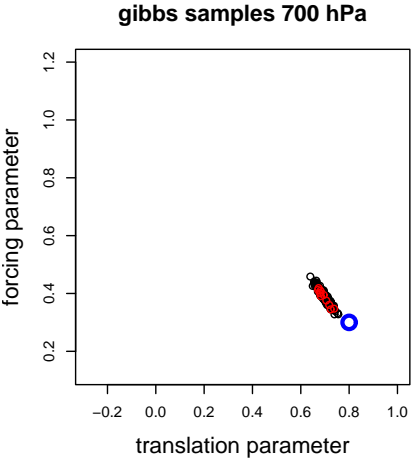
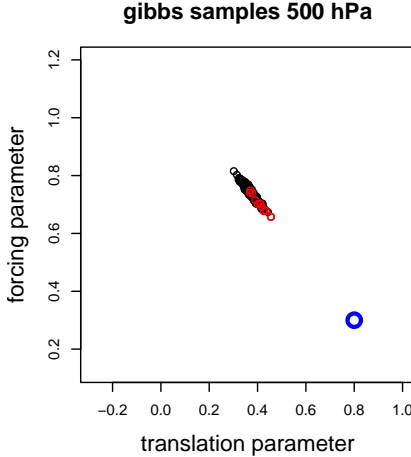
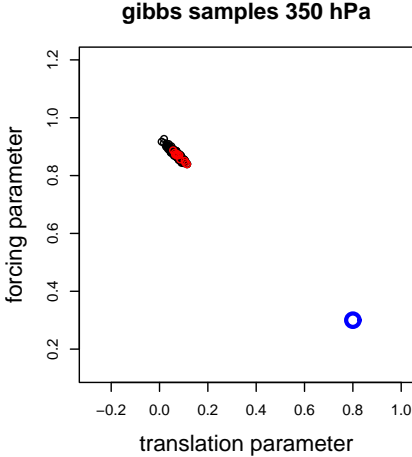
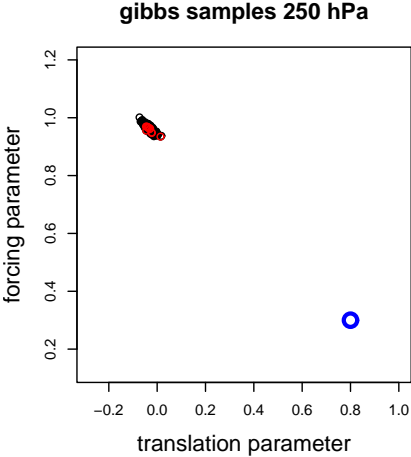
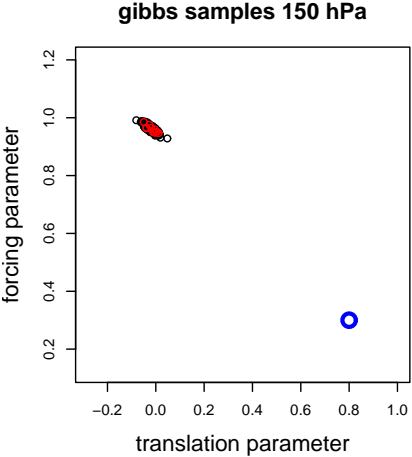


# Gibbs Statistics 250 hPa, time $t_2$

When we use a flat prior it does not work as well. The translation parameter for this level is close to zero and information is not conveyed over time.



# Samples of $f$ versus Samples of $m$



# Full Conditional Means of $m$ and $f$

$$E(m|\cdot) \propto \sum_{k=1}^T \alpha^j(t_{k-1})^T \Sigma_{\eta}^{-1} \alpha^j(t_k) - \sum_{k=1}^T \alpha^j(t_{k-1})^T \Sigma_{\eta}^{-1} \alpha^{j-1}(t_k) \cdot f + \sigma_m^{-2} \mu_m,$$

$$E(f|\cdot) = \left( \sum_{k=1}^T \alpha^{j-1}(t_k)^T \Sigma_{\eta}^{-1} \alpha^{j-1}(t_k) \right)^{-1}.$$

$$\sum_{k=1}^T \alpha^{j-1}(t_k)^T \Sigma_{\eta}^{-1} (\alpha^j(t_k) - \alpha^j(t_{k-1}) \cdot m) + \sigma_f^{-2} \mu_f.$$

- ▶ If  $f \rightarrow 1$  and  $\alpha^j(t_k) \sim \alpha^{j-1}(t_k)$  then  $(\alpha^j(t_k) - \alpha^{j-1}(t_k)f) \rightarrow 0$ .
- ▶ If  $m \rightarrow 0$  then  $E(f|\cdot) \rightarrow 1$ .

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# Future Challenges

- ▶ Apply full advection/diffusion model to output from chemical transport model. To what extent can a simple stochastic model emulate complex dynamics?
- ▶ What impact on parameter estimates and spatial prediction do the sparse sampling have? Will the dynamical model be informative enough to do the spatial interpolation?
- ▶ When we apply the true averaging kernel, how much can we resolve in the vertical?