A Mechanistic Model of Mid-Latitude Decadal Climate Variability (IMAGe T-O-Y Workshop IV)

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North Atlantic Oscillation and Arctic Oscillation



AO

NAO



SST and NAO



SST tripole pattern (Marshall et al. 2001, *Journal of Climate:* Vol. 14, No. 7, pp. 1399–1421)

- Decadal time scale detected in NAO/SST time series
- If real, what dynamics does this signal represent? <u>We will</u> <u>emphasize ocean's dynamical</u> <u>inertia due to eddies</u>
- AGCMs: response to (small) SSTAs is weak and model-dependent
- <u>Nonlinear</u>: small SSTAs – large response??

Coupled QG Model

 Eddy-resolving atmospheric and ocean components, both characterized by vigorous intrinsic variability

 (Thermo-) dynamic coupling via constantdepth oceanic mixed layer with entrainment



Atmospheric circulation



Zonal-jet bimodality in the model



Intra-seasonal oscillations in the atmospheric model



Atmospheric driving of ocean



 <u>Coupled effect</u>: Occupation frequency of atmospheric low-latitude state exhibits (inter)-decadal broad-band periodicity

Oceanic circulation



Eddy effects on O-climatology-I

$$\begin{split} \Psi_1 &= \overline{\Psi}_1 + \Psi'_1 \\ \Psi'_1 &= \Psi'_{1,L} + \Psi'_{1,H} \end{split}$$

$$Q_1 = \nabla^2 \Psi_1 + (f_0^2/g'H)(\Psi_2 - \Psi_1)$$

$$\dot{Q} \sim -J(\Psi_{\mathsf{H}}', Q_{\mathsf{H}}') - J(\Psi_{\mathsf{L}}', Q_{\mathsf{L}}')$$

$$-J(\Psi_{\mathsf{H}}',\,Q_{\mathsf{L}}')-J(\Psi_{\mathsf{L}}',\,Q_{\mathsf{H}}')$$

+ linear terms . . .

Eddy effects on O-climatology-



Eddy effects on O-LFV–I (EPV-flux tendency regressed onto EOF-1 of Ψ_1)



- 10 yr – 5 yr 0 yr + 5 yr + 10 yr Eddy effects on O-LFV–II
Dynamical decomposition into large-scale flow and eddy-flow components, based on parallel integration of the full and "coarsegrained" ocean models (Berloff 2005)

- "Coarse-grained" model forced by randomized spatially-coherent eddy PV fluxes exhibits realistic climatology and variability
- Main eddy effect is <u>rectification</u> of oceanic jet (eddy <u>fluctuations</u> are <u>fundamental</u>)

Dynamics of the oscillation – I





- High Ocean Energy = High-Latitude (HL) O-Jet State
- HL ocean state = A-jet's
 - Low-Latitude (LL) state
- O-Jet stays in HL state for a few years due to O-eddies

Dynamics of the oscillation – II

Oscillation's period
is of about 20 yr in low ocean-drag case and
is of about 10 yr in high ocean-drag case

• Period scales as eddydriven adjustment time



Conceptual model – I

 Fit A-jet position time series from A-only simulations forced by O-states keyed to phases of the oscillation to a stochastic model of the form

$$dx = -V_x dt + \sigma dw$$

[V(x) – polynomial in x]



Conceptual model – II



Conceptual model – III

• "Atmosphere:"
$$dx = -V_x(x, y)dt + \sigma dw$$

• "<u>Ocean</u>:" $\lambda^{-1}=2$ yr, $T_d=5$ yr $\dot{y} = -\lambda y + Ax(t - T_d)$

• <u>Delay</u>: ocean's jet does not "see" the loss of local atmospheric forcing because ocean eddies dominate maintenance of O-jet for as long as T_d

 Atmospheric potential function responds to oceanic changes instantaneously: O-Jet HL state favors A-Jet LL state and vice versa

Conceptual model – IV



Conceptual model – V



Summary

- Mid-latitude climate model involving turbulent oceanic and atmospheric components exhibits inter-decadal coupled oscillation
- Bimodal character of atmospheric LFV is responsible for atmospheric sensitivity to SSTAs
- Ocean responds to changes in occupation frequency of atmospheric regimes with a delay due to ocean eddy effects
- Conceptual toy model was used to illustrate how these two effects lead to the coupled oscillation