Multiscale Interactions and Hierarchical Modeling of Climate Variability

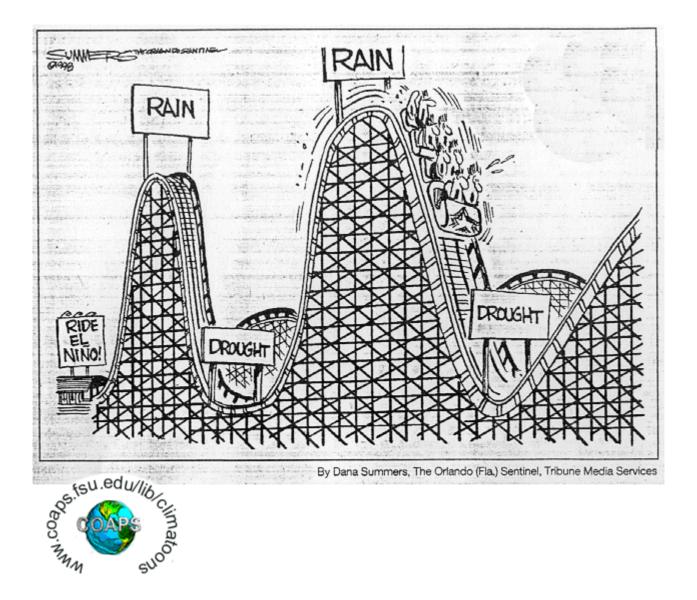
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Questions

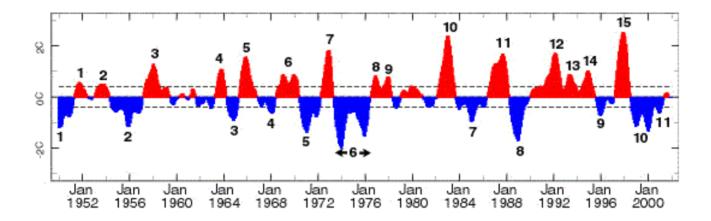
- What role do multi-scale interactions play in climate variability?
- How can we study (and predict) climate variability using a hierarchy of models?

 <u>Strawman</u>: Are simple (and intermediate) models better than complex General Circulations Models?
 OR, When it is appropriate to use simple models?

El Niño-Southern Oscillation



NINO 3.4 index (Kaplan SST data set; base period 1951-1980)



El Niño	Year	 La Niña	Year
10	1982-83	10	1998–2000
11	1986–88	11	2000–01
12	1990–92		
13	1993		
14	1994–95		
15	1997–98		

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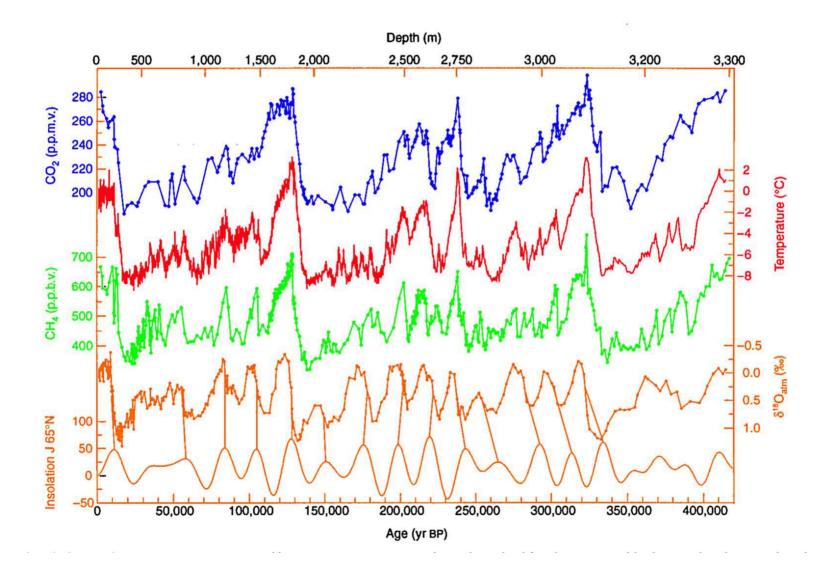
Low frequency variability of Sahel Rainfall

Giannini, Saravanan, and Chang (2003)

3.0 2.0 prcp (standardized) 1.0 0.0 -1.0 -2.0 model observations -3.0 1940 1960 1980 2000

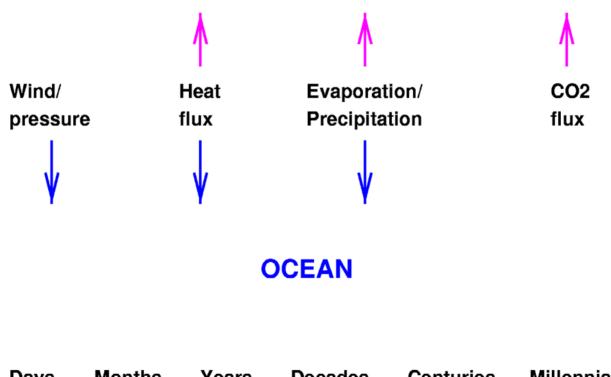
Sahel precipitation - July-September 1930-2000

Vostok core



Ocean-Atmosphere Interaction





Days Months Years Decades Centuries Millennia

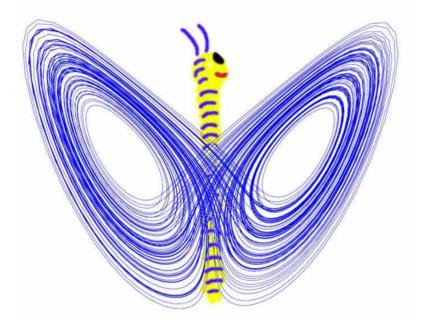
 Days
 Months
 Years
 Decades
 Centuries
 Millennia

 ENSO
 Decadal
 "Climate
 variability
 change"

Multi-scale Interactions

- Often there is no clear separation of scales, i.e., no "spectral gap"
- Interaction of time scales (*global domains*)
 - Flux closures
 - Energy Balance Models
 - Earth Models of Intermediate Complexity (EMICs)
- Interaction of spatial scales (regional domains)
 - "Superparameterizations"
 - Two-way nested models

Deterministic Nonperiodic Flow ("Chaos") Lorenz (1963)



Does the flap of a butterfly's wings in Brazil set off a tornado in Texas? Lorenz (1972)

Superrotation in an Earth-like model Almost intransitive behaviour

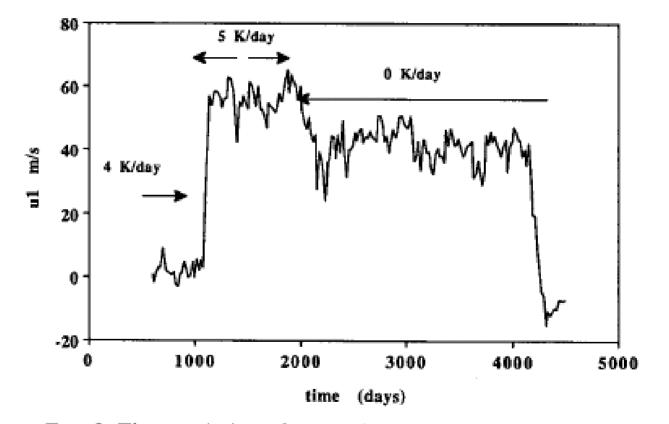


FIG. 3. Time evolution of equatorial $[u_1]$ during a 4500-day integration of the T21 model for varying values of tropical eddy heating rate A_T . The $[u_1]$ values were sampled once every 20 days.

Climatic Predictability Lorenz (1975)

- **Predictability of the First Kind** (*weather prediction*)
 - Effect of uncertainty in initial conditions
- **Predictability of the Second Kind** (*climate prediction*)
 - Response of climate system to boundary conditions
 - Solar radiation
 - Sea surface temperature
 - Soil wetness
 - Sea ice distribution
 - Carbon dioxide concentration

Stochastic climate modelling

- Parameterizing "weather" for climate models
 - Separation in timescales, rather than spatial scales
- Linear models are not very good at weather prediction
 - Timescales of 1-5 days
 - Strong nonlinearity
 - Non-normal PDFs
- Linear models can be quite good at climate prediction
 - Seasonal-to-decadal timescales
 - Weak nonlinearity (the AR(1) model)
 - Approximately normal PDFs

Stochastic coupled models (forward models, not inverse models)

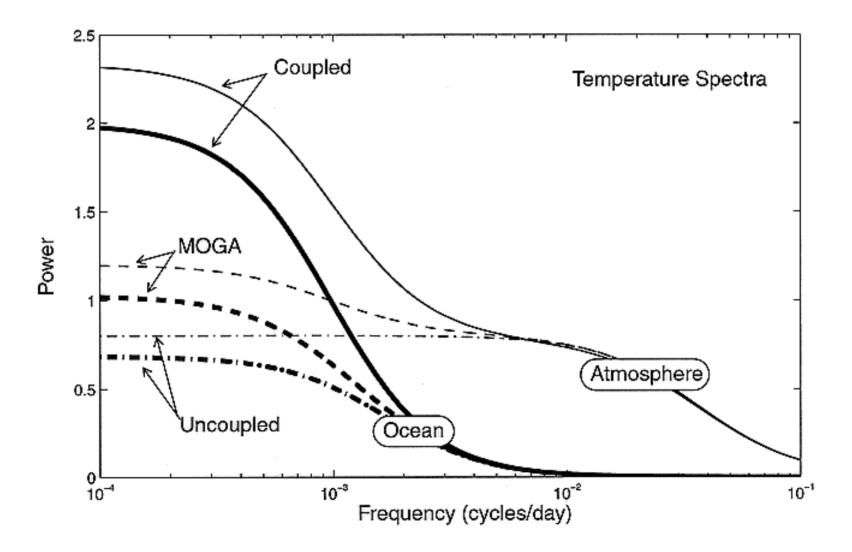
- Stochastic models of atmospheric variability
 - Leith (1975), Branstator (1995), ...
- 0-d stochastic models of oceanic response
 - Hasselmann (1977), ...
- Coupled 0-d models of the atmosphere-ocean system
 - Barsugli & Battisti (1998), ...
- 1-d stochastic models of atmosphere-ocean coupling
 - Saravanan & McWilliams (1998), Chang, Ji, and Saravanan (2001)
 - "Spatial Resonance" (Antarctic Circumpolar Wave, Tropical Atlantic)

0-d climate model (Hasselmann, 1977)

$$\frac{dT_o}{dt} = -\alpha T_o + \varepsilon_a(t)$$

Oceanic Response Atmospheric forcing

0-D Coupled Ocean-Atmosphere Model Barsugli & Battisti (1998)



0-d coupled oscillator

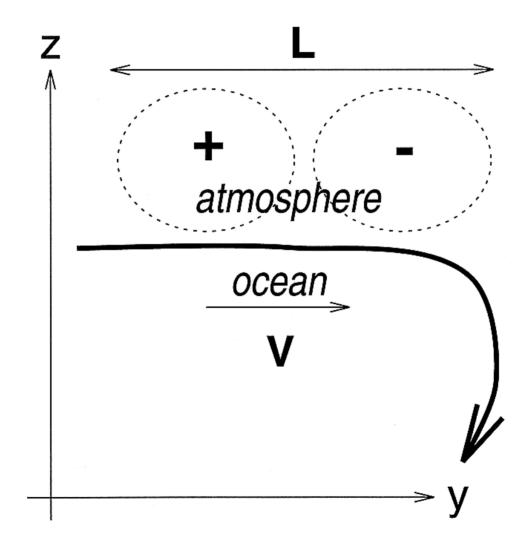
Chang, Saravanan, DelSole, and Wang (2002)

$$\frac{du}{dt} = +fv - \alpha u$$

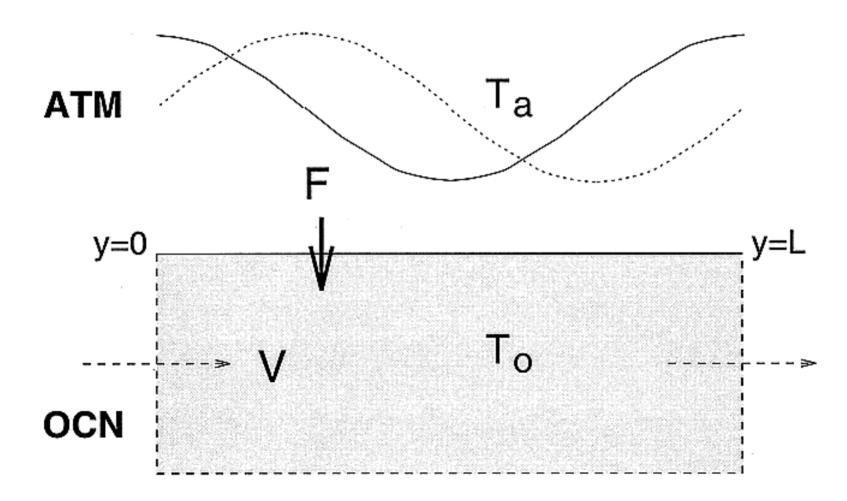
$$\frac{dv}{dt} = -fu - \gamma u$$
Eigenvalues $\lambda_{\pm} = -(\alpha + \gamma) \pm iF$

$$F = \sqrt{f^2 - (\alpha - \gamma)^2}$$

1-d atmosphere-ocean coupling



1-d atmosphere-ocean coupling



1-d advective coupled model

$$\partial_t T_a = -\alpha T_a - \frac{F}{C_a} + \varepsilon_a(y, t),$$

$$\partial_t T_o = -V \partial_y T_o + \frac{F}{C_o},$$

$$F = \kappa (T_a - T_o),$$

$$\frac{\varepsilon_a}{\alpha} = W(t)\sin(2\pi\frac{y}{L}).$$

Boundary condition $T_o(0,t) \equiv 0$

Fourier transform in $t \ldots$

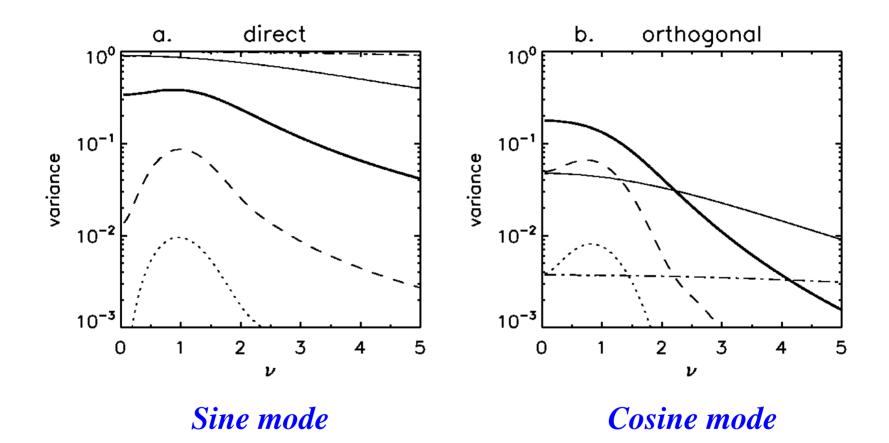
$$\left[1 + \Gamma\left(i\nu + \frac{1}{2\pi}\partial_y\right)\right]\hat{T}_o = \hat{W}(\nu)\sin(2\pi y)$$

$$\Gamma \equiv 2\pi \frac{\lambda_{\text{eff}}^{-1}}{L/V} \propto \frac{\text{damping time-scale}}{\text{advection time-scale}}$$

Laplace transform in y and invert ...

$$\hat{T}_o(y,
u) = rac{\hat{W}}{\Gamma(1+\xi^2)} \left[\xi \sin 2\pi y - \cos 2\pi y + e^{-\xi 2\pi y}
ight],$$
 $\xi \equiv rac{1}{\Gamma} + i
u.$

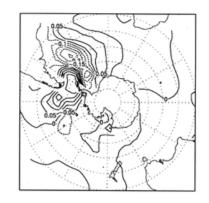
Saravanan & McWilliams (1998)



Oceanic response for $\Gamma = 1/16$, ¹/₄, 1, 4, 16

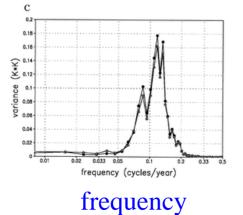
Haarsma, Selten, and Opsteegh (2000)

Antarctic Circumpolar Wave in the ECBILT Model

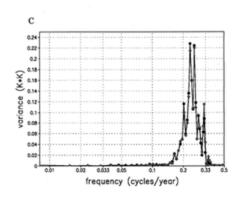


Ocean T (80m-300m)









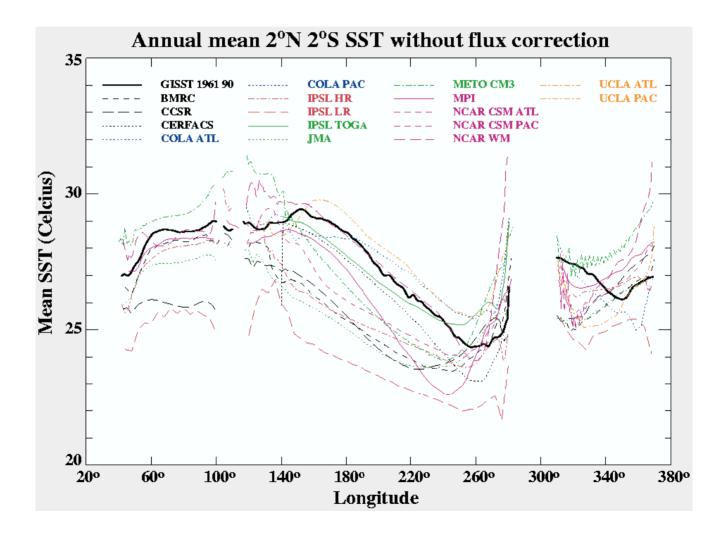
Normal ACC



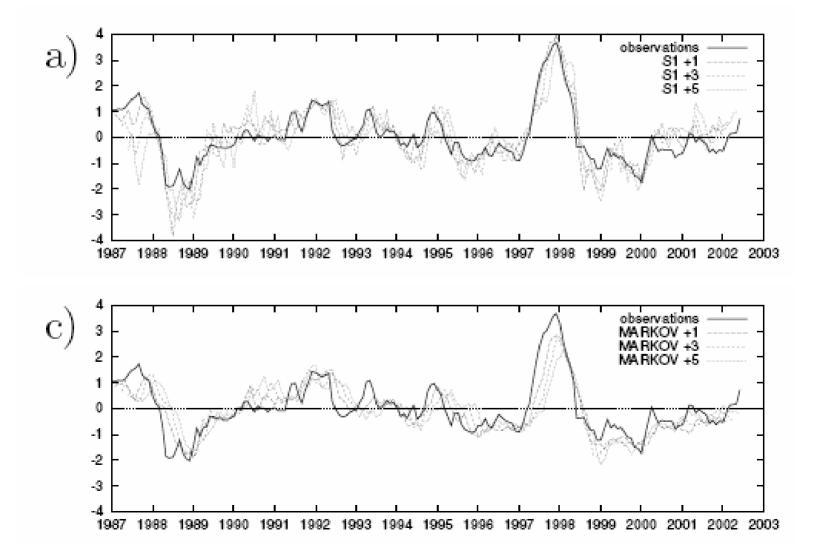
When GCMs have O(1) errors ...

- Simple models can be used to study, and even predict phenomena
- El Niño-Southern Oscillation, Tropical Atlantic
 - Variability
 - GCMs have systematic errors in simulating the mean state and the annual cycle in the Tropical Pacific & Atlantic

Intercomparison of Simulated Equatorial SST (STOIC: Davey et al., 2002)



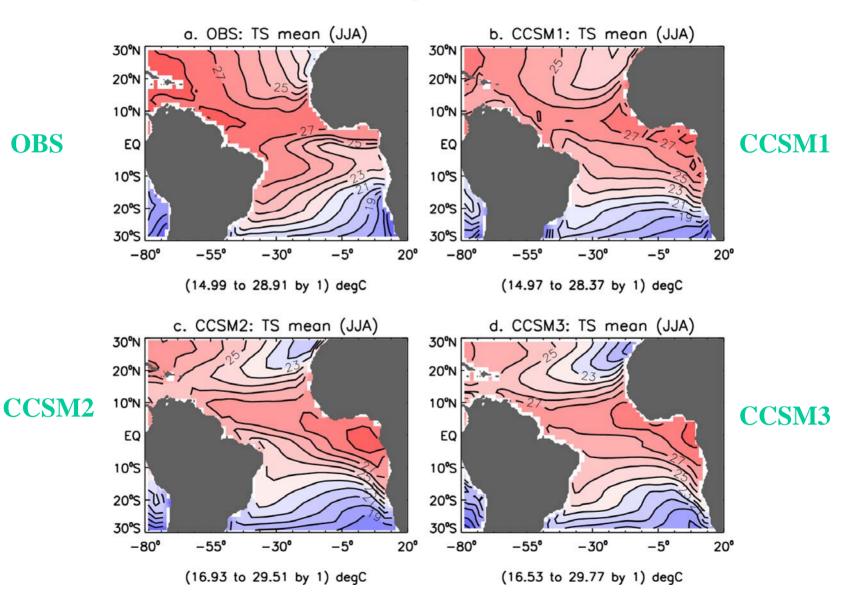
Did the ECMWF seasonal forecast model outperform statistical ENSO forecast models over the last 15 years? (van Oldenborgh et al., 2006)



van Oldenborgh et al., 2006

	monthly Niño3.4			3-monthly Niño3.4	
lead	+1	+3	+5	+0	+3
S1	0.93^{+2}_{-4}	0.87^{+5}_{-8}	0.79^{+9}_{-15}	0.94^{+3}_{-3}	
S2	-	0.88^{+4}_{-8}		0.96^{+2}_{-2}	0.84^{+6}_{-11}
MARKOV	0.91^{+3}_{-4}	0.81^{+7}_{-8}	0.73^{+9}_{-14}		0.78^{+8}_{-9}
CA	0.91^{+3}_{-4}	0.86^{+5}_{-7}	0.78^{+7}_{-10}	0.93^{+2}_{-3}	
STAT		0.79^{+8}_{-12}	0.65^{+12}_{-16}	0.88^{+3}_{-4}	0.67^{+12}_{-15}
CLIPER	0.91^{+3}_{-5}	0.77^{+9}_{-13}	0.61^{+14}_{-16}	0.87^{+5}_{-5}	0.63^{+13}_{-17}
ENSO-CLIPER				0.90^{+3}_{-4}	0.81^{+6}_{-8}
MULTI-MODEL	0.96^{+1}_{-2}	0.90^{+4}_{-5}	0.84^{+6}_{-10}	0.96^{+1}_{-1}	

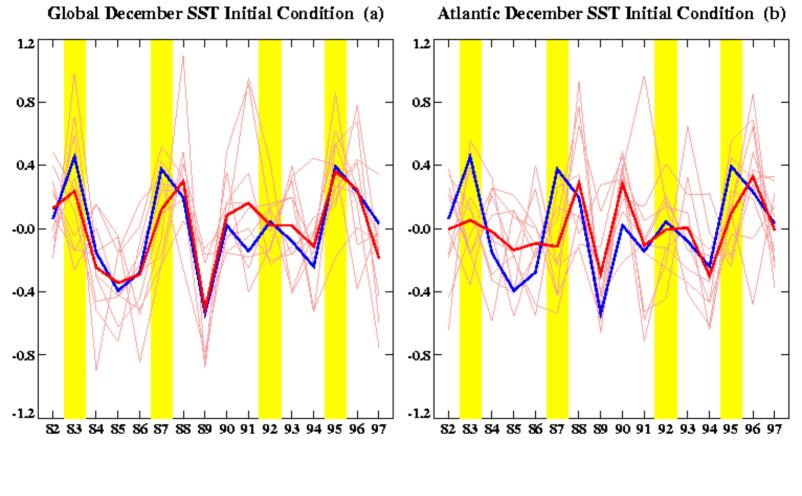
June-July-August mean SST



CCM3-ML integrations

- Control integration
 - CCM3.6.6 + Slab Ocean, with annual-mean mixed layer depth
 - 100 years
- Forecasts with *Global* SST Initial Condition
 - 16 cases (1981-1996), 9 month forecast, 10 member ensemble
 - Observed December monthly-mean initial condition for SST
 - Observed December 15 initial condition for atmosphere
- Forecasts with *Atlantic* SST Initial Conditions
 - Observed SST in Atlantic only (30S 60N)

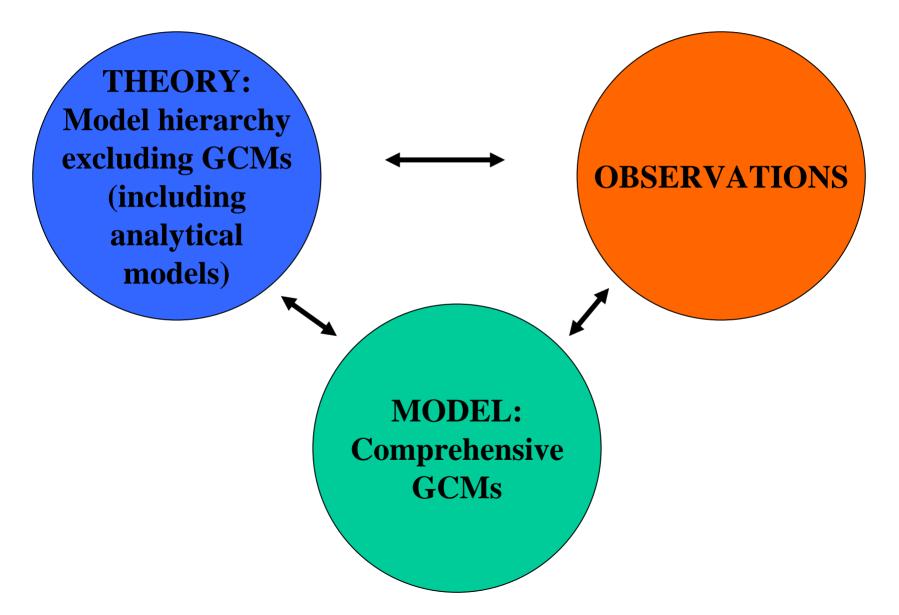
CCM3-ML "Caribbean" predictions (4 month lead time)



observation

prediction

Theory vs. Model vs. Observations

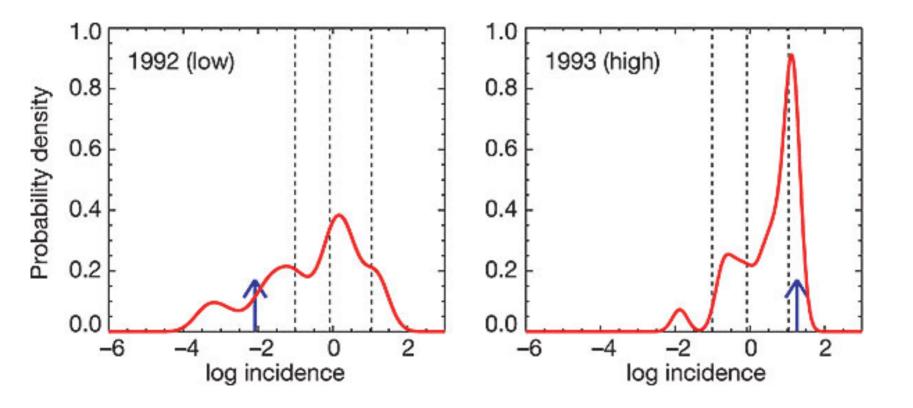


Conclusions

- Utility of model hierarchy
 - "Cheaply" identify mechanisms to test against observations and GCMs
 - Provide the context for GCM behaviour (in climate parameter space)
 - Serve as "scouts" in the frontiers of research
 - Useful for predictions when GCMs have O(1) errors
- It is always possible to construct a GCM that is a superset of the simpler model
 - If a simple model and its corresponding superset GCM disagree, the simple model is more likely to be wrong (as it makes more approximations)

- Parameterizing "weather noise" as a stochastic process can provides some remarkable insights into the nature of climatic variability
 - Red-noise character of climate variability
 - Resonant interactions involving spatial structures
- Is the coupled atmosphere-ocean system subcritical or supercritical (with regard to instability)?
 - In the tropics? In the middle latitudes?
 - Does it matter?
- Forward modelling vs. inverse modelling?

Forecast probability of malaria annual incidence for Botswana (Thomson et al., Nature, 2006)



Mechanistic Modeling

- NSIPP1: NASA Seasonal-to-Interannual Prediction Project
 - 2x2.5 degrees, 34 vertical levels
- NSIPP1/AMIP: Observed SST forcing (1950-1999)
- NSIPP1/ACYC: Annual cycle of SST forcing
- NSIPP1/FBETA: Observed SST, but prescribed "evaporation efficiency" β (ratio of evaporation to potential evaporation)

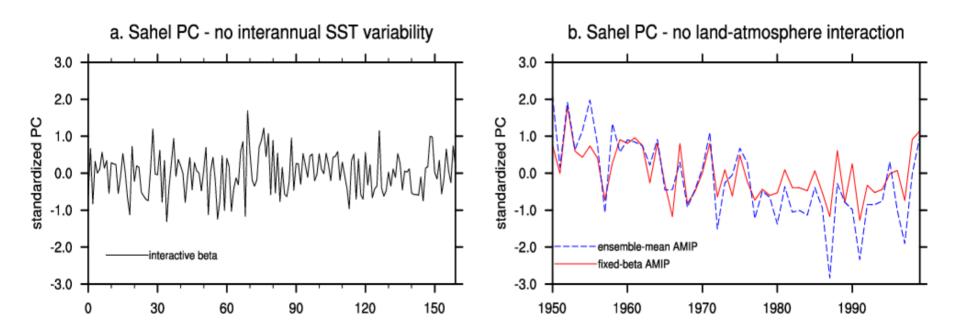
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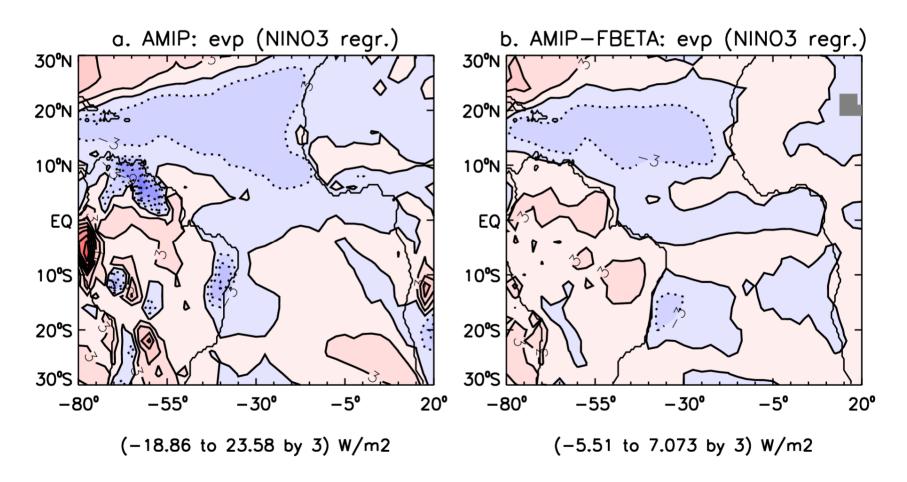
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Sahel precipitation - July-September 1930-2000

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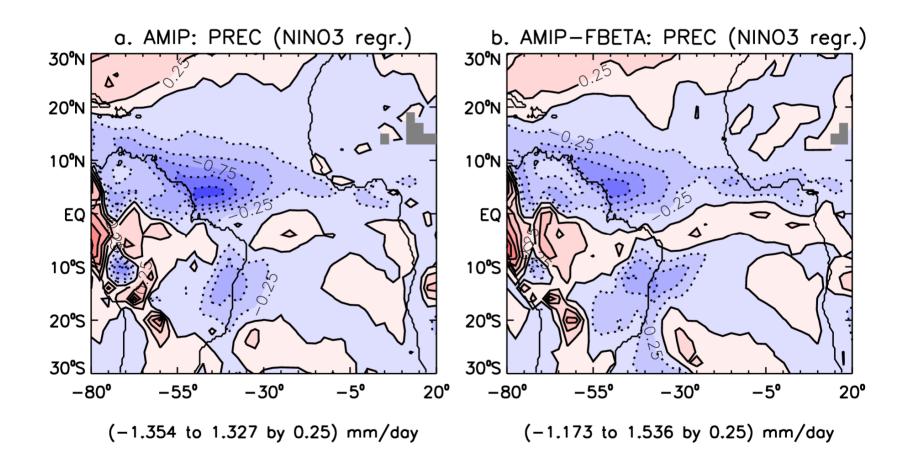
NASA/NSIPP: EVAP regression (DJF)



AMIP

AMIP/Fixed β

NASA/NSIPP: PRECIP regression (DJF)



AMIP/Fixed β

AMIP