Models of the Probability Distribution of Sea-Surface Wind Speeds

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Motivation



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Motivation

Characterisation of wind speed pdfs



Models of the Probability Distribution of Sea-Surface Wind Speeds - p. 2/44

Motivation

- Characterisation of wind speed pdfs
- Empirical models of wind speed pdfs



- Motivation
- Characterisation of wind speed pdfs
- Empirical models of wind speed pdfs
- Mechanistic Model



- Motivation
- Characterisation of wind speed pdfs
- Empirical models of wind speed pdfs
- Mechanistic Model
- Conclusions



 Ocean and atmosphere interact through respective boundary layers, exchanging



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momentum



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 - momentum
 - energy



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- Sea surface winds ("eddy averaged") are an important determinant of turbulence in both boundary layers



- Ocean and atmosphere interact through respective boundary layers, exchanging
 - momentum
 - energy
 - freshwater
 - gases & aerosols
- Sea surface winds ("eddy averaged") are an important determinant of turbulence in both boundary layers
- Turbulence feeds back on surface winds, primarily through surface momentum flux



Motivation: GCMs

Calculation of air/sea fluxes in GCMs requires:



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Calculation of air/sea fluxes in GCMs requires: timestep/gridbox averaged fluxes



 Calculation of air/sea fluxes in GCMs requires: timestep/gridbox averaged fluxes in terms of





Bulk formulae for air/sea fluxes generally have nonlinear dependence on sea-surface wind



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- Averaged surface winds not enough for calculating averaged surface fluxes



- Bulk formulae for air/sea fluxes generally have nonlinear dependence on sea-surface wind
- Averaged surface winds not enough for calculating averaged surface fluxes
- Calculation of averaged fluxes requires full surface wind pdf



Further Complications:



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- Further Complications:
- 1. surface fluxes require surface wind speeds



- Further Complications:
- surface fluxes require surface *wind speeds* **but**



- Further Complications:
- surface fluxes require surface *wind speeds* **but** GCMs give space-time averaged *vector winds*



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⇒ Motivates development of parameterisation of wind speed pdf in terms of vector winds



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- 2. Wind speed pdf arises from vector wind pdf through nonlinear coordinate transformation



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GCMs give space-time averaged vector winds

- \Rightarrow Motivates development of parameterisation of wind speed pdf in terms of vector winds
- 2. Wind speed pdf arises from vector wind pdf through nonlinear coordinate transformation
- ⇒ higher order vector wind moments may affect lower order wind speed moments



Skewness and Kurtosis

Skewness: measure of asymmetry of pdf

skew
$$(x) = \left\langle \left(\frac{x - \langle x \rangle}{\operatorname{std}(x)} \right)^3 \right\rangle$$



Skewness and Kurtosis

Skewness: measure of asymmetry of pdf

$$\operatorname{skew}(x) = \left\langle \left(\frac{x - \langle x \rangle}{\operatorname{std}(x)} \right)^3 \right\rangle$$

Kurtosis: measure of <u>flatness</u> of pdf

skew
$$(x) = \left\langle \left(\frac{x - \langle x \rangle}{\operatorname{std}(x)} \right)^4 \right\rangle - 3$$



Sea-Surface Winds: Notation

Notation:

- **u** vector wind (u, v)
- u along mean wind component
- v cross mean wind component

w wind speed
$$(u^2 + v^2)^{1/2}$$



 Will consider 6-hourly 10m ocean winds from NCEP/NCAR Reanalysis (1948-2005)



Models of the Probability Distribution of Sea-Surface Wind Speeds - p. 8/44

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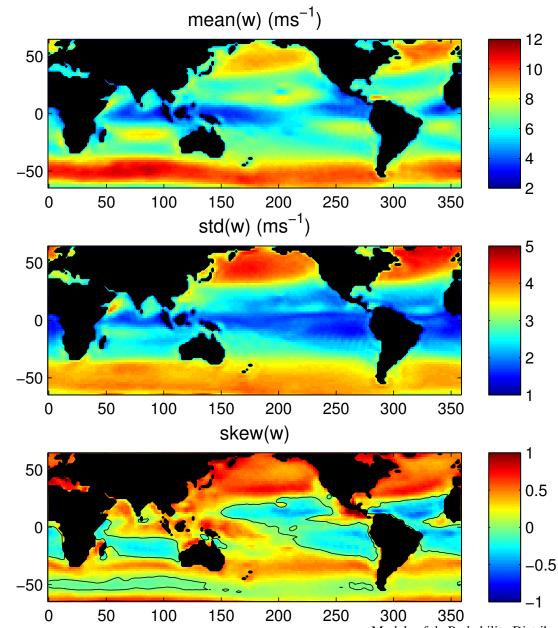
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Sea-Surface Wind Speeds: Moments





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The pdf of wind speed w has traditionally been represented by 2-parameter Weibull distribution:

$$p_w(w) = \begin{cases} \frac{b}{a} \left(\frac{w}{a}\right)^{b-1} \exp\left(-\left(\frac{w}{a}\right)^b\right) & w > 0\\ 0 & w < 0 \end{cases}$$



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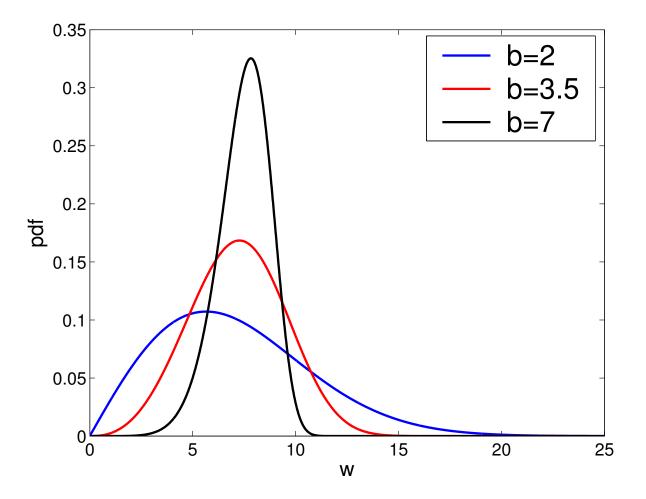
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- a is the <u>scale</u> parameter (pdf centre)
- b is the shape parameter (pdf tilt)
- $p_w(w)$ is unimodal

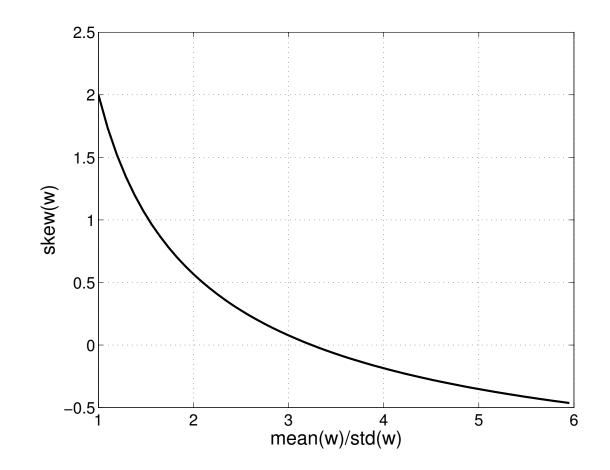


• Weibull pdfs for a = 8





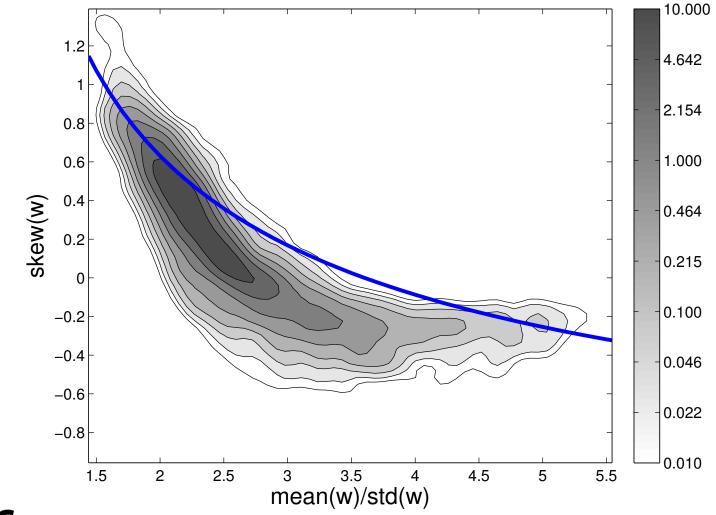
For a Weibull distribution, skew(w) is a decreasing function of mean(w)/std(w)





Wind Speed pdfs: Observed

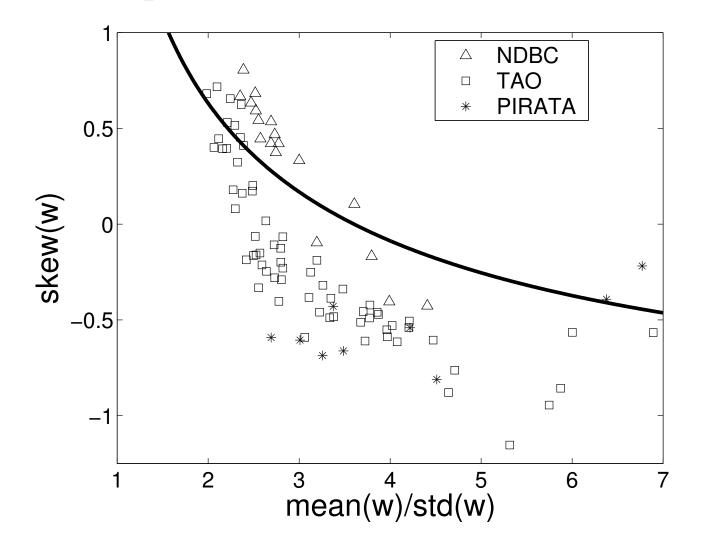
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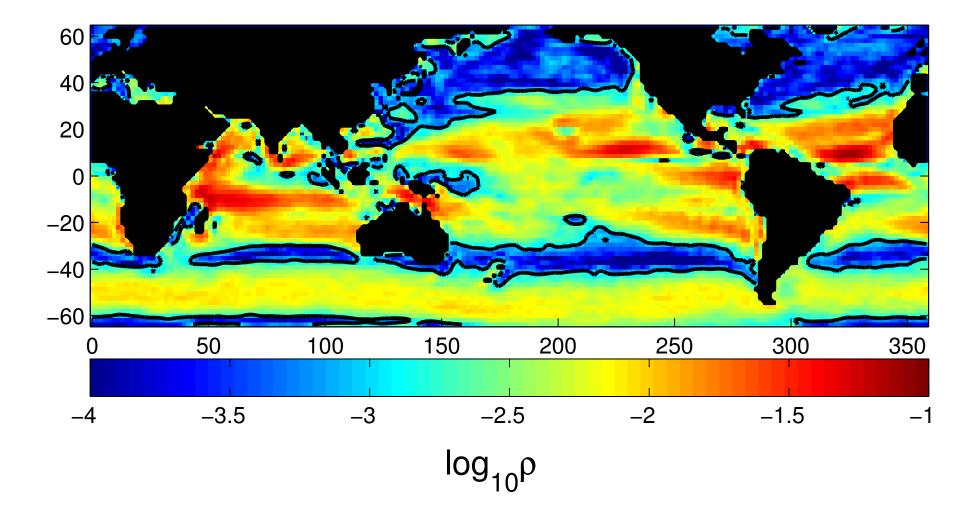
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where:

$$p_w(w)$$
 = observed wind speed pdf
 $q_w(w)$ = best-fit Weibull pdf



Wind Speed pdfs: Relative Entropy





Wind Speed pdfs: Empirical Models

Strategy: systematically construct wind speed pdfs from joint pdf of vector components, $p_{uv}(u, v)$

$$p_w(w) = w \int_{-\pi}^{\pi} p_{uv}(w\cos\theta, w\sin\theta)d\theta$$



Wind Speed pdfs: Empirical Models

Strategy: systematically construct wind speed pdfs from joint pdf of vector components, $p_{uv}(u, v)$

$$p_w(w) = w \int_{-\pi}^{\pi} p_{uv}(w\cos\theta, w\sin\theta)d\theta$$

Approach simplified by assuming u, v independent, so

$$p_{uv}(u,v) = p_u(u)p_v(v)$$



Simplest model assumes isotropic Gaussian fluctuations in vector winds:

$$p_u(u) = \frac{1}{\sqrt{2\pi\sigma^2}} \exp\left(-\frac{(u-\overline{u})^2}{2\sigma^2}\right)$$

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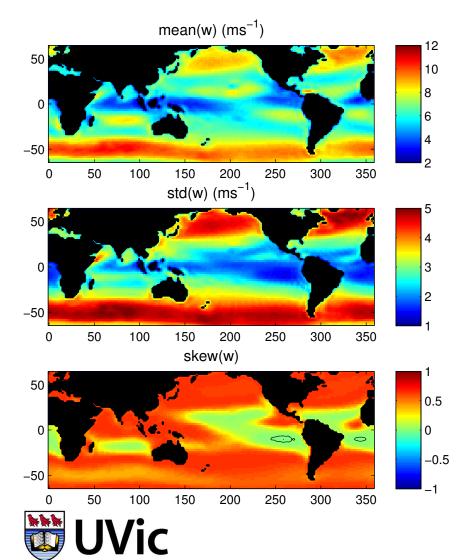
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Integrating over wind direction:

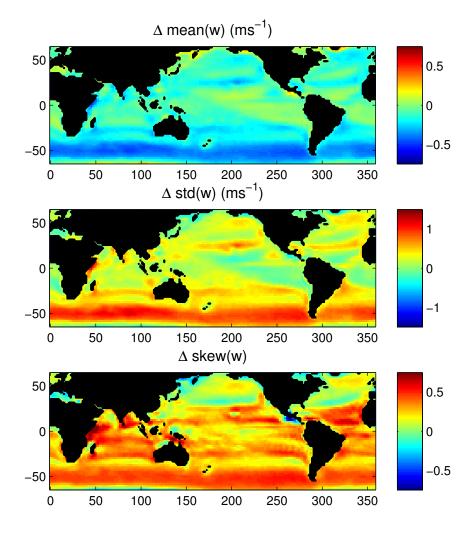
$$p_w(w) = \frac{w}{\sigma^2} \exp\left(-\frac{w^2 + \overline{u}^2}{2\sigma^2}\right) I_0\left(\frac{w\overline{u}}{\sigma^2}\right)$$



Model



Model-Observations



• Isotropic Gaussian model \Rightarrow large biases



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- Relax assumption of isotropy: $\sigma_u = \operatorname{std}(u), \sigma_v = \operatorname{std}(v) \Rightarrow$

$$p_w(w) = \frac{w}{\sigma_u \sigma_v} \exp\left(-\frac{w^2 + \overline{u}^2}{2\sigma_u^2}\right) \times \left\{ I_0\left(\frac{w\overline{u}}{\sigma_u^2}\right) + \sum_{k=1}^{\infty} \left[\frac{w}{\overline{u}}\left(1 - \frac{\sigma_u^2}{\sigma_v^2}\right)\right]^k \frac{\Gamma(k+1/2)}{\sqrt{\pi k!}} I_k\left(\frac{\overline{u}w}{\sigma_u^2}\right) \right\}$$



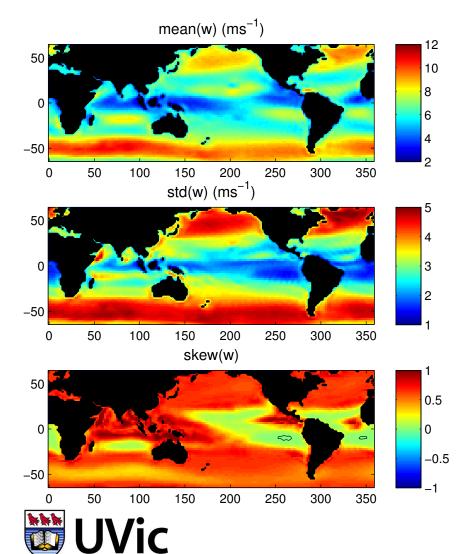
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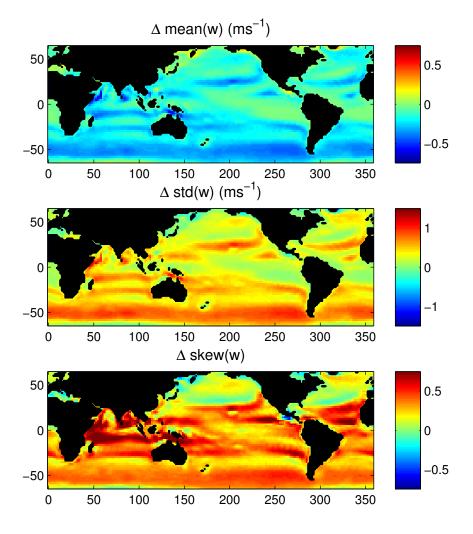
Actually makes approximation somewhat worse



Model



Model-Observations



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Gaussian vector winds unable to model skewness of speed pdfs in tropics and Southern Ocean



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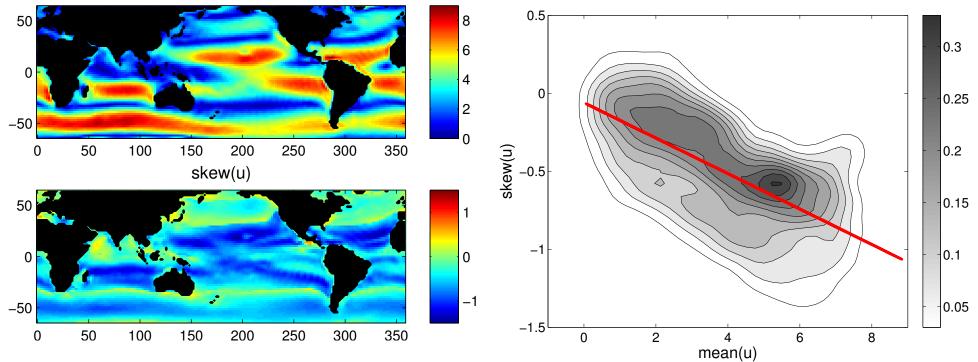
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- Coupling of moments follows from nonlinear dependence of surface drag on surface winds
- Symmetric fluctuations in forcing ⇒ asymmetric response, skewed toward rest



Vector Wind Moments



 $mean(u) (ms^{-1})$

 $skew(u) \simeq (-0.11 \text{ sm}^{-1}) mean(u) - 0.06$



 Need skewed pdf to model u (still assuming v is Gaussian)



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- Question is: how important are details of skewed pdf of u for pdf of w?



Bigaussian pdf

Two half-Gaussians:

$$p_u(u) = \frac{1}{\sqrt{2\pi\sigma^2}} \begin{cases} \exp\left(-\frac{(u-\mu)^2}{2\sigma^2(1-\epsilon)^2}\right) & u < \mu \\ \exp\left(-\frac{(u-\mu)^2}{2\sigma^2(1+\epsilon)^2}\right) & u > \mu \end{cases}$$



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Moments:

$$\begin{aligned} \mathrm{mean}(u) &= \mu + \sqrt{\frac{8}{\pi}} \sigma \epsilon \\ \mathrm{std}(u) &= \sigma \left[1 + \left(3 - \frac{8}{\pi} \right) \epsilon^2 \right]^{1/2} \\ \mathrm{skew}(u) &\simeq \sqrt{\frac{8}{\pi}} \frac{\epsilon}{\mathrm{std}^3(u)} \end{aligned}$$



Centred Gamma pdf

• With
$$z = (u - \overline{u})/\sigma$$
:

$$p_u(u) = \begin{cases} \frac{|\beta|}{\sigma\Gamma(\beta^2)} \left[\beta(z+\beta)\right]^{\beta^2-1} \exp\left[-\beta(z+\beta)\right] & z+\beta > 0\\ 0 & z+\beta < 0 \end{cases}$$



Centred Gamma pdf

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$$mean(u) = \overline{u}$$
$$std(u) = \sigma$$
$$skew(u) = 2/\beta$$



Gram-Charlier Expansion

• With
$$z = (u - \overline{u})/\sigma$$
:

$$p_u(u) = \frac{1}{\sqrt{2\pi\sigma^2}} \left[1 + \frac{\nu}{6} H_3(z) + \frac{\kappa}{24} H_4(z) \right] \exp\left(-\frac{z^2}{2}\right)$$



Gram-Charlier Expansion

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Moments:

$$mean(u) = \overline{u}$$
$$std(u) = \sigma$$
$$skew(u) = \nu$$
$$kurt(u) = \kappa$$



Subject to the constraints:

$$mean(u) = \overline{u}$$
$$std(u) = \sigma$$
$$skew(u) = \nu$$
$$kurt(u) = \kappa$$

find the pdf $p_u(u)$ which maximises the entropy

$$H = -\int p_u(u)\ln p_u(u)du$$



Solution takes the form

$$p_u(u) = \frac{1}{Z} \exp\left(\sum_{i=1}^4 \lambda_i u^i\right)$$



Solution takes the form

$$p_u(u) = \frac{1}{Z} \exp\left(\sum_{i=1}^{4} \lambda_i u^i\right)$$

• Lagrange multipliers $\{\lambda_i\}$ found as solutions to unconstrained dual variational problem



Solution takes the form

$$p_u(u) = \frac{1}{Z} \exp\left(\sum_{i=1}^{4} \lambda_i u^i\right)$$

- Lagrange multipliers $\{\lambda_i\}$ found as solutions to unconstrained dual variational problem
- Maximum entropy pdf is "least biased" among all pdfs with given moments, in a rigorous information theoretic sense



Skew(u) from observations, no kurt(u) information Bigaussian Gram-Charlier 50 50 0 0 -50 -50 100 200 300 100 200 300 0 0 0.5 0.5 -0.5 0 -0.50 Maximum Entropy Centred Gamma 50 50 0 0 -50 -50 100 300 300 200 100 200 0 0 0 0.5 -0.5-0.50 0.5 Vic Modelled skew(w) - observed skew(w)

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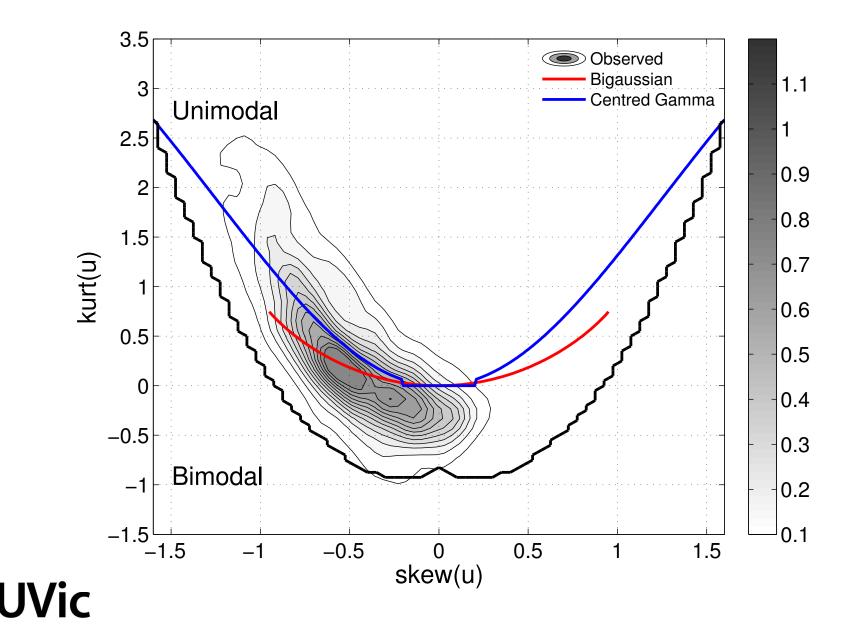
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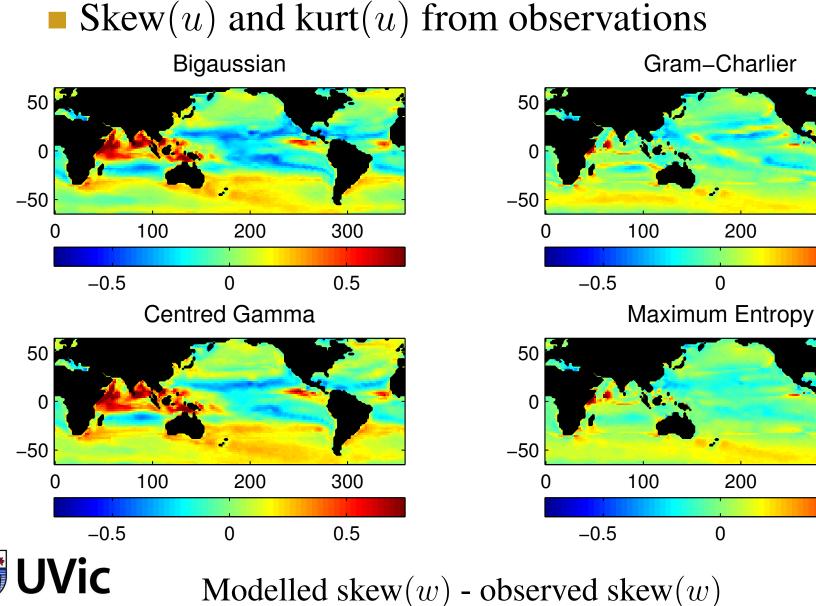
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- In fact, kurtosis of along-mean-wind component non-trivial



$\mathbf{Skew}(u)$ vs. $\mathbf{Kurt}(u)$







Gram-Charlier

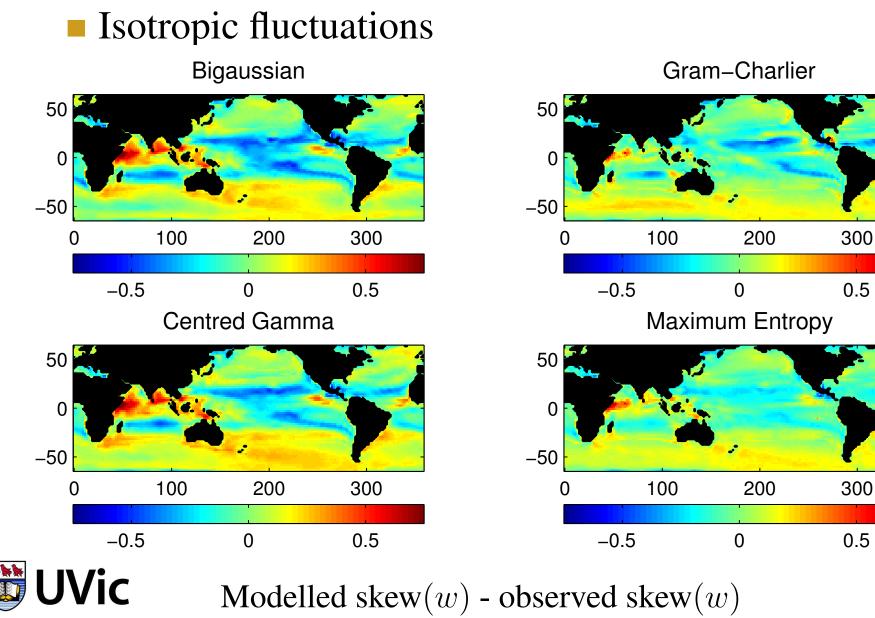
300

0.5

300

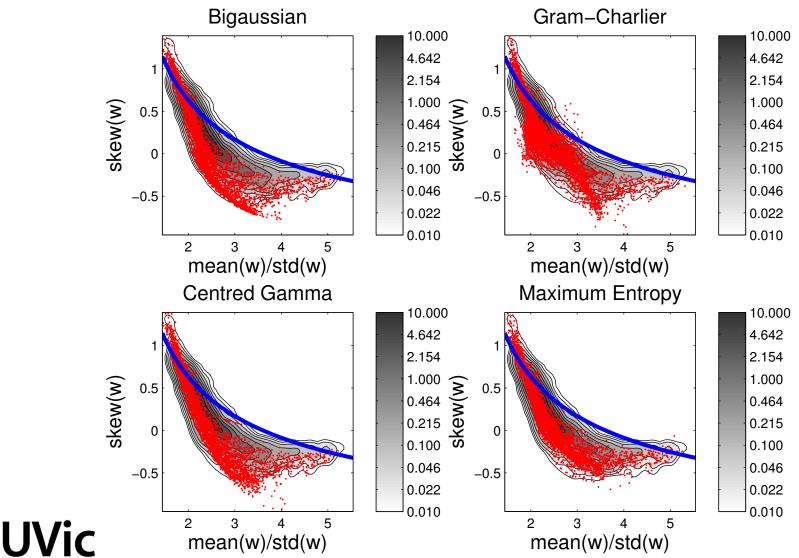
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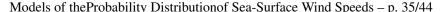
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Relationships between moments





Non-Gaussian structure in vector wind important for accurate simulation of wind speed pdf



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- Parameterisation requires more input information:
 4 moments rather than 2



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- Non-Gaussian structure in vector wind important for accurate simulation of wind speed pdf
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 4 moments rather than 2
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can be parameterised in terms of

 $\operatorname{mean}(u), \operatorname{std}(u)$

Relationship between moments can be found empirically or mechanistically



Mechanistic Model

 Mechanistic model follows from boundary-layer dynamics



Mechanistic Model

- Mechanistic model follows from boundary-layer dynamics
- Horizontal momentum equation:

$$\frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} = -\frac{1}{\rho} \nabla p - f \hat{\mathbf{k}} \times \mathbf{u} - \frac{1}{\rho} \frac{\partial (\rho \overline{\mathbf{u}' u_3'})}{\partial z}$$



• Integrating from z = 0 to z = h



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$$\frac{d\mathbf{u}}{dt} = -\frac{1}{\rho}\nabla p - f\hat{\mathbf{k}} \times \mathbf{u} - \frac{c_d}{h}w\mathbf{u} + \frac{K}{h^2}(\mathbf{U} - \mathbf{u})$$



Mechanistic Model: SDE

Define:

$$\mathbf{\Pi} = -\frac{1}{\rho} \nabla p - f \hat{\mathbf{k}} \times \mathbf{u} + \frac{K}{h^2} \mathbf{U}$$



Mechanistic Model: SDE

Define:

$$\mathbf{\Pi} = -\frac{1}{\rho} \nabla p - f \hat{\mathbf{k}} \times \mathbf{u} + \frac{K}{h^2} \mathbf{U}$$

• Assume fluctuations in Π :

$$\begin{aligned} \Pi_u(t) &= \langle \Pi_u \rangle + \sigma \dot{W}_1(t) \\ \Pi_v(t) &= \sigma \dot{W}_2(t) \end{aligned}$$



Mechanistic Model: SDE

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$$\mathbf{\Pi} = -\frac{1}{\rho} \nabla p - f \hat{\mathbf{k}} \times \mathbf{u} + \frac{K}{h^2} \mathbf{U}$$

Assume fluctuations in Π:

$$\Pi_u(t) = \langle \Pi_u \rangle + \sigma \dot{W}_1(t)$$

$$\Pi_v(t) = \sigma \dot{W}_2(t)$$

Finally, we obtain stochastic differential equation

$$\frac{du}{dt} = \langle \Pi_u \rangle - \frac{c_d}{h} wu - \frac{K}{h^2} u + \sigma \dot{W}_1$$
$$\frac{dv}{dt} = -\frac{c_d}{h} wv - \frac{K}{h^2} v + \sigma \dot{W}_2$$



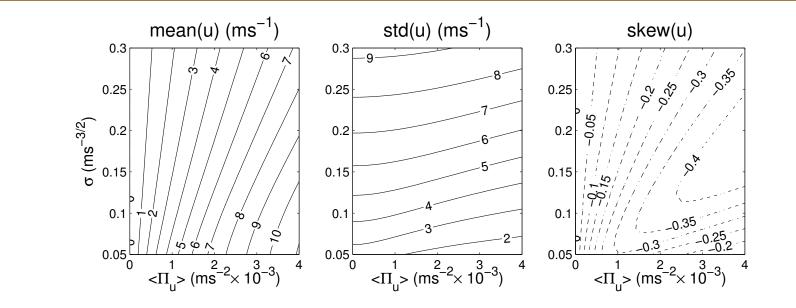
Mechanistic Model: pdf

The stochastic differential equation has an associated Fokker-Planck equation for the stationary pdf, which yields the analytic solution:

$$p_{uv}(u,v) = \mathcal{N}_1 \exp\left(\frac{2}{\sigma^2} \left\{ \langle \Pi_u \rangle \, u - \frac{K}{2h^2} (u^2 + v^2) -\frac{1}{h} \int_0^{\sqrt{u^2 + v^2}} c_d(w') w'^2 \, dw' \right\} \right)$$

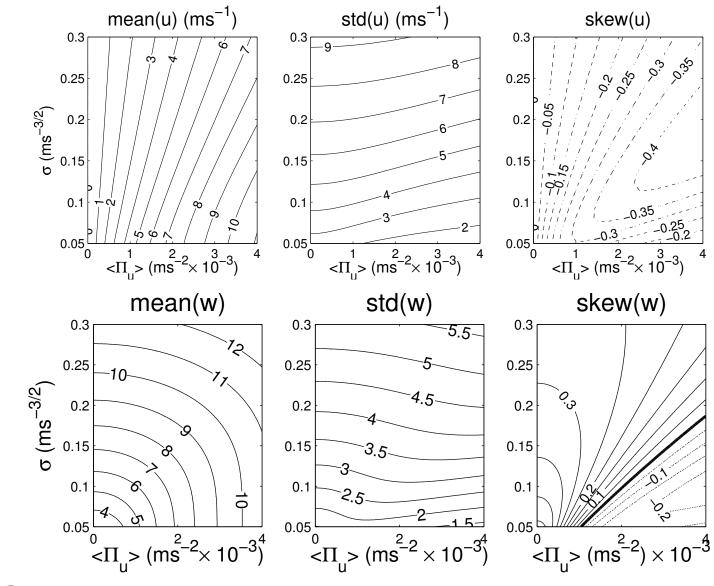


Mechanistic Model: Predictions



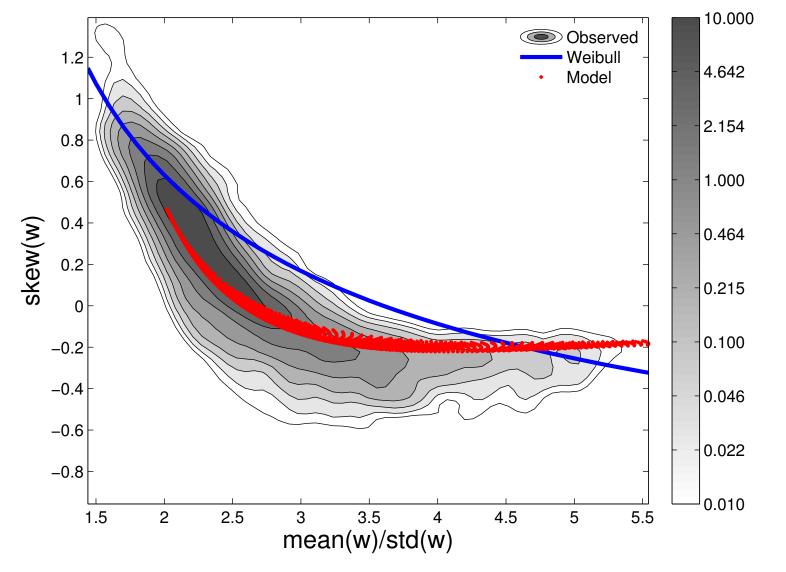


Mechanistic Model: Predictions





Mechanistic Model: Comparison with Observations





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- Qualitative success of model suggests it has captured essential physics: still improvements to be made



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