

Modeling sea ice as a granular material, including dilatancy effect

A person wearing a dark jacket and red pants is walking across a vast, flat, icy landscape. The ground is covered in a layer of ice with some snow patches. The sky is bright and clear, suggesting a sunny day. The person is walking towards the right side of the frame.

Bruno Tremblay

Lamont Doherty Earth Observatory

Outline

- Introduction - Different Approach to Sea Ice Modeling
- Observations
- Granular Model
- Model Against Observations?
- Model Against Observations... continued

Momentum Equation

$$\rho h \frac{du}{dt} = \tau_a + \tau_w - \rho h f k \times u - \rho g h \nabla H + \nabla \cdot \sigma$$

Momentum Equation

$$\rho h \frac{du}{dt} = \tau_a + \tau_w - \rho h f k \times u - \rho g h \nabla H + \nabla \cdot \sigma$$

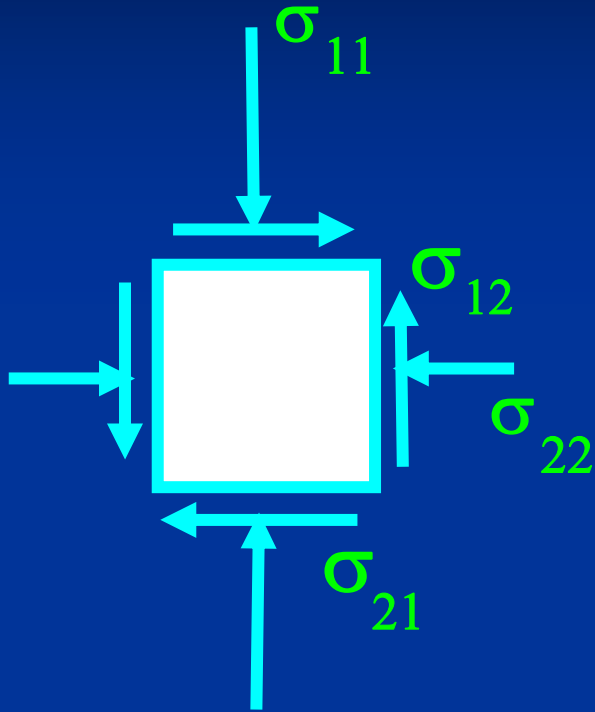
Internal ice stress term



Modeling sea ice

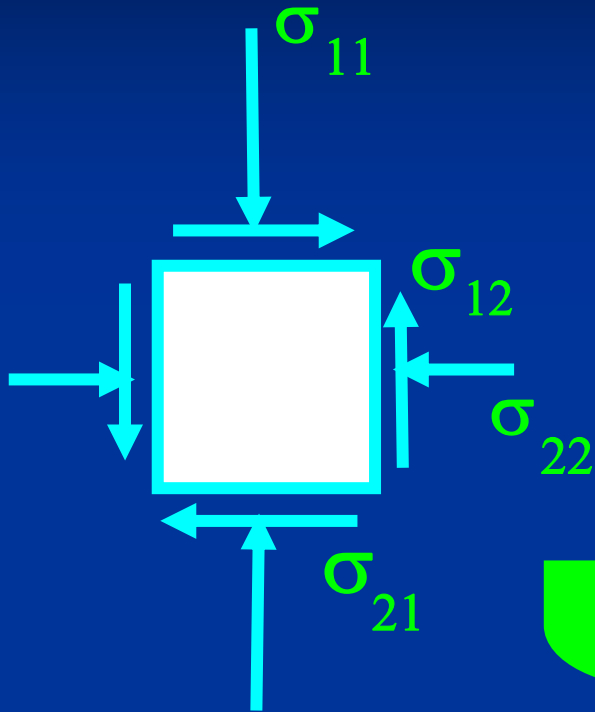
- Deformation
 - Elastic: small, reversible
 - Plastic: large, irreversible
- Elastic-Plastic (EP) [1974]
- Viscous-Plastic (VP) [1979]
- Elastic-Viscous-Plastic (EVP) [1998]
- Granular material model VP [1997]
 - yield curve + flow rule

Stress State



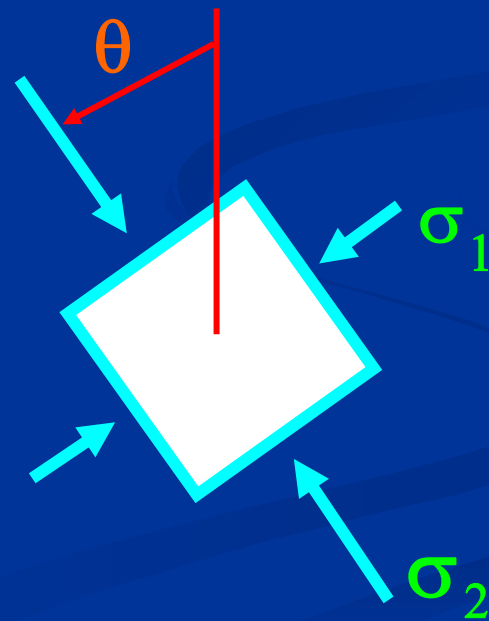
$$\begin{bmatrix} \sigma_{11} & \sigma_{12} \\ \sigma_{21} & \sigma_{22} \end{bmatrix}$$

Stress State



$$\begin{bmatrix} \sigma_{11} & \sigma_{12} \\ \sigma_{21} & \sigma_{22} \end{bmatrix}$$

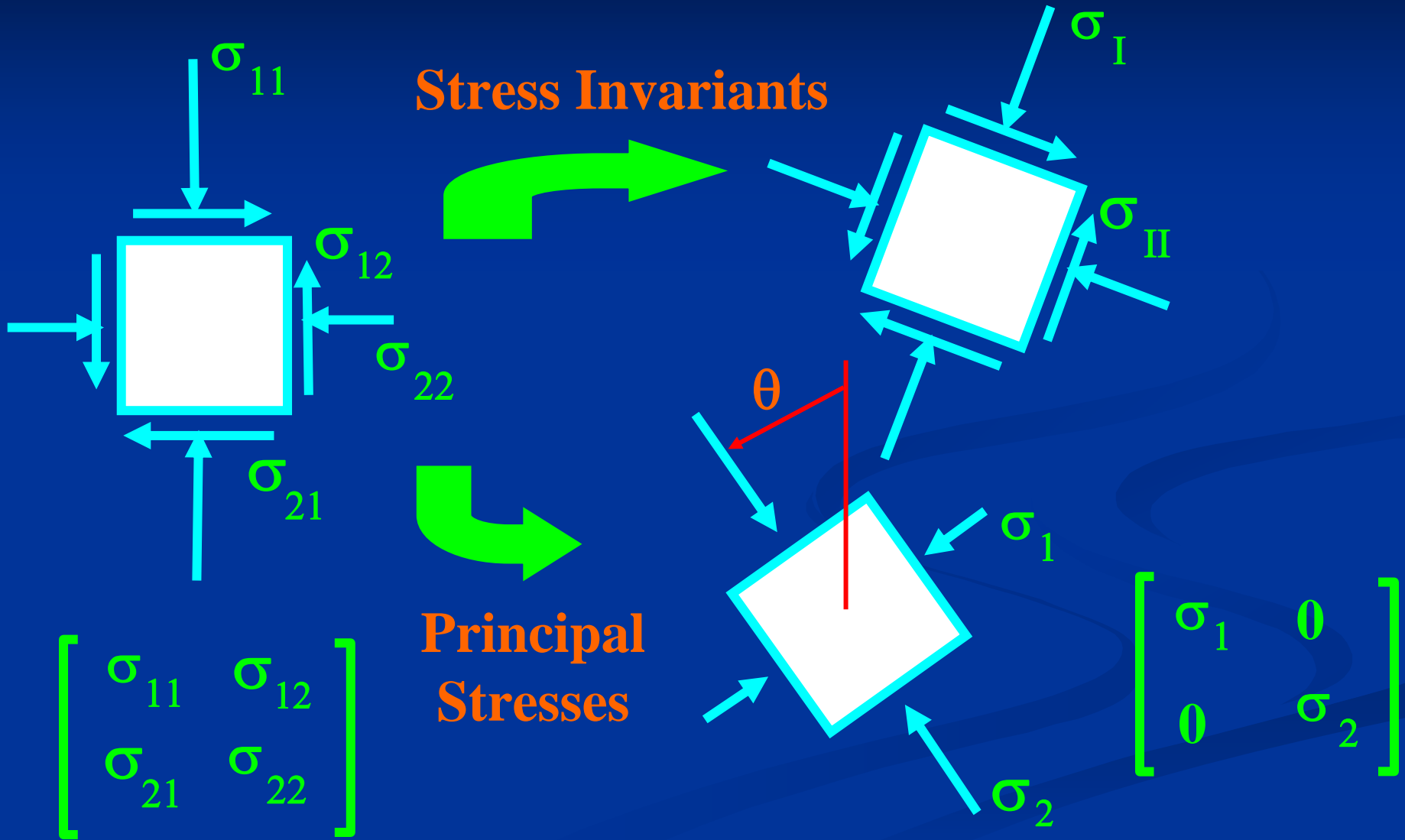
Principal Stresses



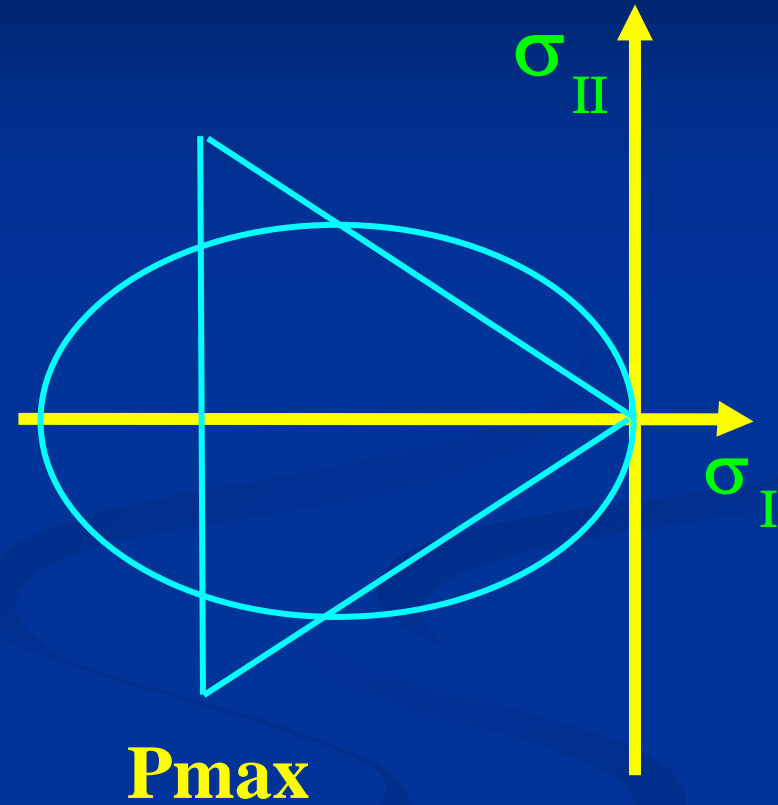
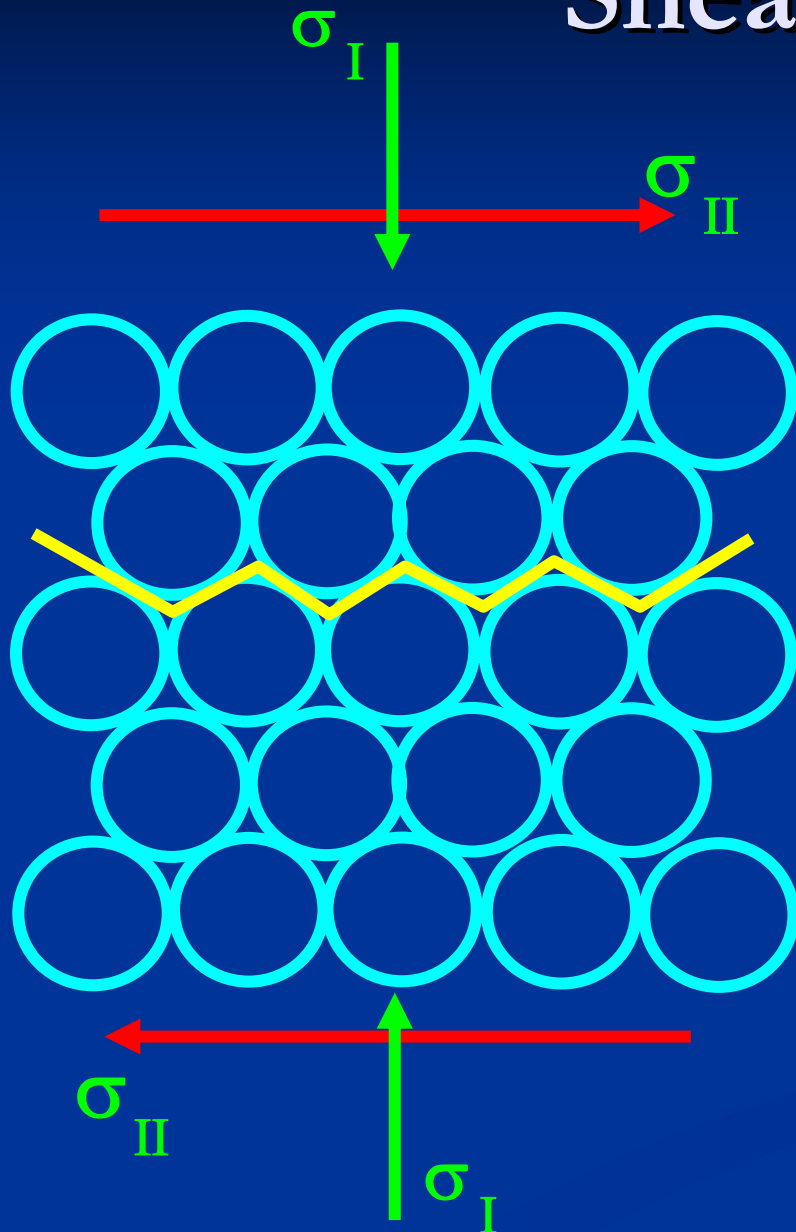
$$\begin{bmatrix} \sigma_1 & 0 \\ 0 & \sigma_2 \end{bmatrix}$$

Stress State

Stress Invariants

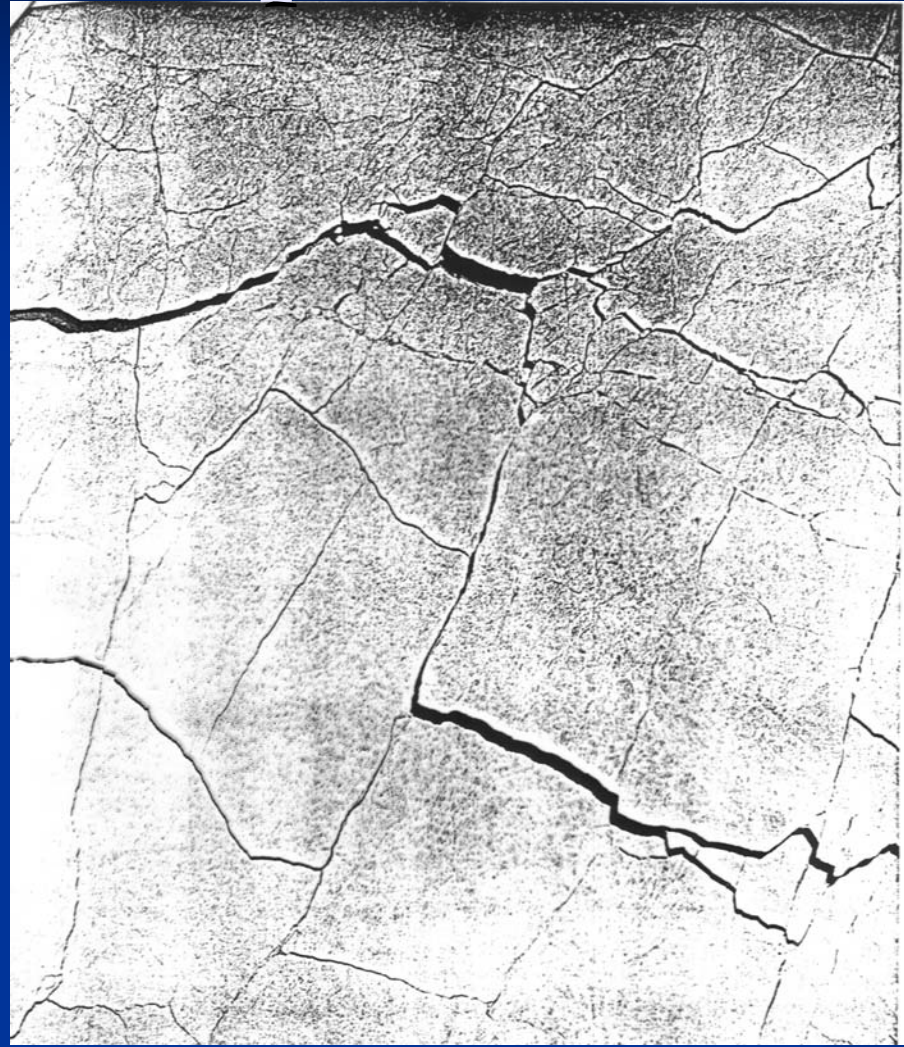


Shear Test



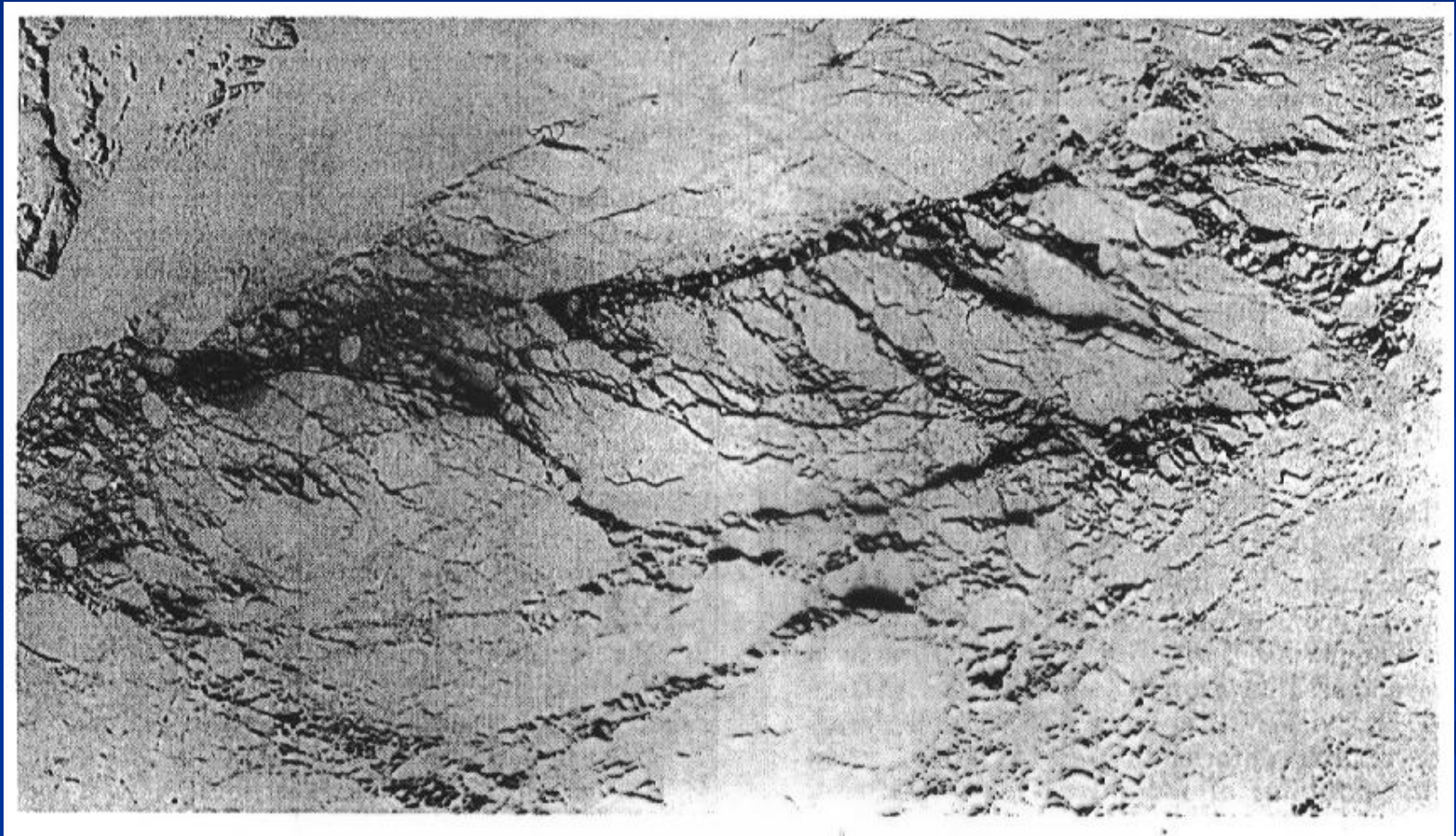
$$P_{\max} = P^* h f(A)$$

Fracture pattern in sea ice



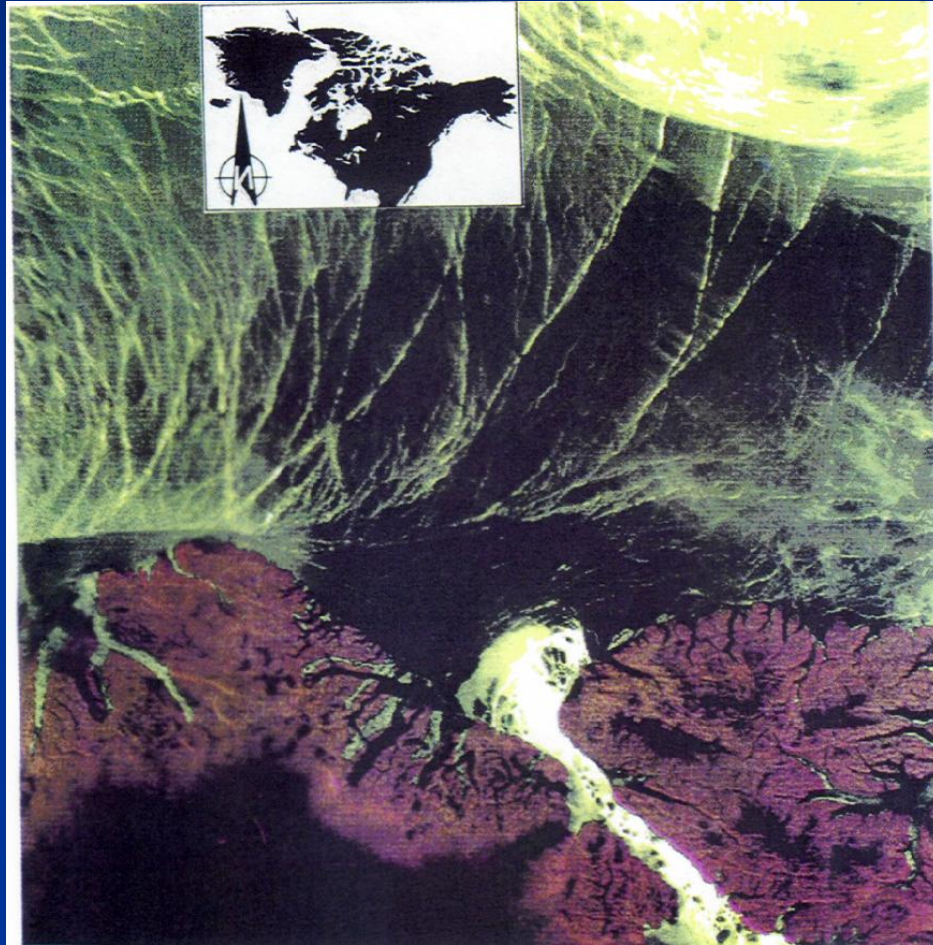
20 km

Fracture pattern in sea ice



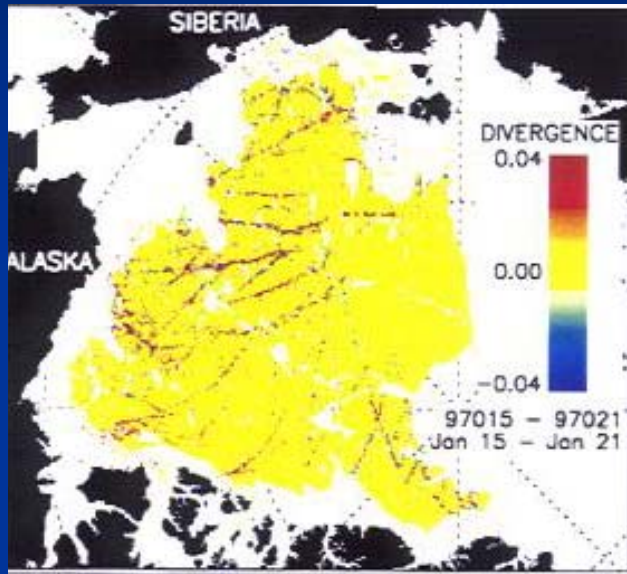
Scale: 200km

Lead pattern north of Greenland



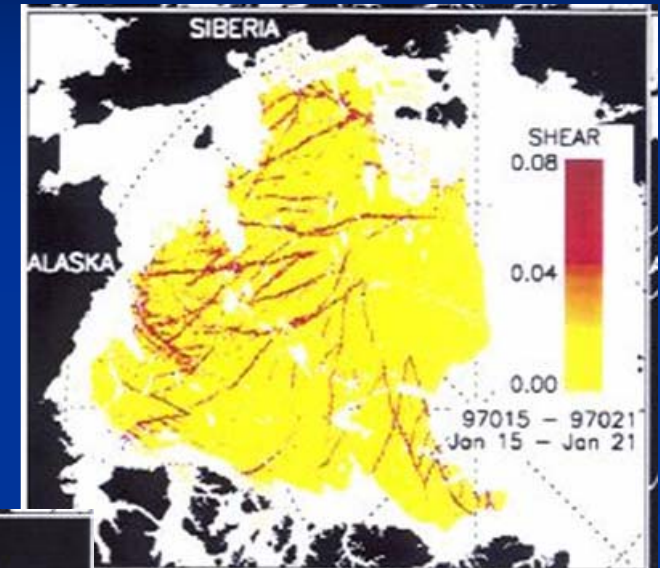
Scale: 500km

Sea ice deformation

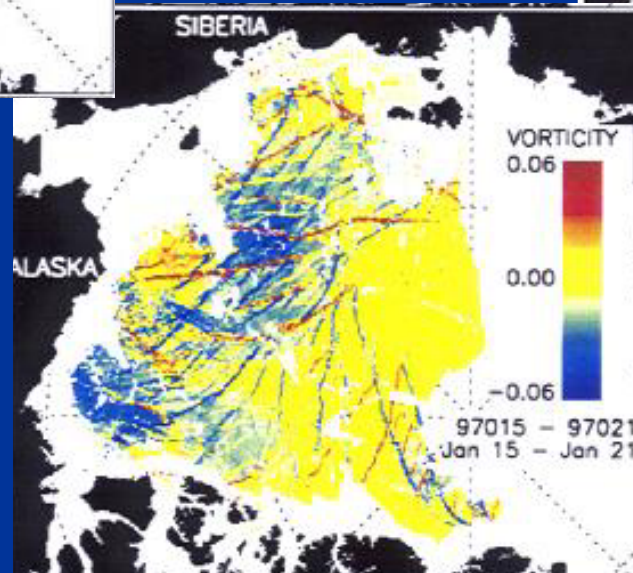


divergence

vorticity

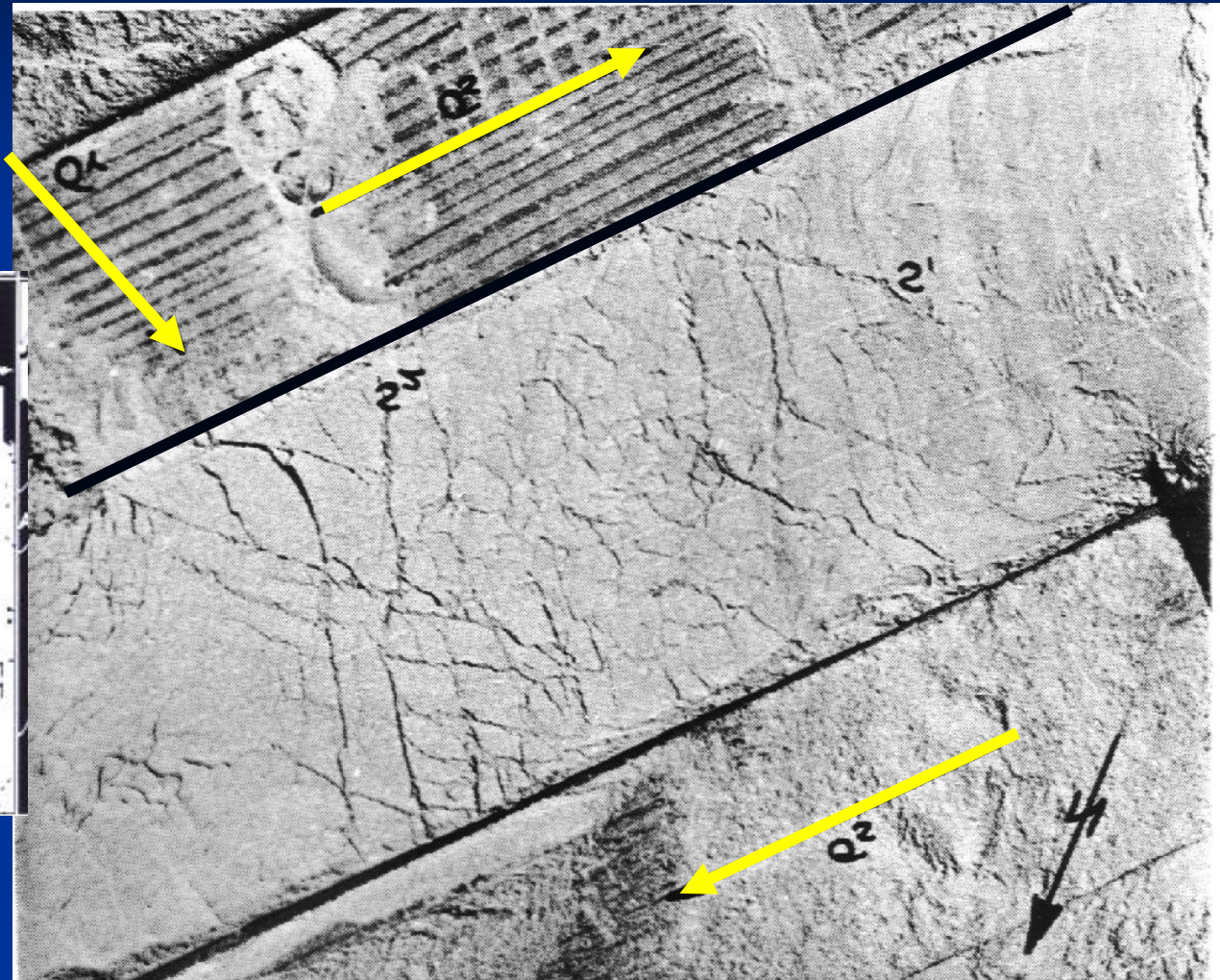
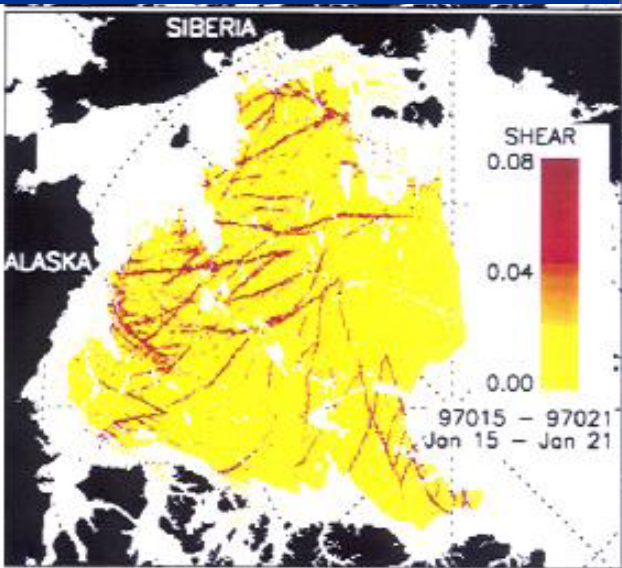


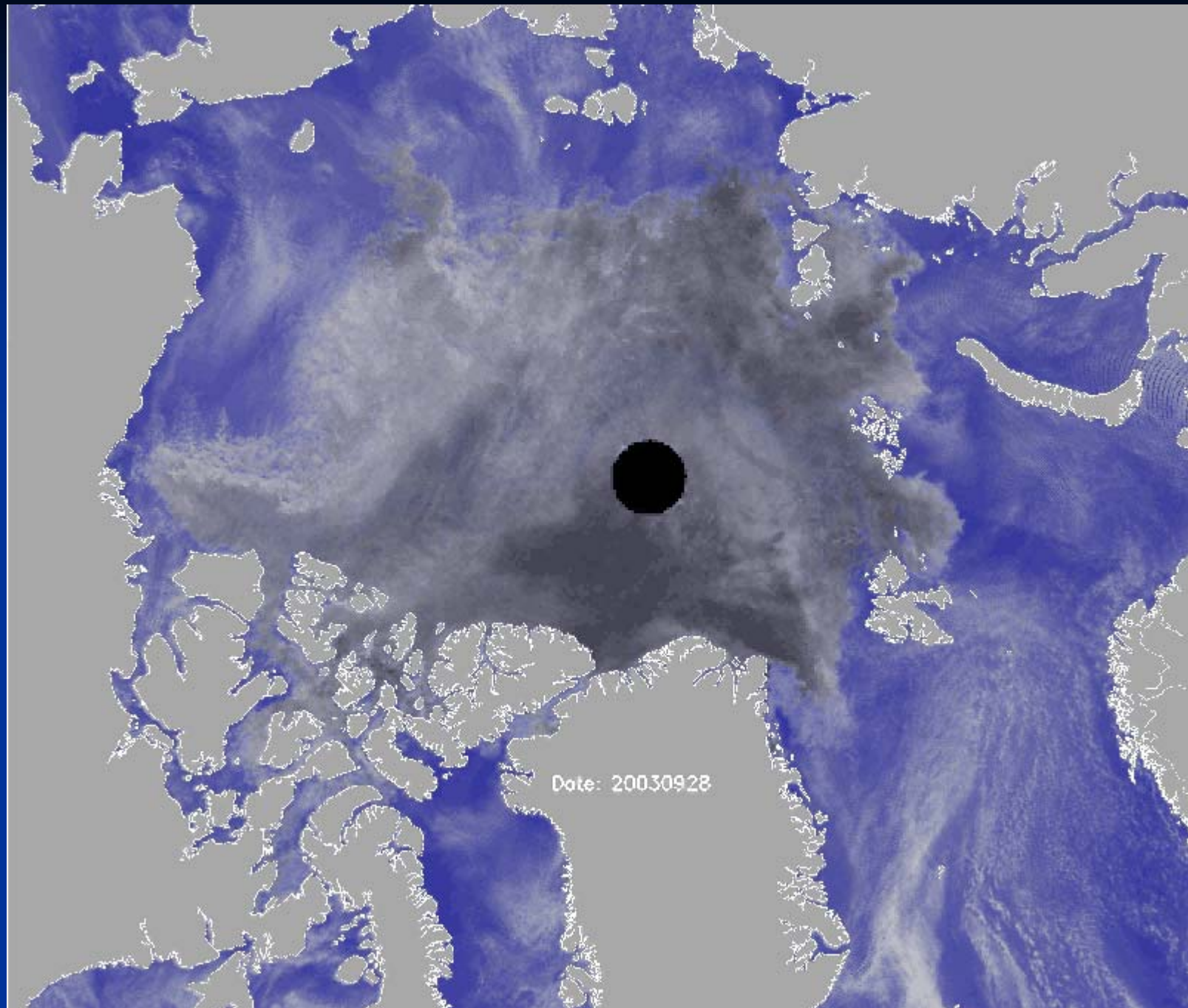
shear



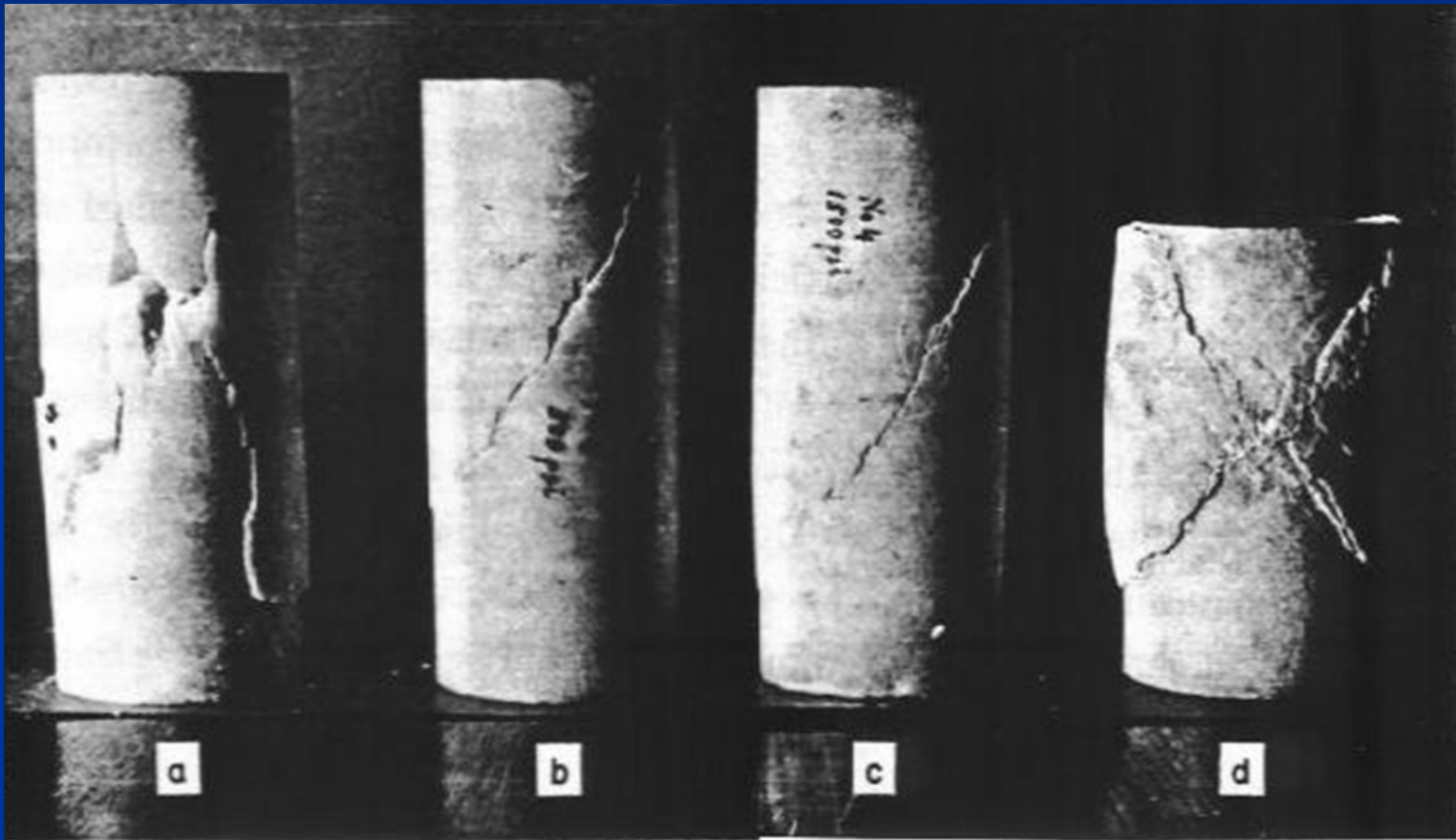
RGPS data, Kwok

Clay under shear deformation

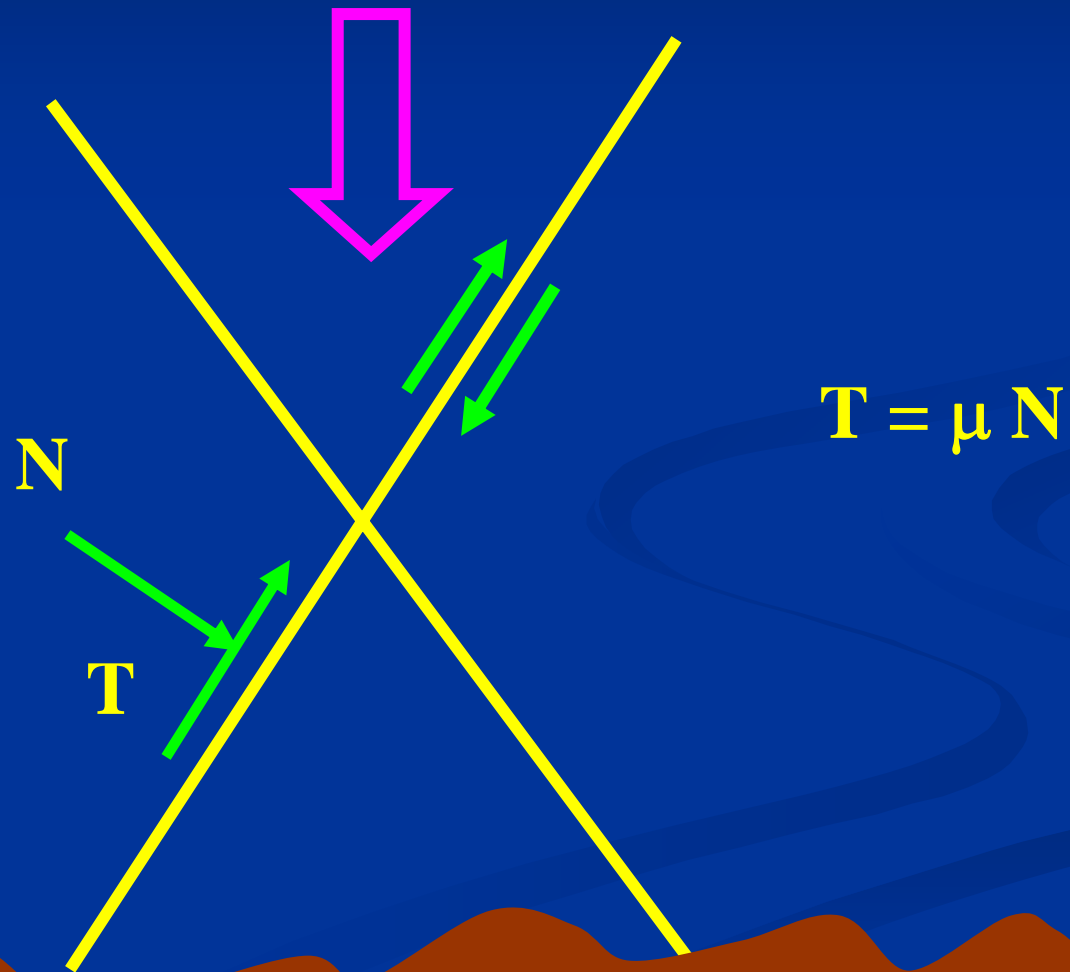




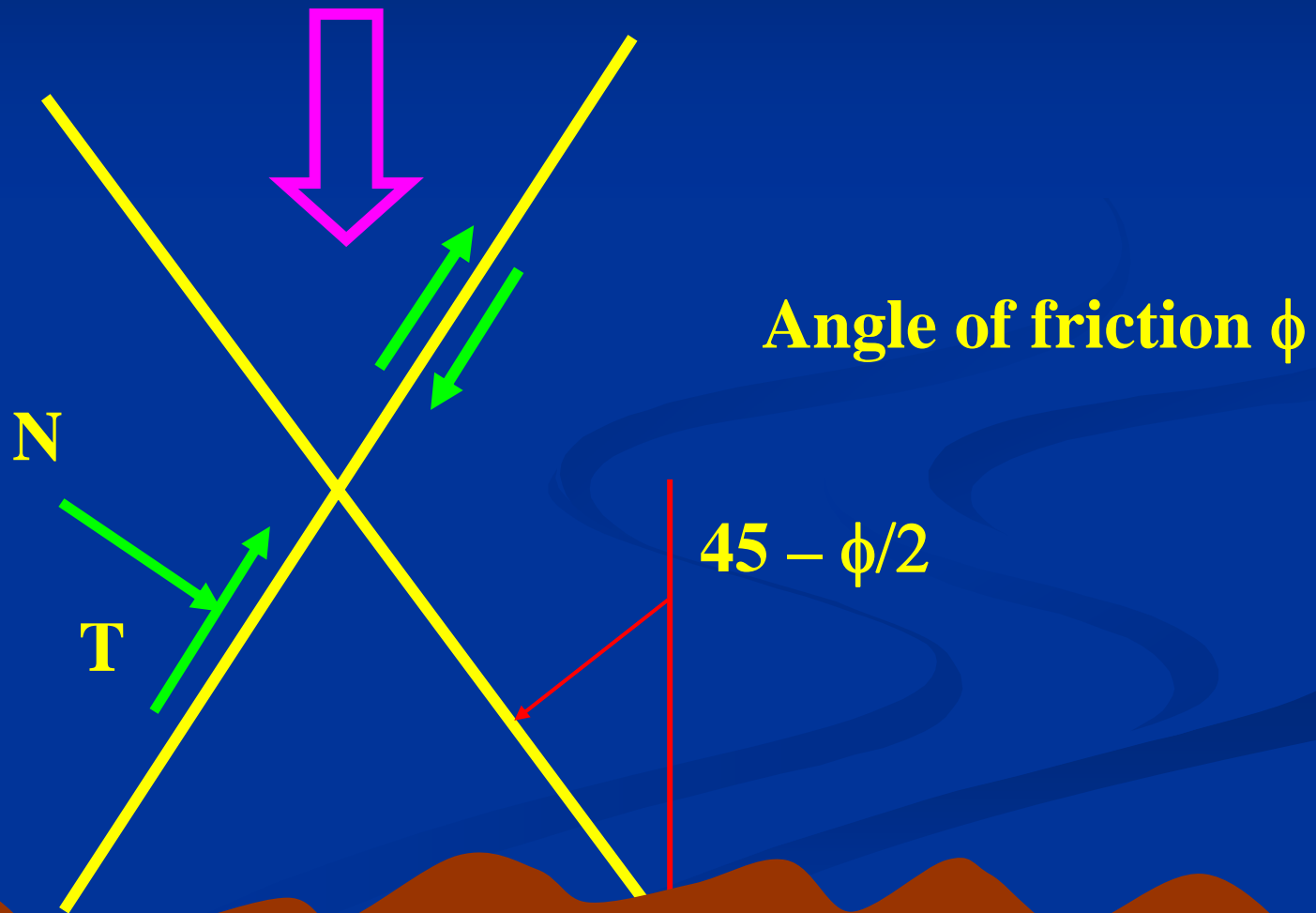
Compression test on concrete



The Model



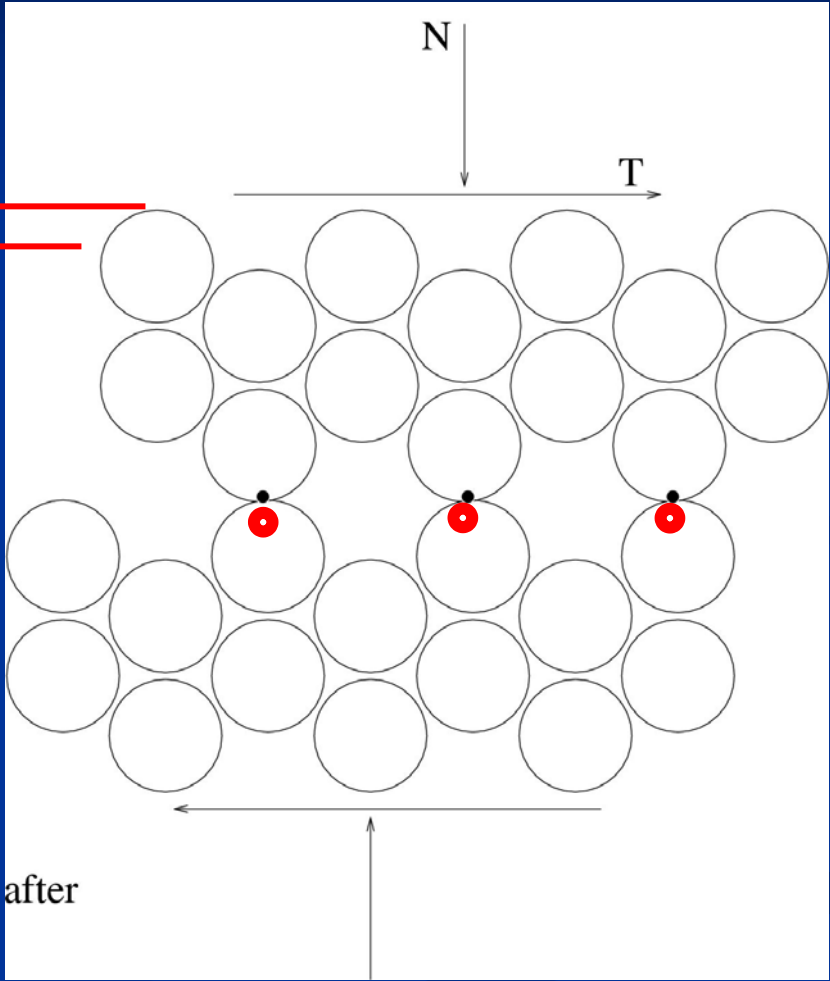
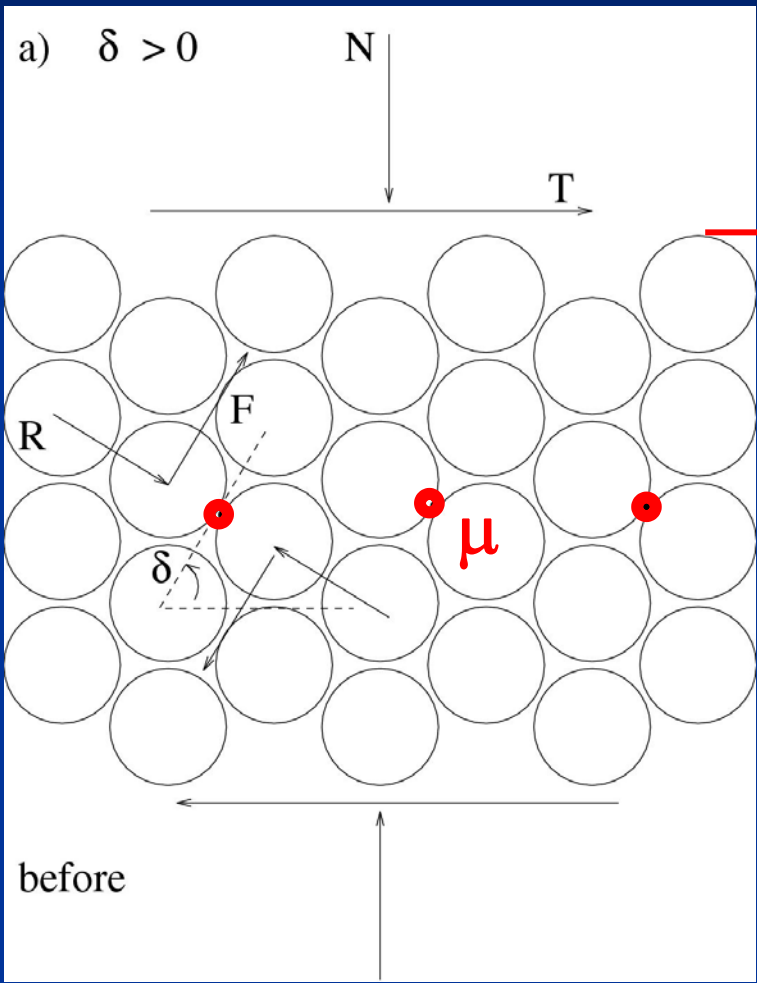
The Model



Summary

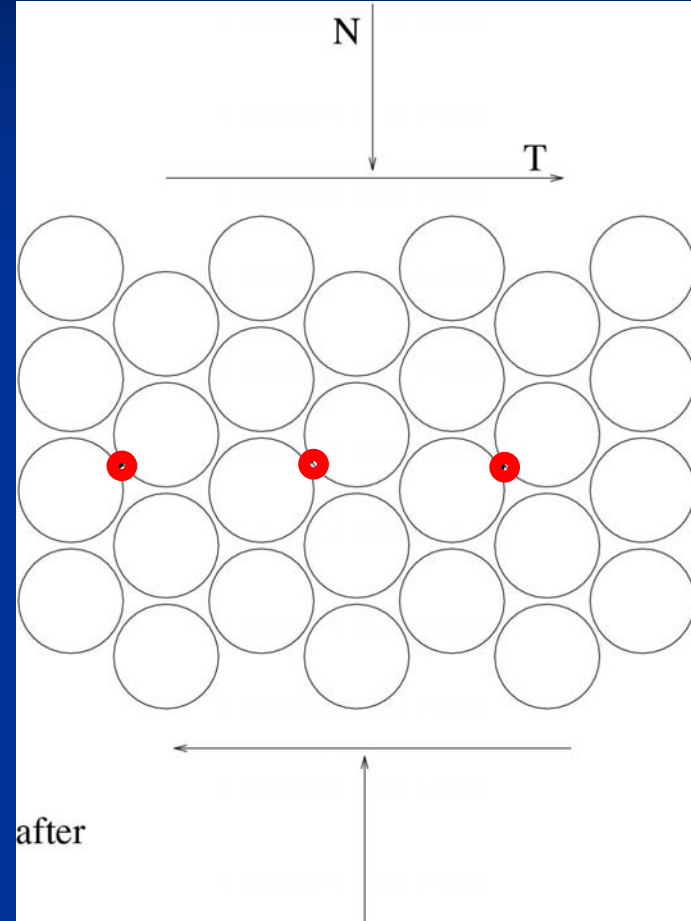
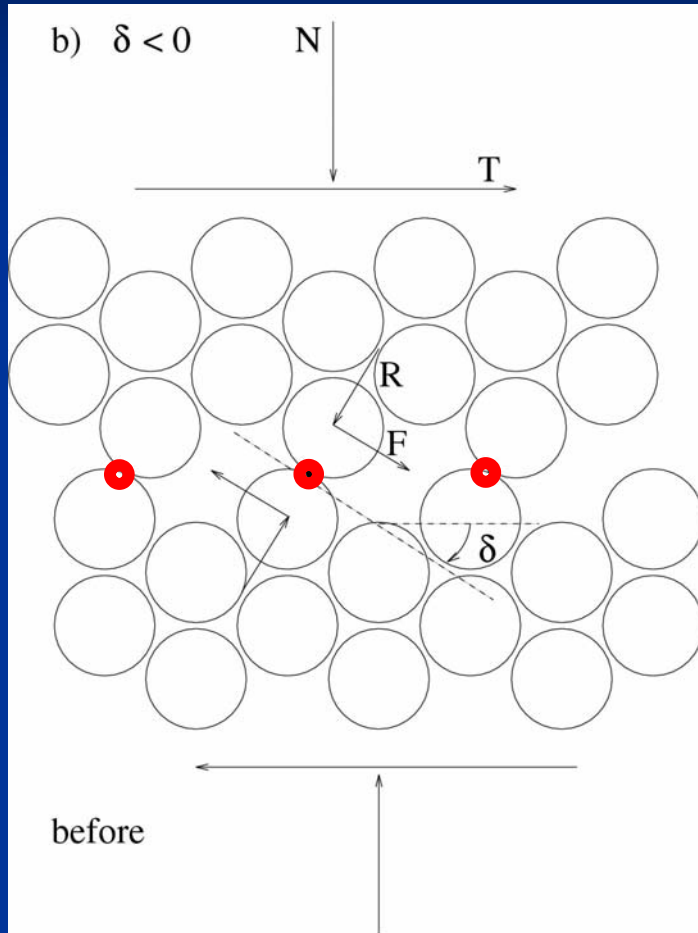
- You identify fracture patterns from satellite pictures
- You calculate the angle of friction from the fracture patterns
- You stick that into the model
- ... and then the model tends to produce to fast ice drifts...

Dilatation δ



$$\phi = \delta + \mu$$

Dilatation



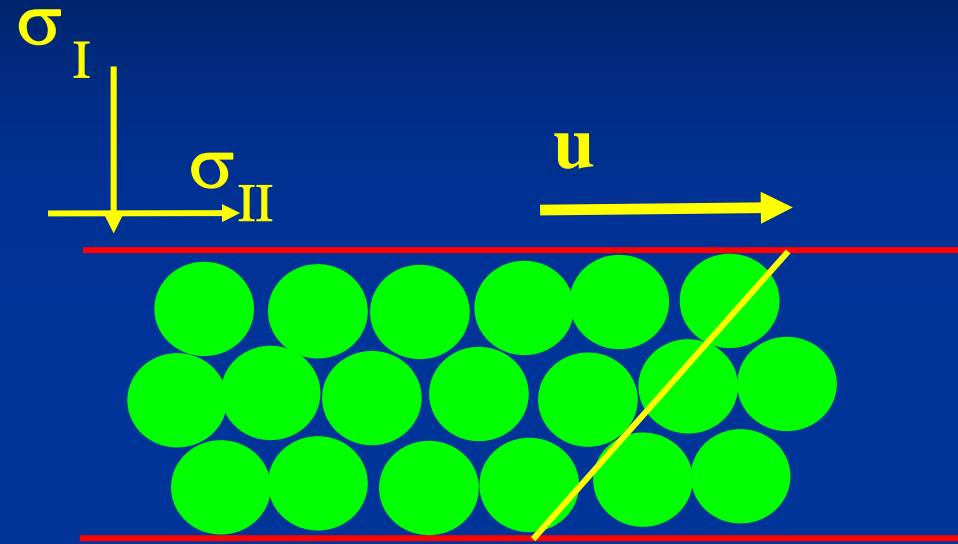
Rheology

$$\sigma_{ij} = -p\delta_{ij} - \eta\dot{\epsilon}_{kk}\delta_{ij} + 2\eta\dot{\epsilon}_{ij}$$

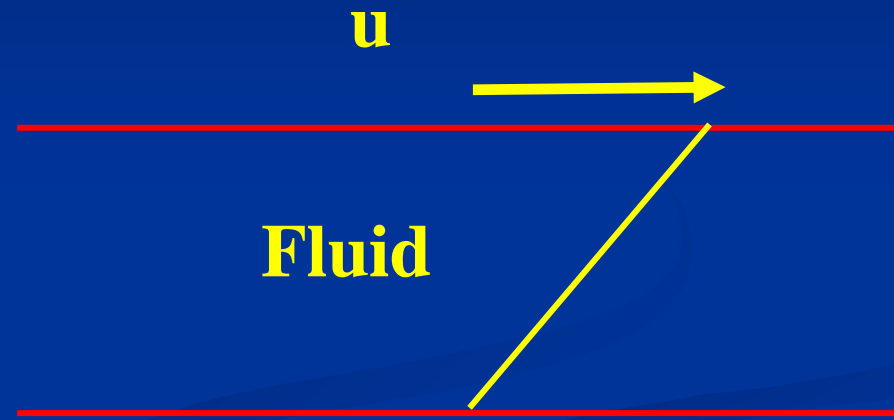
$$\eta = \min\left(\frac{p \sin\phi}{\sqrt{(\dot{\epsilon}_{11} - \dot{\epsilon}_{22})^2 + 4\dot{\epsilon}_{12}^2}}, \eta_{\max}\right)$$

$$0 < p < P^* h \exp[-C(1-A)]$$

Or... in 1-D



$$\sigma_{II} = \mu \frac{\sigma_I}{du/dy}$$

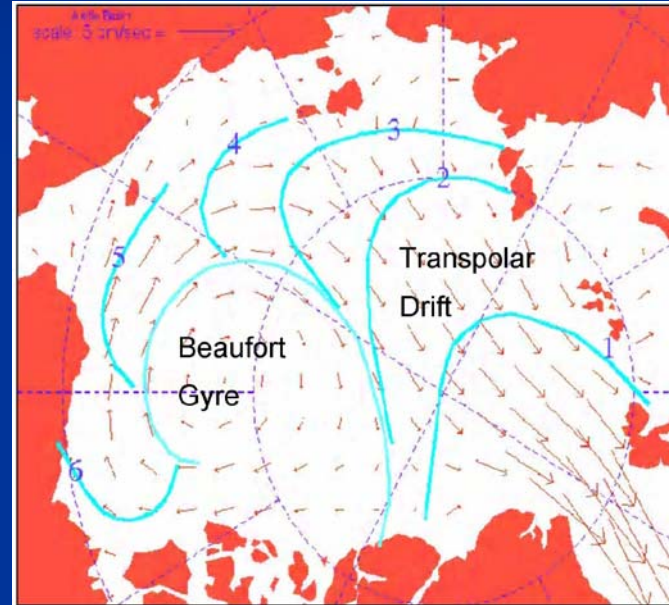
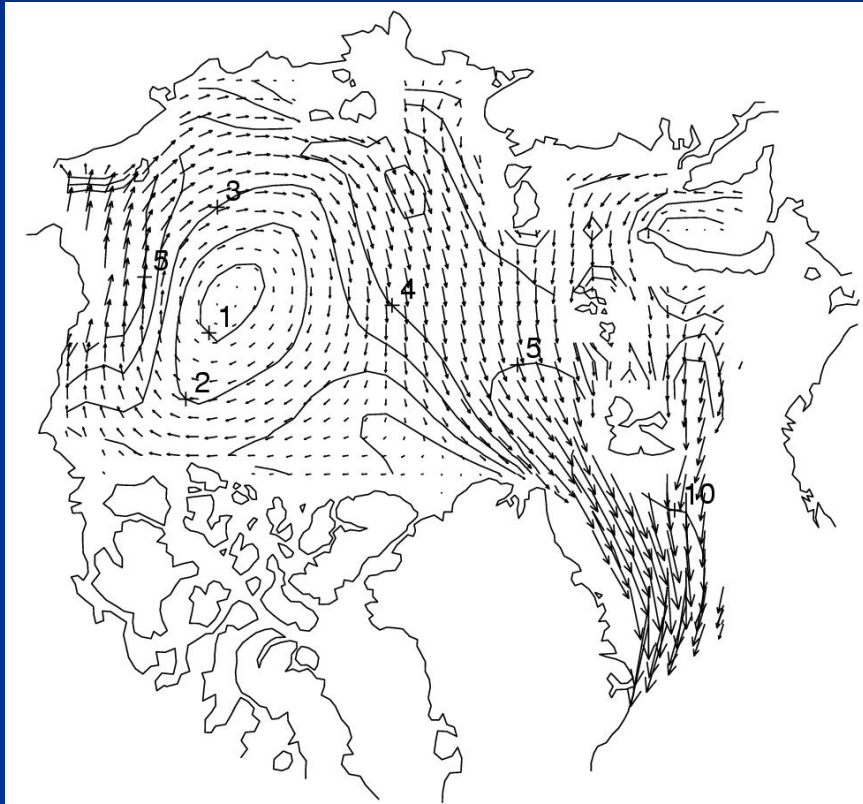


$$\tau = \mu \frac{du}{dy}$$

Comparison with data?

- Sea ice drift
- Sea ice thickness/extent
- Shear lines

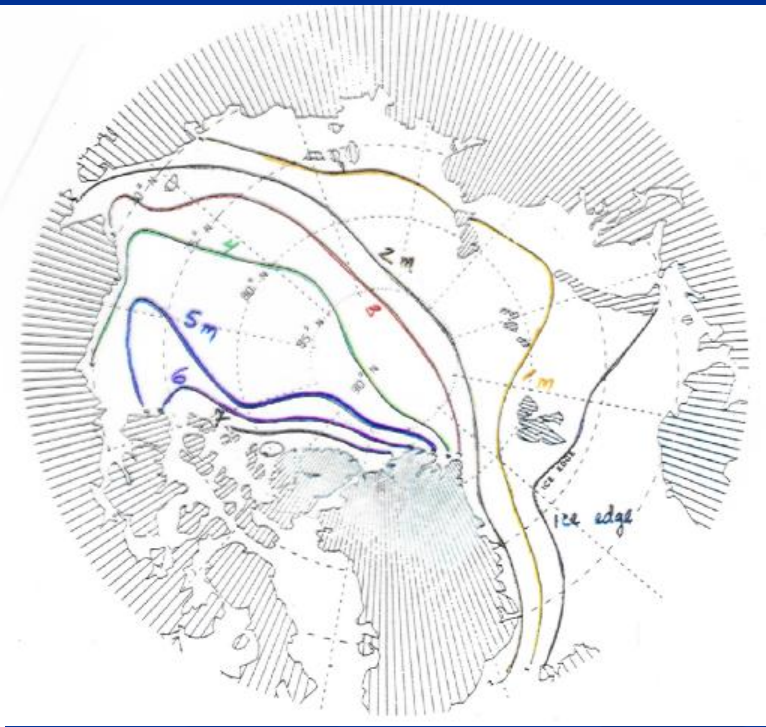
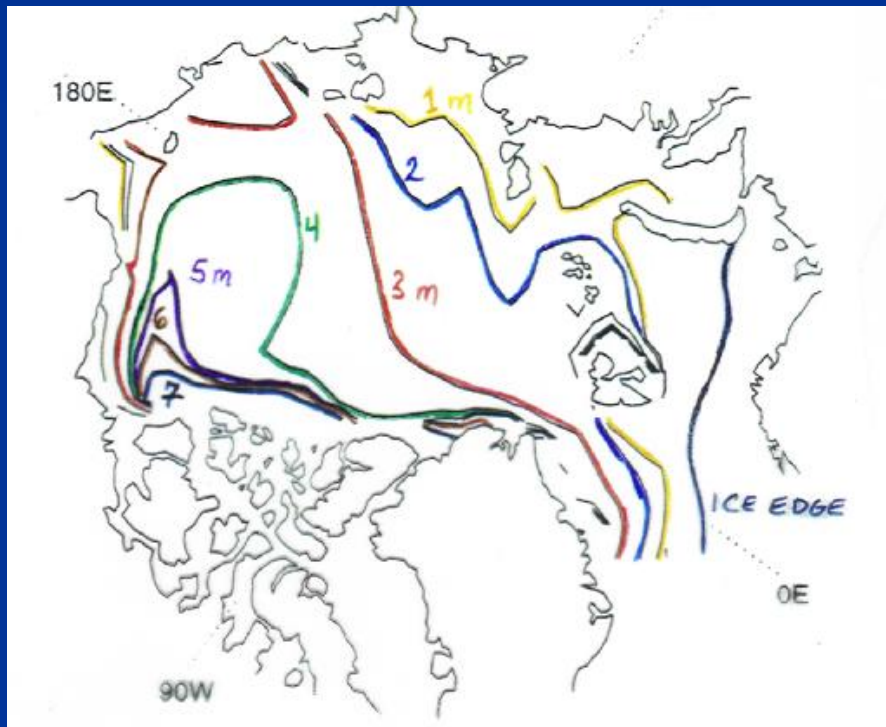
Mean sea ice drift



Sea ice thickness

simulated

observed

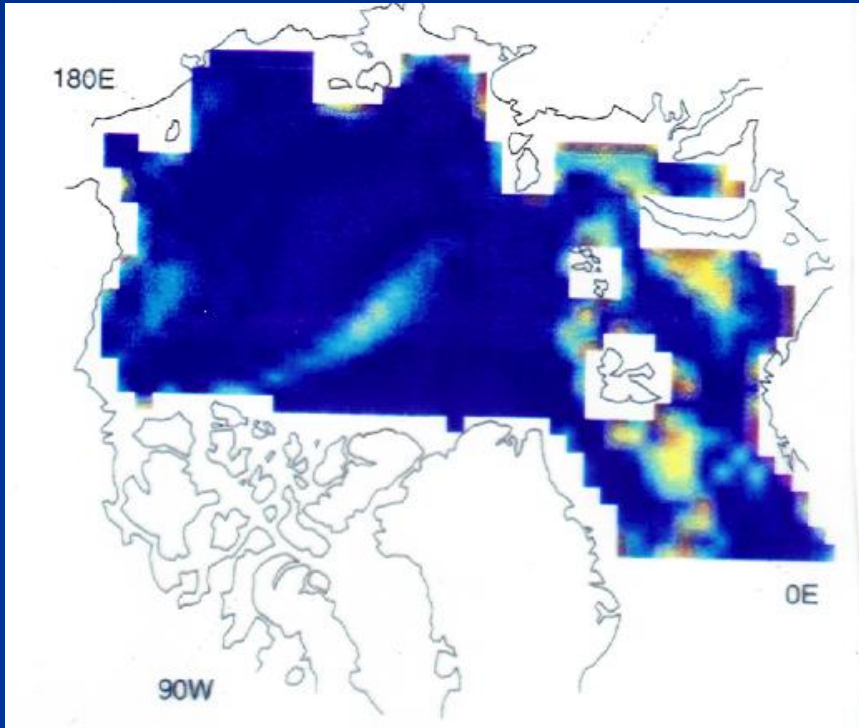


Tremblay Mysak 1997

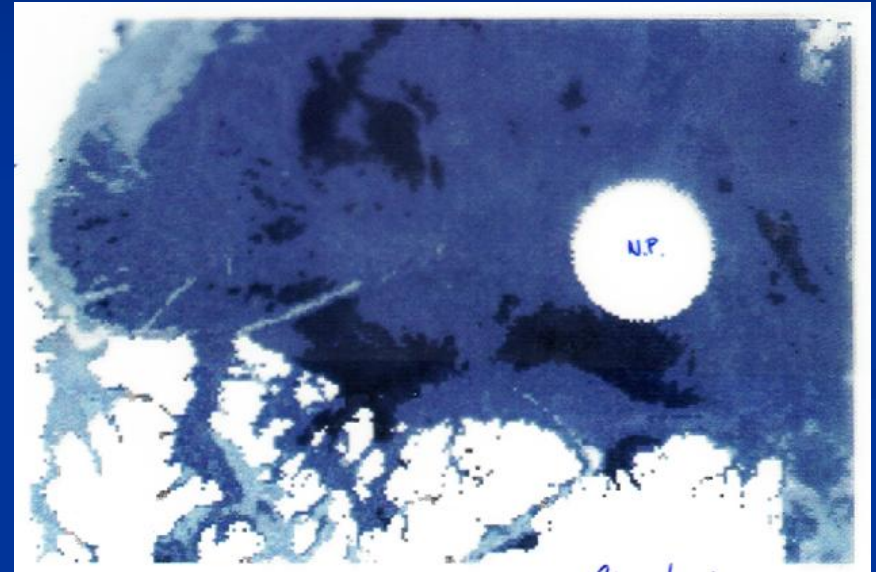
Bourke Garrett, 1986

Shear lines

simulated



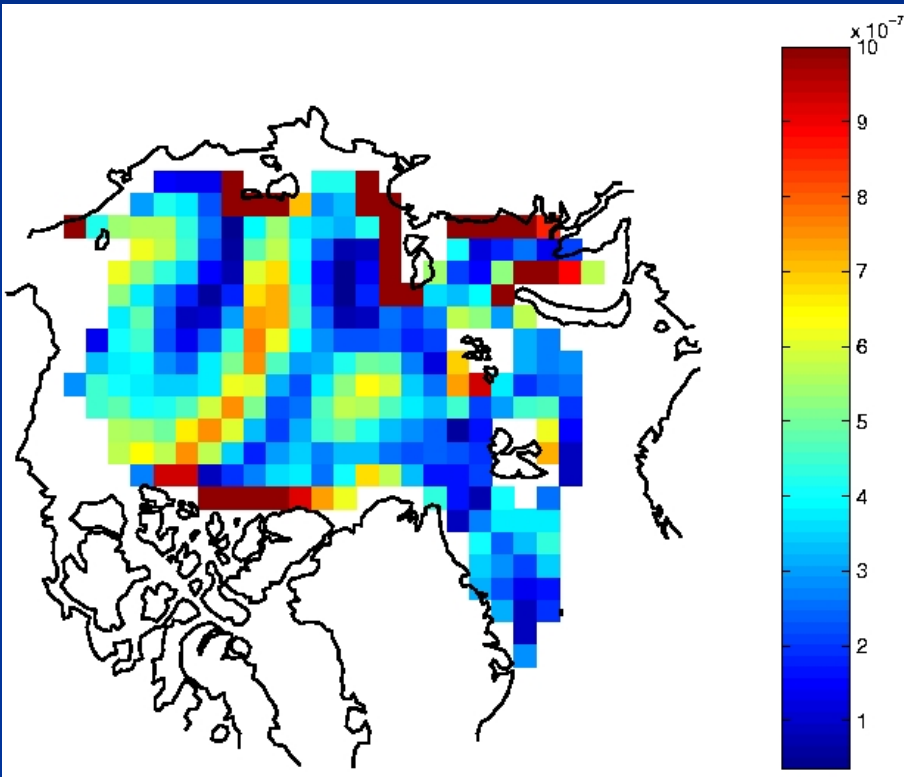
observed



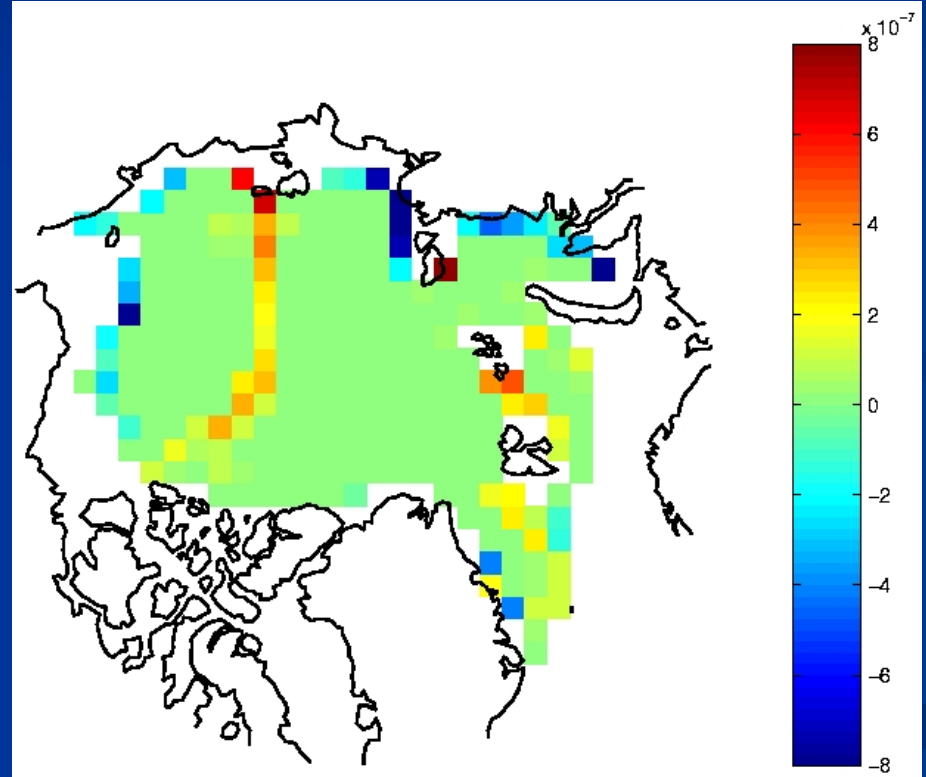
Tremblay Mysak 1997

SSMI passive microwave

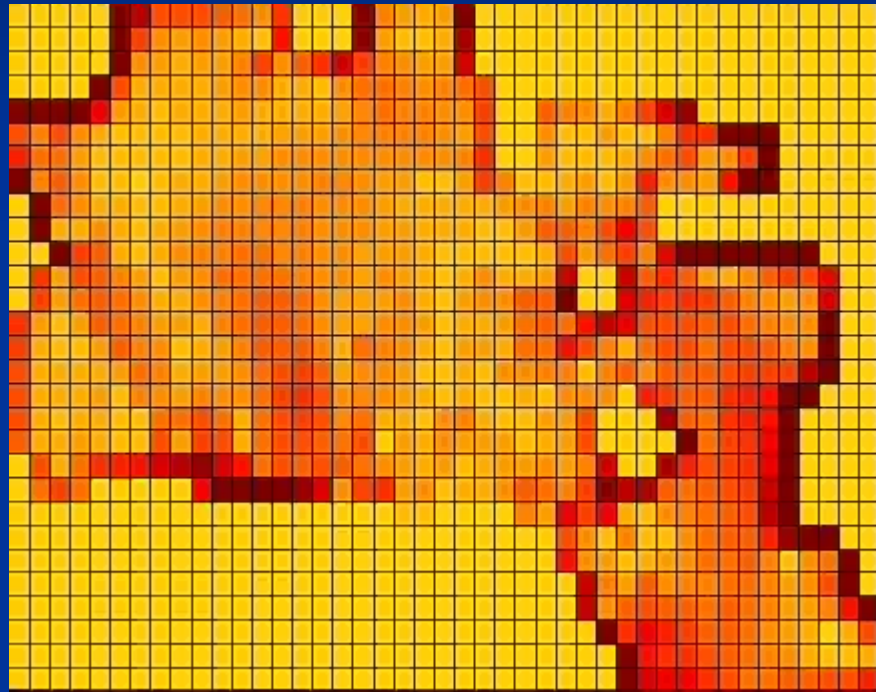
Shear lines



Shear



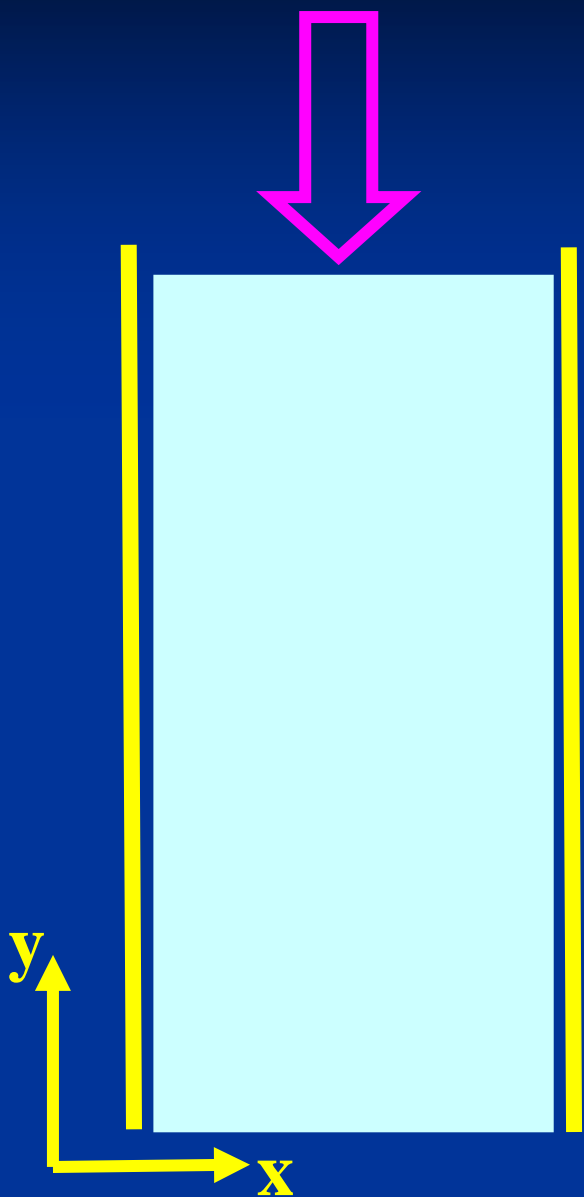
Divergence



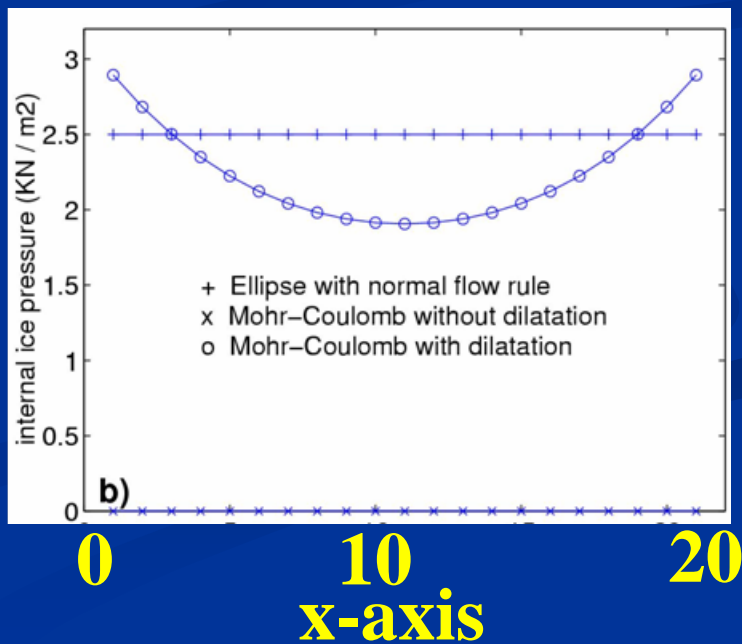
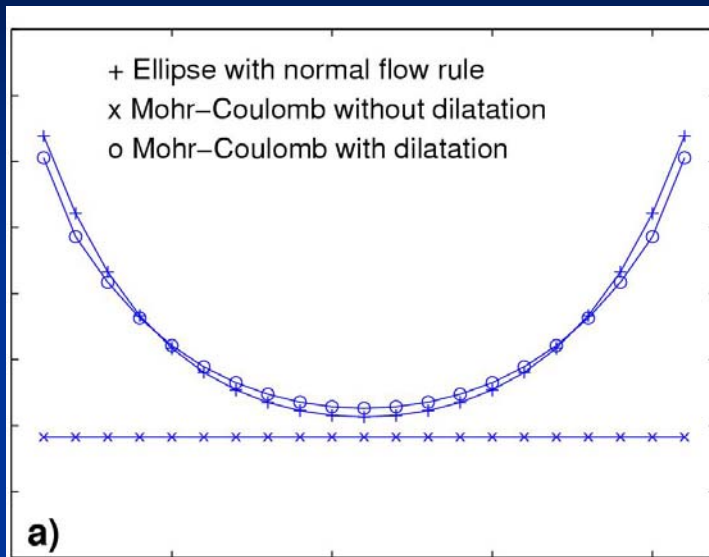
Dilatation effects?

- Flow down a channel
- Arctic simulation

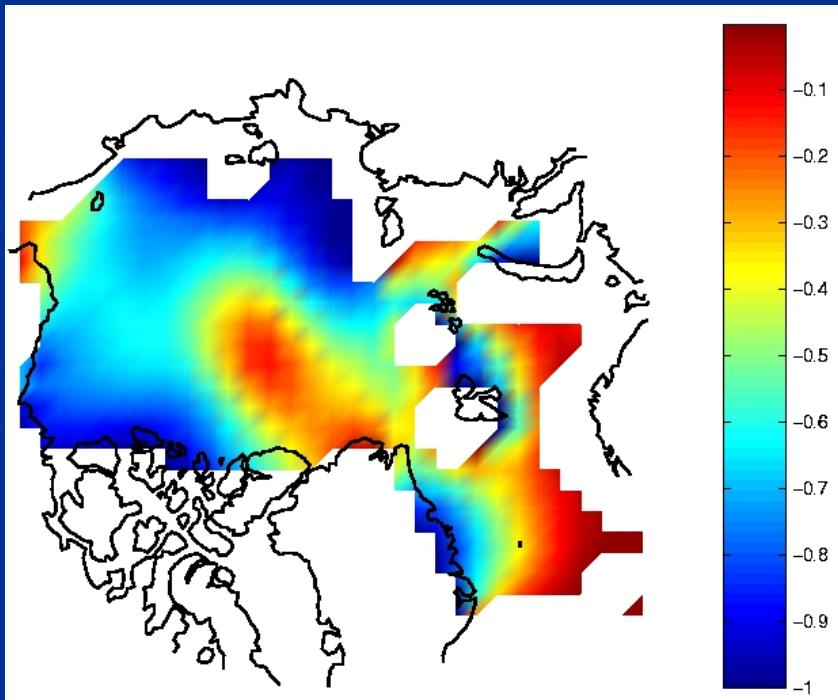
Results



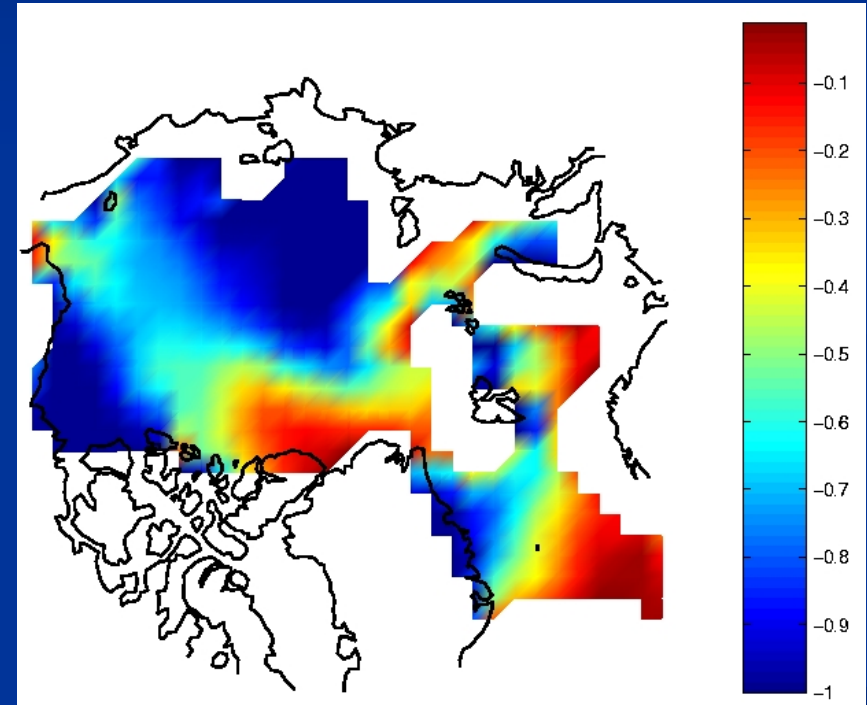
U_{ice} (cm/s) -5
-12
P (kN/m²)



Internal Ice Pressure (normalized)



no dilatation

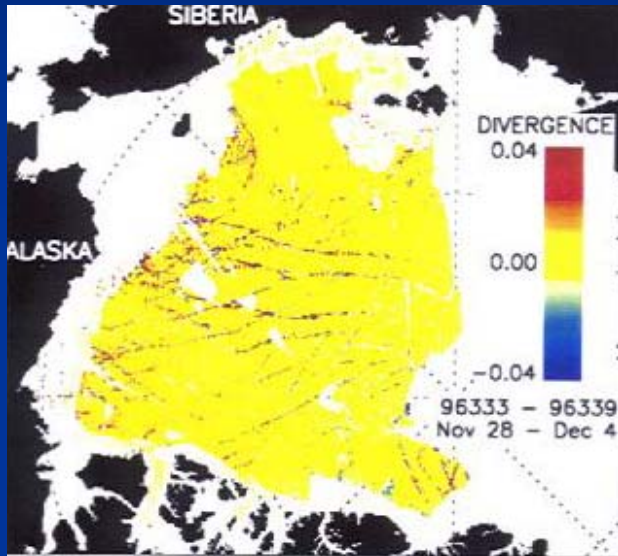


dilatation

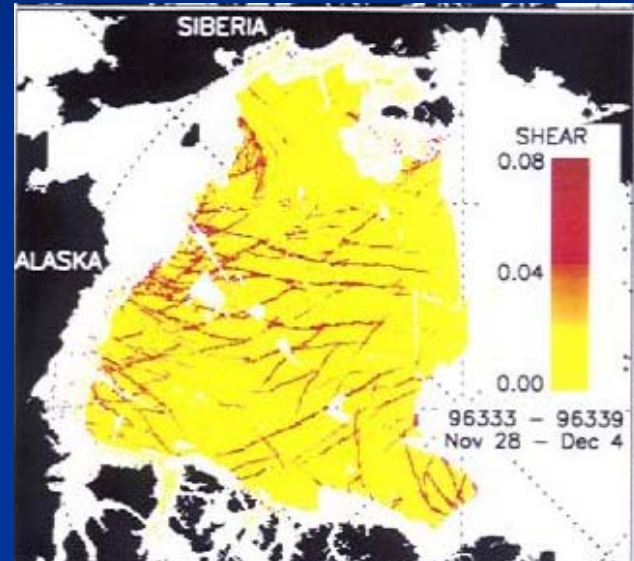
Plastic OR Viscous?

- Plastic OR Viscous?
- Are model iterated until convergence?

Sea ice deformation

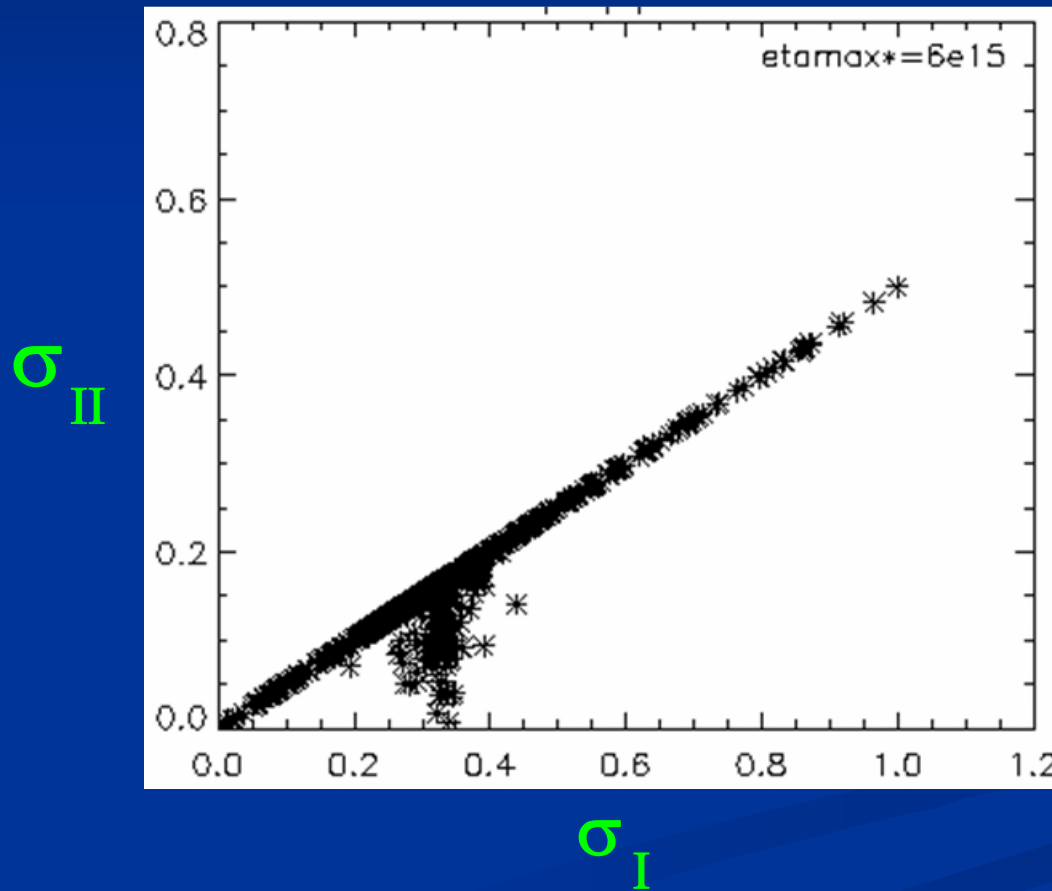


divergence



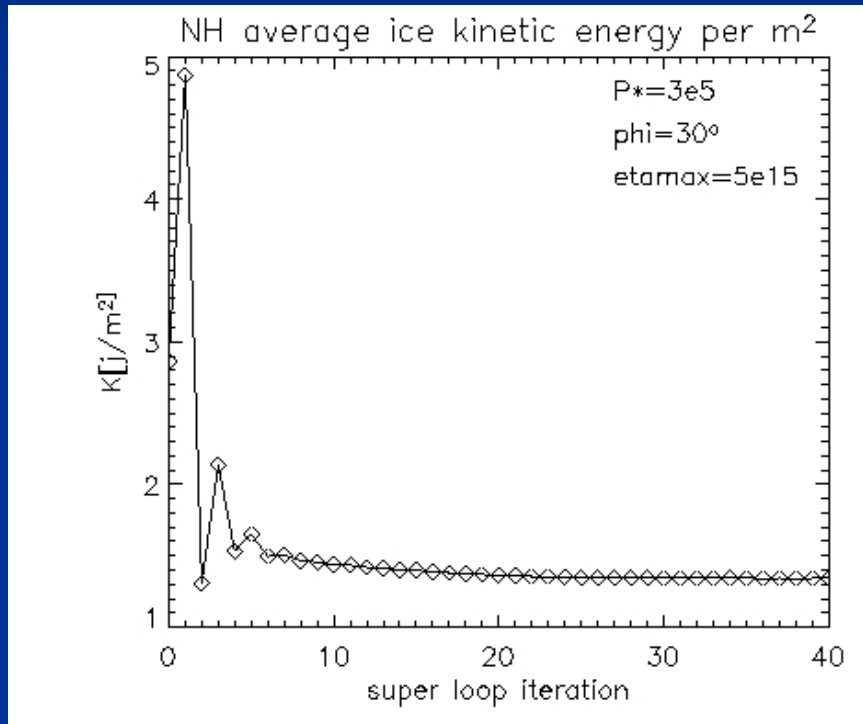
shear

Stress state on the yield curve

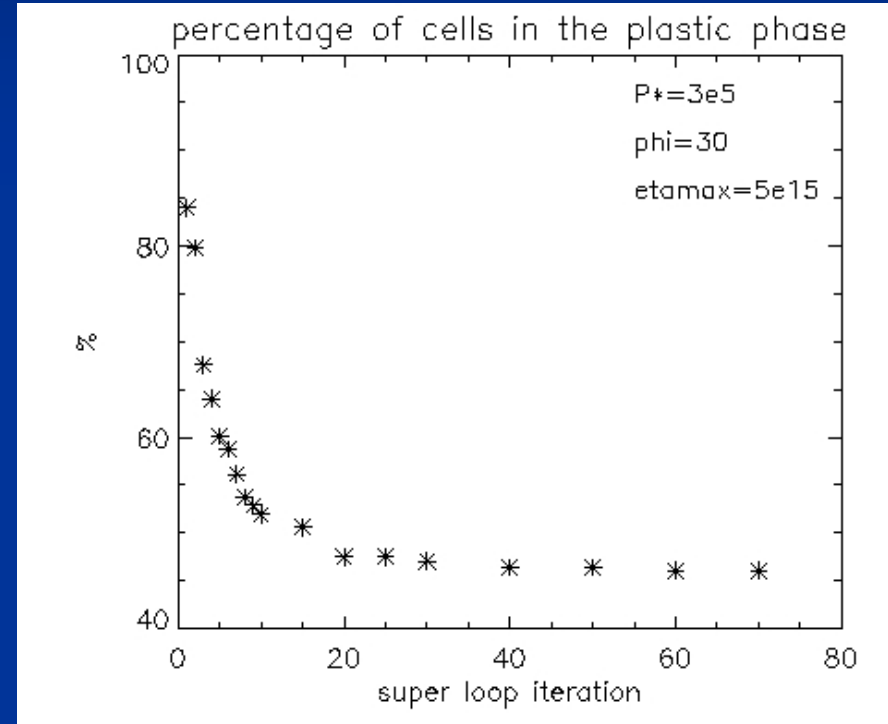


Model integration

Kinetic energy



% plastic phase

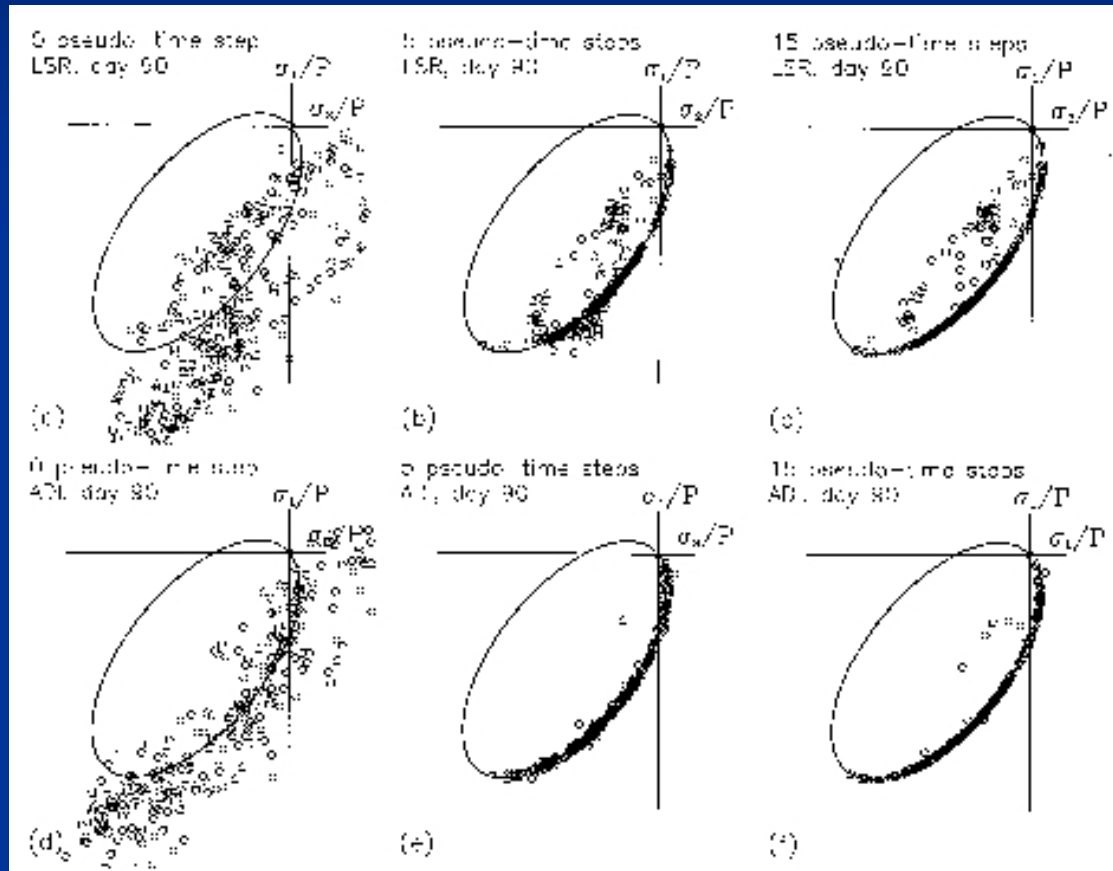


Super loop iteration

Stress states

Line
Successive
Relaxation

Alternating
Direction
Implicit



0

5

15

Pseudo time step

Zhang Rothrock, 2000

Questions

- Should modeler use a larger P^* in their sea ice simulation ? This would lead to more points in the viscous regime as observed...
- What about the mean kinetic energy of the pack then?
- Since most of stress state are observed to be “non-plastic”, should we take more care about how we model the low deformation state? (viscous? elastic?)



Evidence for a higher P^*

- From Model:

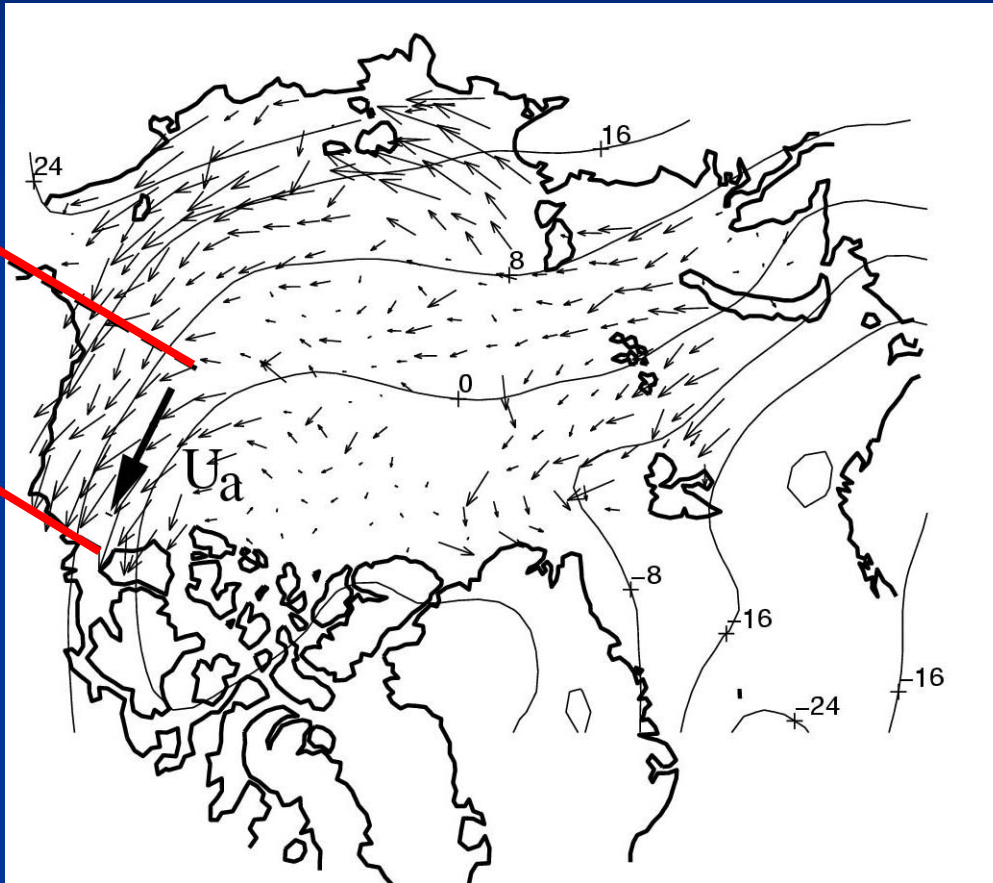
- Kreyscher et al, 1998 $\rightarrow P^* = 15,000 \text{ kN/m}^2$

- Hibler Walsh 1982 $\rightarrow P^* = 27,000 \text{ kN/m}^2$

- From Observations:

Pstar

from satellite/NCEP data



$$F = \tau_a L$$

$$P^*h > \text{or} < F$$

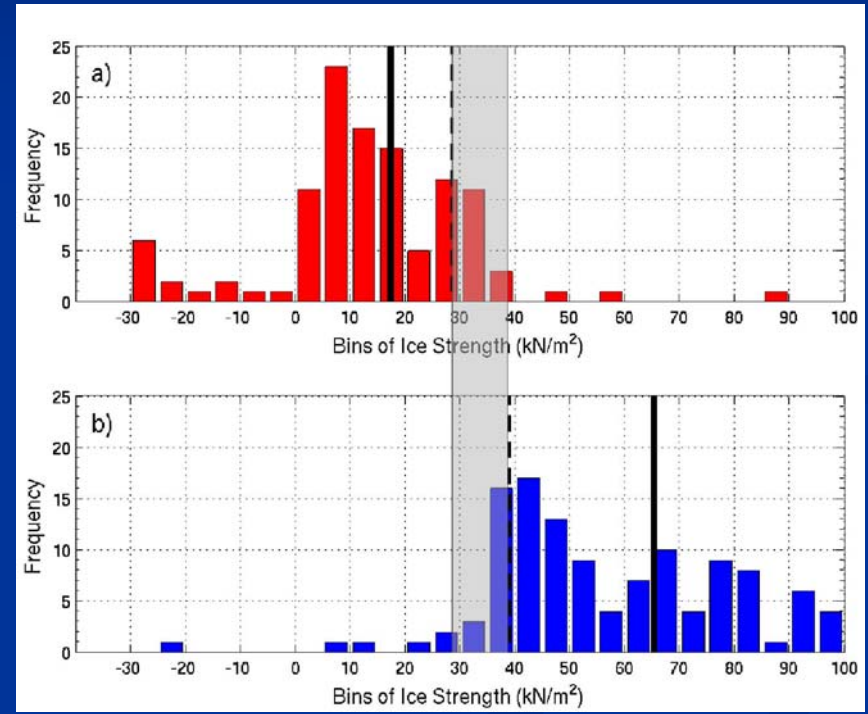
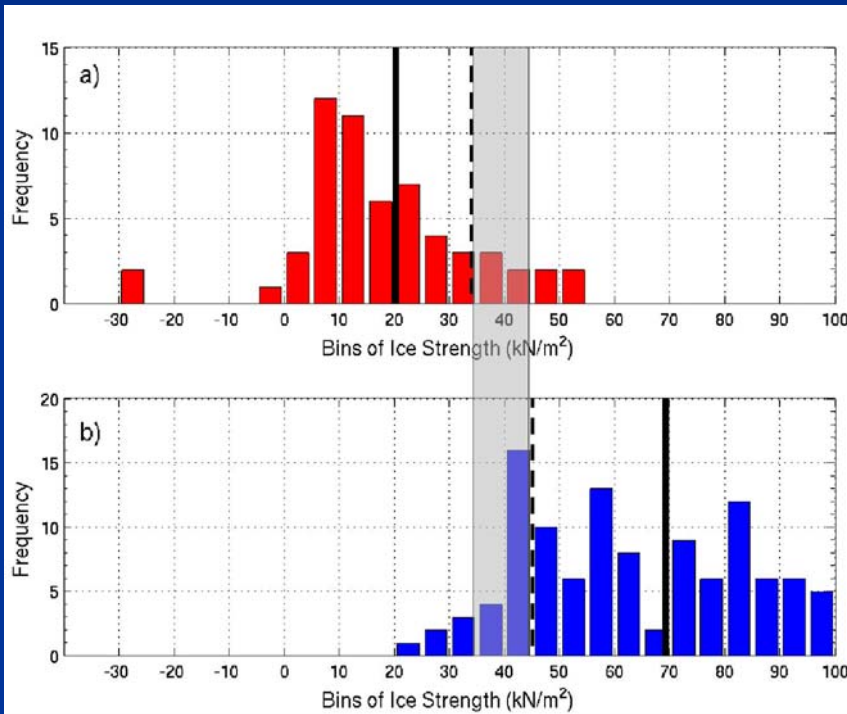
Ice strength estimates

P^*

1992-93

1996-97

frequency



-30

0

100

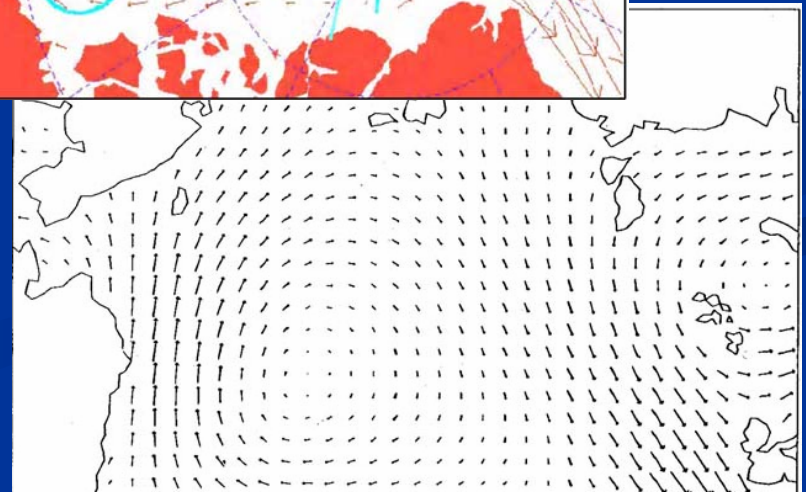
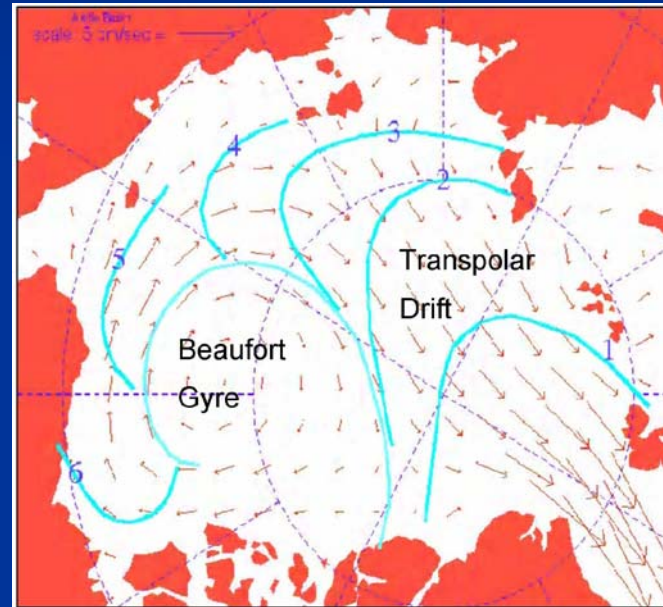
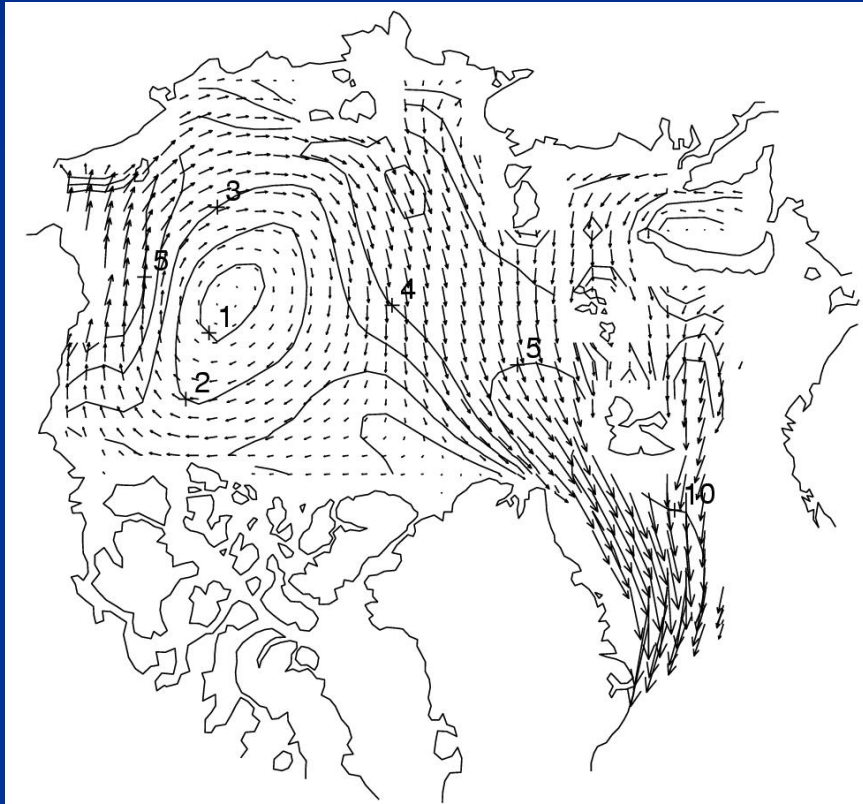
-30

0

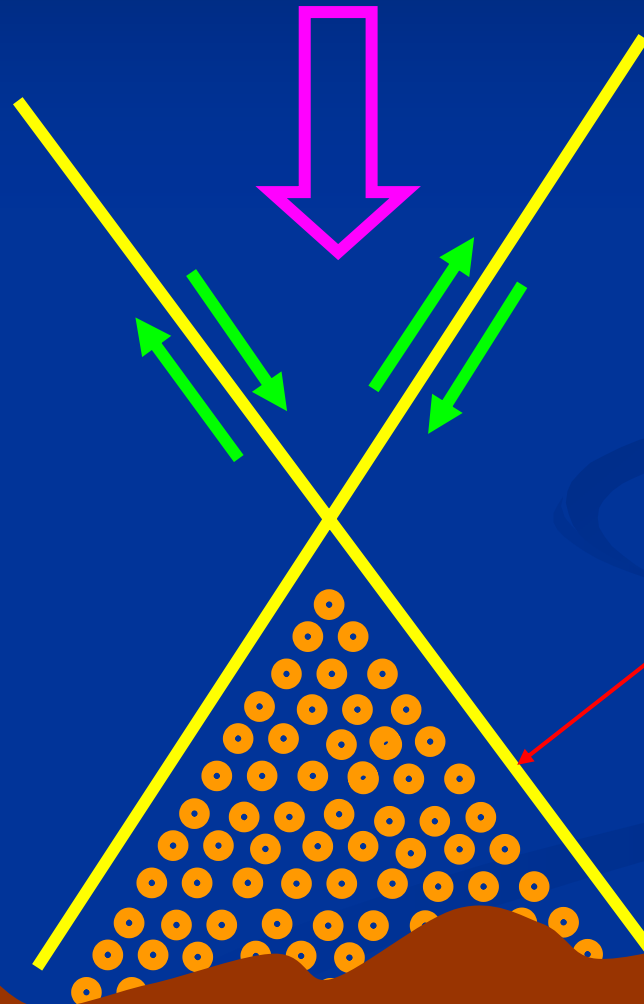
100

P^* (kN/m²)

Mean sea ice drift



The Model



Angle of friction ϕ

$$45 - \phi/2$$