

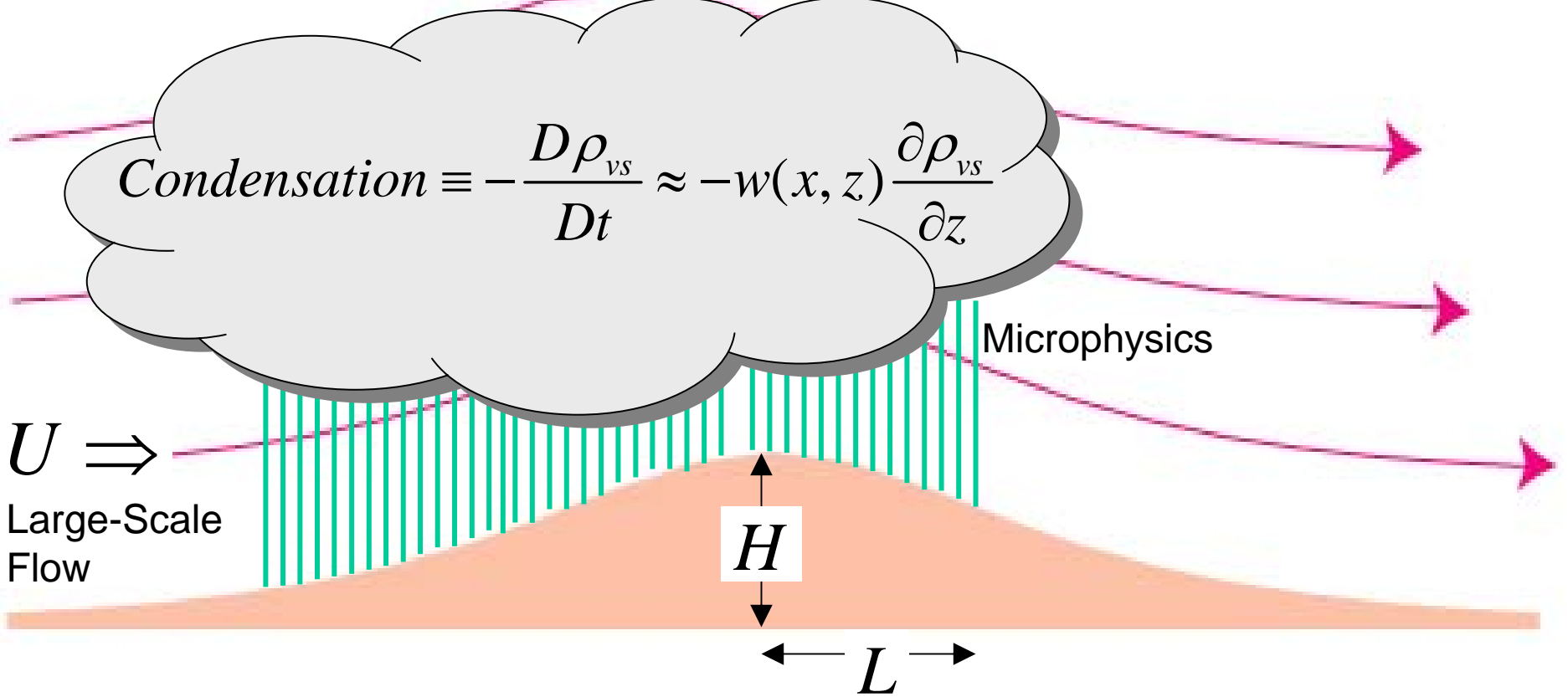
Orographic Precipitation I: Observation and Theory Overview

Richard Rotunno

National Center for Atmospheric Research , USA

*“To explain the amount and distribution of orographic rain ...requires consideration of aspects of meteorology on three different scales. First, there are the **large-scale synoptic factors** which determine the characteristics of the air mass which crosses the hills, its wind speed and direction, its stability and humidity. Second there is the **dynamics of the air motion over and around the hill or hills** with which we are concerned; this determines to what depth and through what layer of the air mass is lifted. Thirdly, there is the **microphysics of the cloud and rain**, which determines whether the water which is condensed as cloud will reach the ground as rain or snow, or whether it will be merely re-evaporated on the leeward side.”*

J. S. Sawyer (1956)



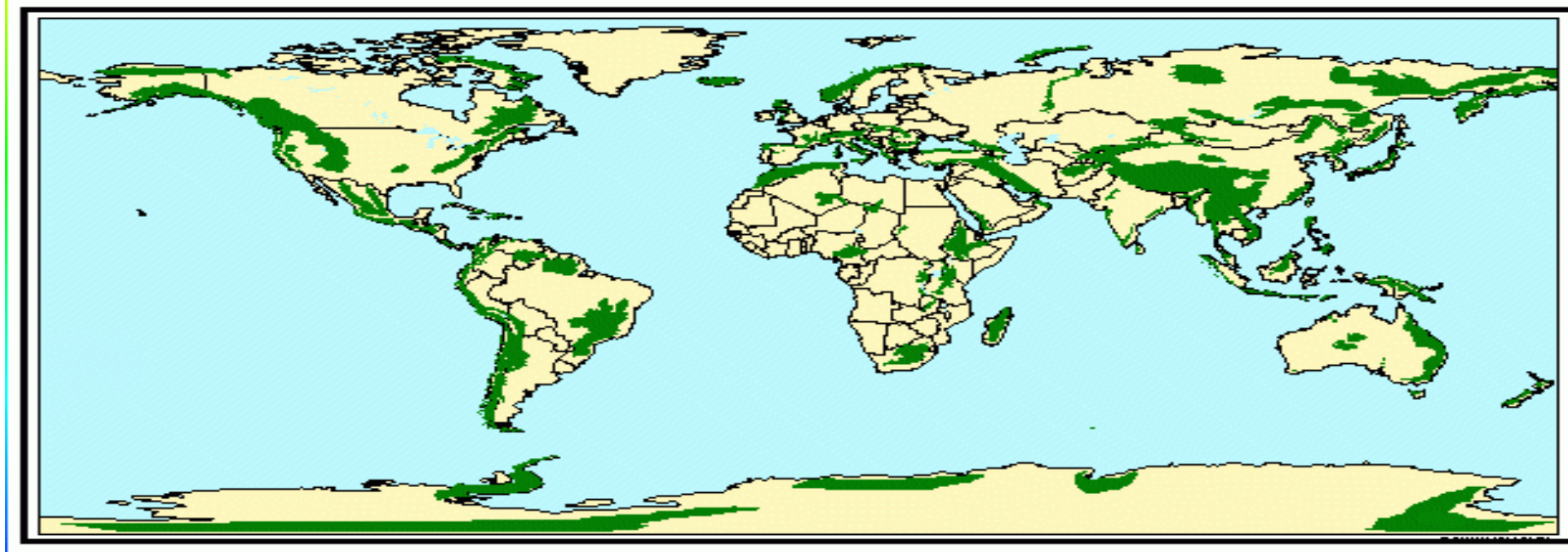
$$w = w(H, L, U, \text{Stability}, \text{Coriolis}, \text{3D Effects})$$

w = vertical velocity

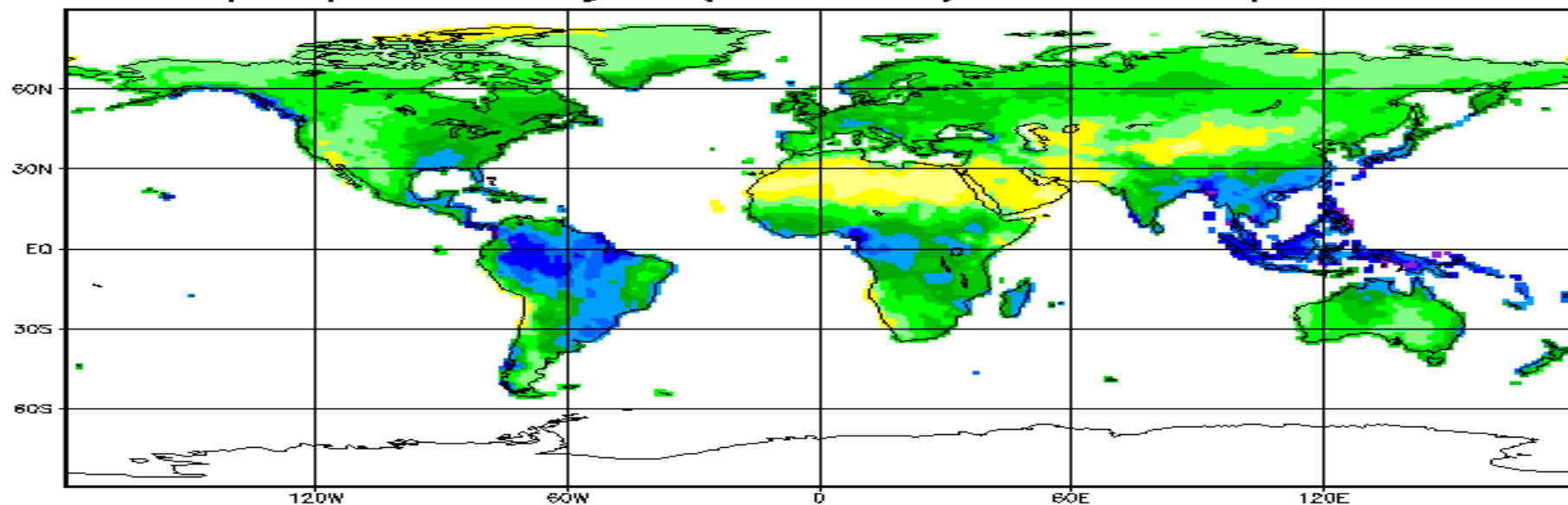
ρ_{vs} = saturation vapor density

...large-scale synoptic factors...

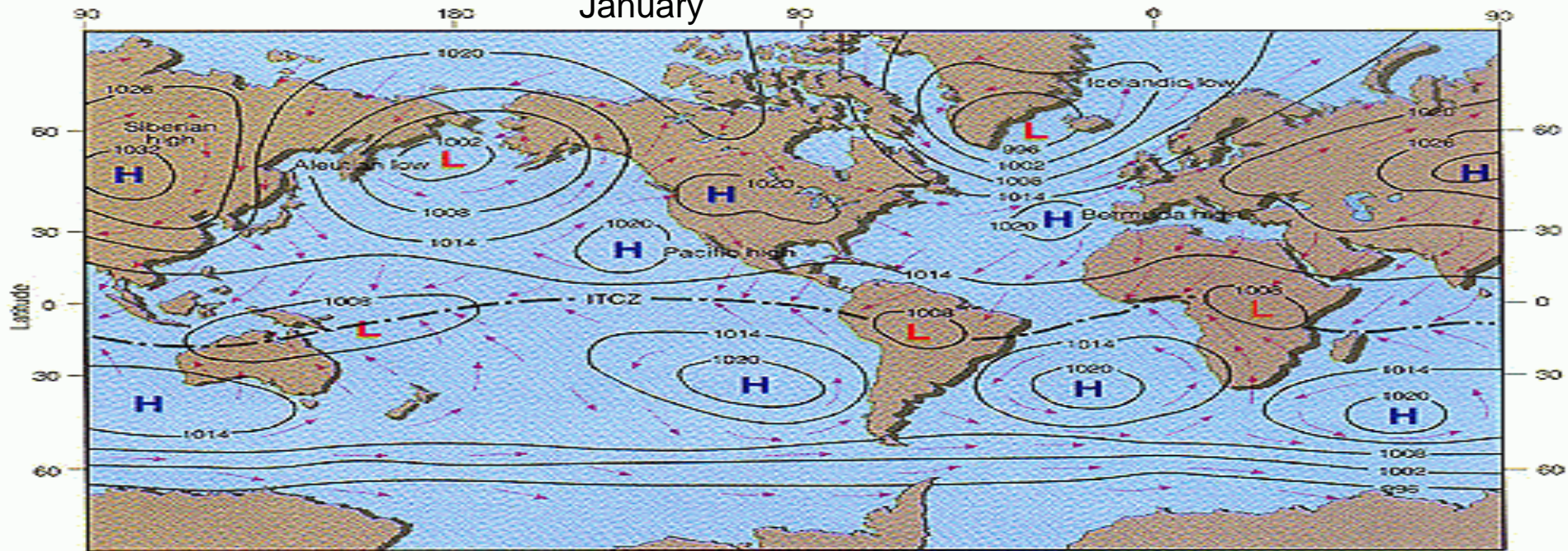
Orography (green > 1km)



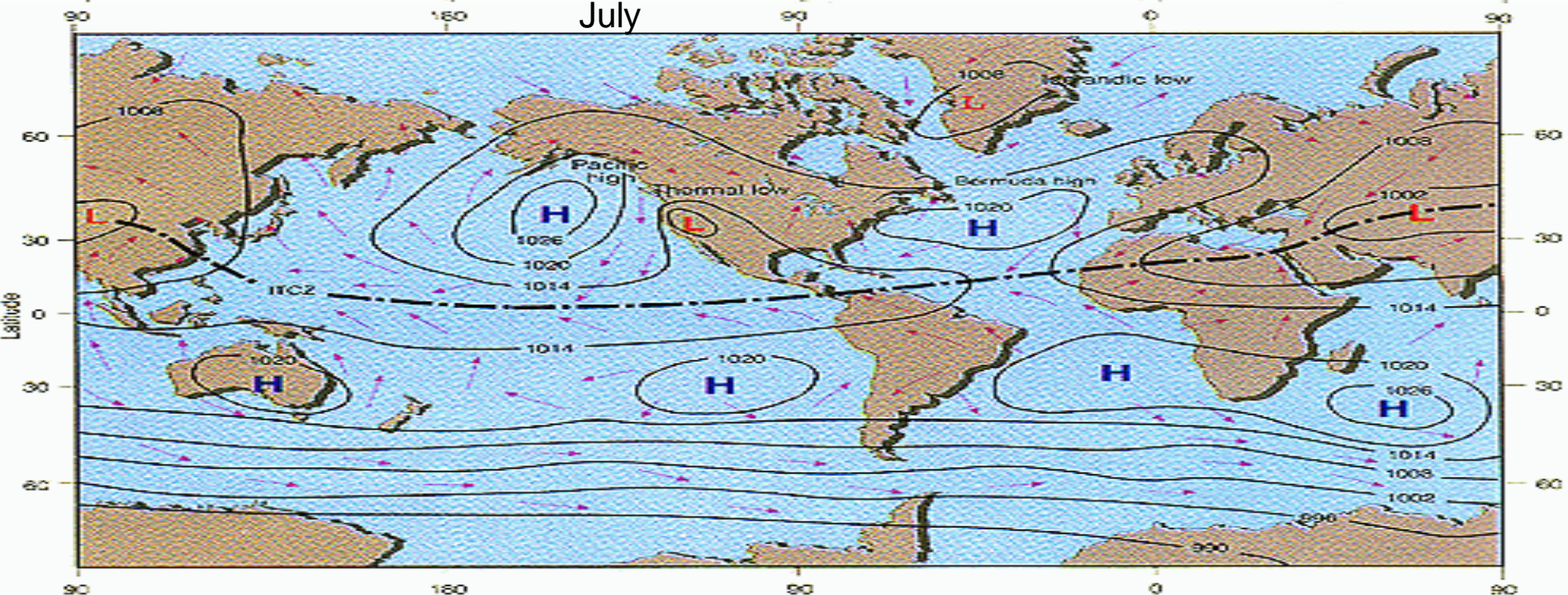
GPCC Monitoring Product Gauge-Based Analysis 1.0 degree precipitation for year (Jan - Dec) 2001 in mm/month



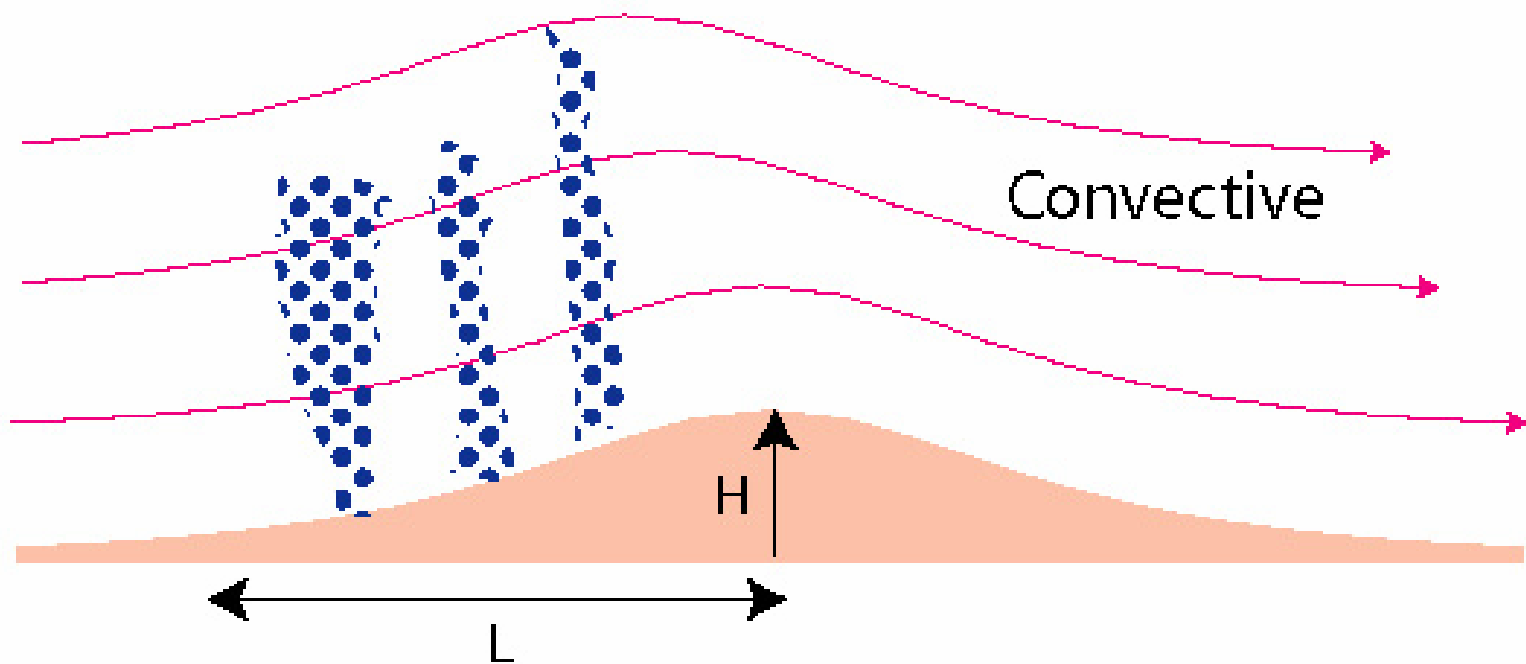
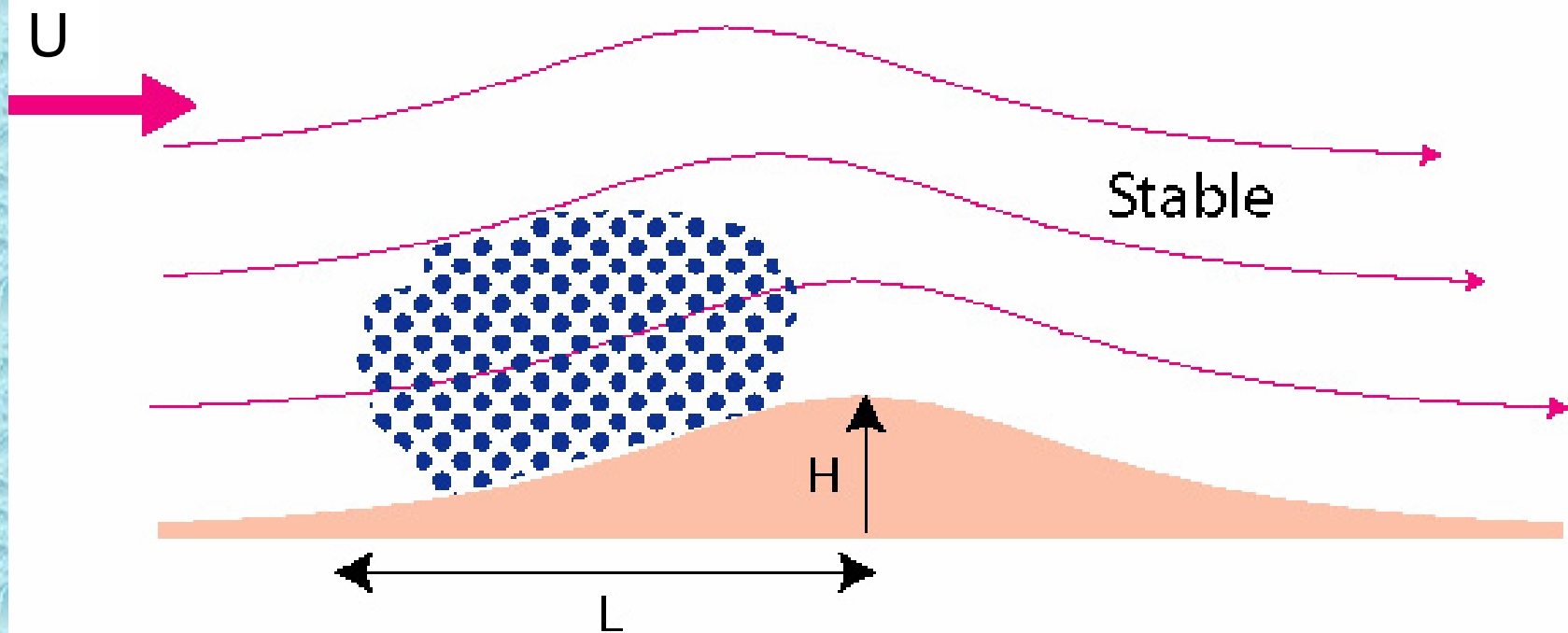
January



July

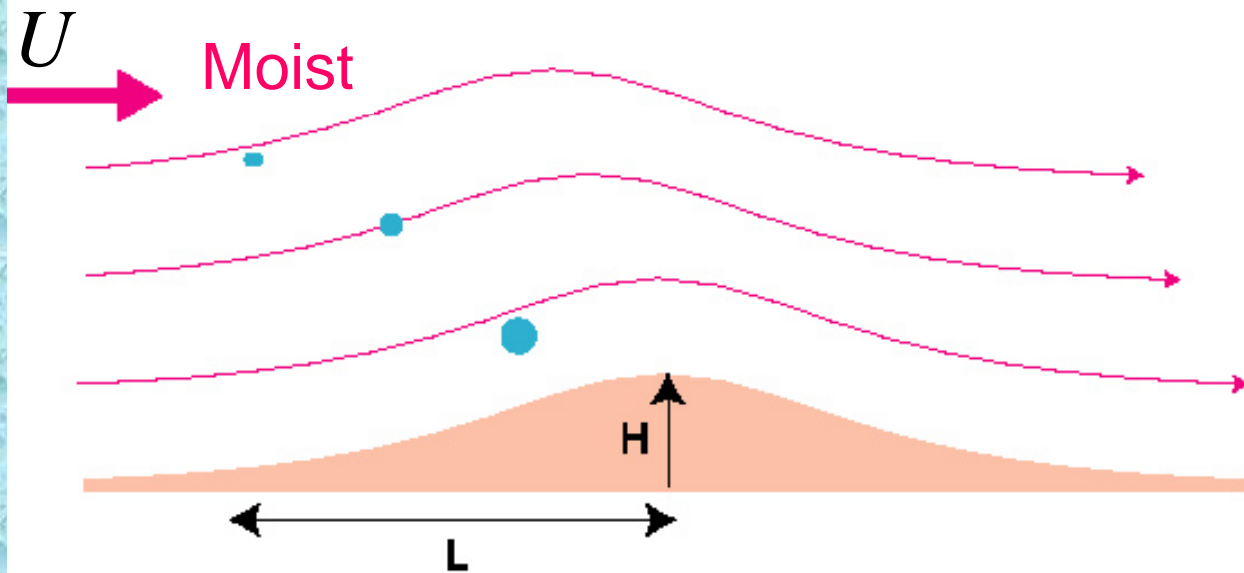


*...dynamics of the air motion
over and around the hill...*





...microphysics of the cloud and rain...



$$\tau_{microphysics} \ll \tau_{airflow}$$

$$1000s \ll \frac{L}{U}$$

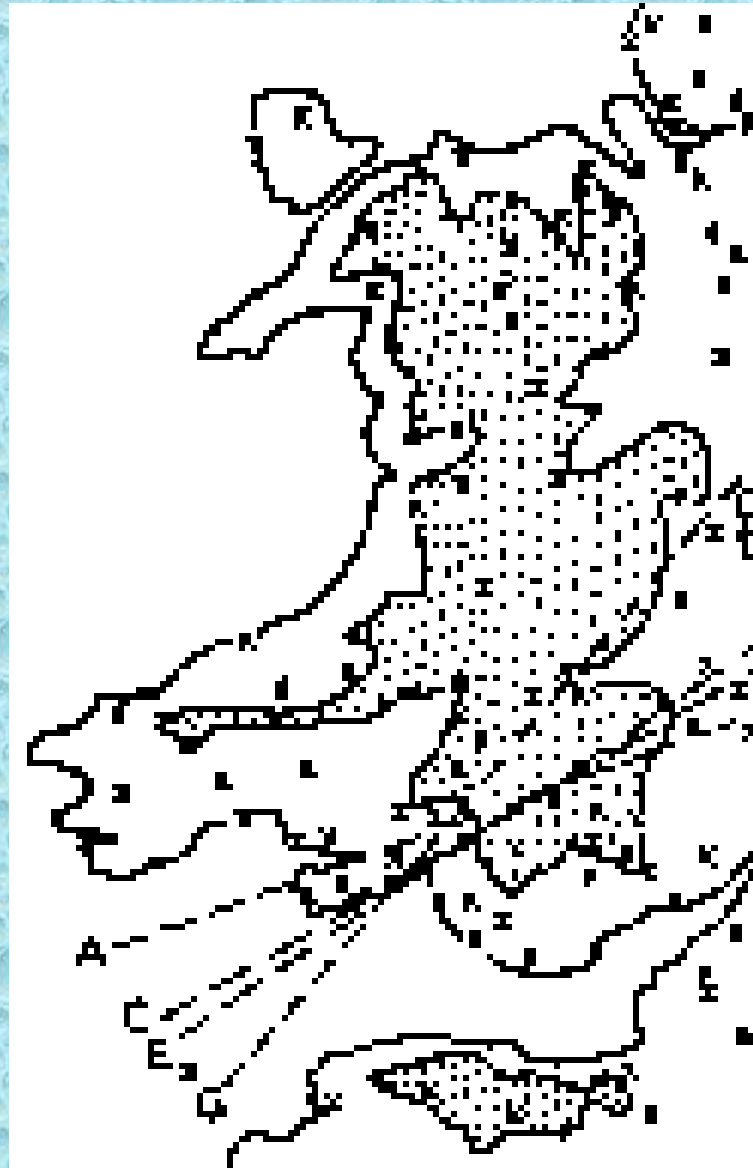
$$U = 10m/s$$

L	L/U
100km	10000s
10km	1000s



A Few Examples

Wales , UK (stippled > 150 m)



100 km

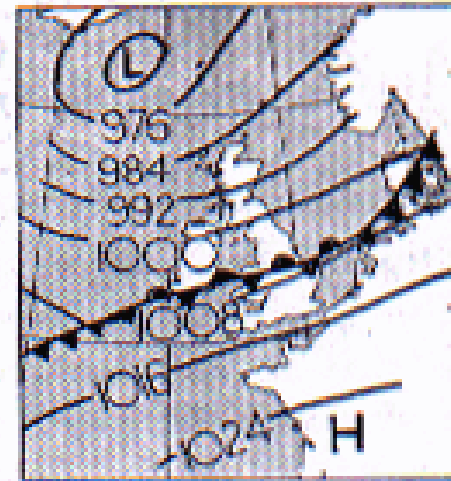


28th MARCH 1953

RAINFALL (INCHES)



Synoptic chart 1800 G.M.T



Valentia 1500 GMT.

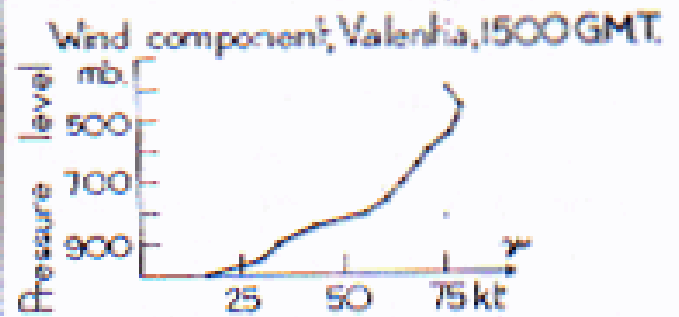
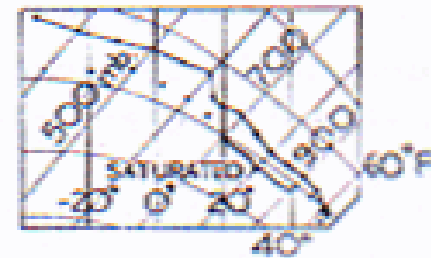


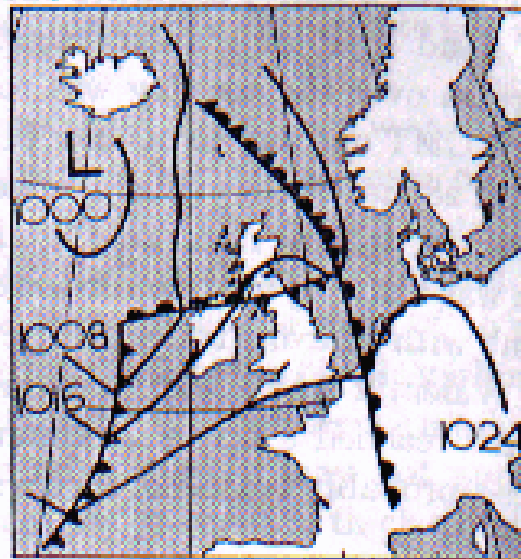
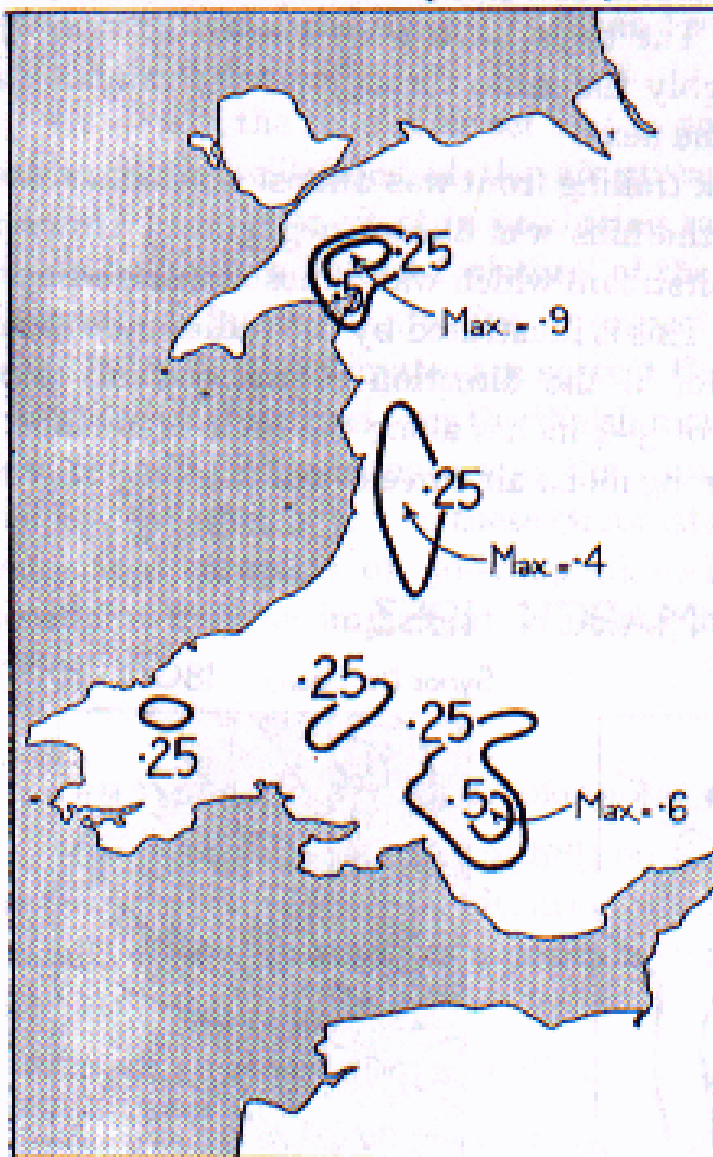
Fig. 2

Sawyer (1956)

17th. JUNE 1954

RAINFALL (INCHES)

Synoptic chart 1800 GMT.



Camborne 1400 GMT.

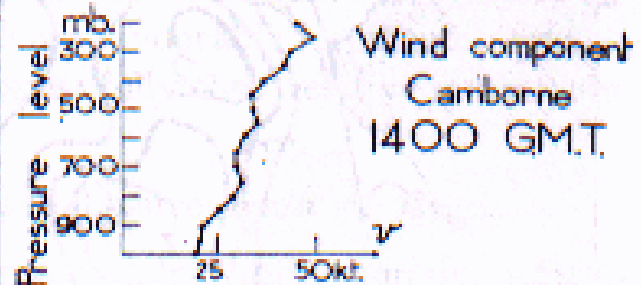
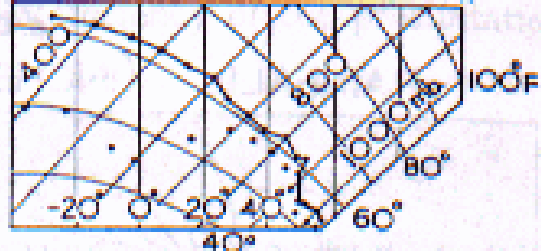
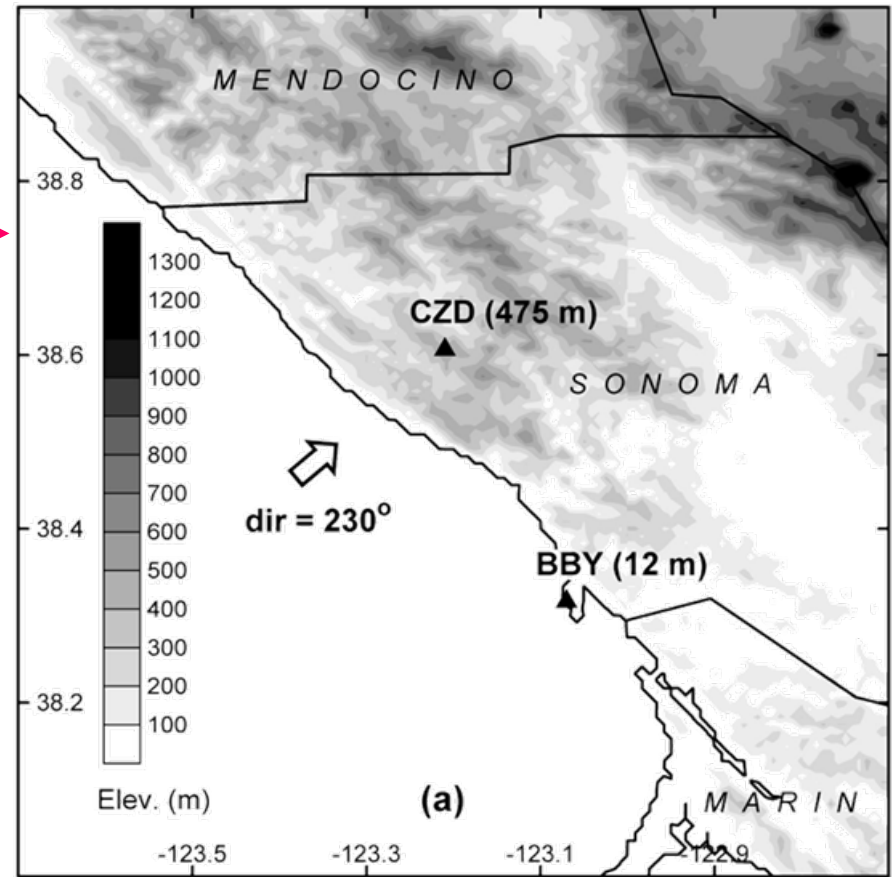
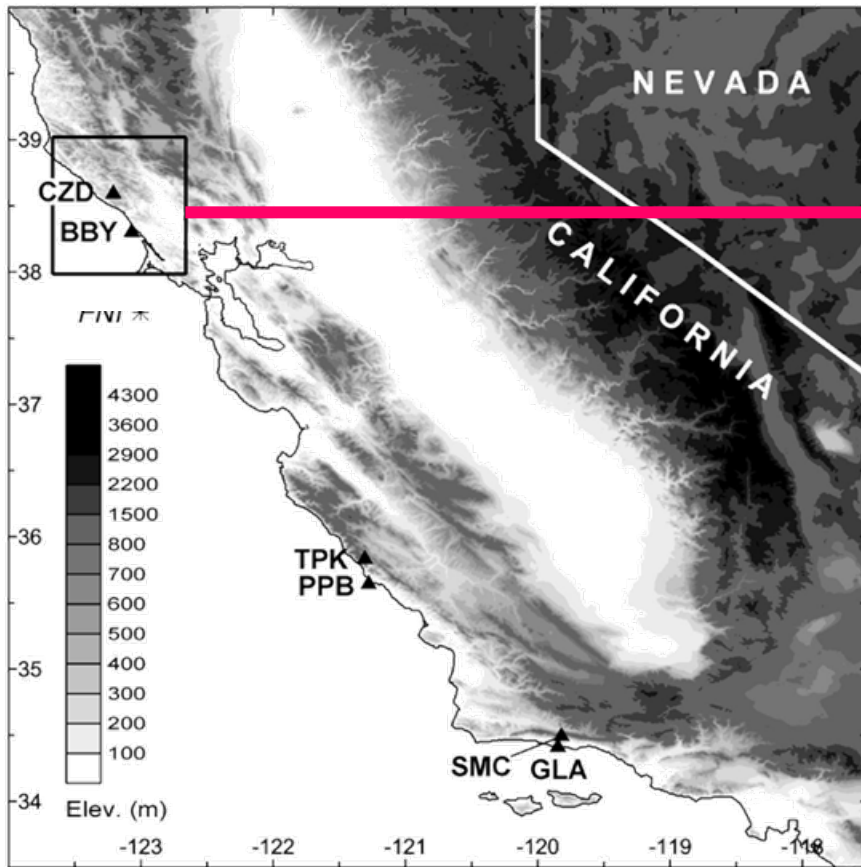
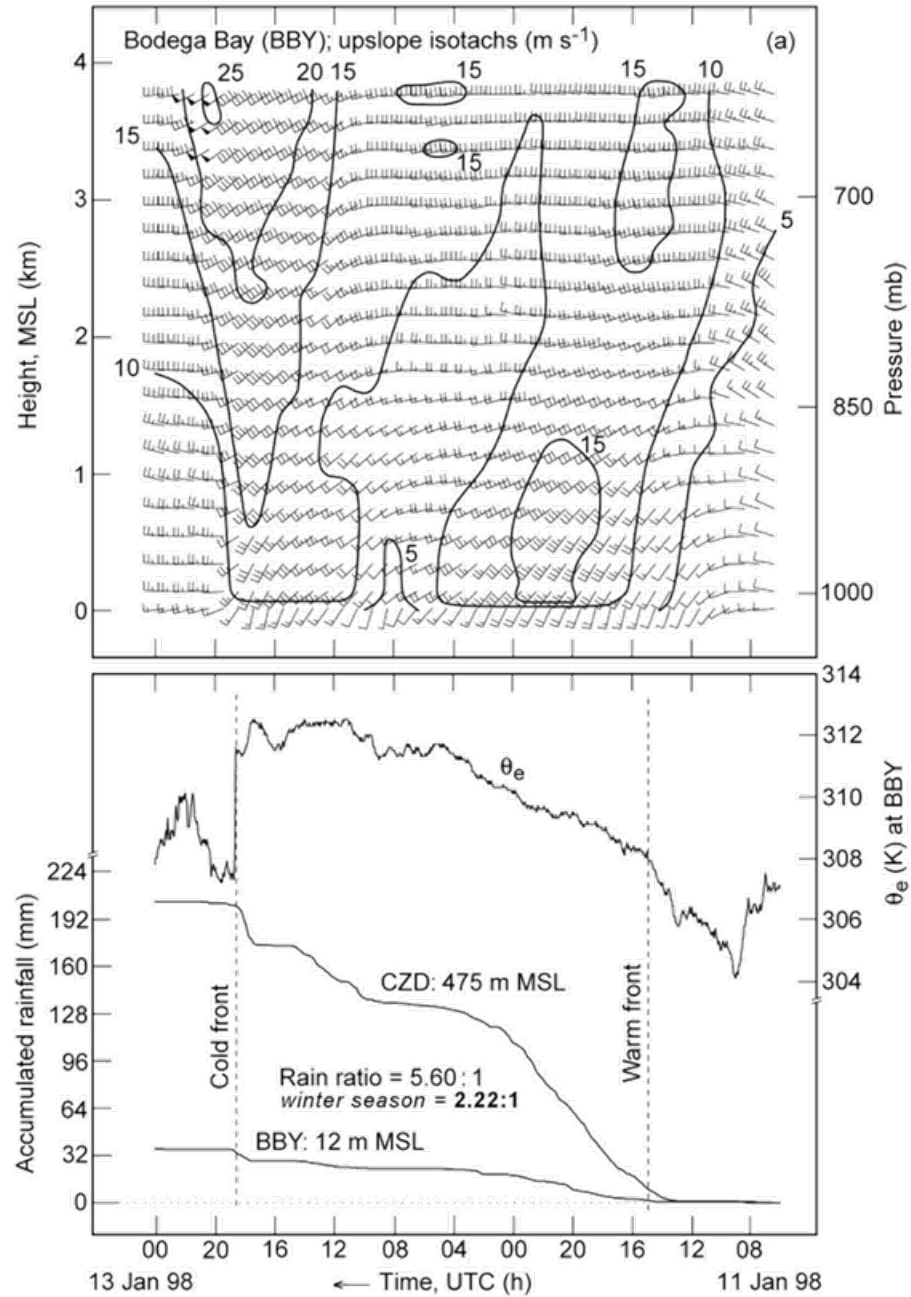


Fig. 3

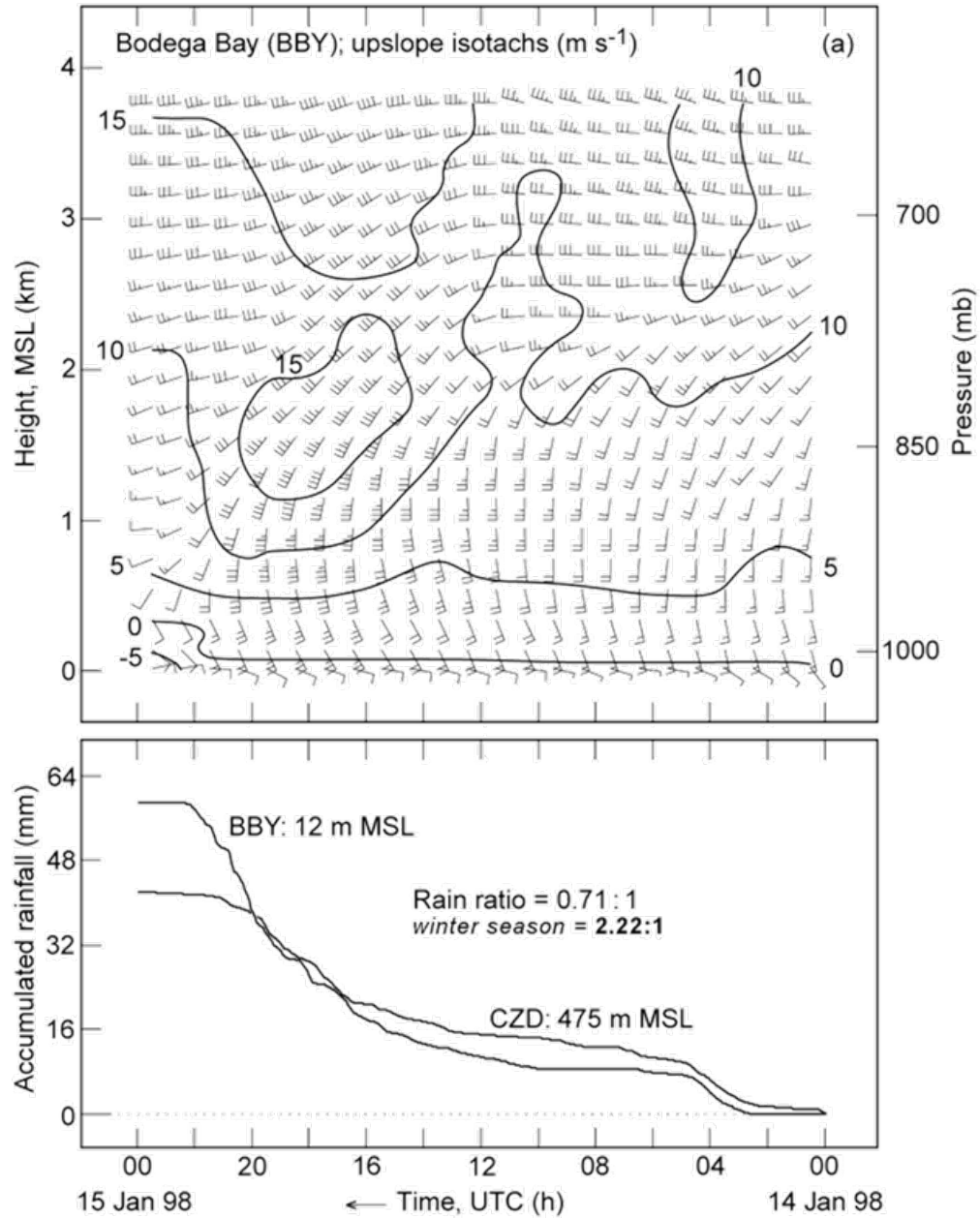
Sawyer (1956)

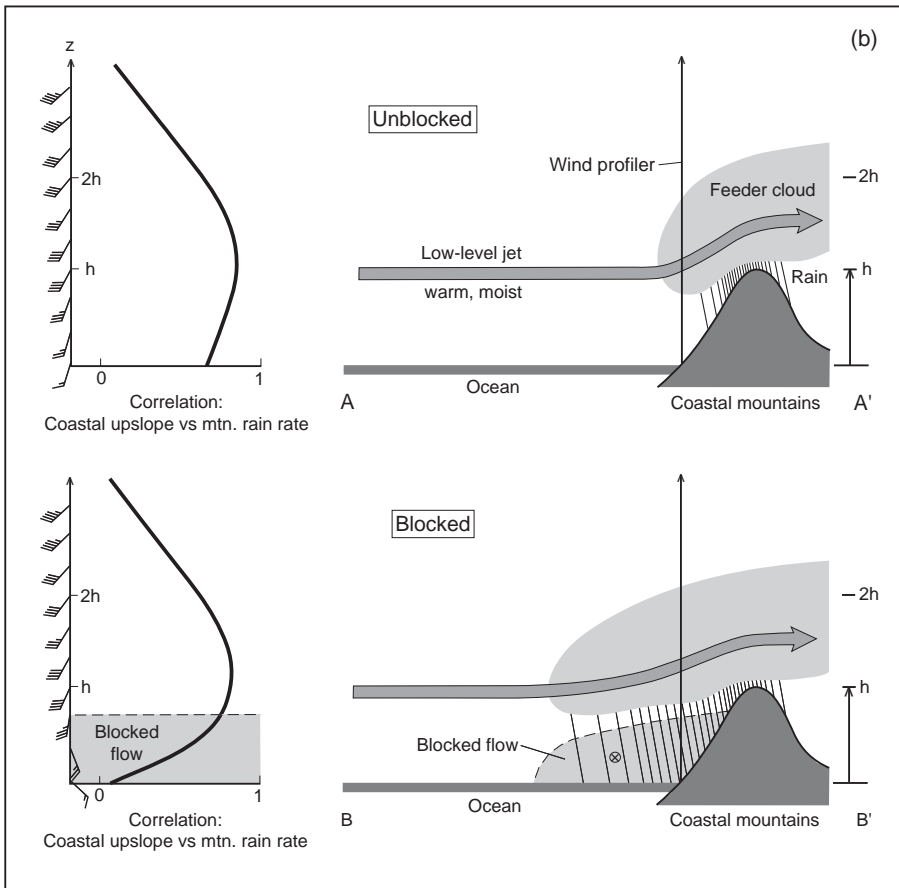
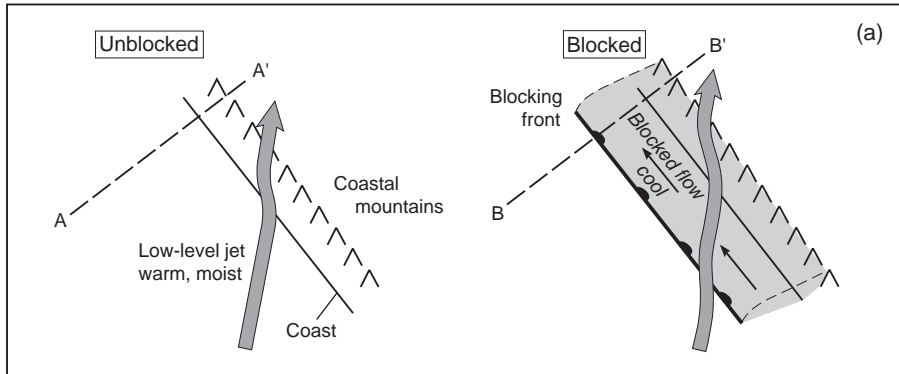


Unblocked



Blocked





➤ Rain rate in coastal mtns directly linked to upslope flow at coast.

➤ In blocked flow, near-sfc winds do not provide useful rain-rate info.

063/046 900804 EOSAT TM-4 741 93210015

E0151

H155-581

E020

E0251

E030 Z

1 0 4 4 5 5 M

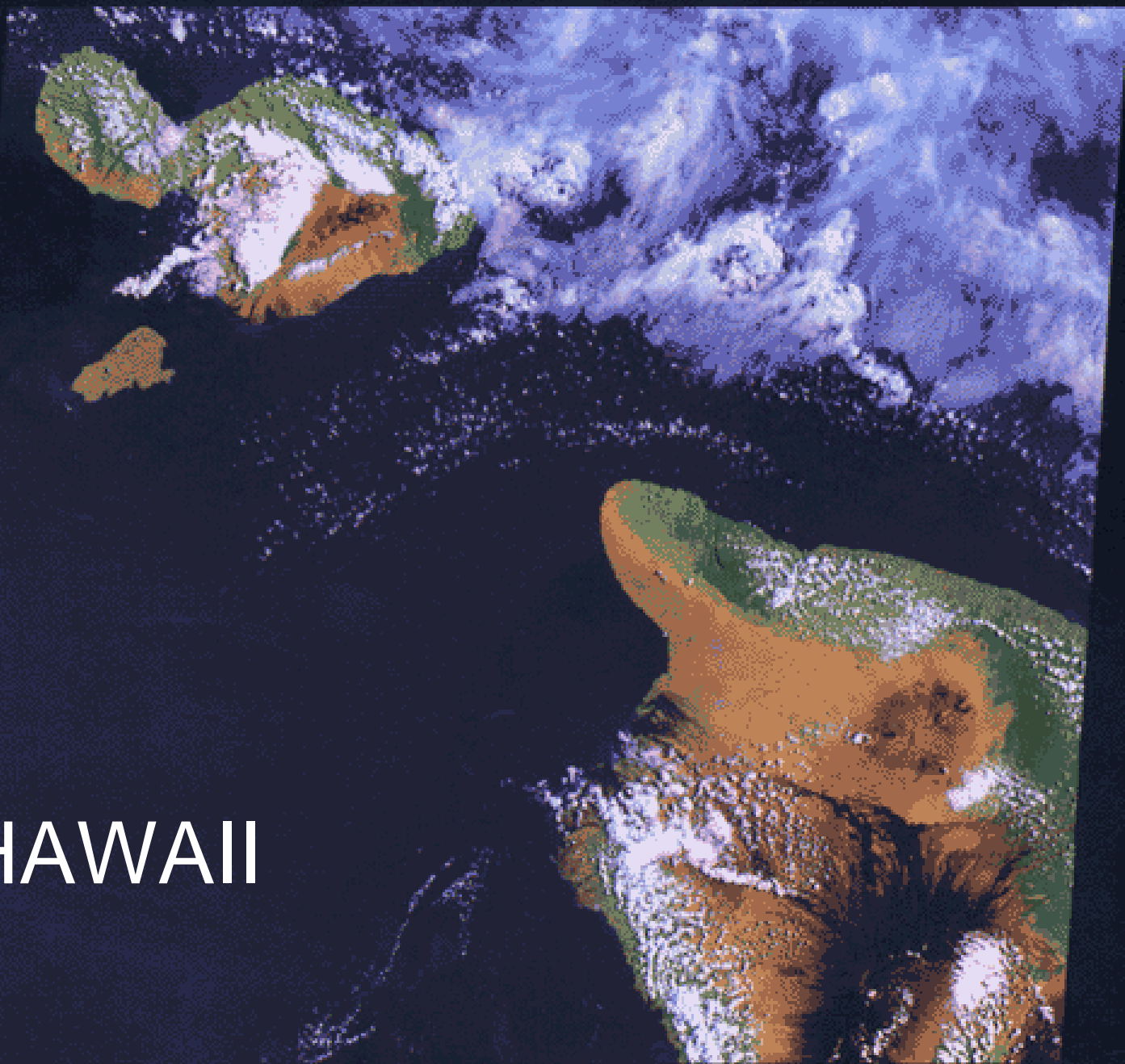
1 0 4 4 5 5 M

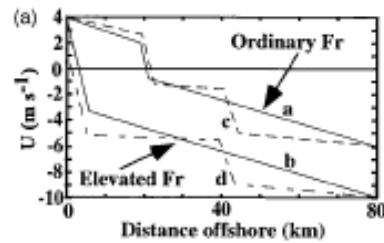
1 0 4 4 5 5 M

1 0 4 4 5 5 M

1 0 4 4 5 5 M

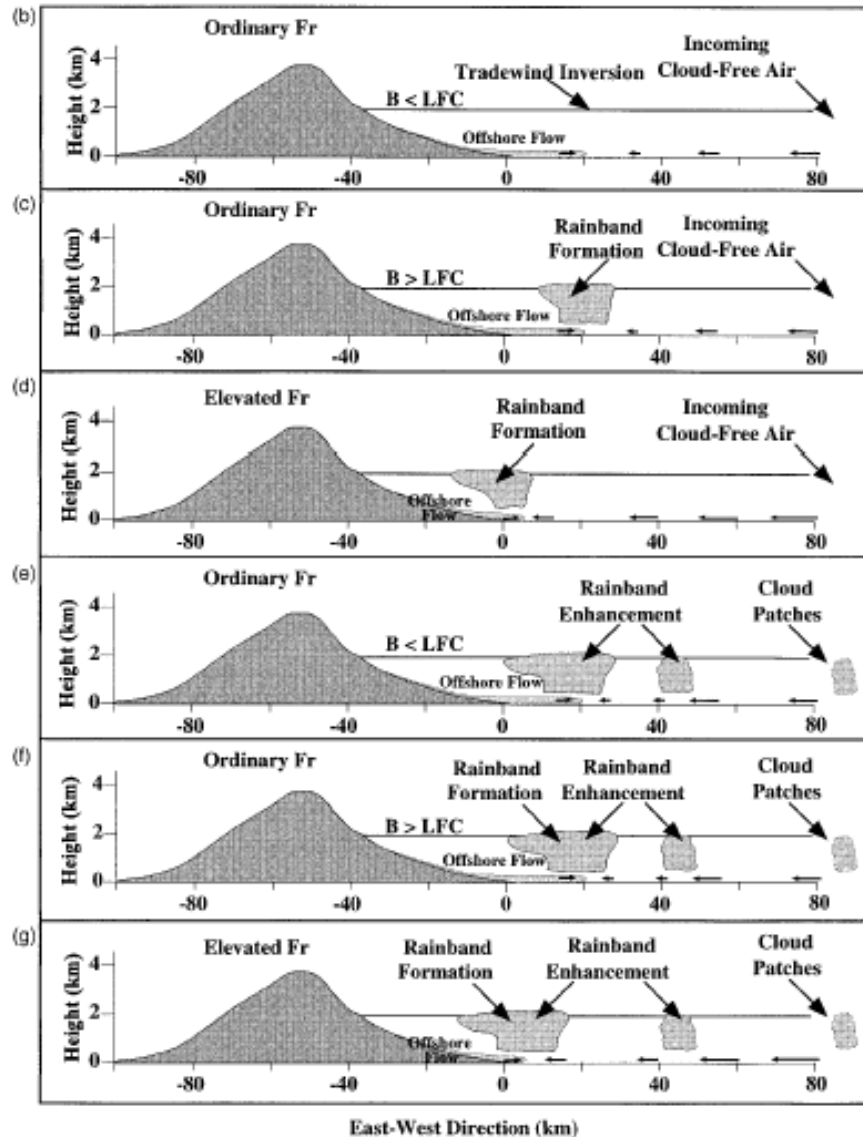
HAWAII





HAWAII

$$B = \frac{U}{N} \quad Fr = \frac{U}{NH}$$



Wang, Rauber, Ochs
And Carbone (2000)

Oahu Flood, Hawaii 1974

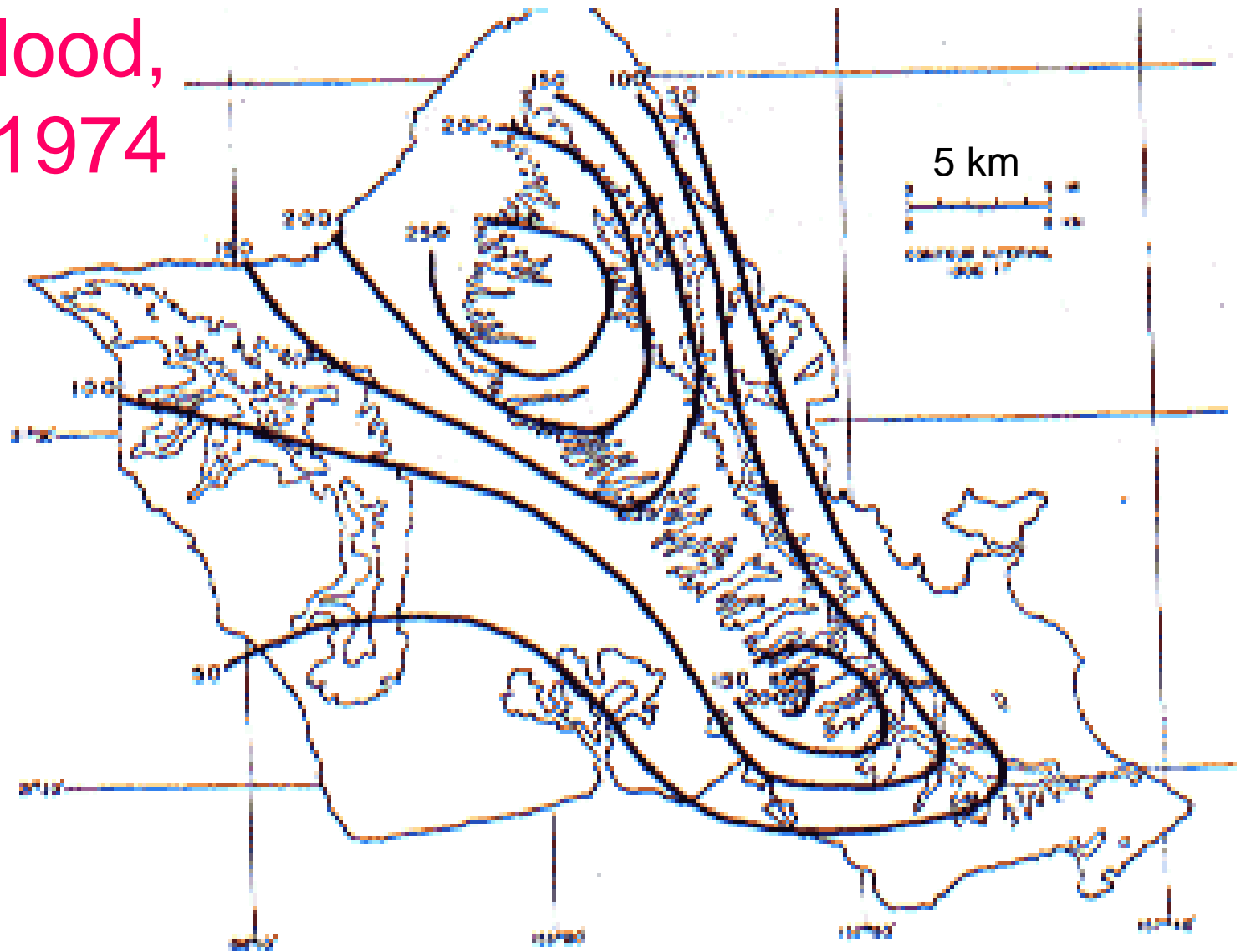
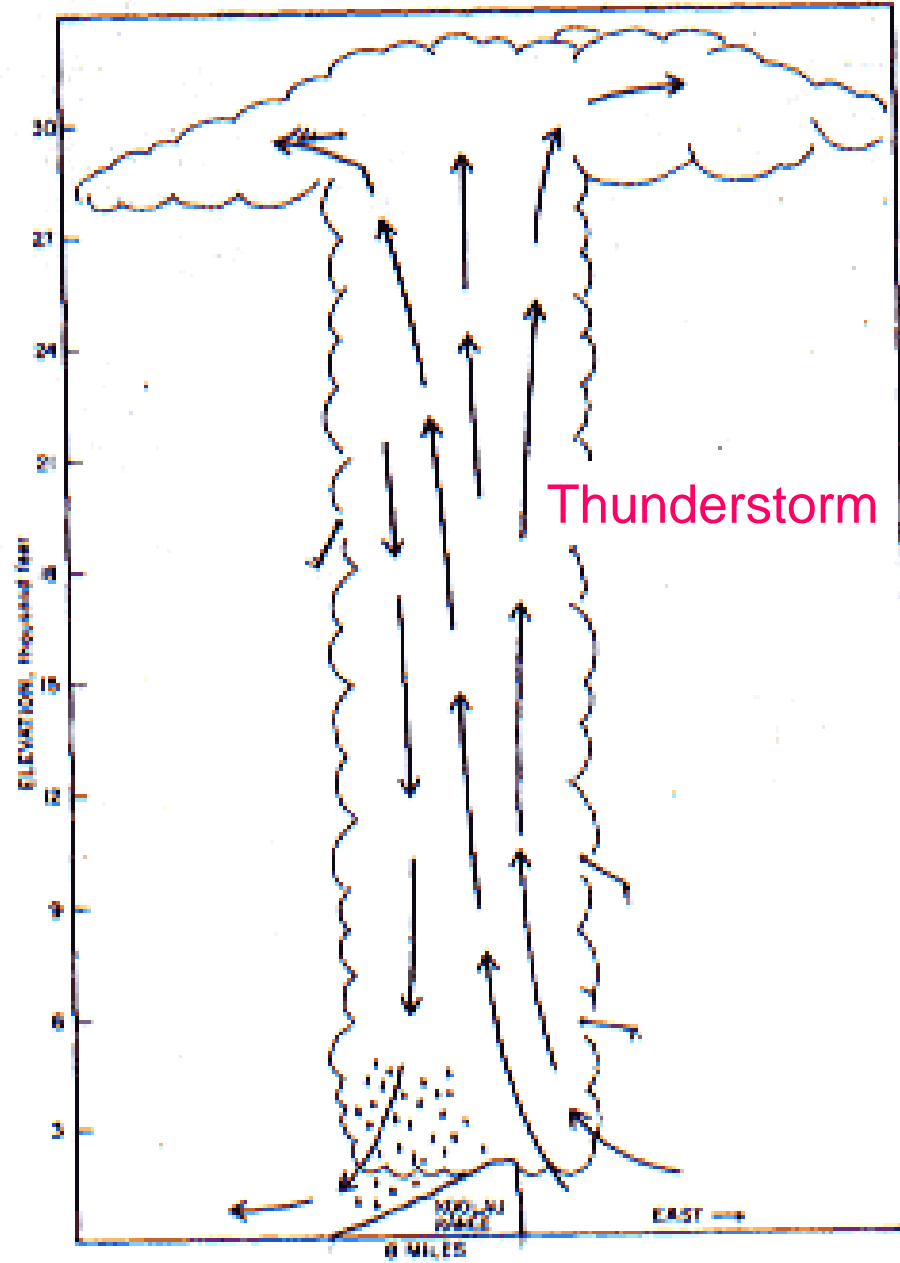


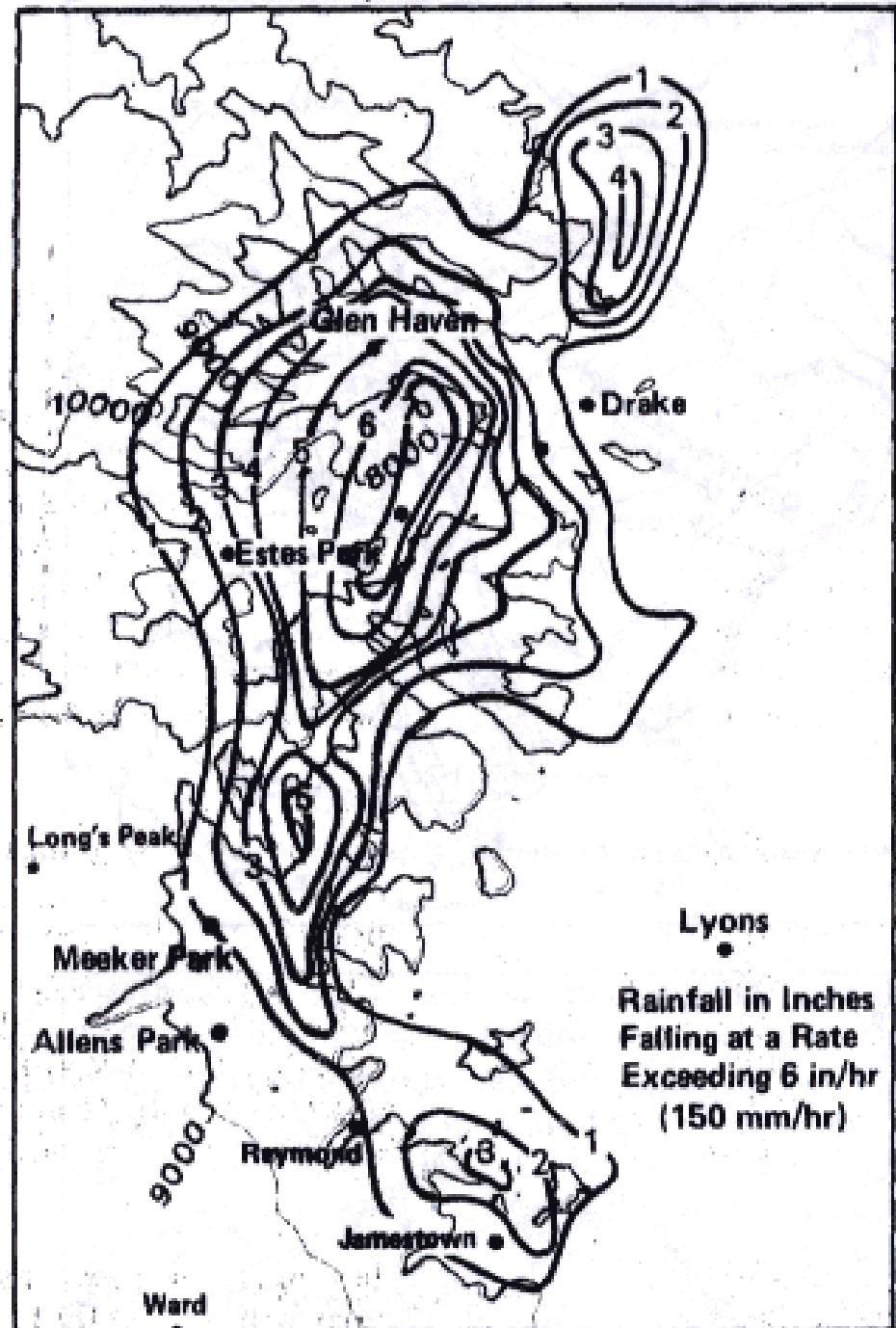
FIG. 7. Analyzed storm isohyets for period 0200 to 1500 hours on 19 April 1974. Isohyets are in millimeters.

$$R_{\max} \approx 70 \text{ mm/h}$$

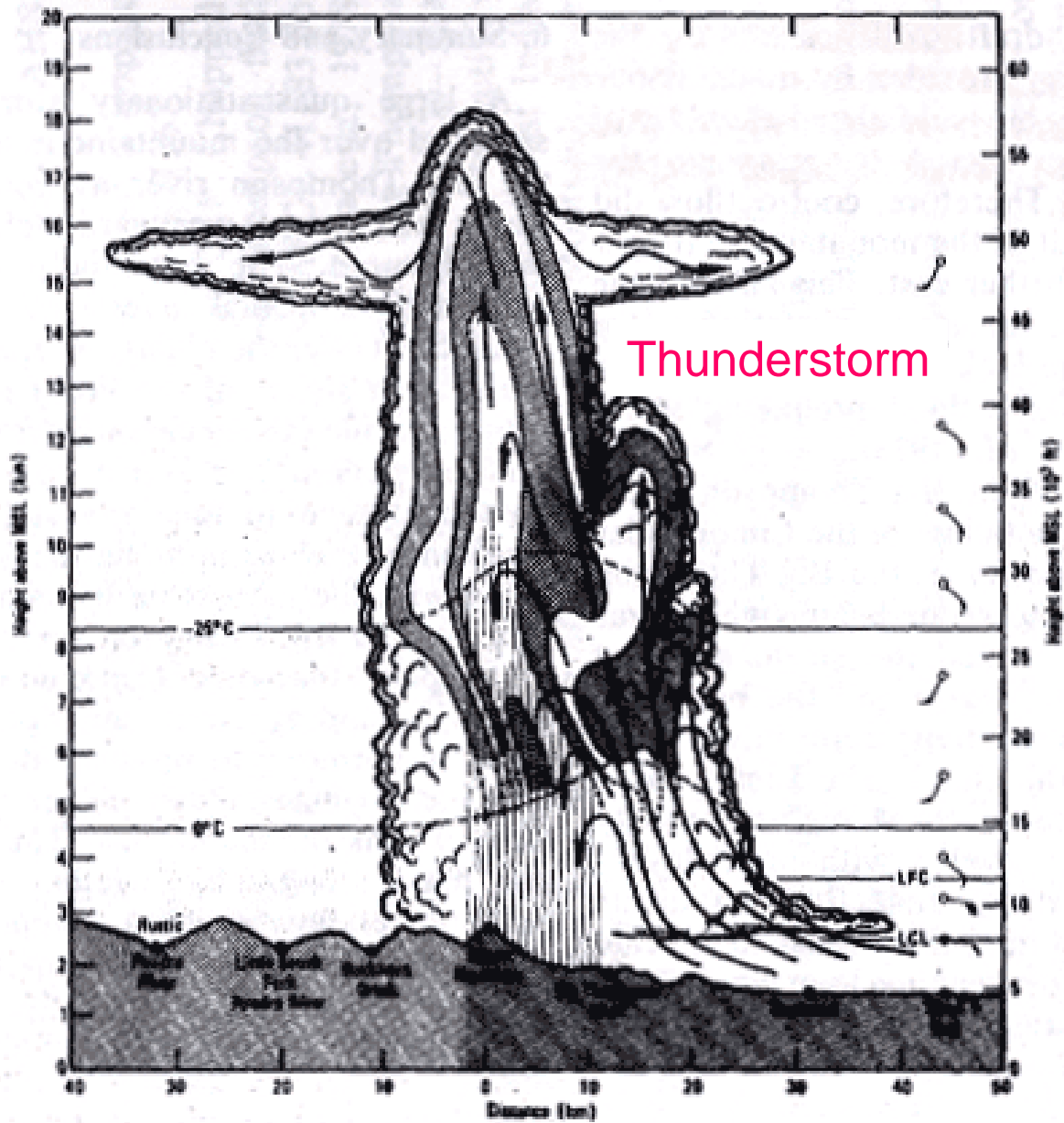


Schroeder (1977)

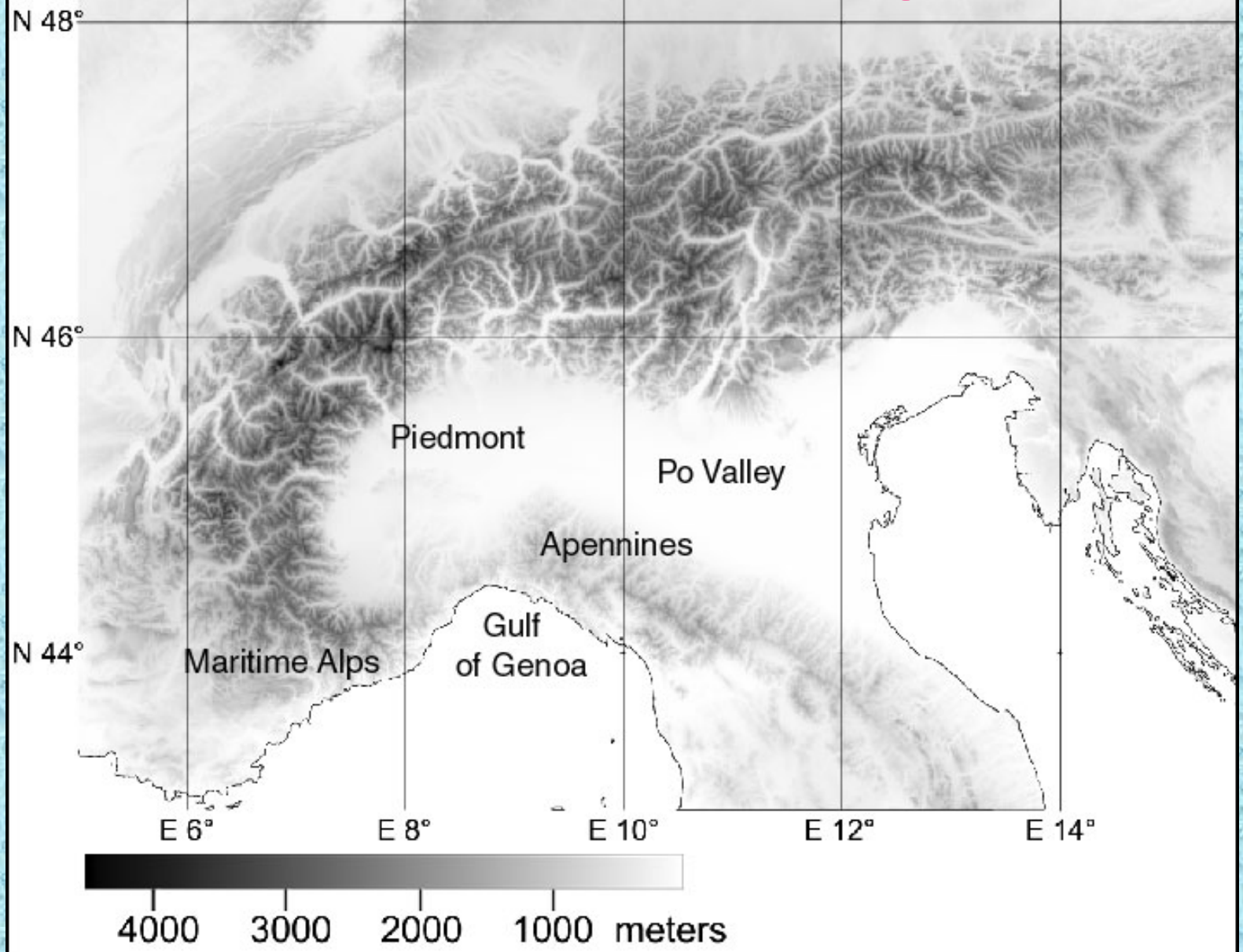
Big Thompson Flood Colorado, 1976



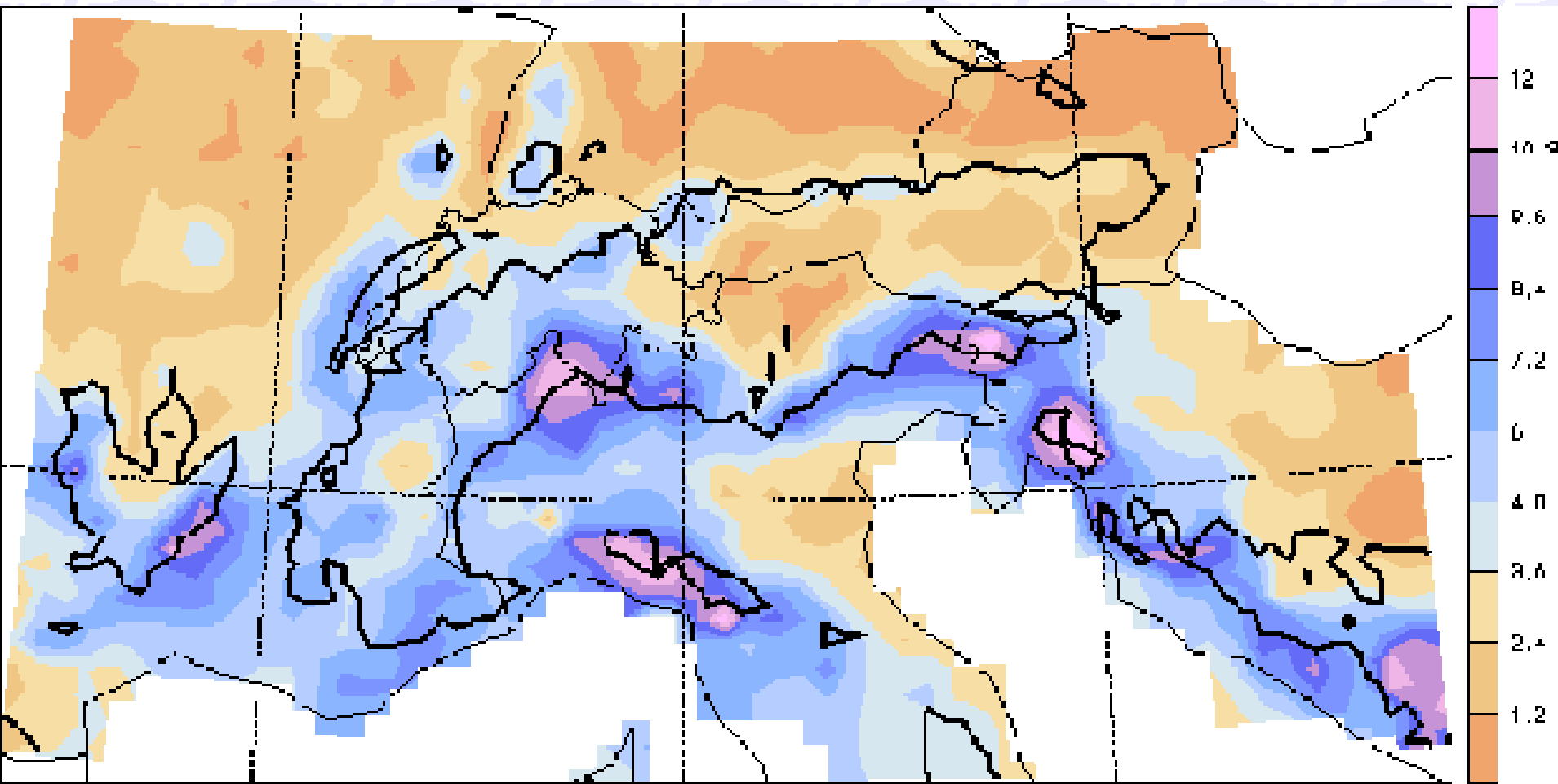
Caracena et al. (1979)



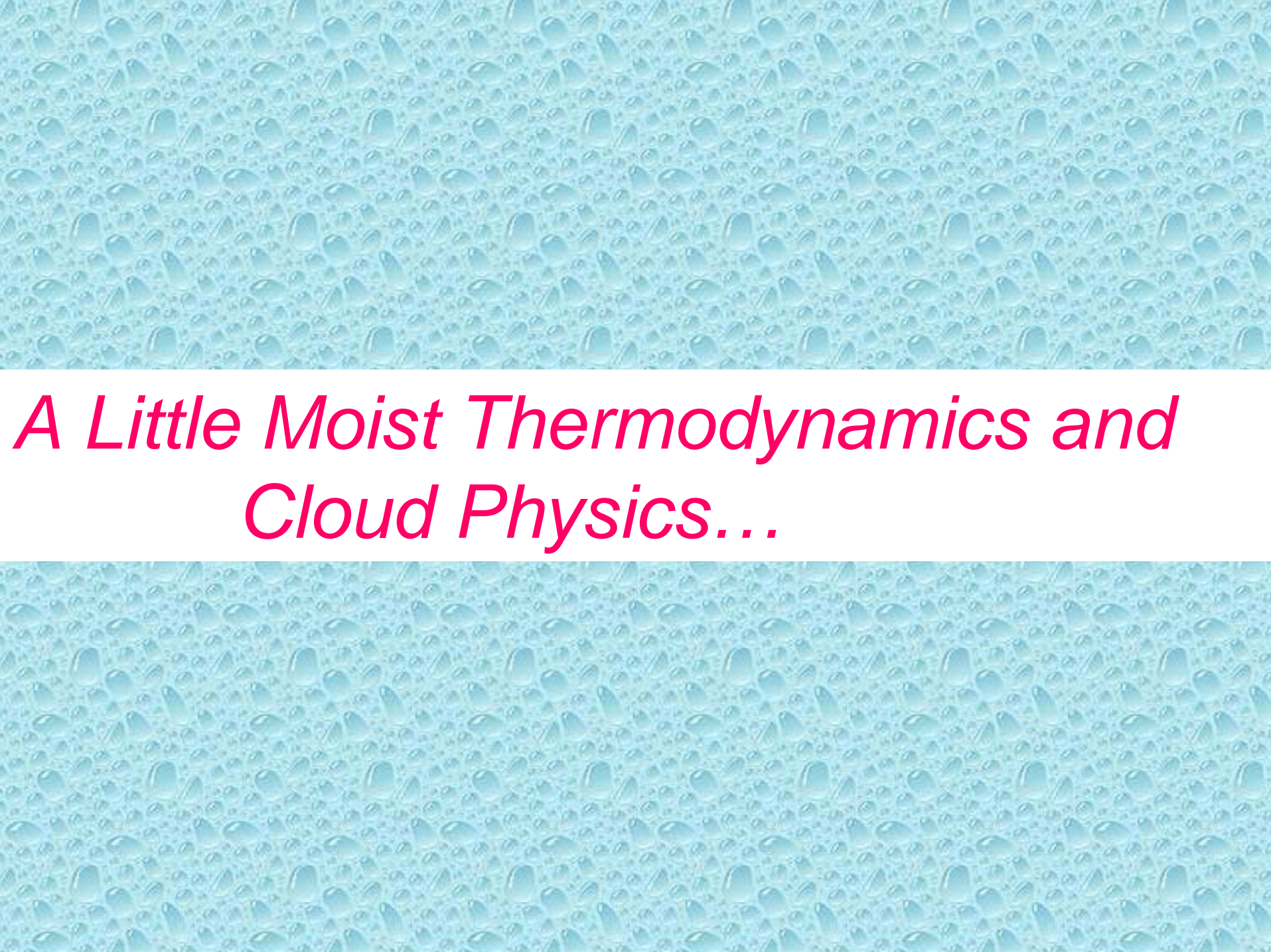
Many-Scaled Topography



Alpine Precipitation Climatology

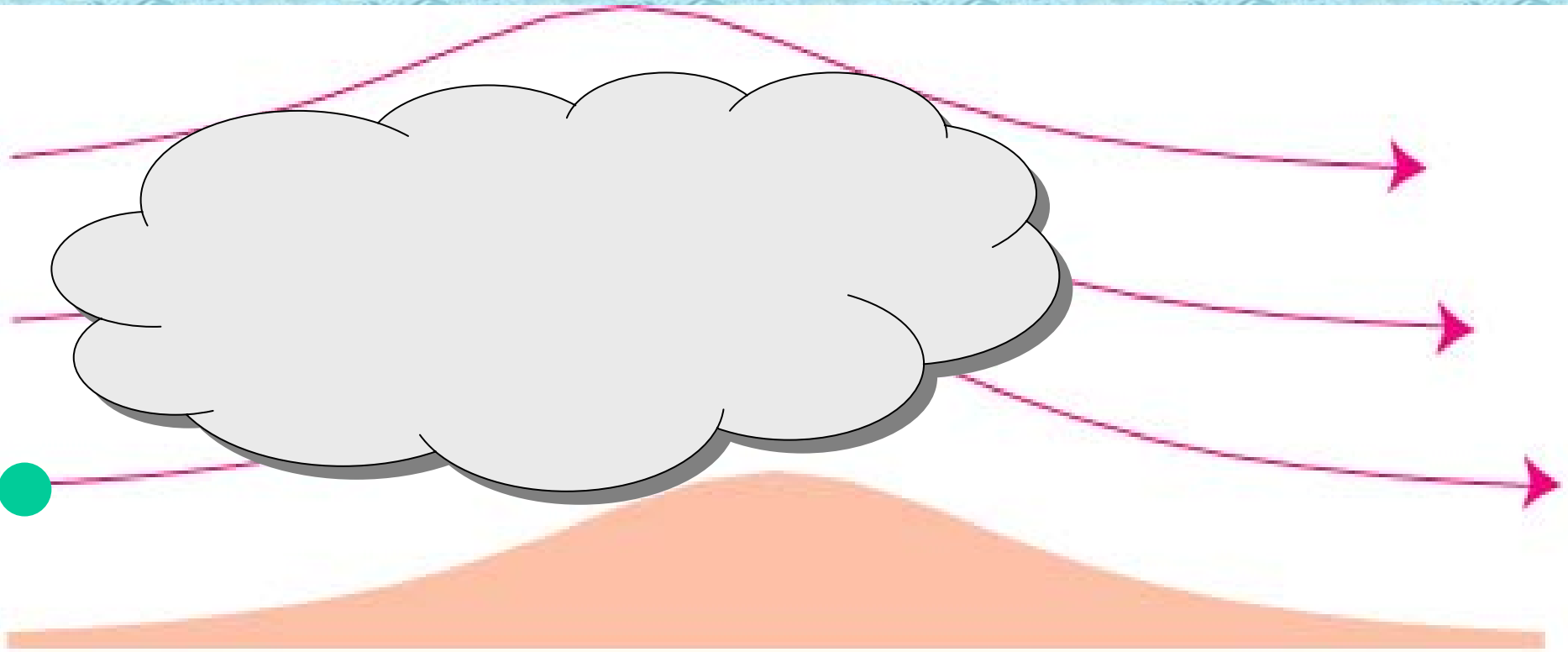


Frei and Schär (1998)



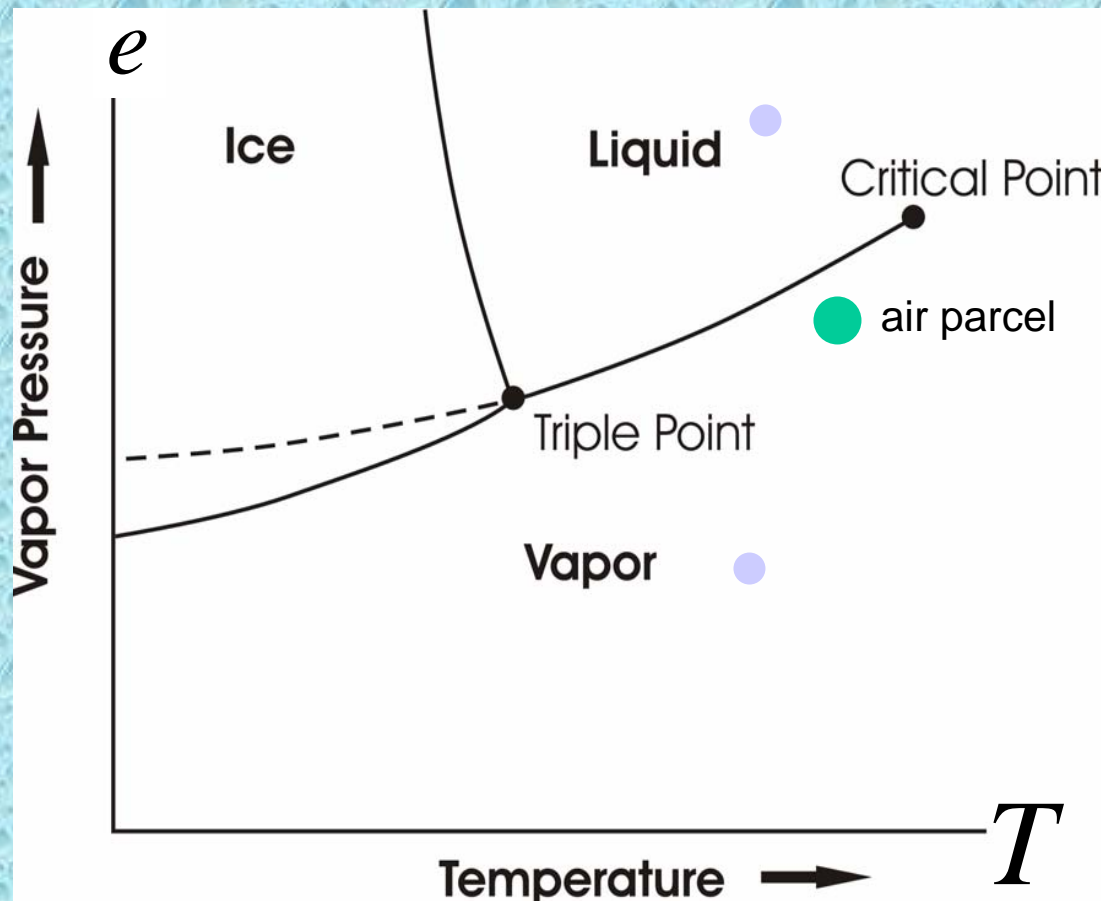
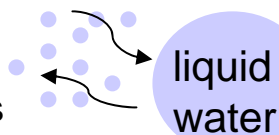
*A Little Moist Thermodynamics and
Cloud Physics...*

Condensation

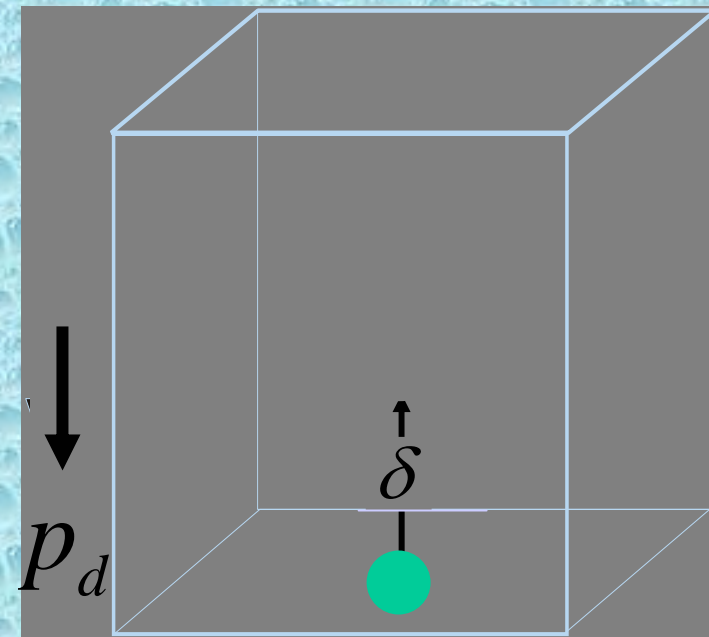


Phase Equilibria

water molecules



Saturation thru Adiabatic Cooling



$$\frac{\Delta T}{T} = \frac{c_p}{R_d} \frac{\Delta p_d}{p_d} < 0$$

gas law \rightarrow

$$\Delta e = \Delta p_d q_v R_v / R_d$$

p_d dry air pressure q_v water vapor mixing ratio

Conventional borderline
between cloud drops and
raindrops

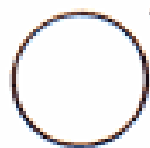
$$r = 100$$
$$V = 70$$

Large cloud drop

$$r = 50^3$$
$$n = 10^3$$
$$V = 27$$

Typical condensation nucleus

$$r = 0.1$$
$$n = 10^6$$
$$V = 0.0001$$



Typical cloud drop

$$r = 10$$
$$n = 10^6$$
$$V = 1$$

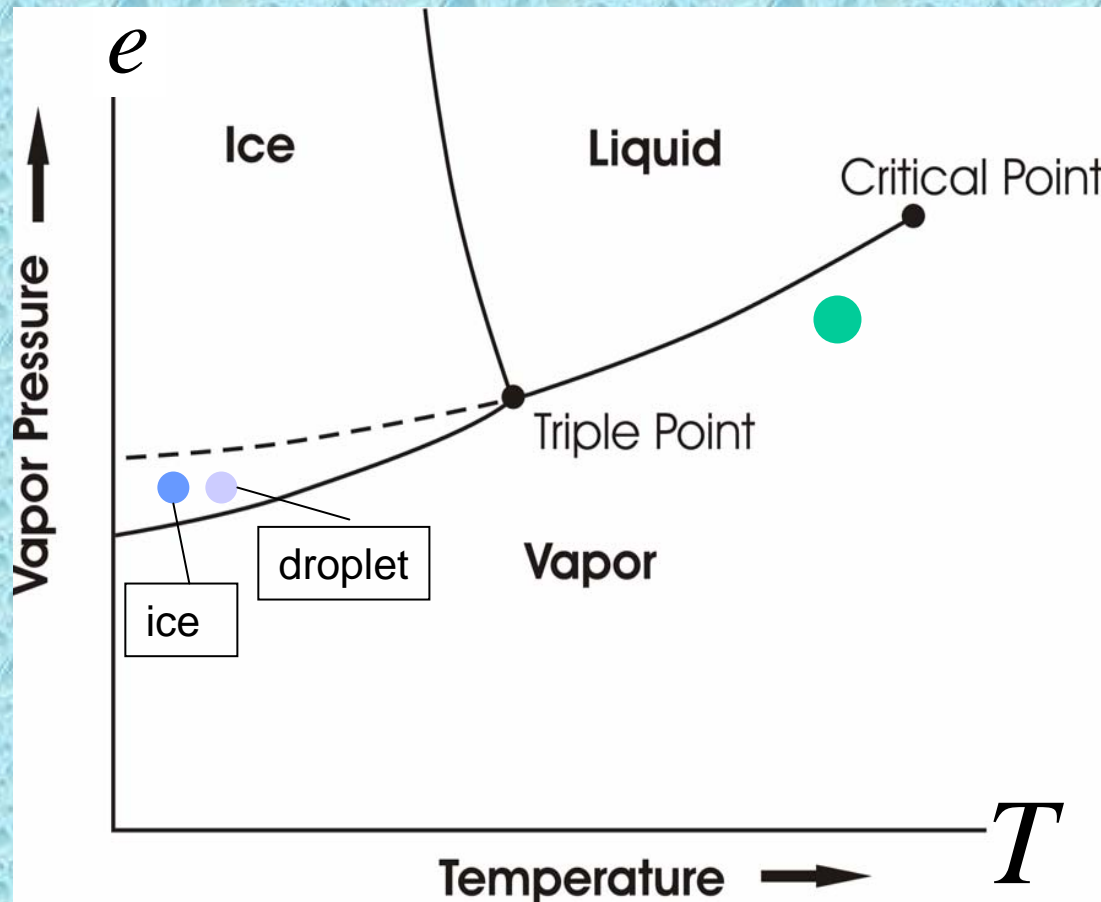
Key:

r - radius in microns
 n - number per liter
 V - terminal velocity
in centimeters
per second

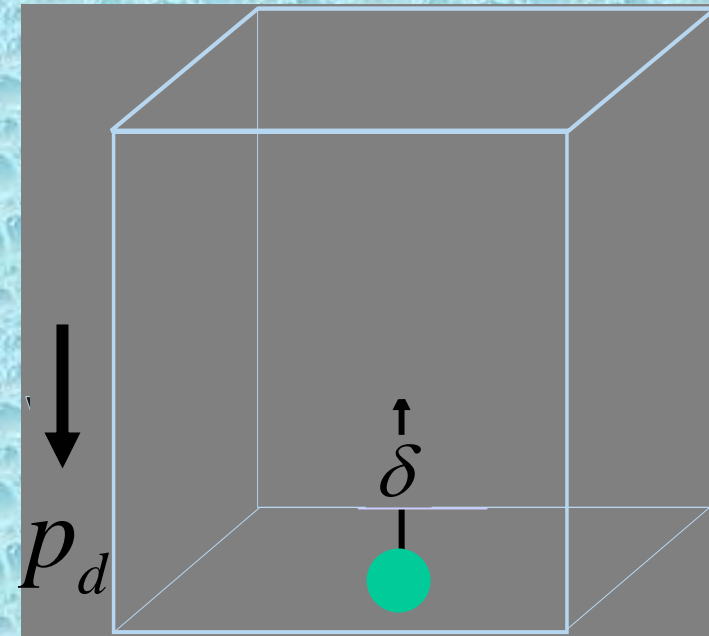
Typical raindrop $r=1000$, $n=1$, $V=650$

McDonald (1958)

Phase Equilibria



Saturation thru
Adiabatic Cooling



$$\frac{\Delta T}{T} = \frac{c_p}{R_d} \frac{\Delta p_d}{p_d} < 0$$

gas law \rightarrow

$$\Delta e = \Delta p_d q_v R_v / R_d$$

p_d dry air pressure q_v water vapor mixing ratio

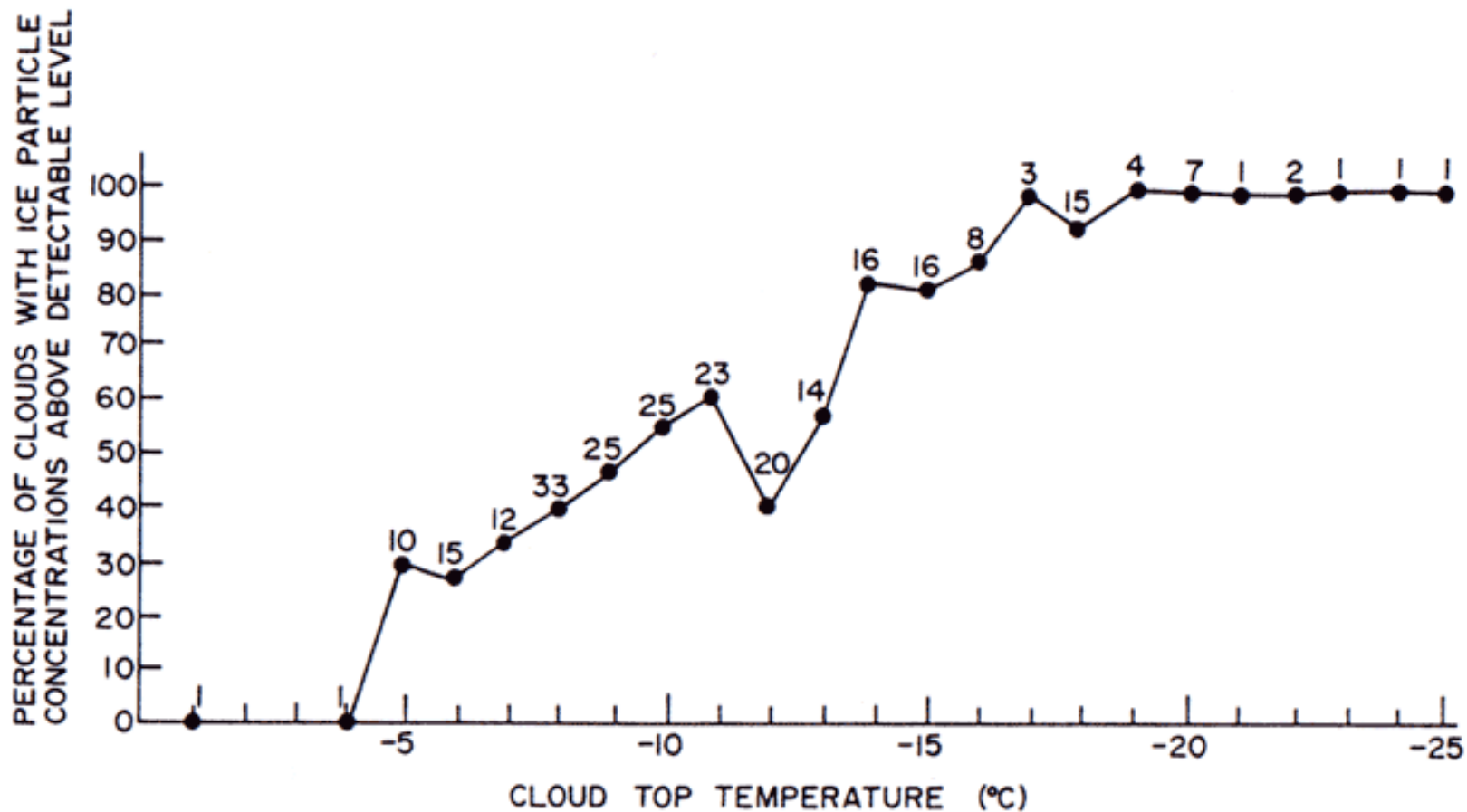
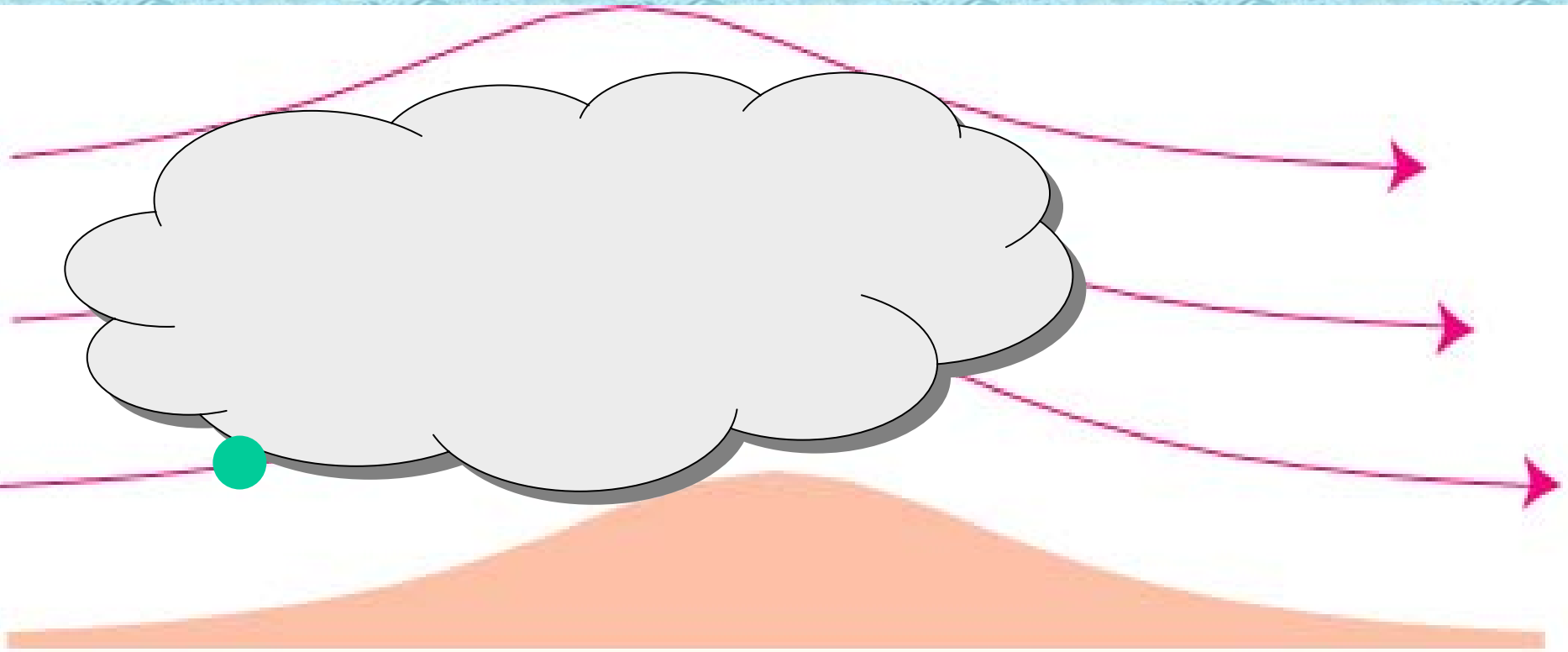


Fig. 4.4 Percentage chance of ice being detected in clouds as a function of the cloud top temperature. Results are based on field observations of 30 orographic cloud systems. The number above each point is the total number of cloud samples for that temperature. [From Proc. Amer. Met. Soc. 1st National Conf. on Weather Modification, Albany, N. Y., 1968, p. 306; Quart. J. Roy. Met. Soc., 96, 487 (1970); Proc. Intern. Conf. on Weather Modification, Canberra, Australia, 1971, p. 5.]

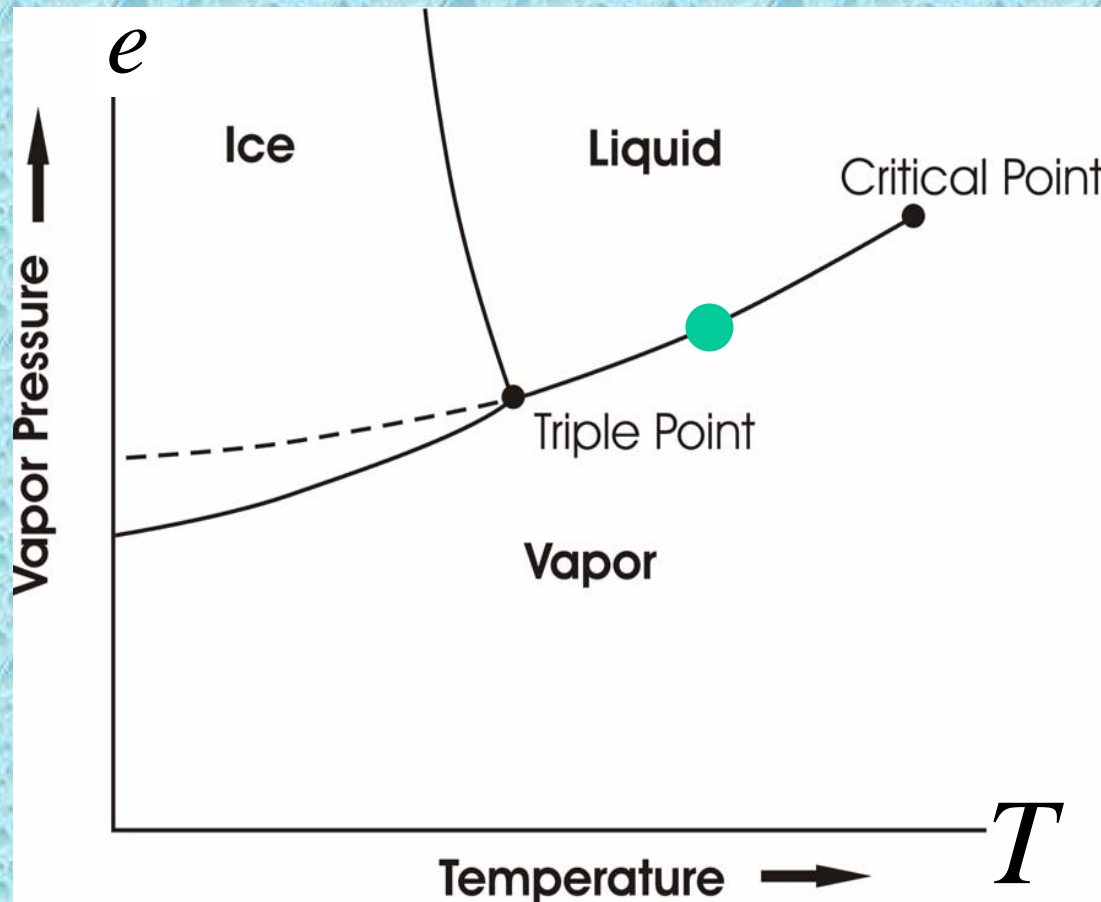
Substance	Crystal lattice dimension		Temperature to nucleate ice (°C)	Comments
	<i>a</i> axis (Å)	<i>c</i> axis (Å)		
Pure substances				
Ice	4.52	7.36	0	—
AgI	4.58	7.49	-4	Insoluble
PbI ₂	4.54	6.86	-6	Slightly soluble
CuS	3.80	16.43	-7	Insoluble
CuO	4.65	5.11	-7	Insoluble
HgI ₂	4.36	12.34	-8	Insoluble
Ag ₂ S	4.20	9.50	-8	Insoluble
CdI ₂	4.24	6.84	-12	Soluble
I ₂	4.78	9.77	-12	Soluble
Minerals				
Vaterite	4.12	8.56	-7	(Silicate)
Kaolinite	5.16	7.38	-9	
Volcanic ash	—	—	-13	
Halloysite	5.16	10.1	-13	
Vermiculite	5.34	28.9	-15	
Cinnabar	4.14	9.49	-16	
Organic materials				
Testosterone	14.73	11.01	-2	(Bacteria in leaf mold)
Chloesterol	14.0	37.8	-2	
Metaldehyde	—	—	-5	
β-Naphthol	8.09	17.8	-8.5	
Phloroglucinol	—	—	-9.4	
Bacterium	—	—	-2.6	
<i>Pseudomonas</i>				
<i>Syringae</i>				

Rogers and Yau (1991)

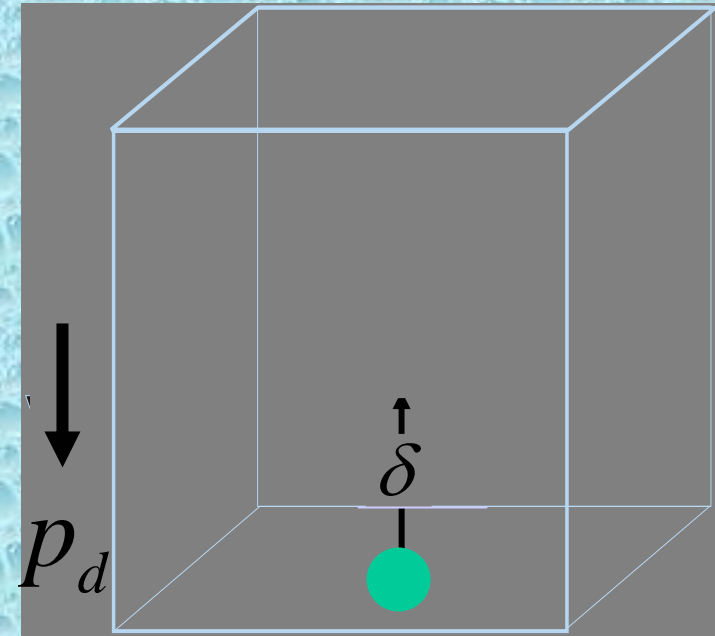
Condensation



Phase Equilibria



Latent Heating Offsets Adiabatic Cooling



$$\frac{\Delta T}{T} = \frac{c_p}{R_d} \frac{\Delta p_d}{p_d} - \frac{L}{T c_p} \Delta q_{vs}$$

gas law \rightarrow

$$q_{vs} = e_s / p_d$$

p_d dry air pressure q_v water vapor mixing ratio

Orographic Precipitation Summary

- Large – Scale (Wind, Humidity, Stability)
- Mesoscale dynamics of orographic air flow
- Microphysics