



Statistical Fluctuations in Convective Forcing Computed from Big-Domain Cloud-Resolving Model Simulations

Glenn Shutts

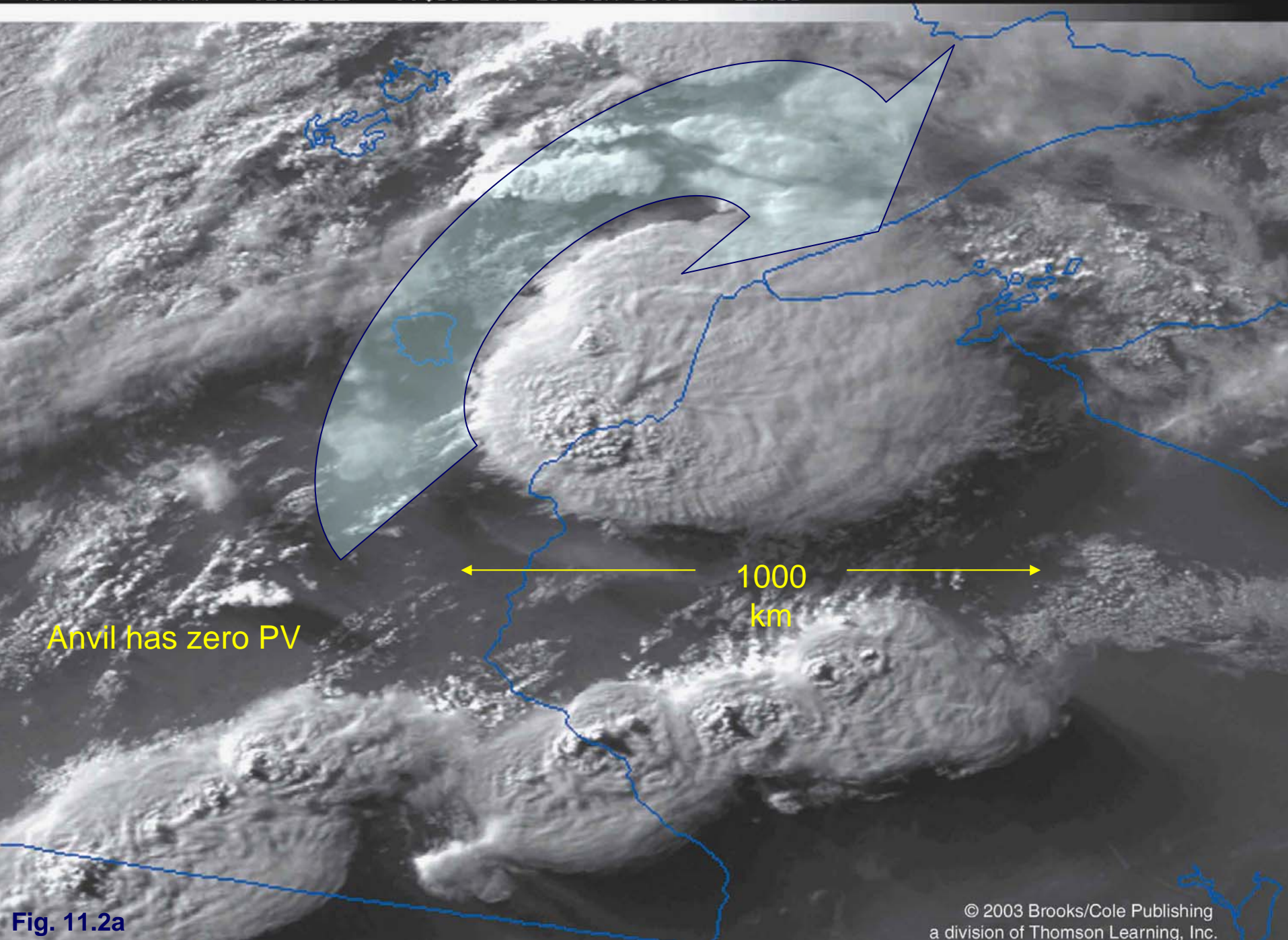
November 2 2005

- Peter Bechtold, Tim Palmer
- Rachael Stratton, Tom Allen
- Roy Kershaw, Leon Hermanson

Why ?

- ‘Stochastic physics’ – lack of EPS spread associated with missing statistical fluctuations from parametrized sub-filter scale phenomena
- under forecasting of low frequency variability in the Tropics e.g. Madden-Julian oscillation
- realism of convection parametrization
- Lin and Neelin (2002) use a stochastic variant of a convective parametrization and find :
“substantial impact on atmospheric variability in the tropics, including intraseasonal and longer time scales”

- Convection is a 'multi-scale phenomenon' (updraughts ~ 100 m to stratiform anvils of 100's kms)
- Horizontal autocorrelation length scale ~ 130 km; time scale ~ 5.5 hours (Ricciardulli and Sardeshmukh, 2002)
- Global NWP models contain < 1 Cb per gridbox.
- This is no **ensemble of deep clouds** within each model grid column
- Parametrization assumes deterministic relationship between ensemble-average effect of convection and instantaneous vertical profile



Anvil has zero PV

1000
km

Fig. 11.2a

- Non-interacting ‘point clouds’
- Individual cloud mass fluxes (m) follow Boltzmann-like exponential pdf with mean $\langle m \rangle$
- total mass flux (M) of a cloud ensemble the following pdf :

$$P(N_*) = \sqrt{\frac{\langle N \rangle}{N_*}} e^{-\langle N \rangle + N_*} \cdot \underbrace{I_1}_{\text{Modified Bessel function}} \left[2\sqrt{\langle N \rangle N_*} \right] + \delta(N_*) e^{-\langle N \rangle}$$

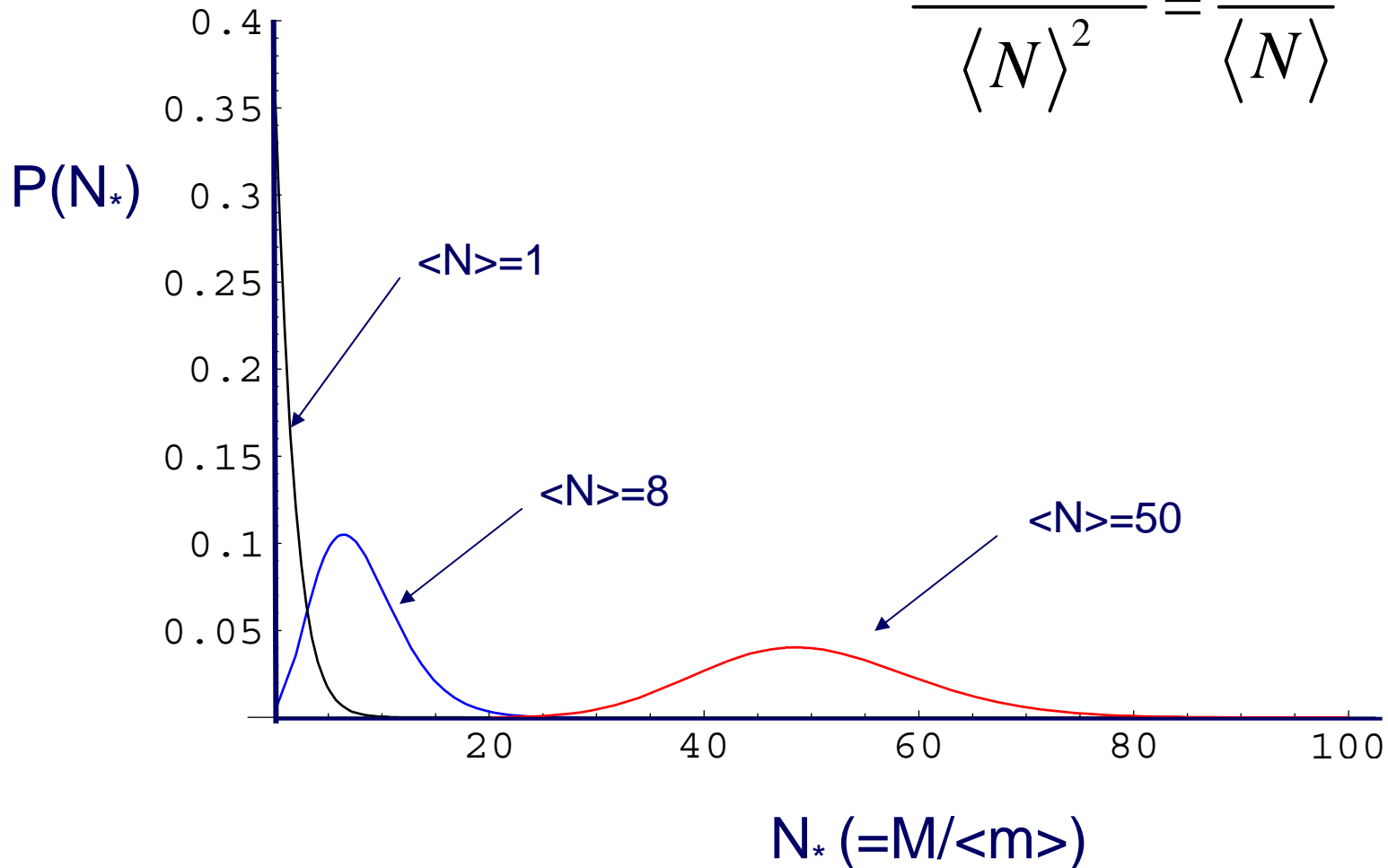
Modified Bessel
function

where $N_* = M / \langle m \rangle$ and $\langle N \rangle$ is the ensemble-mean number of clouds per finite area

Pdf of normalized convective mass flux $P(N_*)$



$$\frac{Var(N_*)}{\langle N \rangle^2} = \frac{2}{\langle N \rangle}$$



$$P(N_*) = \langle N \rangle e^{-\langle N \rangle + N_*} + \underbrace{\delta(N_*) e^{-\langle N \rangle}}_{\substack{\text{Probability of no} \\ \text{clouds} \\ e^{-\langle N \rangle}}}$$

Exponential mass flux PDF

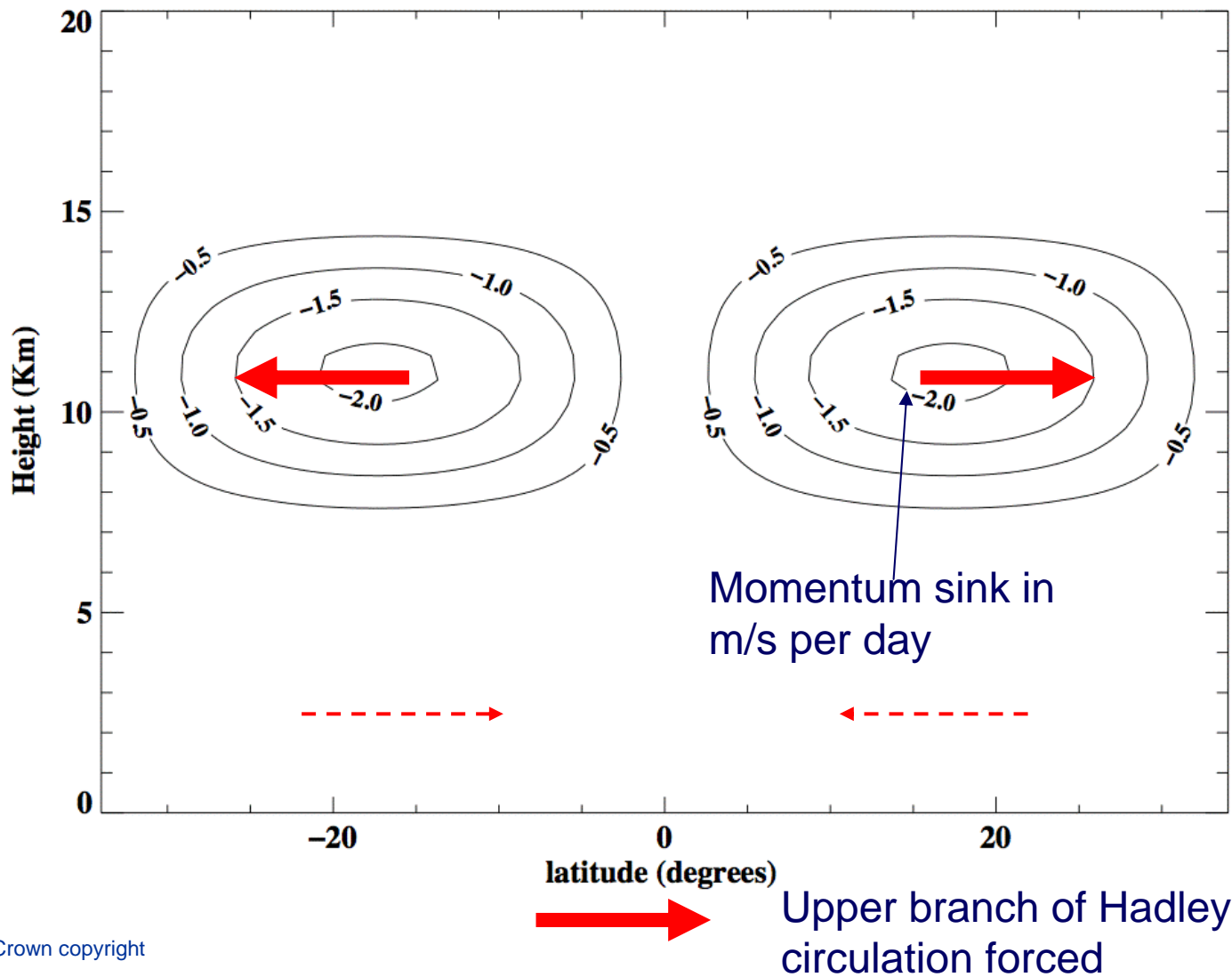
Suggests that the convective precipitation rate should have exponential PDF

Cloud-resolving model to calibrate model error due to convection

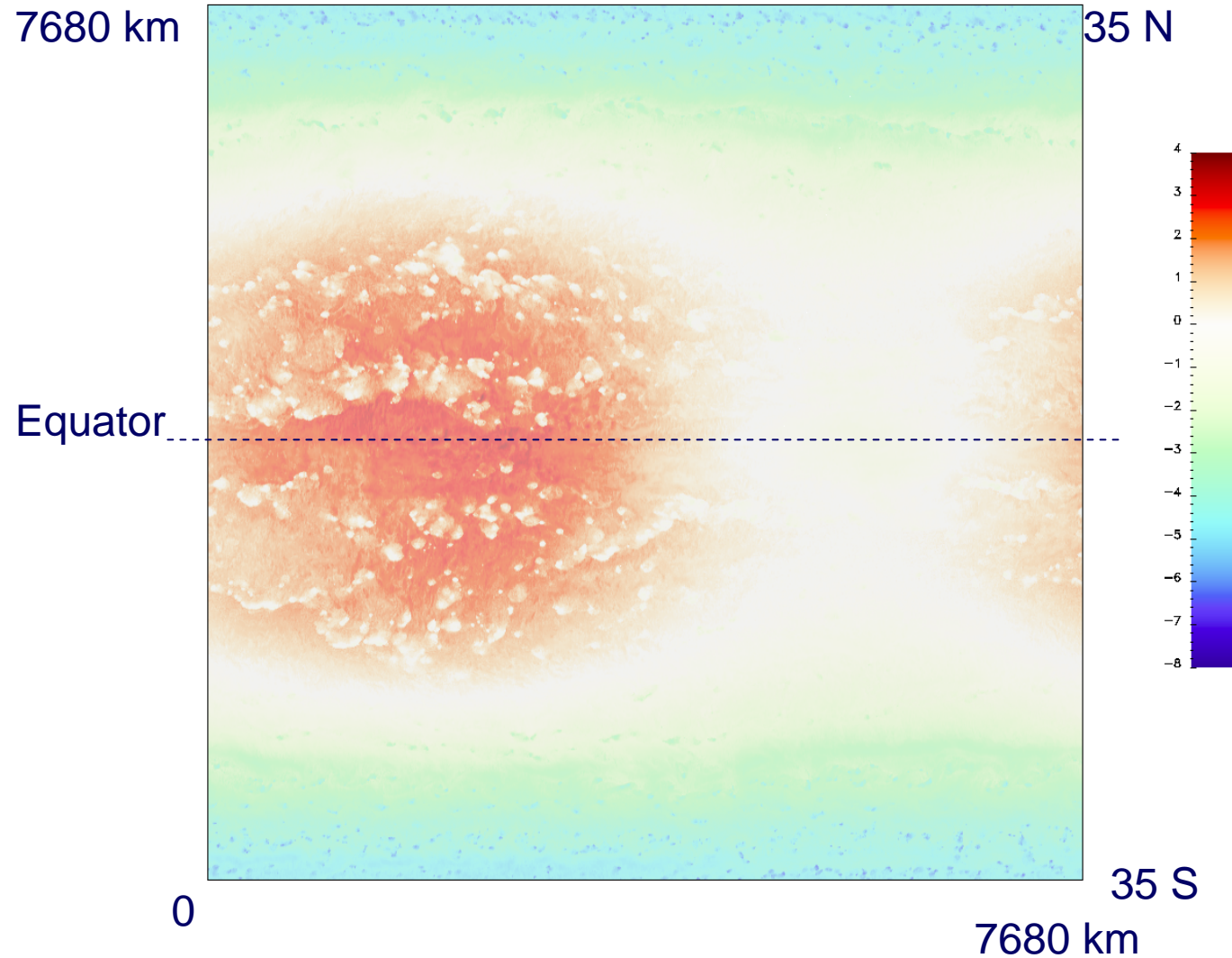


- Met Office LEM configured as an equatorial beta-plane with SST variation, imposed cooling and 'Trade Wind forcing function'
- Domain 7680 x 7680 x 30 km. 50 levels dx= 2km
dy=10 km. Latitude range +/- 35 degs
- Simulations produce cloud 'super-clusters', squall lines, double ITCZ and tropical cyclones
- **Compute convective temperature and vorticity forcing seen by an NWP model grid and characterize its statistical properties**

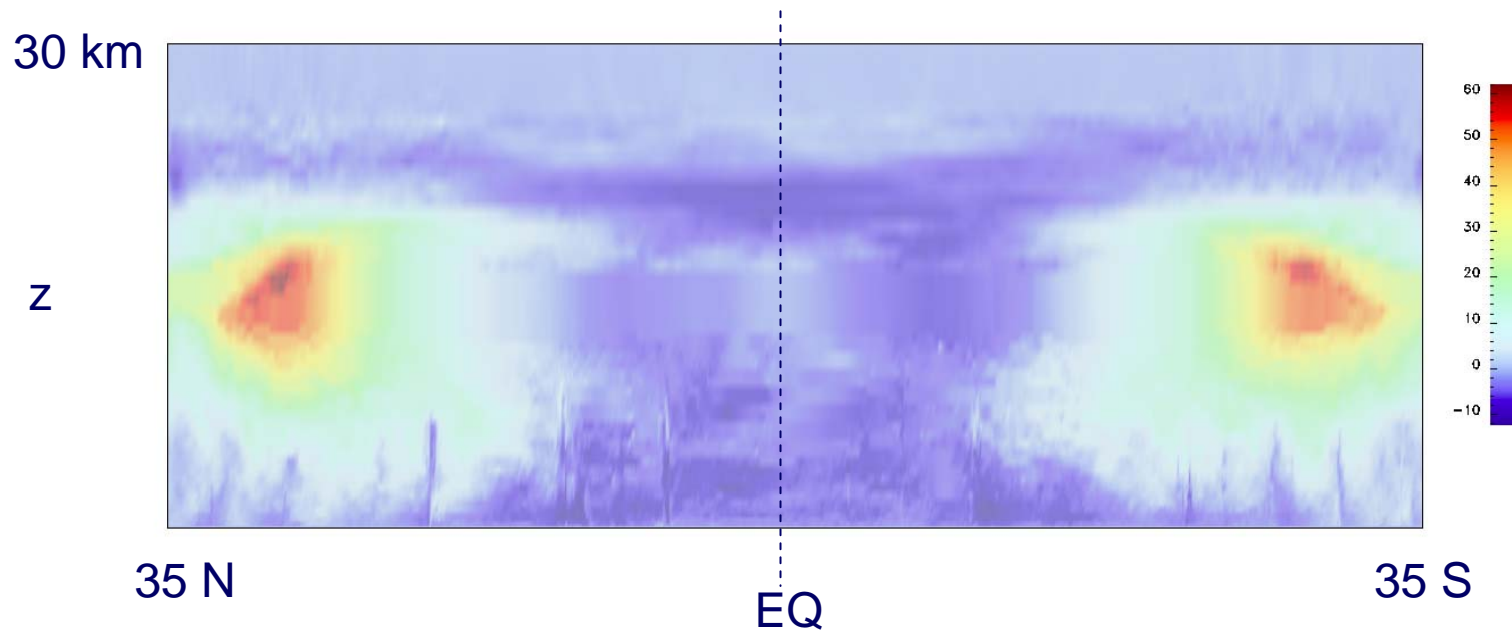
Driving the sub-tropical surface easterlies



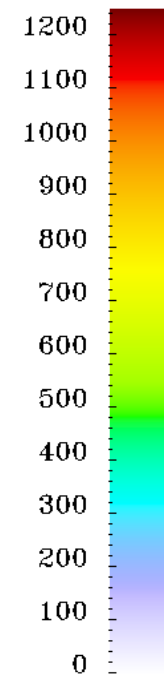
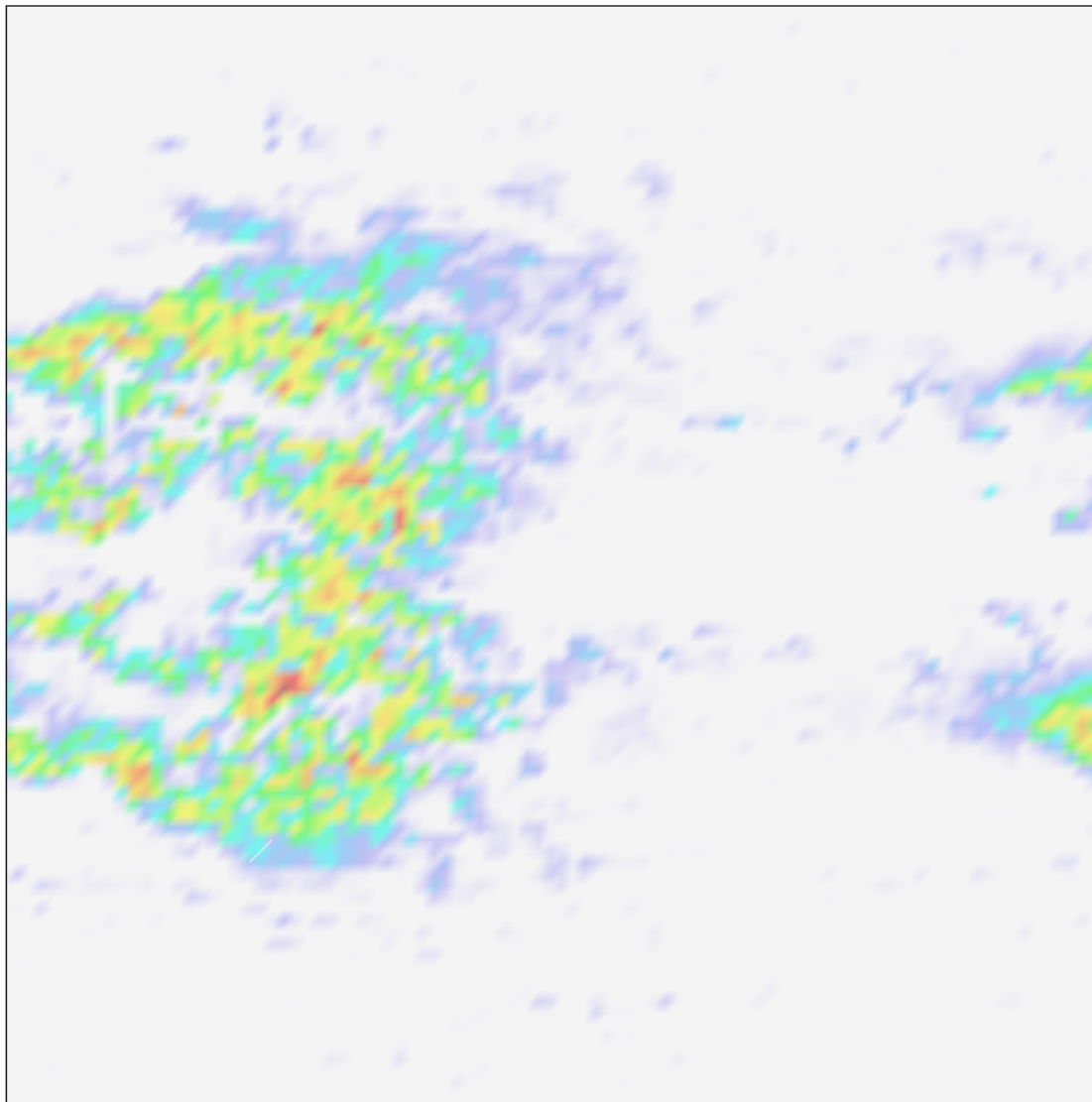
Surface Temperature (z=78 m) at day 4.5



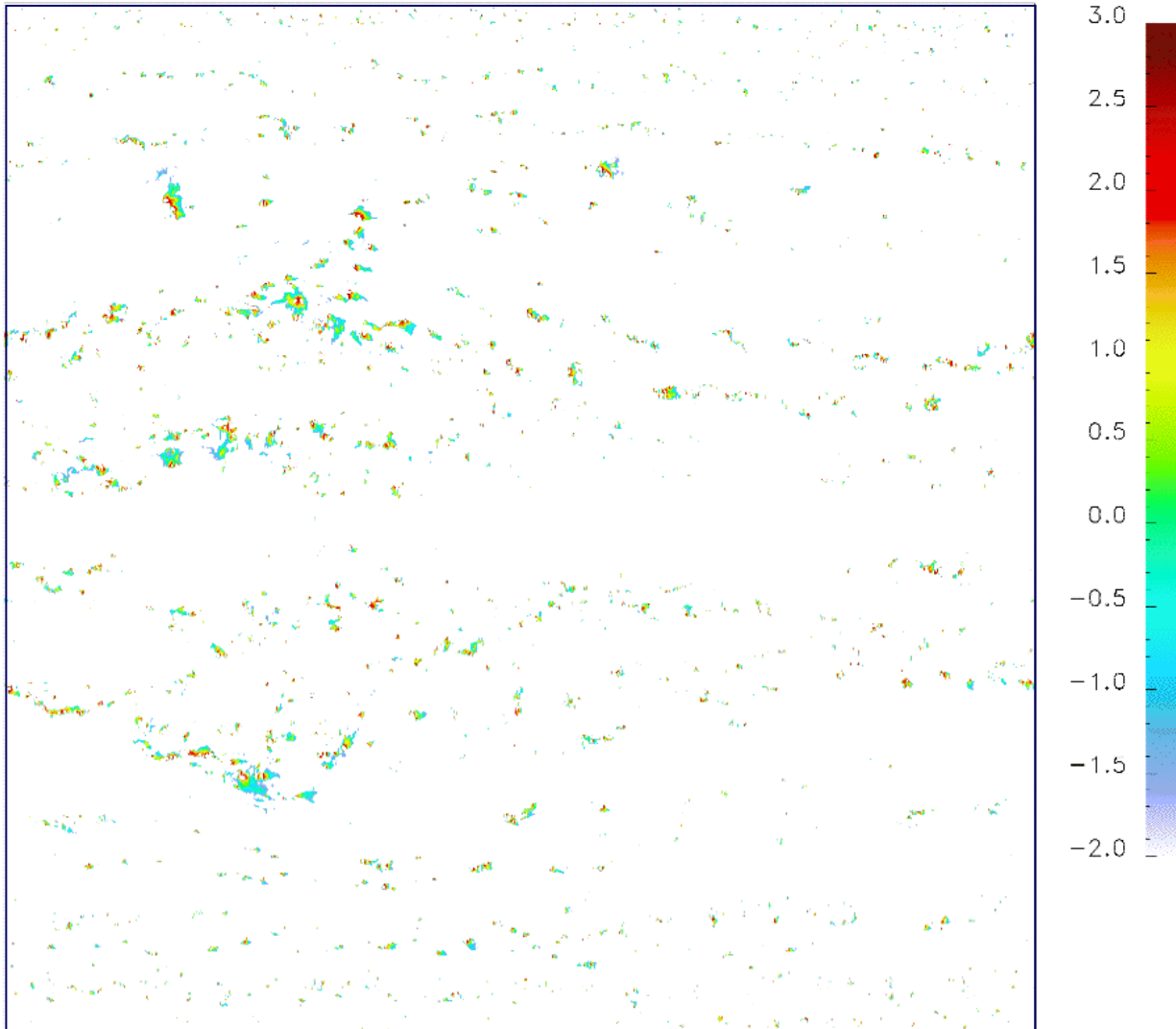
Latitude-height section of zonal wind



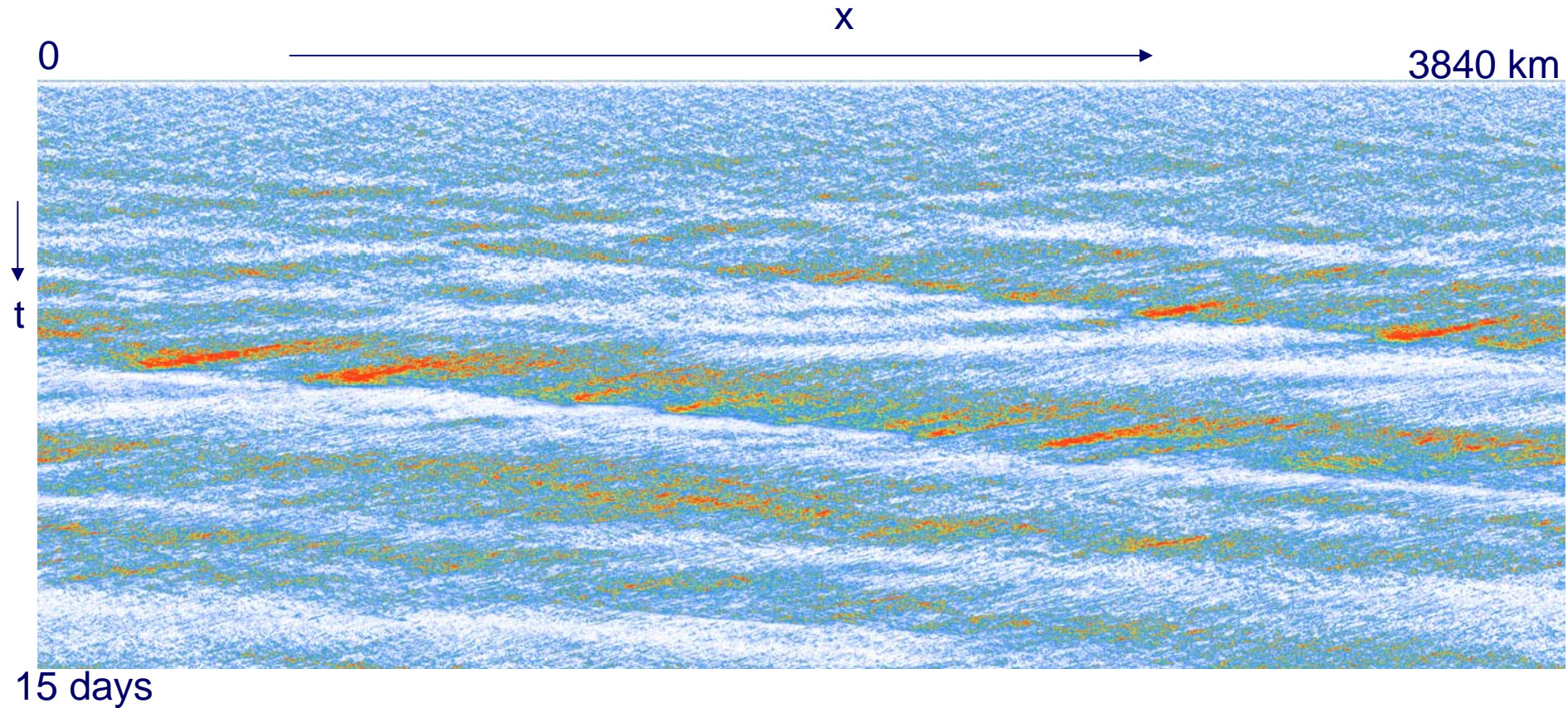
Convective Available Potential Energy



Surface rainfall rate at day 4.5

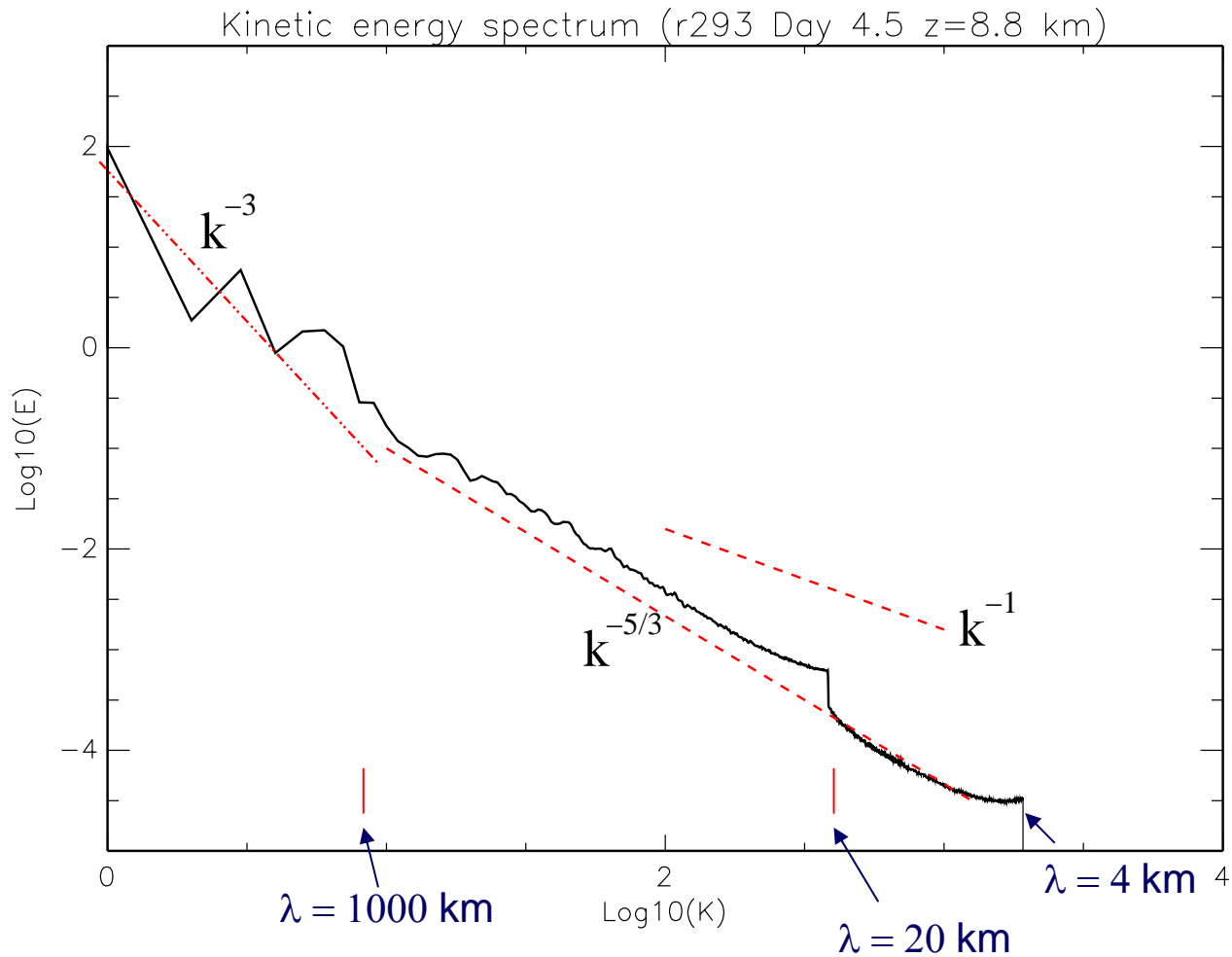


Hovmuller Diagram of rainfall rate from anisotropic grid run



Red is > 50 mm/day

Kinetic Energy spectrum at day 4.5



$$\frac{\partial \theta}{\partial t} = -\mathbf{V} \cdot \nabla \theta + Q$$

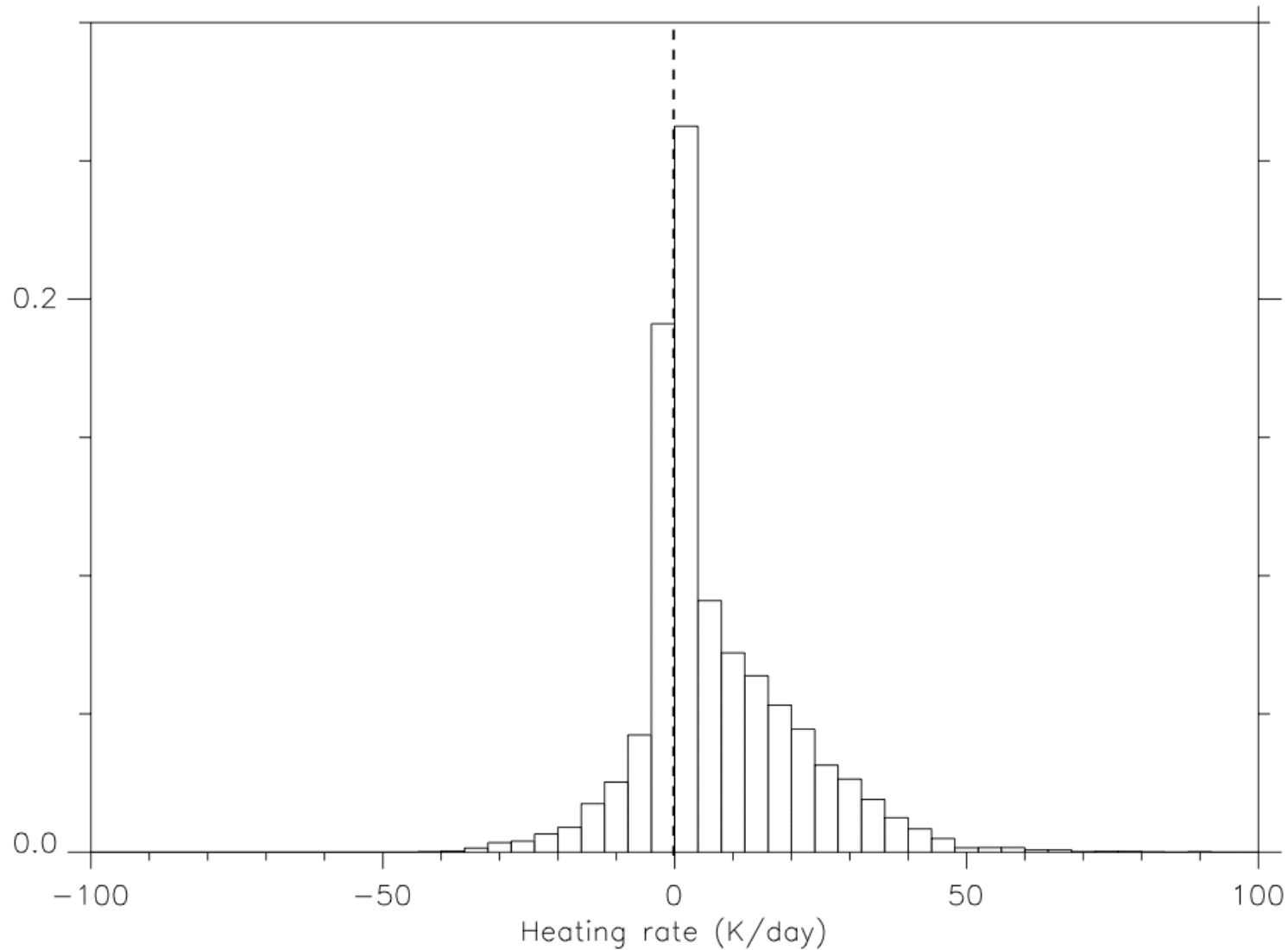
Let overbar denote average over a coarse grid box, then:

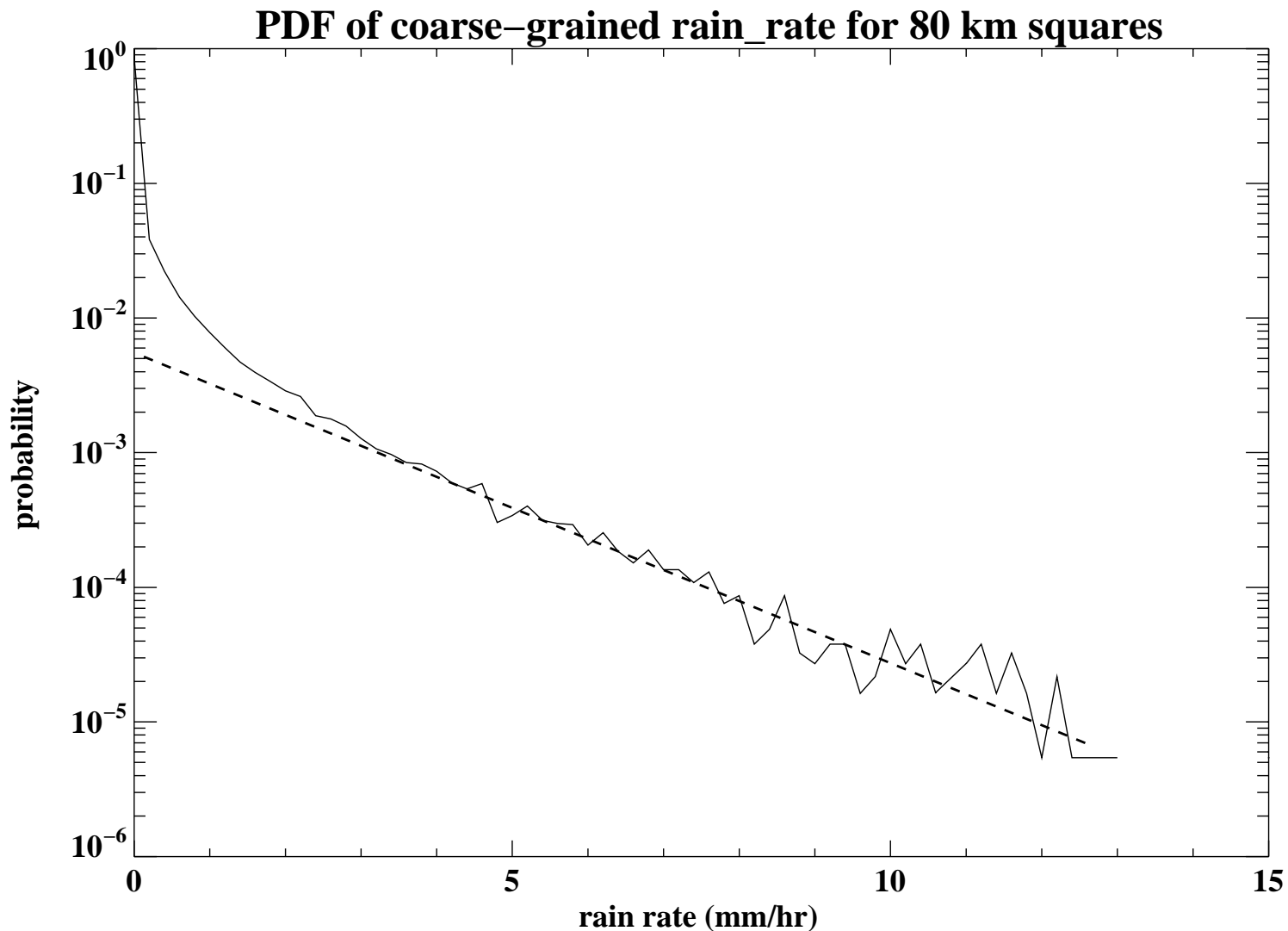
$$\frac{\partial \bar{\theta}}{\partial t} = -\overline{(\mathbf{V} \cdot \nabla \theta)} + \bar{Q}$$

$$\frac{\partial \bar{\theta}}{\partial t} + \bar{\mathbf{V}} \cdot \nabla \bar{\theta} = \underbrace{\bar{\mathbf{V}} \cdot \nabla \bar{\theta} - \overline{(\mathbf{V} \cdot \nabla \theta)}}_{\text{Parametrized + resolved heating}} + \bar{Q} = \tilde{Q}$$

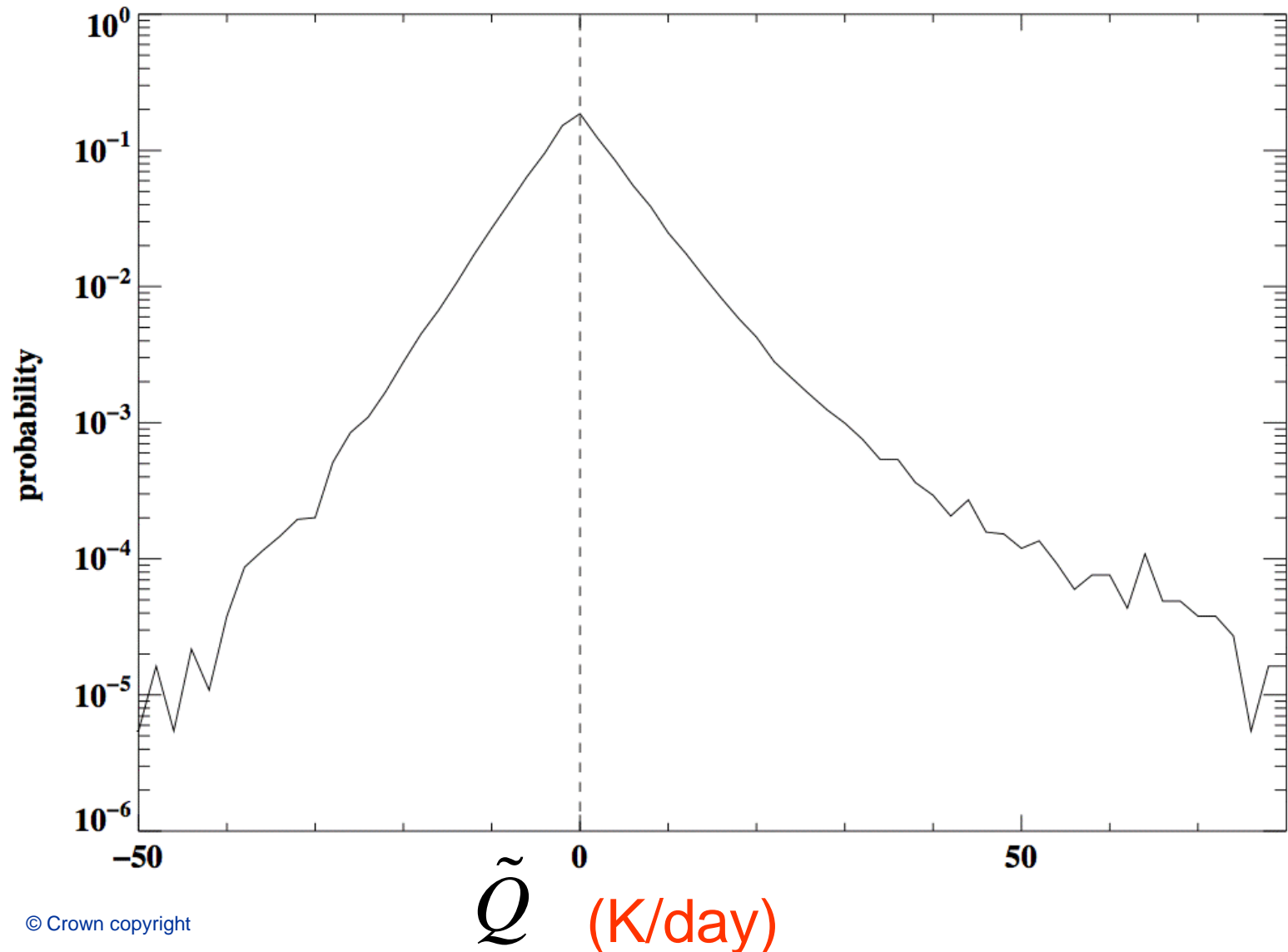
Parametrized + resolved heating

Histogram of diabatic heating coarse-grained to an 80 km grid at $z=9.4$ km





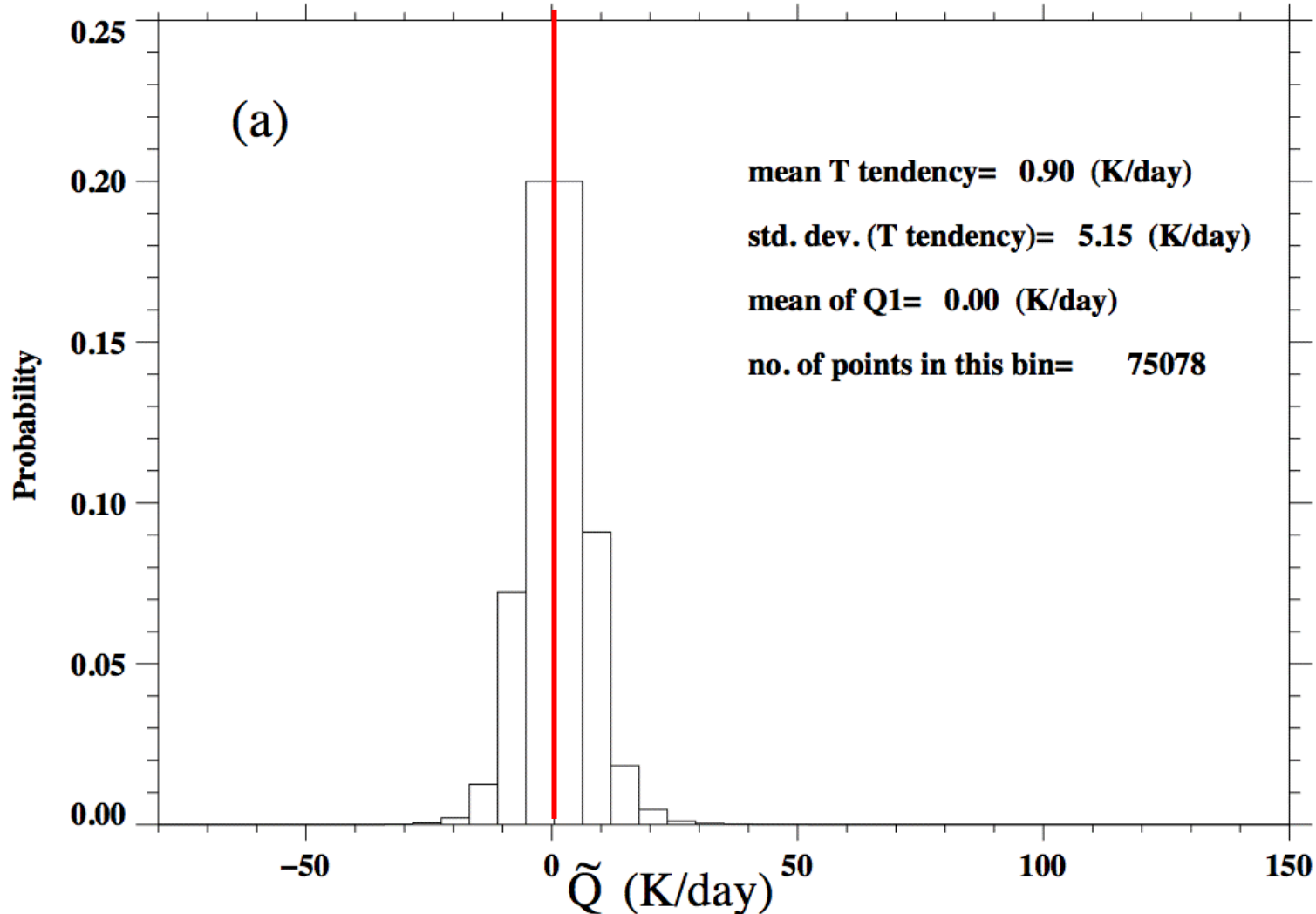
PDF of effective T forcing as seen from an 80 km grid



- take the coarse-grained CRM fields and feed them into a convective parametrization scheme (Bechtold et al, 2001) $\rightarrow Q1$ - the convective warming rate
- at any model level, bin the effective temperature tendency \tilde{Q} according to different ranges of $Q1$
- see how the variance of effective temperature tendency depends on $Q1$
- use knowledge of variance dependence to calibrate 'stochastic physics' schemes based on **multiplicative noise**

Pdf of effective CRM dT/dt for which $-0.1 < Q1 < 0.1 \text{K/day}$

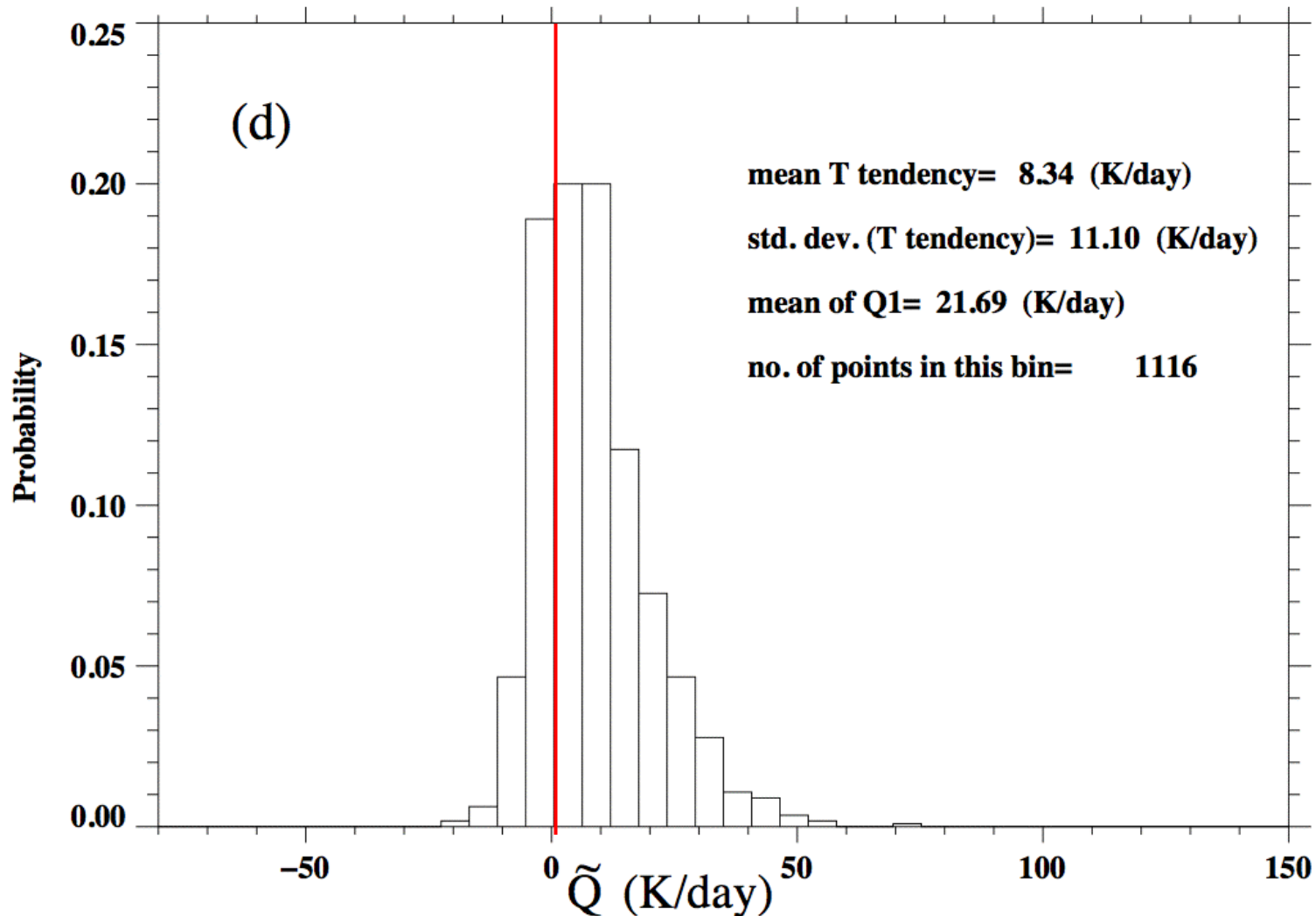
Coarse box size= 120 km



Pdf of effective CRM dT/dt for which $18 < Q1 < 27$ K/day

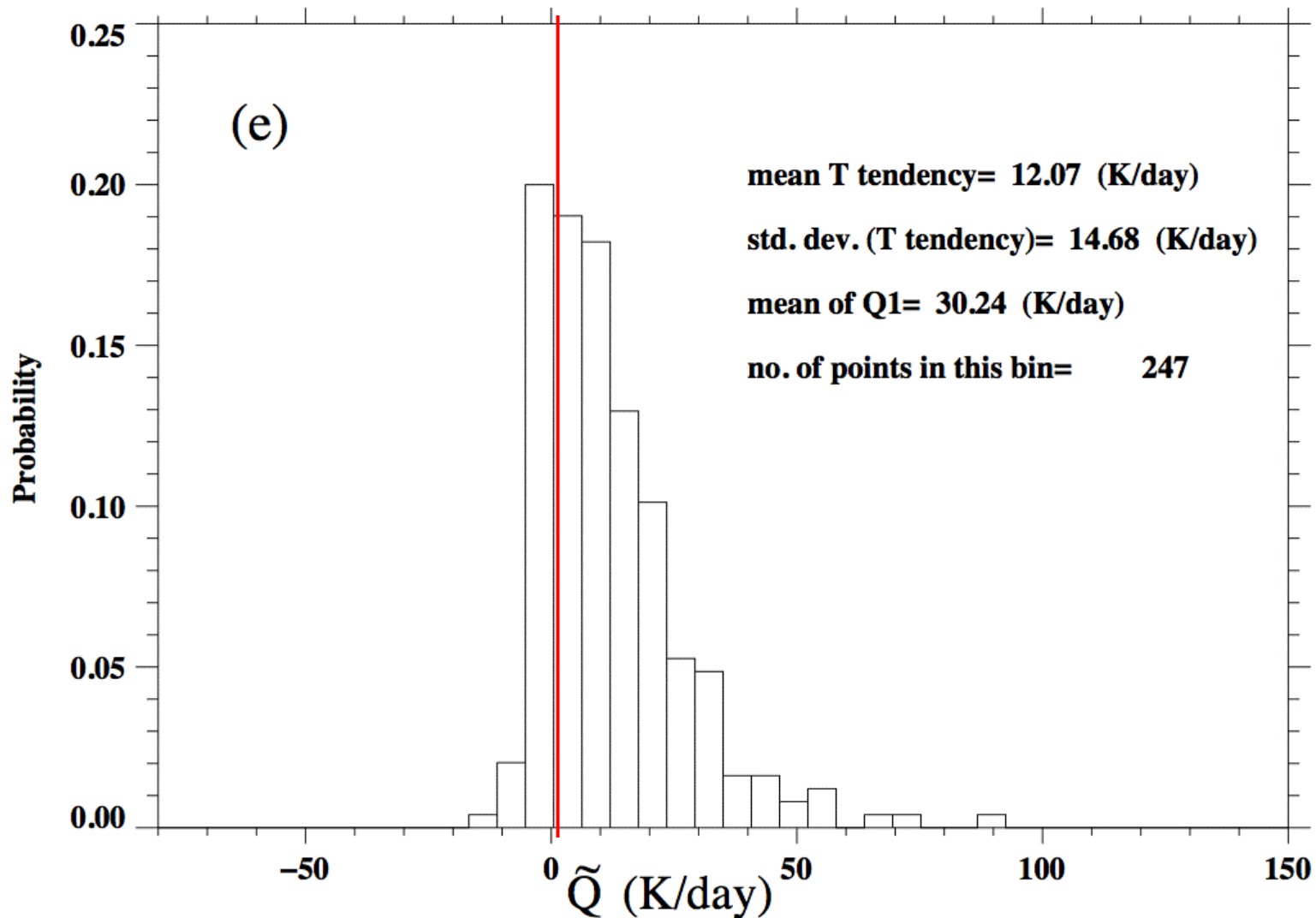


Coarse box size= 120 km



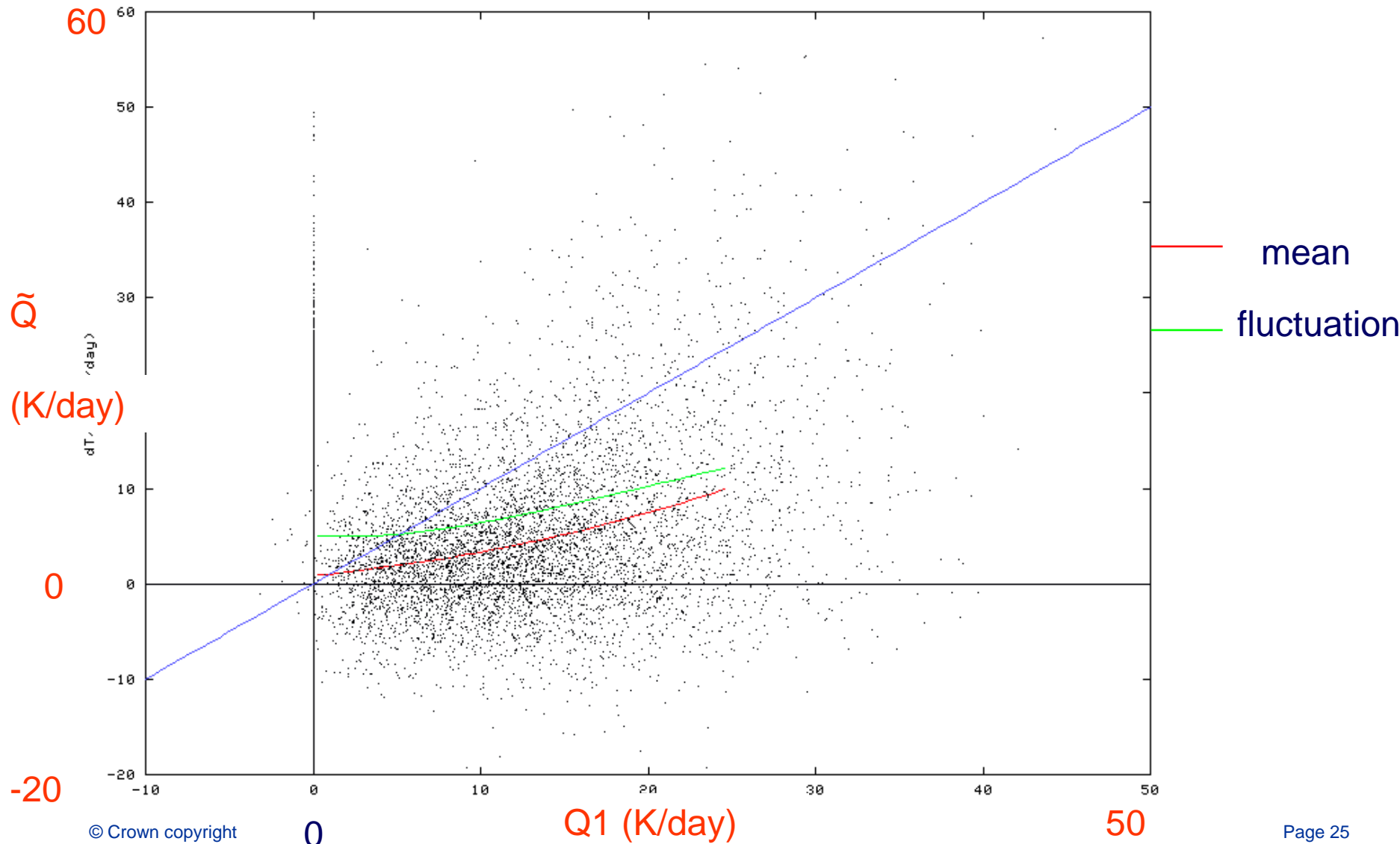
Pdf of effective CRM dT/dt for which $27 < Q1 < 36$ K/day

Coarse box size= 120 km



Scatter plot of effective T forcing computed from CRM versus Q1

Coarse box size= 120 km

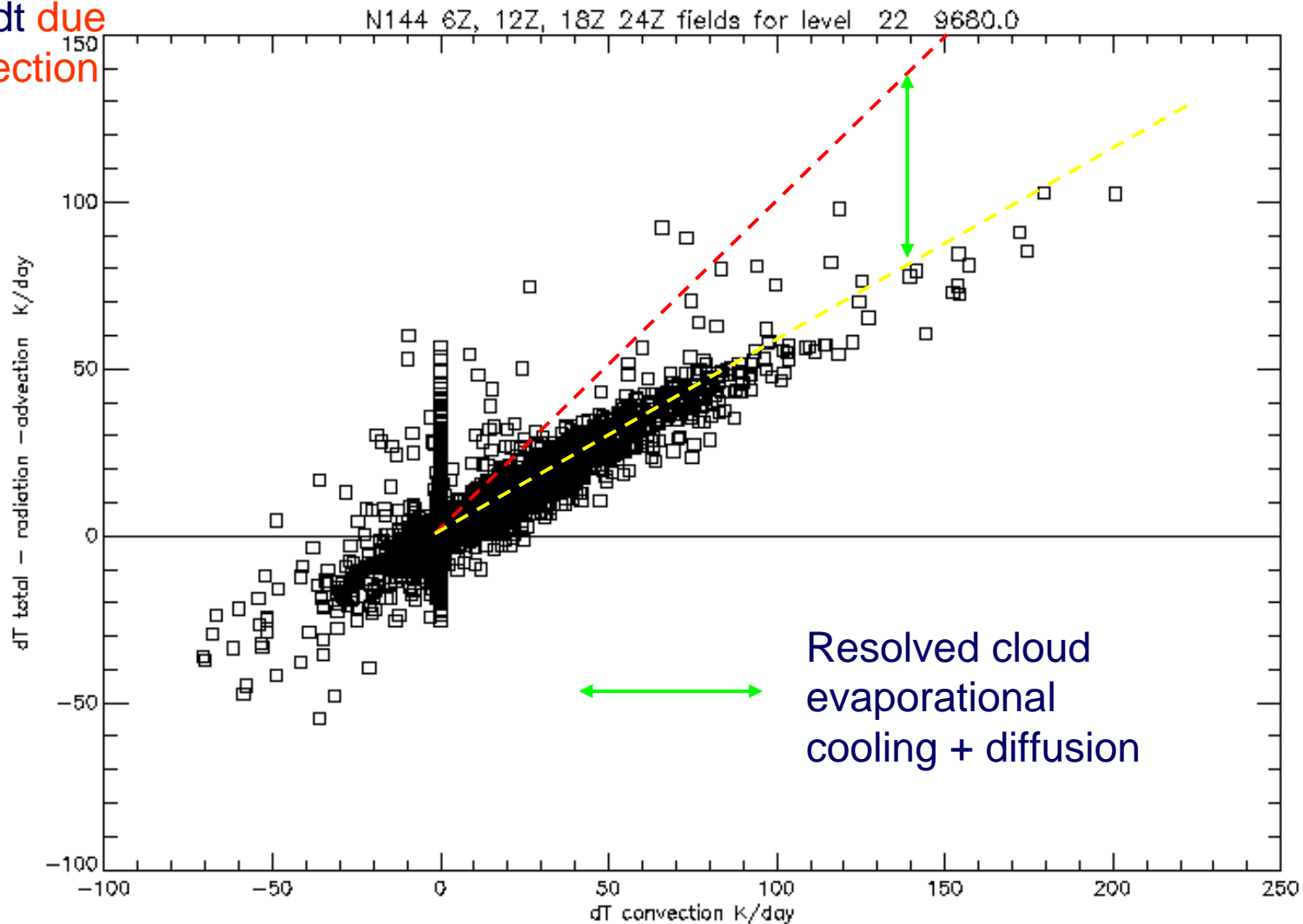


Scatter plot UM (dT/dt – radiation – advection) versus Q1 from aquaplanet run.

Provided by Rachel Stratton

Resolution= N144, z=9.68 km

Net dT/dt due to convection



Resolved cloud evaporational cooling + diffusion

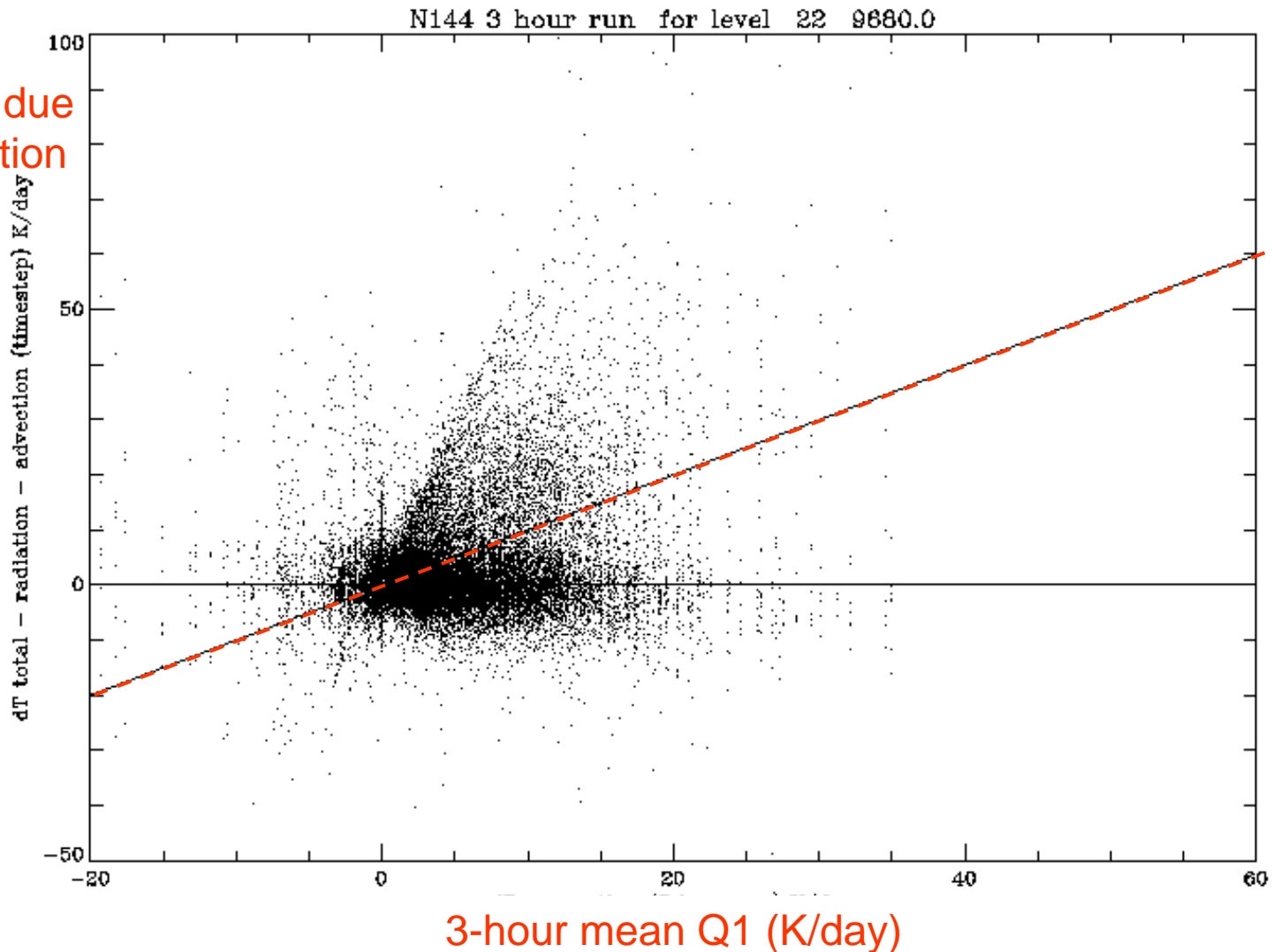
Scatter plot UM (dT/dt – radiation – advection) versus Q1 from aquaplanet run.



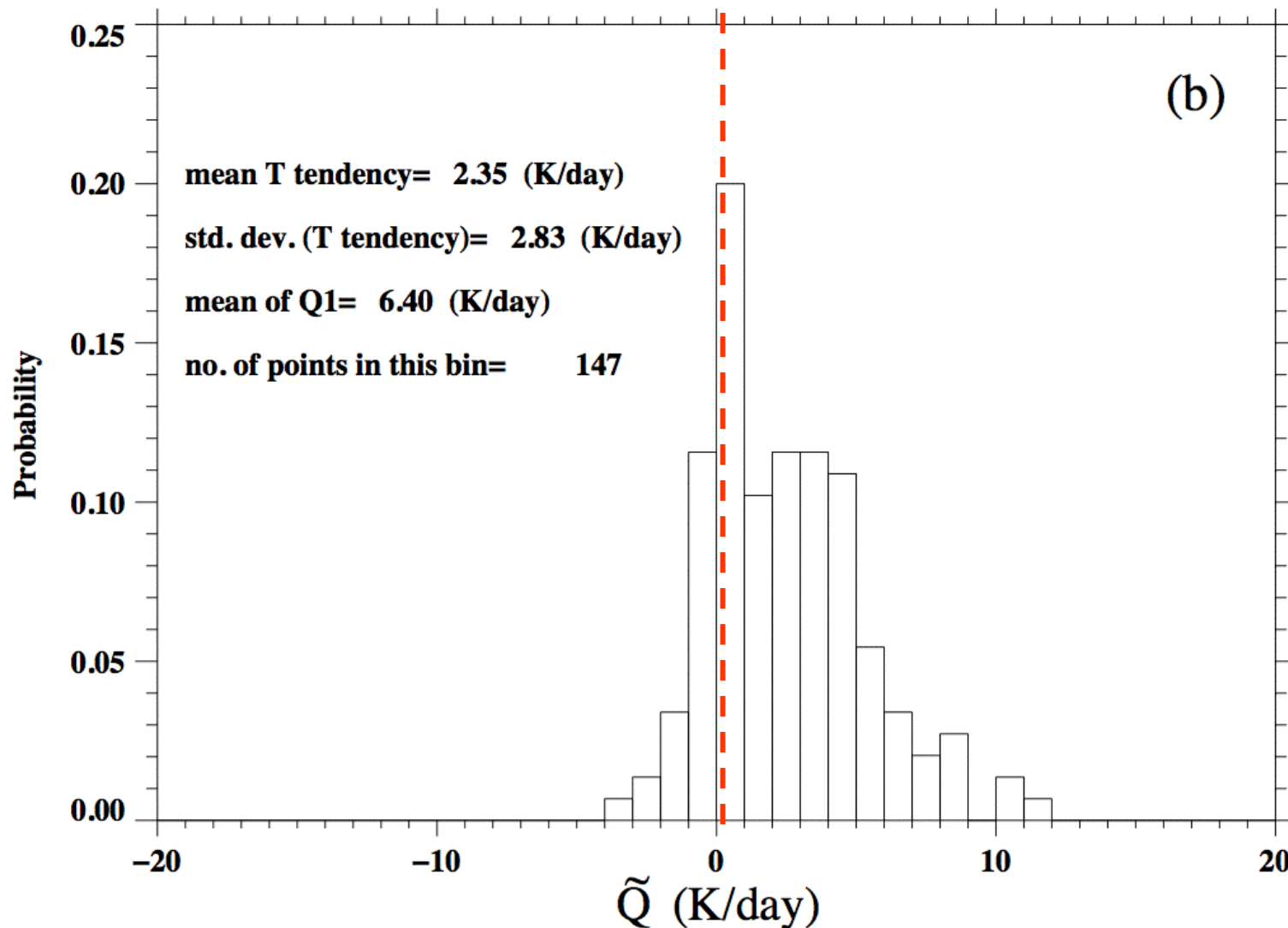
Provided by Rachel Stratton

Resolution= N144, z=9.68 km

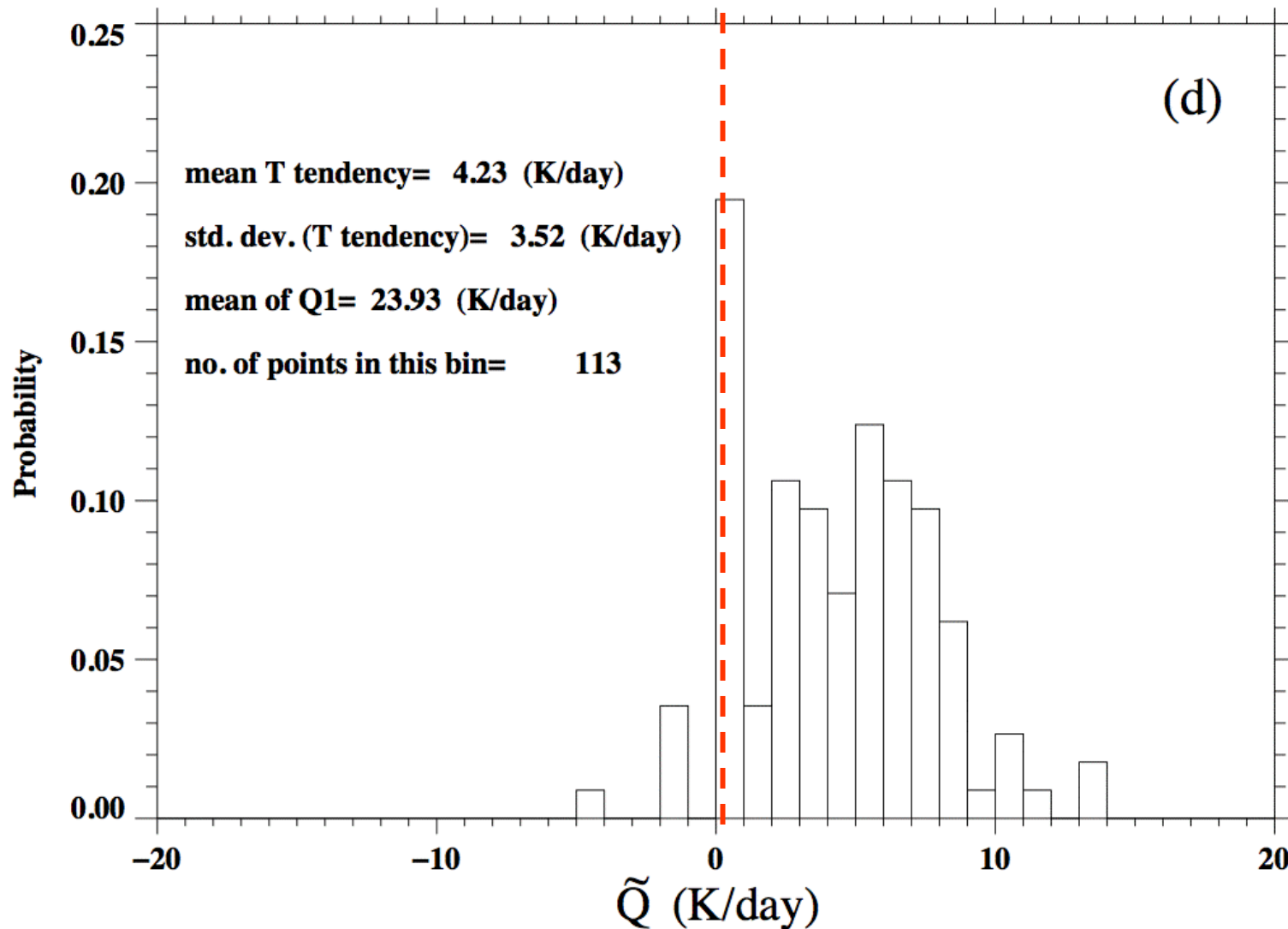
Net dT/dt due
to convection
(K/day)



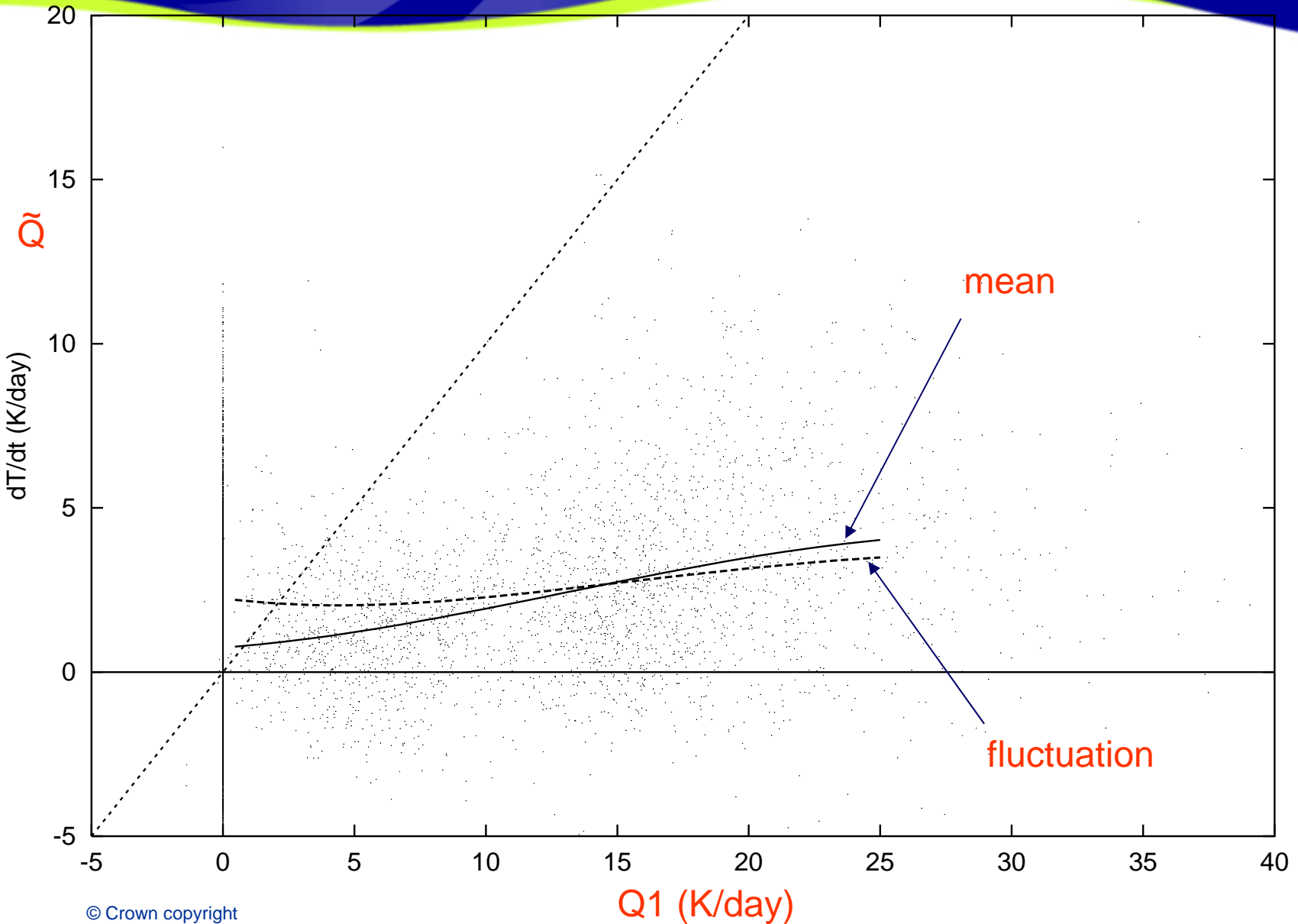
Pdf of effective CRM on 320 km grid dT/dt for which $0 < Q1 < 10$ K/day



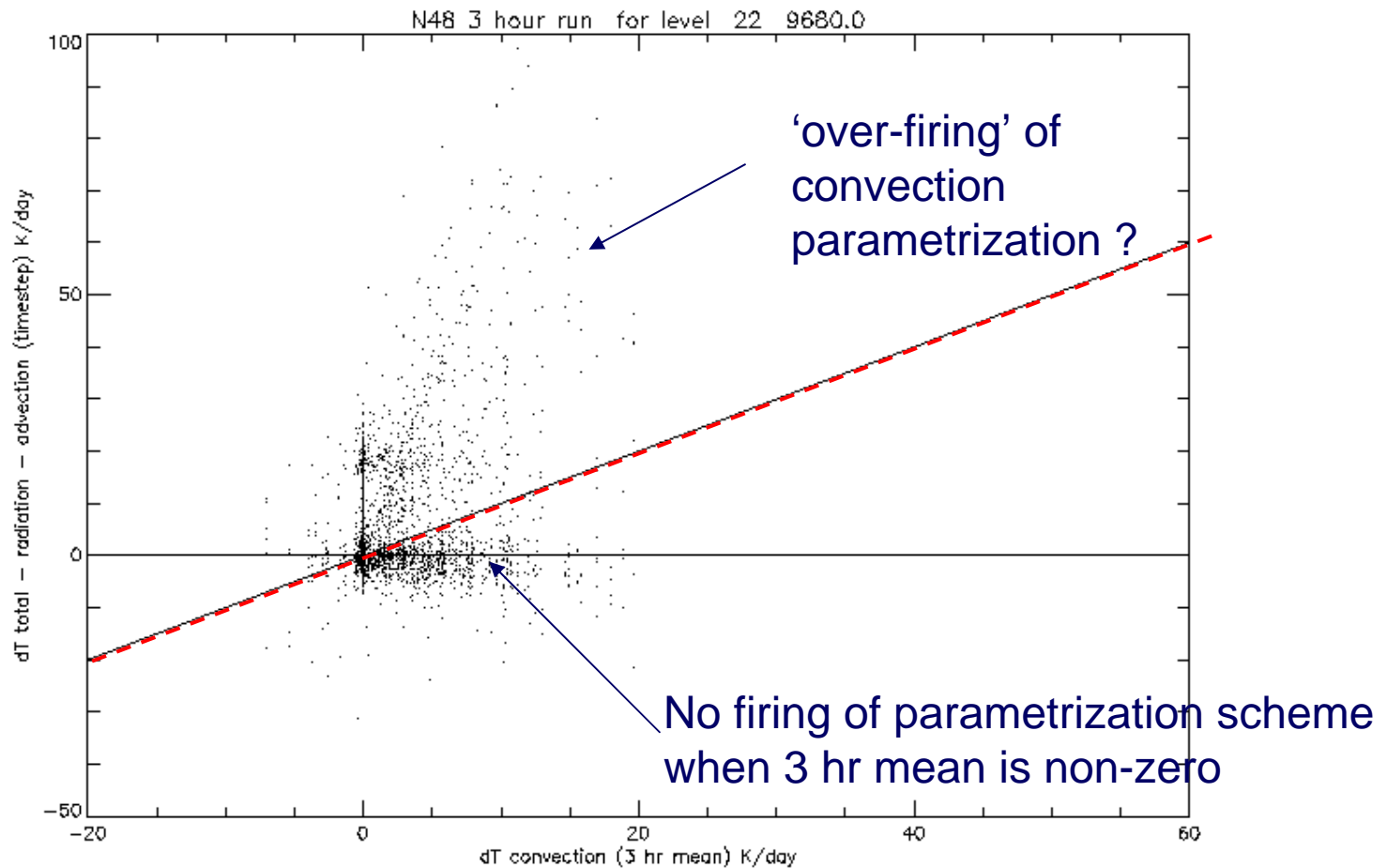
Pdf of effective CRM on 320 km grid dT/dt for which $20 < Q1 < 40$ K/day



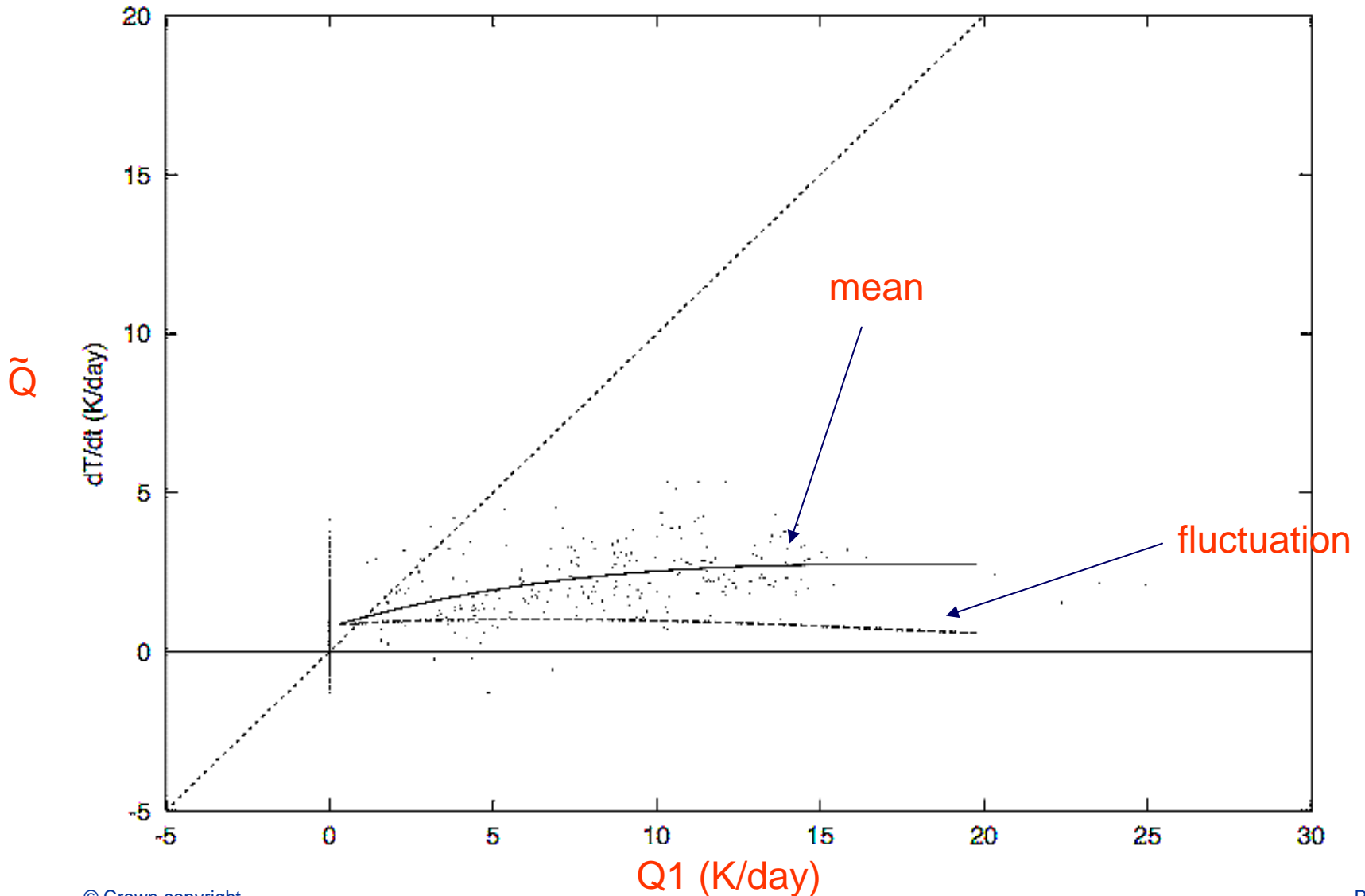
Scatter plot of effective dT/dt for 320 km grid versus Q_1



Scatter plot UM (dT/dt – radiation – advection) versus 3-hr mean Q1 at N48



Scatter plot of effective dT/dt for 960 km grid versus Q_1



- use a cellular automaton to prescribe time-evolving patterns of convective forcing (Palmer, 1997)
- Cellular Automaton Stochastic Backscatter (CASB) scheme has been used in ECMWF EPS and climate runs to inject KE back into flow. (Shutts, 2005)
- CASB only concerned with convective momentum tendencies

Living cells are **red**

Dead cells are white

0	0	0	0	0	0
0	0	0	0	0	0
1	2	3	3	2	1
1	1	2	2	1	1
1	2	3	3	2	1
0	0	0	0	0	0

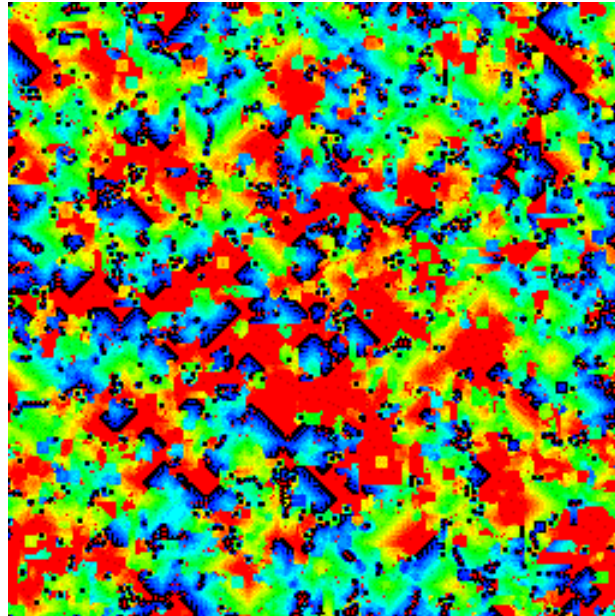


0	0	0	0	0	0
0	1	2	2	1	0
0	2	3	3	2	0
0	3	5	5	3	0
0	2	3	3	2	0
0	1	2	2	1	0

Rules:

Survival: 2 or 3 living neighbours

Birth: 3 living neighbours



- Blue cells are young
- Red cells are old or dead

- estimate the total rate of generation of sub-grid scale kinetic energy due to numerical and physical sources (D)
- define time-evolving patterns with a cellular automaton (CA)
- compute a field of vorticity increments using the CA pattern and scaling the amplitude with \sqrt{D}
- streamfunction tendency due to upscale energy transfer given by:

$$\frac{\partial \psi}{\partial t} = \alpha \cdot \Delta s \cdot \Psi(\lambda, \phi, t) \sqrt{\Delta \tau \cdot D} / \Delta \tau$$

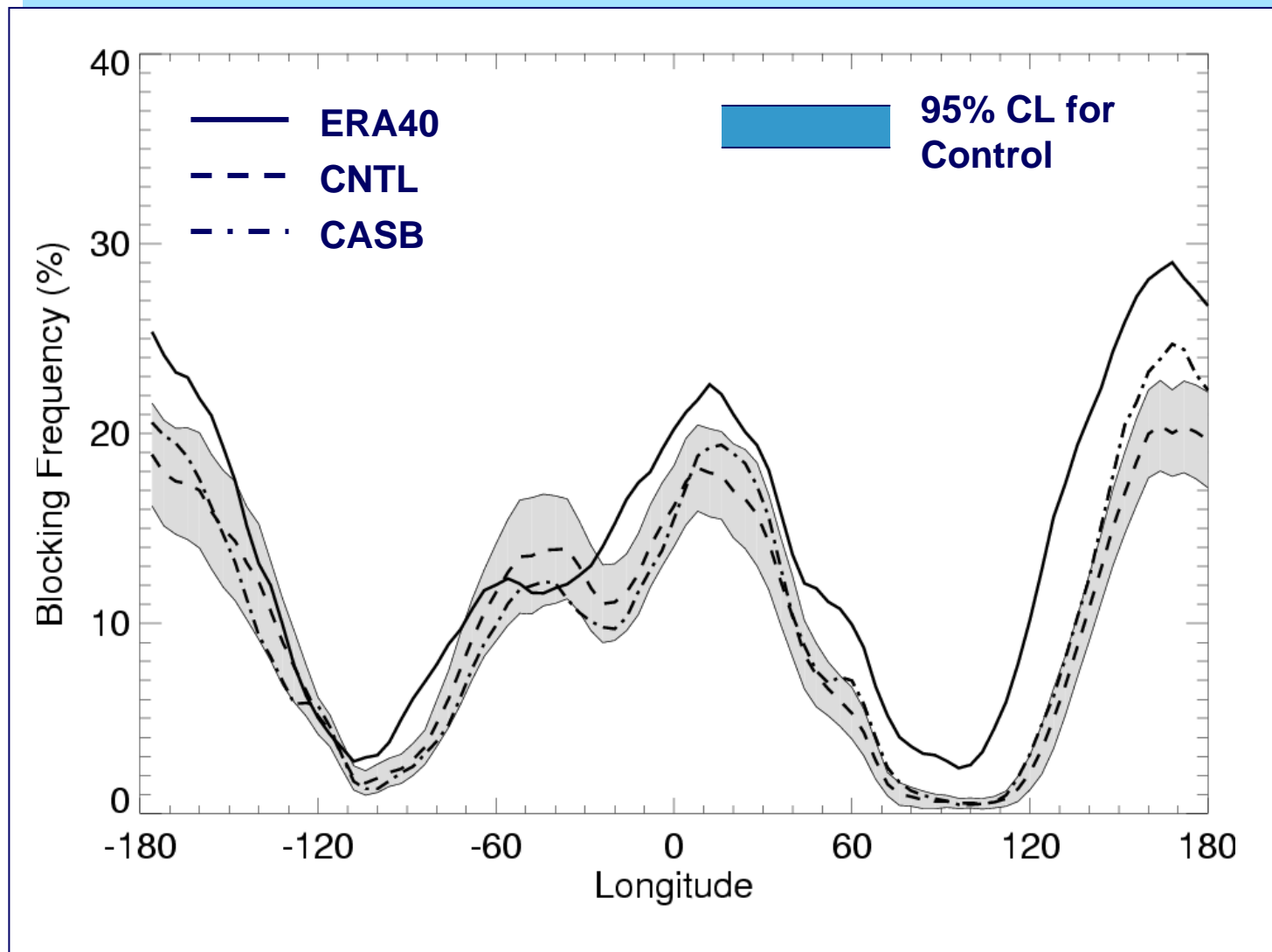
α = dimensionless parameter determining backscatter fraction

Δs = CA gridlength Ψ = CA pattern amplitude (function of lat. /lon.)

$\Delta \tau$ = time interval between CA states

Impact of Stochastic backscatter on Atmospheric Blocking (Dec-Mar 1962-2001)

Work by Thomas Jung



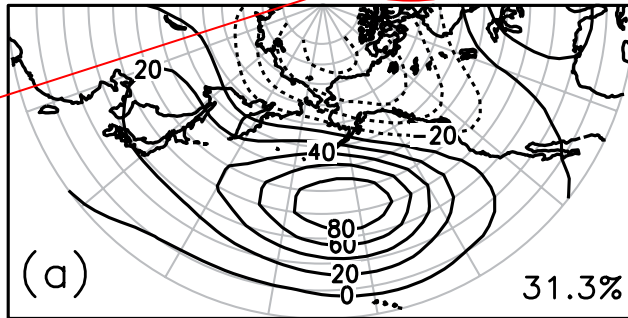
Cluster centroids of wintertime Z500 anomalies

From Jung, Shutts and Palmer (2005)

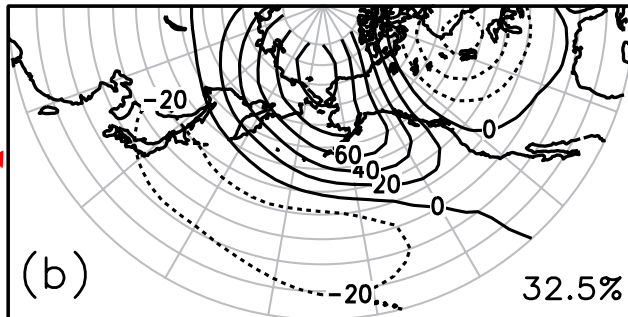


For winters in the period 1962-2001

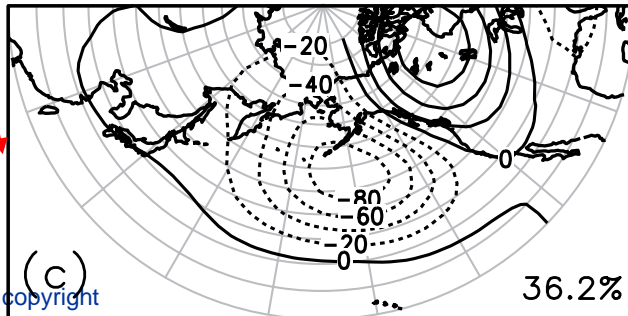
Cluster #1 (ERA40)



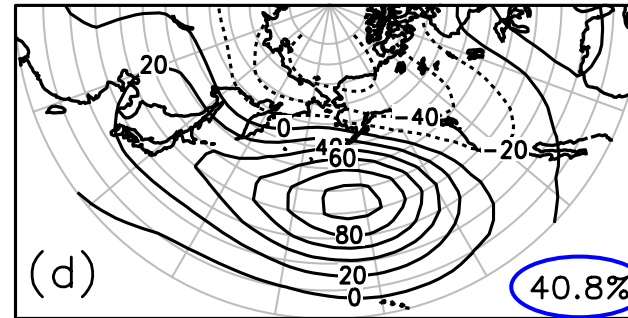
Cluster #2 (ERA40)



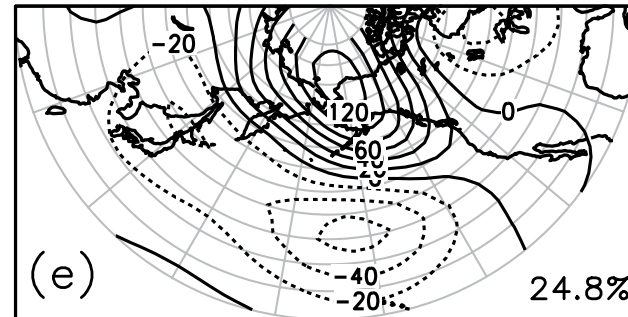
Cluster #3 (ERA40)



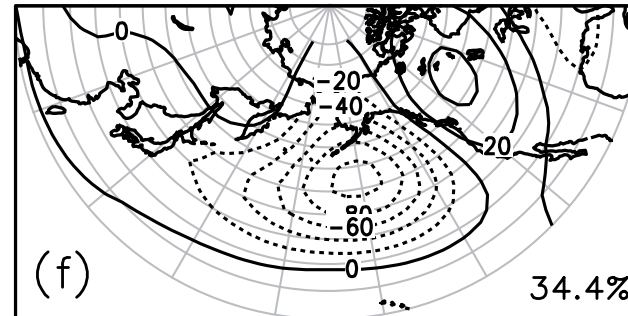
Cluster #1 (CNTL)



Cluster #2 (CNTL)



Cluster #3 (CNTL)



ERA 40
'truth'

T95

Units= m

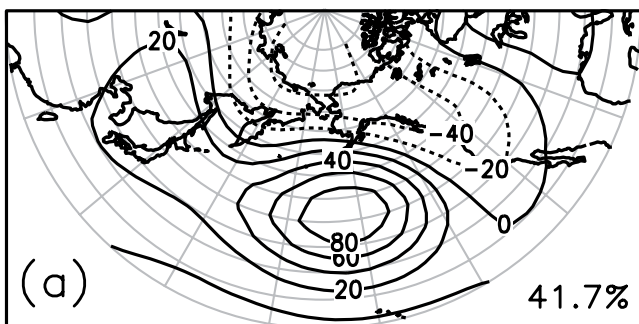
frequency

control

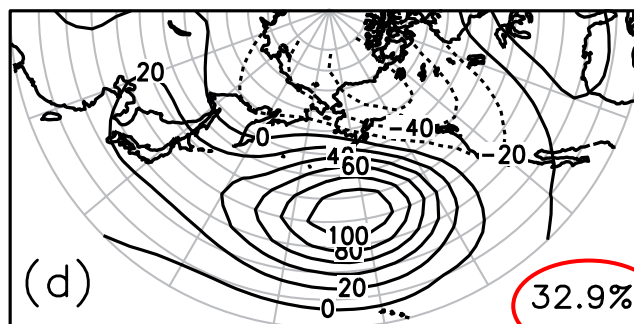
Stochastically-perturbed cluster centroids

Op.
stochastic
physics
(SSP)

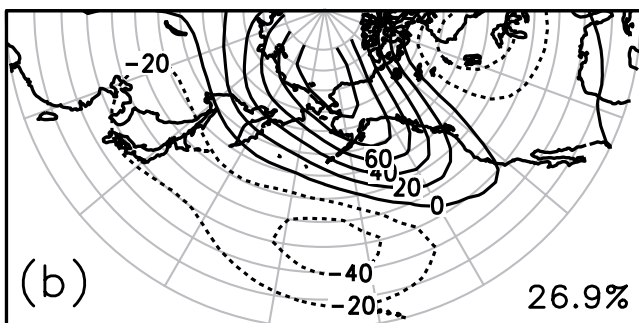
Cluster #1 (SSP)



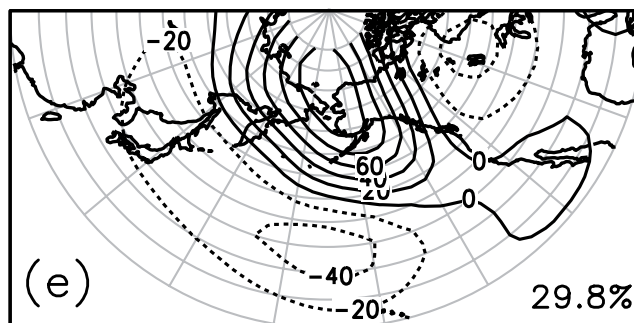
Cluster #1 (CASB)



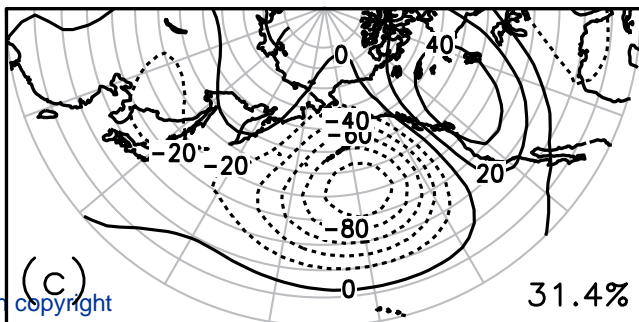
Cluster #2 (SSP)



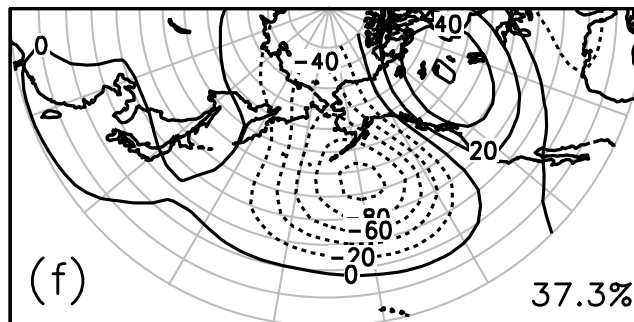
Cluster #2 (CASB)



Cluster #3 (SSP)



Cluster #3 (CASB)



CASB

Closer to
ERA 40 than
CNTL or SSP

- idealized tropical circulation can be numerically-simulated with a CRM
- using a coarse-graining approach the effective cloud forcing can be computed
- PDF of latent heating on an 80 km grid shows cooling as well as warming due to evaporation
- PDFs of Q can be conditioned on the parametrized convective tendency Q_1
- Variance of Q increases with Q_1 (although at 960 km apparently not !)
- Calibrate stochastic convective parametrizations with PDF info from cloud-resolving model simulations