Parameterizing Submesoscale Eddies for Global Ocean Models

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Typical Ocean Stratification Permits

cale and SubMesoscale (Boccaletti et al., 2006)



Mes



Typical Ocean Stratification Permits

Mes cale and SubMesoscale (Boccaletti et al., 2006)

Vertical fluxes are Submesoscale

And tend to restratify Horizontal fluxes are Mesoscale and tend to stir



Having a Mixed Layer Counts! The vertical buoyancy flux in the ML (<w'b'>) without diurnal cycle is notless than with cycle (ML)



Having a Mixed Layer Counts! The vertical buoyancy flux in the ML (<w'b'>) without diurnal cycle is less than with cycle (ML)



Having a Mixed Layer Counts! The vertical buoyancy flux in the ML (<w'b'>) without diurnal cycle is 2x less than with cycle (ML)



Having a Mixed Layer Counts! The vertical buoyancy flux in the ML (<w'b'>) without diurnal cycle is 2x less than with cycle (ML)



Having a Mixed Layer Counts! The vertical buoyancy flux in the ML (<w'b'>) without diurnal cycle is 3x less than with cycle (ML)



Having a Mixed Layer Counts! The vertical buoyancy flux in the ML (<w'b'>) without diurnal cycle is 4x less than with cycle (ML)





Simple Spindown



Plus, Diurnal Cycle and KPP





Simple Spindown



Plus, Diurnal Cycle and KPP





Simple Spindown



Plus, Diurnal Cycle and KPP





Simple Spindown



Plus, Diurnal Cycle and KPP







Simple Spindown







Simple Spindown

Plus, Diurnal Cycle and KPP

Plus, Diurnal Cycle and KPP

Simple Spindown

Plus, Diurnal Cycle and KPP

Simple Spindown

Plus, Diurnal Cycle and KPP

y (km)

The Parameterization:

 $\Psi = \frac{C_e H^2 \mu(z)}{|\mathbf{f}|} \nabla \overline{b} \times \mathbf{\hat{z}}$

$$\mu(z) = \left[1 - \left(\frac{2z}{H} + 1\right)^2\right] \left[1 + \frac{5}{21}\left(\frac{2z}{H} + 1\right)^2\right]$$

Fox-Kemper et al. (08) The horizontal fluxes are downgradient:

$$\overline{\mathbf{u}_{\mathbf{H}}'b'} = -\frac{C_e H^2 \mu(z) \frac{\partial b}{\partial z}}{|f|} \nabla_H \overline{b}$$

Vertical fluxes always upward to restratify with correct extraction rate of potential energy:

$$\overline{w'b'} = \frac{C_e H^2 \mu(z)}{|f|} |\nabla \overline{b}|^2$$

It works for Prototype Sims:

Red: No Diurnal Blue: With Diurnal 10¹ 10⁰ w'b' $\overline{\overline{b}}_y$ ^{Sul}→10⁻², Ð 10⁻³ 10⁻² 10⁻³ 10⁰ 10^{-1} C H² M² Ifl⁻¹ 10⁻¹ C H² M² IfI⁻¹ 10⁻² 10⁰ 10¹ 10^{1}

>2 orders of magnitude!

Circles: Balanced Initial Cond. Squares: Unbalanced Initial Cond.

Works In 'Real' Models!

HIM/GOLD and CCSM/POP

Soon to be in MITgcm & MOM

0.1 POP mle BMFLI POP control -BMFLI POP control -BMFLI POP control -BMFLI POP control -Levitus Difference in Time-Mean Mixed Layer Depth (m) Change of Time-Mean Boundary Layer Depth in POP

© RMS: 16m->8m, Skew 2.4->0.6

Summary I:

Submesoscale features, and mixed layer eddies in particular, exhibit large vertical fluxes of buoyancy often ignored in climate models.

A parameterization of mixed layer eddy fluxes as an overturning streamfunction is proposed. The magnitude comes from extraction rate of potential energy.

Many observations are consistent, and model biases are reduced. Biogeochemical effects are likely, as vertical fluxes and mixed layer depth are changed.

In HIM & CCSM, soon to be in MITgcm and MOM.

4 Papers so far... fox-kemper.com/research

Param gives same scaling, but...

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-500 (III) Hade D D D 1d -1500 -2000 100 150 50 Domain Distance (km) -50 5d Depth (m) -1000 -1500 -2000 50 100 150 Domain Distance (km) -500 30d Depth (m) -1000 -150 -2000 50 100 150 Domain Distance (km) -50 Depth (m) -1000 50d -1500 -2000 50 100 150 Domain Distance (km)

Param gives same scaling, but...

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Vertical structure is different...

Jones & Marshall 97

100 150 Domain Distance (km)

50

-2000

4d16h from 2d parameterization 0 -500 -1000 -1500 -2000 50 100 150 200 0 y (km)

30d04h from 2d parameterization

49d24h from 2d parameterization 0 -500 -1000 -1500 -2000 50 100 150 200 0 y (km)

Consider SQG Eady + Mix

Possibly 3 growing modes and 3 decaying.

Consider SQG Eady + Cool 0-100m

Consider SQG Eady + Cool 100-200m

Consider SQG Eady + Cool 200-300m

Consider SQG Eady + Cool 300-400m

Consider SQG Eady + Cool 400-500m

Summary II:

A parameterization of mixed layer eddy fluxes as an overturning streamfunction is proposed.

The parameterization reduces model bias in GOLD and POP.

However, difficulties arise in parameterization of submesoscale features if restratification isolates reduced stratification away from boundaries.

Preliminary testing/linear instability analysis reveals that extending submesoscale may be relatively easy, but preventing mesoscale double-counting will not.

4 Papers so far... fox-kemper.com/research

The Global Parameterization:

 $\Psi = \frac{C_e H^2 \mu(z)}{\nabla \bar{b} \times \hat{z}}$

$$\mu(z) = \left[1 - \left(\frac{2z}{H} + 1\right)^2\right] \left[1 + \frac{5}{21}\left(\frac{2z}{H} + 1\right)^2\right]$$

Account for equator by going to subinertial ML approx (Young 94)

$$\Psi = \frac{C_e H^2 \mu(z)}{\sqrt{f^2 + \tau^{-2}}} \nabla \overline{b} \times \mathbf{\hat{z}}$$

Account for coarse res. by assuming

$$E_b(k) \sim k^{-2} \to \mathbf{\Psi} = \left[\frac{L_f}{\Delta x}\right] \frac{C_e H^2 \mu(z)}{\sqrt{f^2 + \tau^{-2}}} \nabla \overline{b} \times \mathbf{\hat{z}}$$

Obs. reveal $L_f \sim R_d$

Estimate of Vert. Heat Flux from satellite data agrees

Changes To Mixing Layer Depth in Eddy-Resolving Southern Ocean Model

Changes To Mixing Layer Depth in Eddy-Resolving Southern Ocean Model

 $N^{2}(z^{b})$

