

A Systematic Multi-Scale Framework for Meteorological Modelling II: *Moist Flows*

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Thanks to ...

Motivation

Moist Aero-Thermodynamics

Meso-convective interactions revisited

Narrow columns

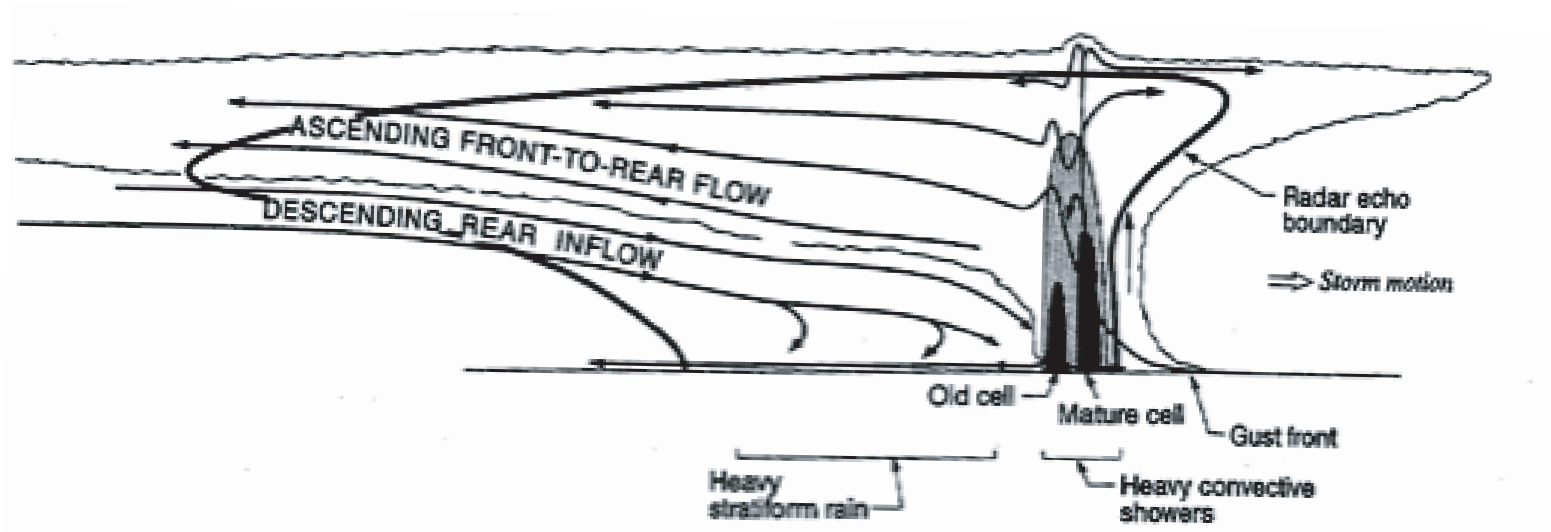
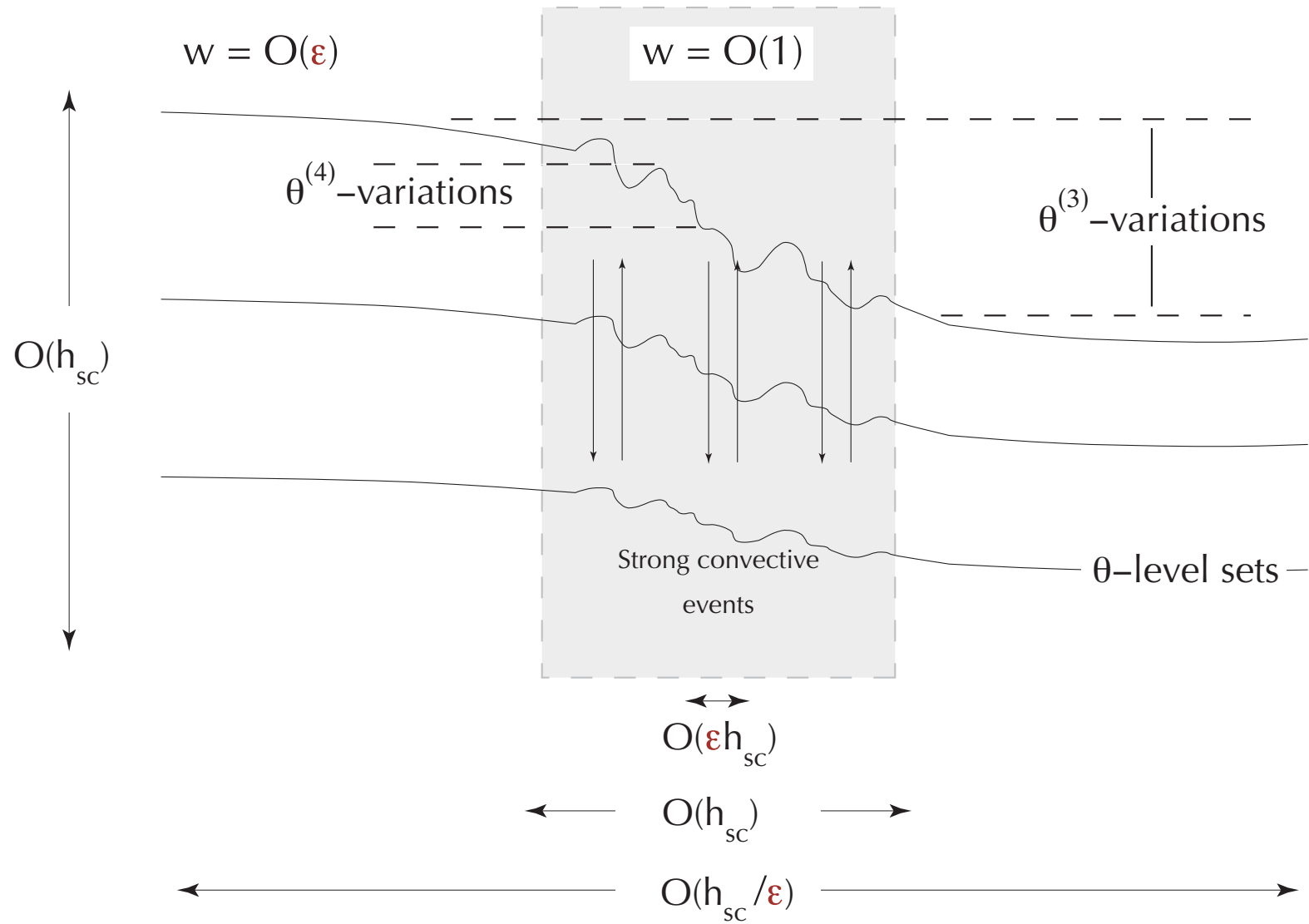


FIG. 1. Conceptual model of a multicell squall line with trailing stratiform precipitation. The storm is viewed in a cross section perpendicular to the convective line (adapted from Houze et al. 1989).

(from: Pandya & Durran (1996))

Motivation and Background



Motivation and Background

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Mass, momentum, energy balances

$$\rho_t + \nabla_{\parallel} \cdot (\rho \mathbf{u}) + (\rho w)_z = 0$$

$$\mathbf{u}_t + \mathbf{u} \cdot \nabla_{\parallel} \mathbf{u} + w \mathbf{u}_z + \frac{1}{\mathbf{Ro}_B} (\boldsymbol{\Omega} \times \mathbf{v})_{\parallel} + \frac{1}{\mathbf{M}^2} \frac{1}{\rho} \nabla_{\parallel} p = D_u$$

$$w_t + \mathbf{u} \cdot \nabla_{\parallel} w + w w_z + \frac{1}{\mathbf{Ro}_B} (\boldsymbol{\Omega} \times \mathbf{v})_{\perp} + \frac{1}{\mathbf{M}^2} \frac{1}{\rho} p_z = D_w - \frac{1}{\overline{\mathbf{Fr}}^2}$$

$$\theta_t + \mathbf{u} \cdot \nabla_{\parallel} \theta + w \theta_z = D_{\theta} + S_{\theta}$$

Moisture balances

$$q_{v,t} + \mathbf{u} \cdot \nabla_{\parallel} q_v + w q_{v,z} = (C_{ev} - C_d) + D_{qv}$$

$$q_{c,t} + \mathbf{u} \cdot \nabla_{\parallel} q_c + w q_{c,z} = (C_d - C_{ac} - C_{cr}) + D_{qc}$$

$$q_{r,t} + \mathbf{u} \cdot \nabla_{\parallel} q_r + w q_{r,z} + \frac{1}{\rho} (\rho q_r V_T)_z = (C_{ac} + C_{cr} - C_{ev}) + D_{qr}.$$

$$\rho \theta = p^{\frac{1}{\gamma}} \quad S_{\theta} = \frac{\gamma - 1}{\gamma} \frac{\theta}{p} L^* (C_d - C_{ev}).$$

$$q_{vs} = \frac{q_{vs}^*}{p} \exp \left(A^* \frac{T - T_0^*}{T_1^* + (T - T_0^*)} \right)$$

$$T(\theta, p) = \theta p^{\frac{\gamma-1}{\gamma}}.$$

$$C_d = C_d^* (q_v - q_{vs}) H(q_c) (q_c + q_{cn}^*)$$

$$C_{ev} = C_{ev}^* \frac{p}{\rho} (q_{vs} - q_v) H(q_r) q_r^{1/2}$$

$$C_{cr} = C_{cr}^* q_c q_r$$

$$C_{ac} = C_{ac}^* \max(0, q_c - q_c^*)$$

Distinguished Limits

$$q_{vs}^* = 0.021 = \epsilon^2 q_{vs}^{**}, \quad A^*, L^* = (16, 30) = \frac{1}{\epsilon} (A^{**}, L^{**}), \quad T_0^* = 1, \quad \dots$$

Multilayer Difficulty

$$q_{vs} = \epsilon^2 \frac{q_{vs}^{**}}{p} \exp \left(\frac{A^{**}}{\epsilon} \frac{T-1}{1 + (T-1) - \epsilon T_1^{**(1)}} \right), \quad T = \theta p^{\frac{\gamma-1}{\gamma}}$$

Newtonian Limit*

$$\frac{\gamma-1}{\gamma} = \epsilon \Gamma \quad \Rightarrow \quad T = \theta \left(1 + \epsilon \Gamma \ln p + O(\epsilon^2) \right)$$

Weak stratification**

$$\theta = 1 + \epsilon^2 \Theta_2(z) + O(\epsilon^3)$$

⇓

$$\underline{q_{vs} = \epsilon^2 q_{vs}^{**} \exp \left(- [A^{**}\Gamma - 1] z \right) (1 + O(\epsilon)) .}$$

* I. Newton (long ago), Crighton (2000), Lipps, Hemler (1982), Bannon (1995)

** Majda, Klein (2003)

$$(q_v, q_c, q_r) = \varepsilon^2 (\hat{q}_v, \hat{q}_c, \hat{q}_r)$$

$$\begin{aligned} \hat{q}_{v,t} + \mathbf{u} \cdot \nabla_{\parallel} \hat{q}_v + w \hat{q}_{v,z} &= -\frac{1}{\varepsilon^n} \hat{C}_d + \hat{C}_{ev} + D_{\hat{q}_v} \\ \hat{q}_{c,t} + \mathbf{u} \cdot \nabla_{\parallel} \hat{q}_c + w \hat{q}_{c,z} &= +\frac{1}{\varepsilon^n} \hat{C}_d - \frac{1}{\varepsilon} \hat{C}_{cr} - \hat{C}_{ac} + D_{\hat{q}_c} \\ \hat{q}_{r,t} + \mathbf{u} \cdot \nabla_{\parallel} \hat{q}_r + w \hat{q}_{r,z} + \frac{1}{\rho} (\hat{V}_T \rho \hat{q}_r)_z &= \frac{1}{\varepsilon} \hat{C}_{cr} - \hat{C}_{ev} + \hat{C}_{ac} + D_{\hat{q}_r} \end{aligned}$$

$$\hat{C}_d = \underline{C_d^{**} (\hat{q}_v - \hat{q}_{vs}) H(\hat{q}_c) (\hat{q}_c + \varepsilon q_{cn}^{**})}$$

$$\hat{C}_{ev} = C_{ev}^{**} (\hat{q}_{vs} - \hat{q}_v) H(\hat{q}_r) \hat{q}_r^{\frac{1}{2}}$$

$$\hat{C}_{cr} = C_{cr}^{**} \hat{q}_c \hat{q}_r$$

$$\hat{C}_{ac} = C_{ac}^{**} \max(0, \hat{q}_c - \varepsilon q_c^{**}) .$$

$$\underline{\hat{S}_\theta = \varepsilon^2 \Gamma L^{**} \frac{\theta}{p} \left[\frac{1}{\varepsilon^n} \hat{C}_d - \hat{C}_{ev} \right]}$$

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$$\mathbf{U}(\mathbf{x}, z, t; \epsilon) = \sum_i \epsilon^i \mathbf{U}^{(i)}(\mathbf{x}, \boldsymbol{\xi}, z, t),$$

$$\boldsymbol{\xi} = \epsilon \mathbf{x}$$

$$\mathbf{x} = \frac{\mathbf{x}'}{10 \text{ km}}$$

Mesoscale–Subsynoptic interactions

Mesoscale dynamics

$$\bar{\mathbf{u}}_t + \nabla_{\xi} \pi' = -\partial_z \left(\frac{\overline{S''_{\theta} \mathbf{u}}}{d\Theta_2/dz} \right),$$

$$\theta'_t + \overline{w'} \frac{d\Theta_2}{dz} = \overline{S'_{\theta}},$$

$$\partial_z \pi' = \theta',$$

$$\rho_0 \nabla_{\xi} \cdot \bar{\mathbf{u}} + \partial_z (\rho_0 \overline{w'}) = 0.$$

Convective scale dynamics

Anelastic* for near-moist adiabatic stratification

WTG otherwise

$\Rightarrow S''_{\theta}$

* essentially

Mesoscale–Convective interactions

Sublinear growth conditions

Mass:

$$\nabla_x \cdot (\rho \mathbf{u})^{(0)} + (\rho w)_z^{(0)} = 0$$

$$\nabla_x \cdot (\rho \mathbf{u})^{(1)} + (\rho w)_z^{(1)} = -\nabla_\xi \cdot (\rho \mathbf{u})^{(0)}$$

Horizontal momentum:

$$\nabla_x p^{(0)} = 0$$

$$\nabla_x p^{(j+1)} + \nabla_\xi p^{(j)} = 0 \quad (j \in \{0, 1, 2\})$$

$$(\rho \mathbf{u})_t^{(0)} + \nabla_x \cdot (\rho \mathbf{u} \circ \mathbf{u})^{(0)} + (\rho w \mathbf{u})_z^{(0)} + \nabla_x p^{(4)} + \nabla_\xi p^{(3)} = \mathbf{D}_{\rho \mathbf{u}}^{(0)}.$$

Vertical momentum:

$$\partial_z p^{(j)} = -\rho^{(j)} \quad (j \in \{0, 1, 2, 3\})$$

Sublinear growth conditions

Potential temperature:

$$\theta = 1 + \epsilon^2 \Theta_2(z) + \epsilon^3 \underline{\Theta^{(3)}(\boldsymbol{\xi}, z, t)} + \epsilon^3 \theta^{(4)}(\boldsymbol{x}, \boldsymbol{\xi}, z, t) \dots$$

$$S_\theta^{(0)} = 0,$$

$$S_\theta^{(1)} = 0,$$

$$w^{(0)} \frac{d\Theta_2}{dz} = S_\theta^{(2)},$$

$$\theta_t^{(3)} + w^{(0)} \theta_z^{(3)} + w^{(1)} \frac{d\Theta_2}{dz} = S_\theta^{(3)},$$

Meso-convective interactions

Mesoscale wave dynamics

$$\bar{\mathbf{u}}_t + \nabla_{\boldsymbol{\xi}} \pi^{(3)} = -\partial_z (\overline{w\mathbf{u}}) ,$$

$$\Theta_t^{(3)} + \overline{w^{(1)}} \frac{d\Theta_2}{dz} = \overline{S_{\theta}^{(3)}} ,$$

$$\partial_z \pi^{(3)} = \Theta^{(3)} ,$$

$$\rho_0 \nabla_{\boldsymbol{\xi}} \cdot \bar{\mathbf{u}} + \partial_z \left(\rho_0 \overline{w^{(1)}} \right) = 0 .$$

where

$$\pi^{(3)} = \frac{p^{(3)}}{\rho_0(z)}$$

Moisture transport

Water vapor:

$$C_d^{(i)} = 0 \quad (i = -(n+2), \dots, 1)$$

$$q_v^{(2)} + \mathbf{u}^{(0)} \cdot \nabla_{\mathbf{x}} q_v^{(2)} + w^{(0)} q_v^{(2)} = -C_d^{(2)} + C_{ev}^{(2)}.$$

$$q_v^{(3)} + \mathbf{u}^{(0)} \cdot \nabla_{\mathbf{x}} q_v^{(3)} + \mathbf{u}^{(1)} \cdot \nabla_{\mathbf{x}} q_v^{(2)} + w^{(0)} q_v^{(3)} + w^{(1)} q_v^{(2)} = -C_d^{(3)} + C_{ev}^{(3)}.$$

Cloud water:

$$C_{cr}^{(1)} = C_{cr}^{**} q_c^{(2)} q_r^{(2)} = 0,$$

$$q_c^{(2)} + \mathbf{u}^{(0)} \cdot \nabla_{\mathbf{x}} q_c^{(2)} + w^{(0)} q_c^{(2)} = C_d^{(2)} - C_{cr}^{(2)}.$$

Rain water:

$$q_r^{(2)} + \mathbf{u}^{(0)} \cdot \nabla_{\mathbf{x}} q_r^{(2)} + w^{(0)} q_r^{(2)} + \frac{1}{\rho_0} \left(\rho_0 V_T q_r^{(2)} \right)_z = C_{cr}^{(2)} - C_{ev}^{(2)}.$$

Nonlinear control of convective motions

Recall that

$$w^{(0)} \frac{d\Theta_2}{dz} = S_\theta^{(2)}$$

$$\theta_t^{(3)} + w^{(0)} \theta_z^{(3)} + w^{(1)} \frac{d\Theta_2}{dz} = S_\theta^{(3)}$$

Source terms

$$S_\theta^{(2)} = \frac{\Gamma L^{**}}{p_0(z)} \left(-H_{\geq}(q_v - q_{vs}) w^{(0)} \frac{dq_{vs}^{(2)}}{dz} + H_{>}(q_{vs} - q_v) \mathbf{C}_{ev}^{(2)} \right)$$

$$S_\theta^{(3)} = \frac{\Gamma L^{**}}{p_0(z)} \left(-H_{\geq}(q_v - q_{vs}) \left(w^{(0)} \frac{dq_{vs}^{(3)}}{dz} + w^{(1)} \frac{dq_{vs}^{(2)}}{dz} \right) + H_{>}(q_{vs} - q_v) \mathbf{C}_{ev}^{(3)} \right) + \hat{S}_\theta^{(3)}$$

Nonlinear control of convective motions / **undersaturated**

Strongly undersaturated Air ($q_{\text{vs}}^{(2)} - q_{\text{v}}^{(2)} > 0$)

$$w^{(0)} = -\frac{\Gamma L^{**} C_{\text{ev}}^{**}}{(p_0 d\Theta_2/dz)(z)} \left(q_{\text{vs}}^{(2)}(z) - q_{\text{v}}^{(2)} \right) q_{\text{r}}^{(2)\frac{1}{2}} \quad \text{for} \quad q_{\text{v}}^{(2)} < q_{\text{vs}}^{(2)}(z).$$

Weakly undersaturated Air ($q_{\text{vs}}^{(2)}(z) - q_{\text{v}}^{(2)} \equiv 0$ **but** $q_{\text{vs}}^{(3)}(z) - q_{\text{v}}^{(3)} > 0$)

$$w^{(0)} = 0$$

$$w^{(1)} = \left(-\frac{\Gamma L^{**} C_{\text{ev}}^{**}}{p_0} \left(q_{\text{vs}}^{(3)}(z) - q_{\text{v}}^{(3)} \right) q_{\text{r}}^{(2)\frac{1}{2}} + \left(\hat{S}_{\theta}^{(3)} - \theta_t^{(3)} \right) \right) \left(\frac{d\Theta_2}{dz} \right)^{-1}.$$

Nonlinear control of convective motions / saturated

$$w^{(0)} \frac{d\Theta_2}{dz} = -w^{(0)} \frac{\Gamma L^{**}}{p_0(z)} \frac{dq_{vs}^{(2)}}{dz}$$

$$\underline{\frac{d\Theta_2}{dz} = -\frac{\Gamma L^{**}}{p_0} \frac{dq_{vs}^{(2)}}{dz}} \quad \text{or} \quad w^{(0)} \equiv 0$$

Precipitating clouds: rapid cloudwater collection, ($C_{cr}^{(1)} = C_{cr}^{} \underline{q_c^{(2)}} q_r^{(2)} = 0$)**

$$C_d^{(2)} - C_{cr}^{(2)} = -w^{(0)} \frac{dq_{vs}^{(2)}}{dz} - C_{cr}^{**} q_c^{(3)} q_r^{(2)1/2} = 0 \quad \text{or} \quad w^{(0)} = -\frac{C_{cr}^{**} q_c^{(3)} q_r^{(2)1/2}}{dq_{vs}^{(2)}/dz}.$$

$$\underline{w^{(0)} \geq 0}$$

So, how fast do we go upward, then??

Source terms in third order θ -equation

$$S_{\theta}^{(3)} = \frac{\Gamma L^{**}}{p_0} \mathbf{C}_d^{(3)} + \hat{S}_{\theta}^{(3)}$$

$$\mathbf{C}_d^{(3)} = -w^{(1)} \frac{dq_{vs}^{(2)}}{dz} - w^{(0)} \frac{dq_{vs}^{(3)}}{dz}$$

Third order θ -equation

$$\theta_t^{(3)} + w^{(0)} \theta_z^{(3)} + \underline{w^{(1)} \frac{d\Theta_2}{dz}} = -\frac{\Gamma L^{**}}{p_0} \left(\underline{w^{(1)} \frac{dq_{vs}^{(2)}}{dz}} + w^{(0)} \frac{dq_{vs}^{(3)}}{dz} \right) + \hat{S}_{\theta}^{(3)}$$

$$\Theta_t^{(3)} + w^{(0)} \Theta_z^{(3)} = -\frac{\Gamma L^{**}}{p_0} w^{(0)} \frac{dq_{vs}^{(3)}}{dz} + \hat{S}_{\theta}^{(3)}$$

Coupling with mesoscale waves

$$\underline{w_*^{(0)} = \left(\hat{S}_{\theta}^{(3)} - \Theta^{(3)}_t \right) \left(\Theta^{(3)}_z + \frac{\Gamma L^{**}}{p_0} \frac{dq_{vs}^{(3)}}{dz} \right)^{-1}} \quad w^{(0)} = H \left(w_*^{(0)} \right)$$

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$$\mathbf{U}(\mathbf{x}, z, t; \epsilon) = \sum_i \epsilon^i \mathbf{U}^{(i)}(\eta, \mathbf{x}, z, \tau)$$

$$\eta = \mathbf{x}/\epsilon$$

$$\tau = t/\epsilon$$

$$\mathbf{x} = \frac{\mathbf{x}'}{10 \text{ km}}, \quad t = \frac{t'}{20 \text{ min}}$$

Onset of Deep Convection

Mesoscale Dynamics

$$\mathbf{u}_\tau + \nabla_{\mathbf{x}} \pi' = 0$$

$$\bar{w}_\tau + \pi'_z = \bar{\theta}'$$

$$\bar{\theta}'_\tau + \bar{w} \frac{d\Theta_2}{dz} = \frac{\Gamma L^{**}}{p_0} \bar{C}''$$

$$\rho_0 \nabla_{\mathbf{x}} \cdot \mathbf{u} + (\rho_0 \bar{w})_z = 0$$

Bulk Microscale Column Dynamics

$$\left(\partial_\tau + \mathbf{u}^{(0)} \cdot \nabla_\eta \right) \tilde{w} = \tilde{\theta}'$$

$$\tilde{\theta}'_\tau + \tilde{w} \frac{d\Theta_2}{dz} = \frac{\Gamma L^{**}}{p_0} \tilde{C}''.$$

Coupling to Moisture Physics

$$C'' = H(q'_c) C''_d + [1 - H(q'_c)] C''_{\text{ev}}$$

Onset of Deep Convection

Saturated Air

$$C_d''' = C_d^{**} \underline{\delta q_v^{(n^*)}} q_c' = \underline{- \left[\left(\widetilde{w^{(0)}} + \overline{w^{(0)}} \right) \frac{dq_{vs}''}{dz} - D_{qv}'' \right]}$$

$$\left(\partial_\tau + \mathbf{u}^{(0)} \cdot \nabla_\eta \right) q_c' = H(q_c') C_d'' - C_{cr}^{**} q_r'' q_c'$$

$$\left(\partial_\tau + \mathbf{u}^{(0)} \cdot \nabla_\eta \right) q_r'' = 0$$

Undersaturated Air

$$C_{ev}'' = -C_{ev}^{**} (q_{vs}''(z) - q_v'') q_r''^{1/2}$$

$$\left(\partial_\tau + \mathbf{u}^{(0)} \cdot \nabla_\eta \right) q_v'' = 0$$

$$\left(\partial_\tau + \mathbf{u}^{(0)} \cdot \nabla_\eta \right) q_r'' = 0$$

Conclusions

