

Zonal jet formation and Equatorial Super-rotation in the Destabilization of Short Mixed Rossby-Gravity Waves

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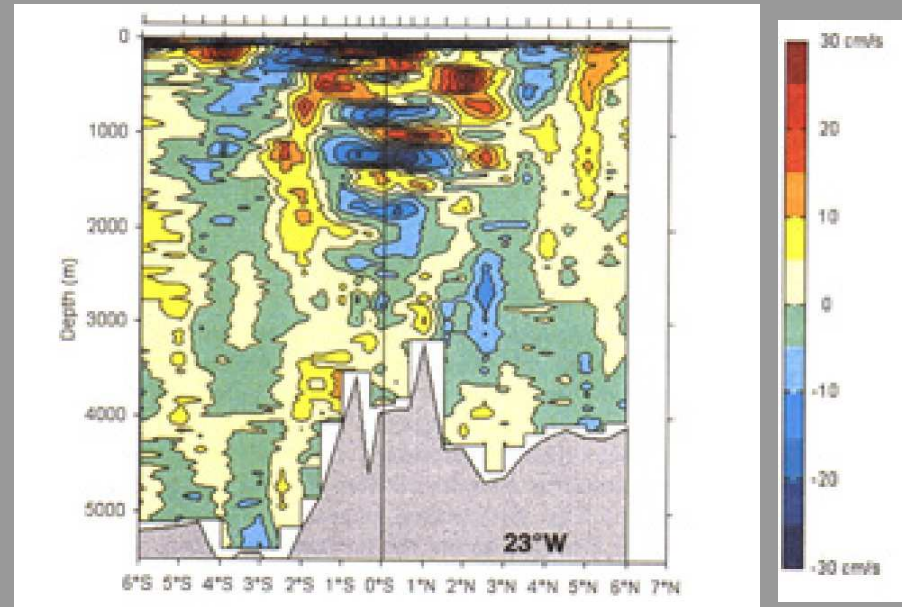
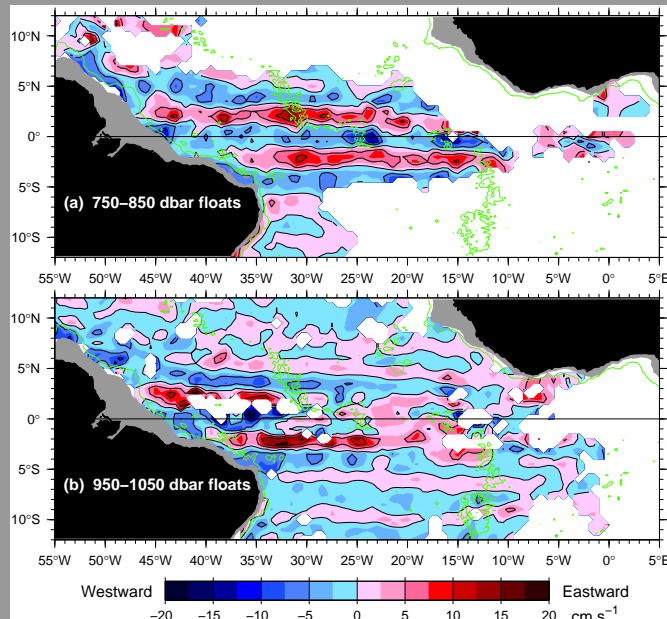
Turbulent Theory and Modeling Workshop; Boulder, Co

February 27th - 29th, 2008

OUTLINE

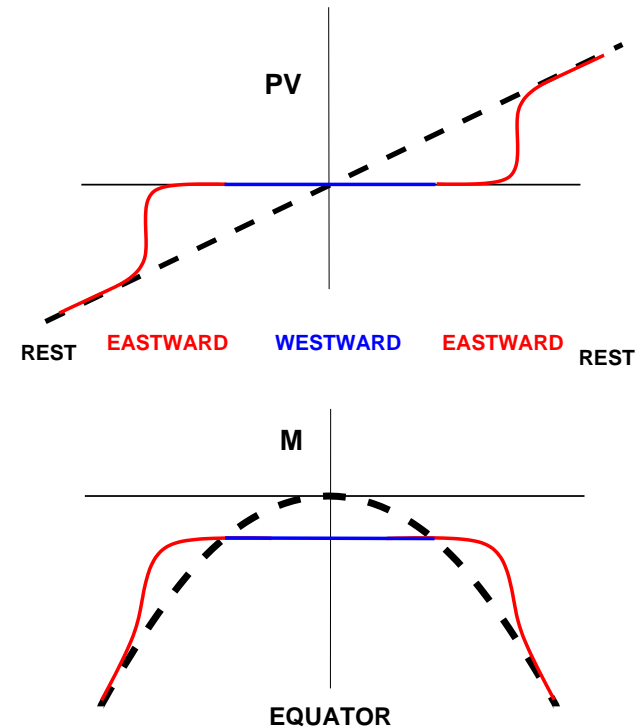
1. Equatorial zonal jets
2. Destabilization of short westward MRG wave
3. Inertial Instability and PV Homogenization
4. Super-rotation and the non-traditional Coriolis force

Gouriou et al. 2001 \Rightarrow
 – instantaneous merid. sect. 23° W
 – strongly barotropic jets



\Leftarrow Ollitraut et al. 2006
 – zonal currents near 1000 m depth
 – multiple, equally spaced jets

- Westward flow at equator and eastward jets at $\pm 2^\circ$ latitude (Gouriou et al. 2001)
- often corresponds to angular momentum M (and hence PV) homogenized about equator surrounded by PV barriers.

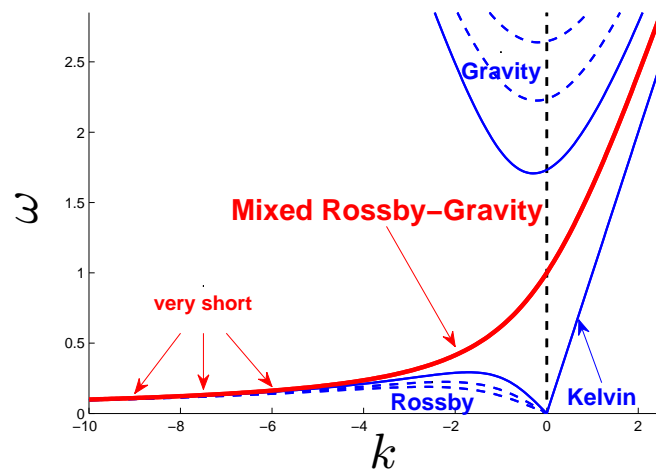


- Propose mechanism for generating jets by barotropic instability of short equatorial waves
- Also explains “equatorial deep jets” (d’Orgeville et al. (2006), Hua et al. (2008))

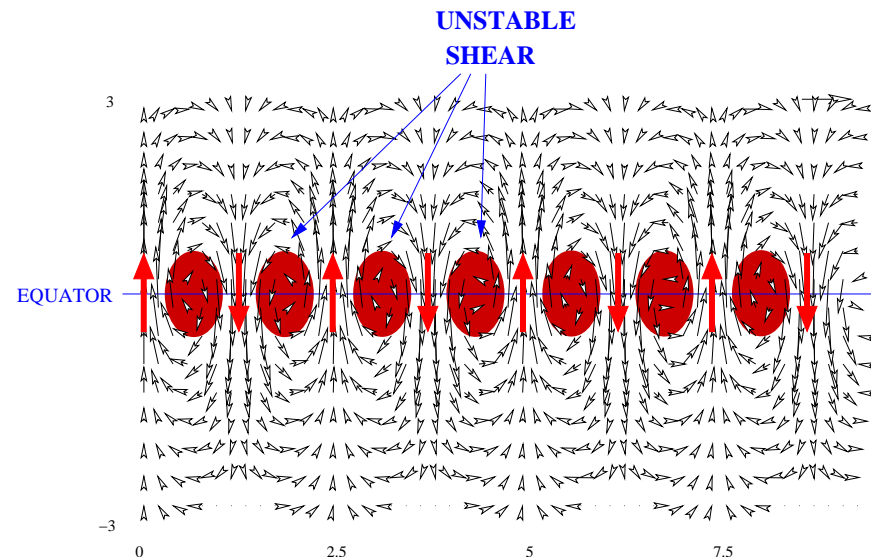
2. Destabilization of short westward MRG wave

MRG wave in large negative k limit: $\omega \sim -k^{-1}$

DISPERSION DIAGRAM FOR
EQUATORIALLY TRAPPED WAVES



MRG WAVE VELOCITY FIELD



- Cases considered: $-6 < k < -16$
- Oceanic vertical mode 1 \Rightarrow 50 - 200 day period, $1.2^\circ - 3^\circ$ wavelength.

Linear Theory in $k \ll -1$ limit:

Scale length and time like $(x, y) = k^{-1}(\xi, \eta)$ and $t = (kV_0)^{-1}\tau$.

Velocity components of basic state wave:

$$V \sim V_0 \exp\left(\frac{-\eta^2}{2k^2}\right) \cos(z) \cos(\xi),$$

$$U \sim k^{-1}V_0 \frac{\eta}{k^2} \exp\left(\frac{-\eta^2}{2k^2}\right) \cos(z) \sin(\xi),$$

$$W \sim k^{-2}V_0 \left(\frac{-\eta^2}{k^2}\right) \exp\left(\frac{-\eta^2}{2k^2}\right) \sin(z) \cos(\xi),$$

$$Z \sim -kV_0 \exp\left(\frac{-\eta^2}{2k^2}\right) \cos(z) \cos(\xi),$$

Look for **horizontally non-divergent** perturbations $u' = -\psi'_\eta$, $v' = \psi'_\xi$.

Perturbation vorticity equation is approximately

$$\tilde{\zeta}'_\tau + V\tilde{\zeta}'_\eta + k^{-1}Z_\xi u' = \nabla^2\psi'_\tau + \cos(z) \exp\left(\frac{-\eta^2}{2k^2}\right) \cos(\xi) (\nabla^2\psi'_\eta + \psi'_\eta) = 0$$

Next order in k^{-1} : time dep. of MRG wave, advection by U , and β effect.

Case 1: η small, $\exp\left(\frac{-\eta^2}{2k^2}\right) \approx 1$

Special case of problem of Gill (1974) for barotropic Rossby wave.

Approximate solution by truncated series:

$$\psi' = \text{Re} \left[e^{\mu \cos(z)\tau} e^{il\eta} (\psi'_{-1} e^{-i\xi} + \psi'_0 + \psi'_1 e^{i\xi}) \right]$$

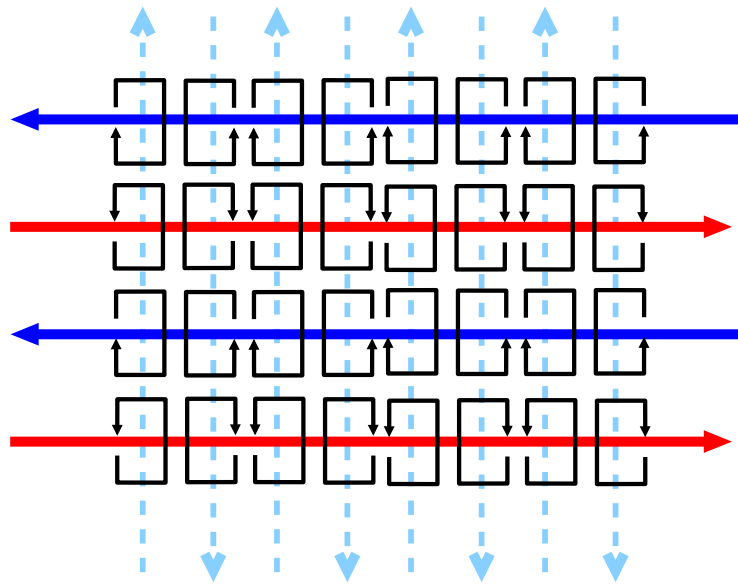
ψ'_0 is zonal jet, $\psi'_{\pm 1}$ zonally short wave.

Instability if eigenvalue

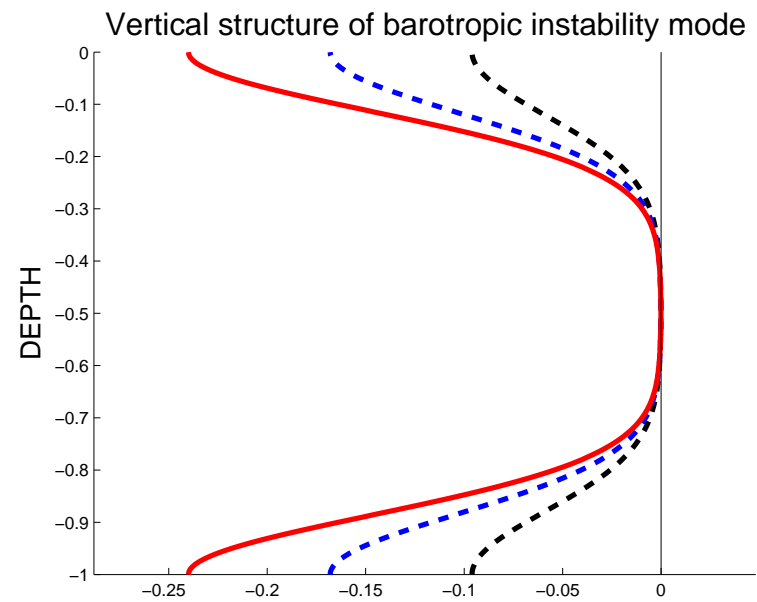
$$\mu = \frac{l}{\sqrt{2}} \sqrt{\frac{1-l^2}{l^2+1}} > 0$$

\Rightarrow Instability if $|l| < 1 \Rightarrow$ perturbation of larger scale than original wave (expected on grounds of simultaneous conservation of energy and enstrophy). Fastest growth rate for $l = \sqrt{\sqrt{2} - 1} \approx 0.64$.

Horizontal structure



Vertical structure



Zonal jets alternating in direction in latitude, and small scale cells stirring within jet centres.

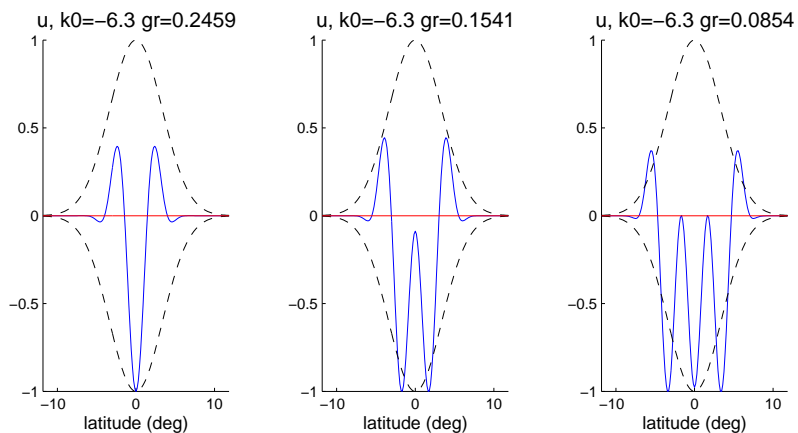
Case 2: Solution valid on $-\infty < \eta < \infty$

Again look for truncated series solution:

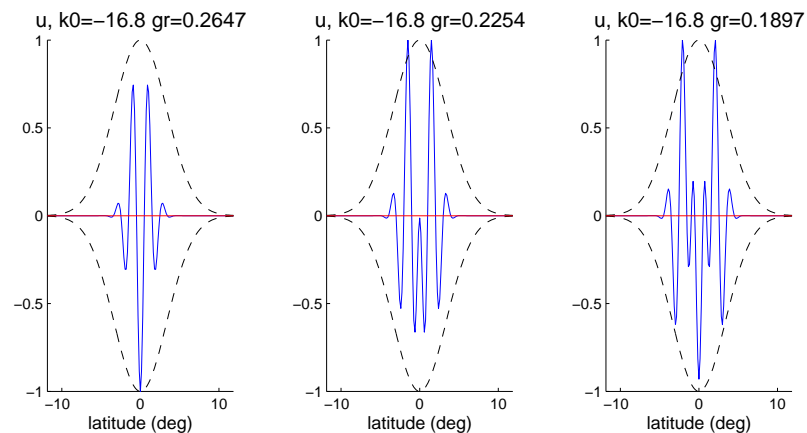
$$\psi' = \text{Re} \left[e^{\mu \cos(z)\tau} (\psi'_{-1}(\eta)e^{-i\xi} + \psi'_0(\eta) + \psi'_1(\eta)e^{i\xi}) \right]$$

Solve numerically using finite difference method. Structure of $\psi'_0(\eta)$ for fastest growing eigenmodes:

$k = -6.3$



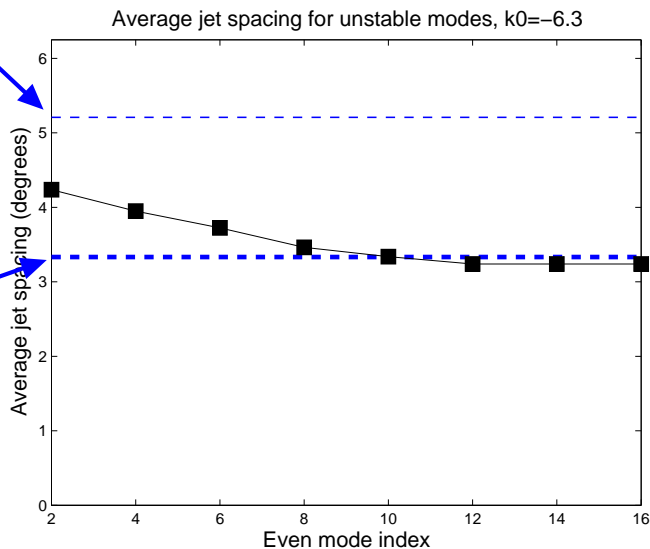
$k = -16.8$



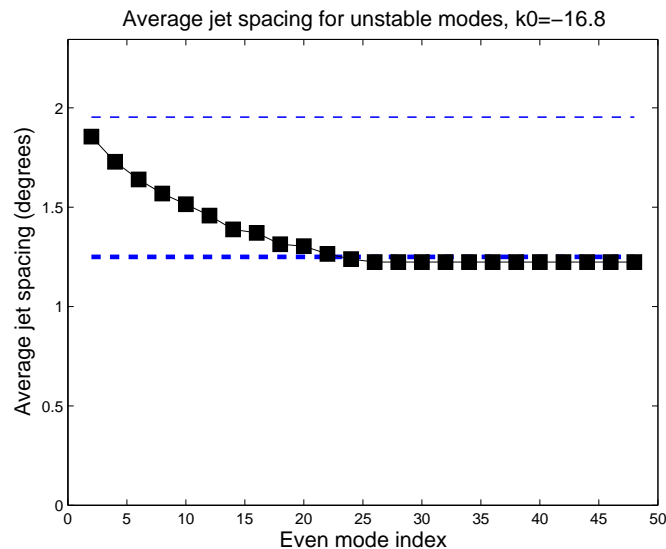
Average jet spacing:

$V = \text{const}$

MRG wavelength



$k = -6.3$



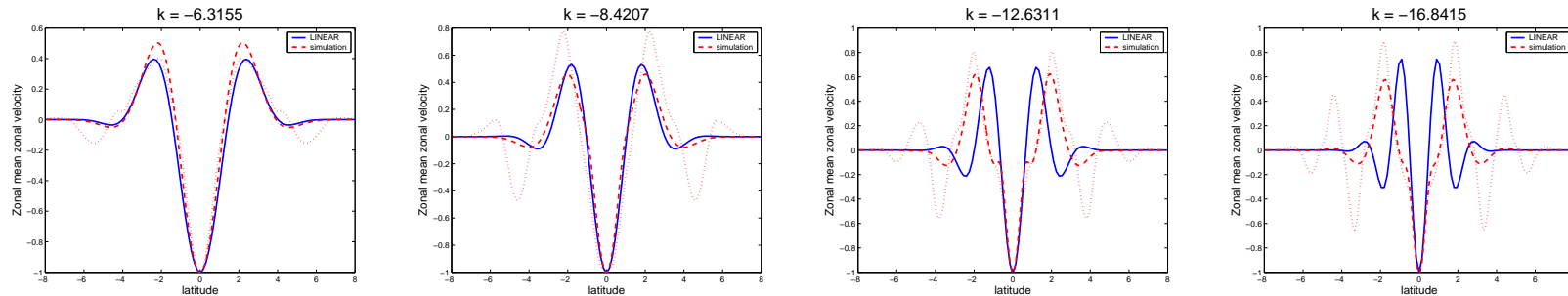
$k = -16.8$

- Jet spacing wider than wavelength of MRG wave
- Spacing less than for plane barotropic wave.

Simulations at $0.1^\circ \times 0.1^\circ$, 100-200 vertical levels using ROMS model:

Amplitude 0.36 cm, $k = -6.3$ (3.3° wavelength)

Comparison with simulations

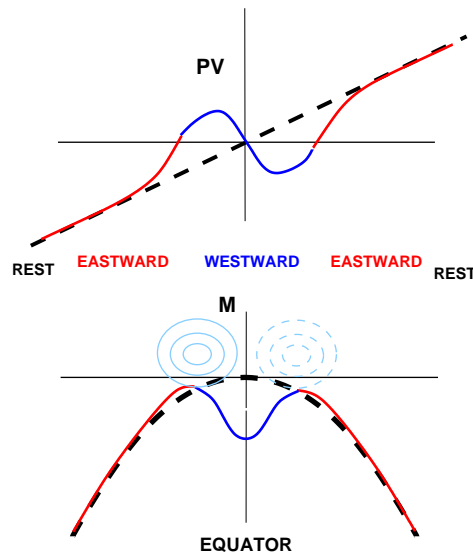


- Simulations **even** (in latitude) because odd modes are **inertially unstable**
- Equatorial jet always **westward**. Probably due to mixing of planetary angular momentum by short modes (which are strongest at equator) biasing simulations towards westward flow at equator.
- Extra-equatorial jet positioning poleward of in most unstable linear mode : **barotropic stability?**

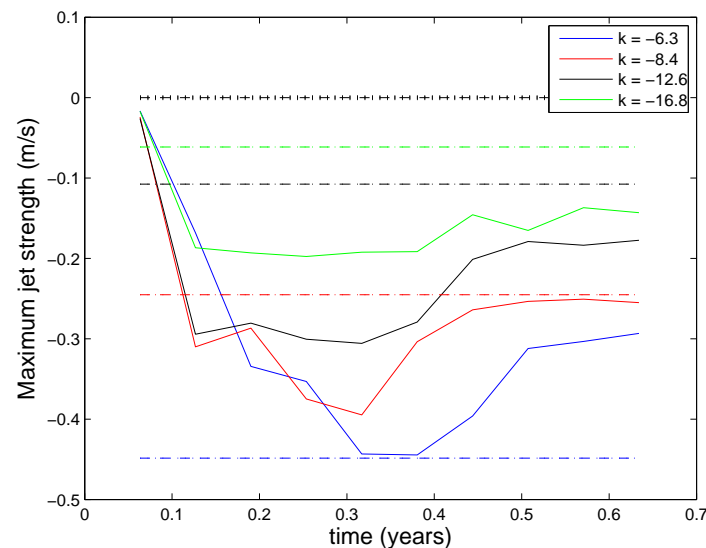
3. Inertial Instability and PV Homogenization

- inertial instability occurs when angular momentum increases poleward, i.e. if $f(PV) < 0$
- linear mode will have curvature greater than β at some moment, at which point flow will become inertially unstable

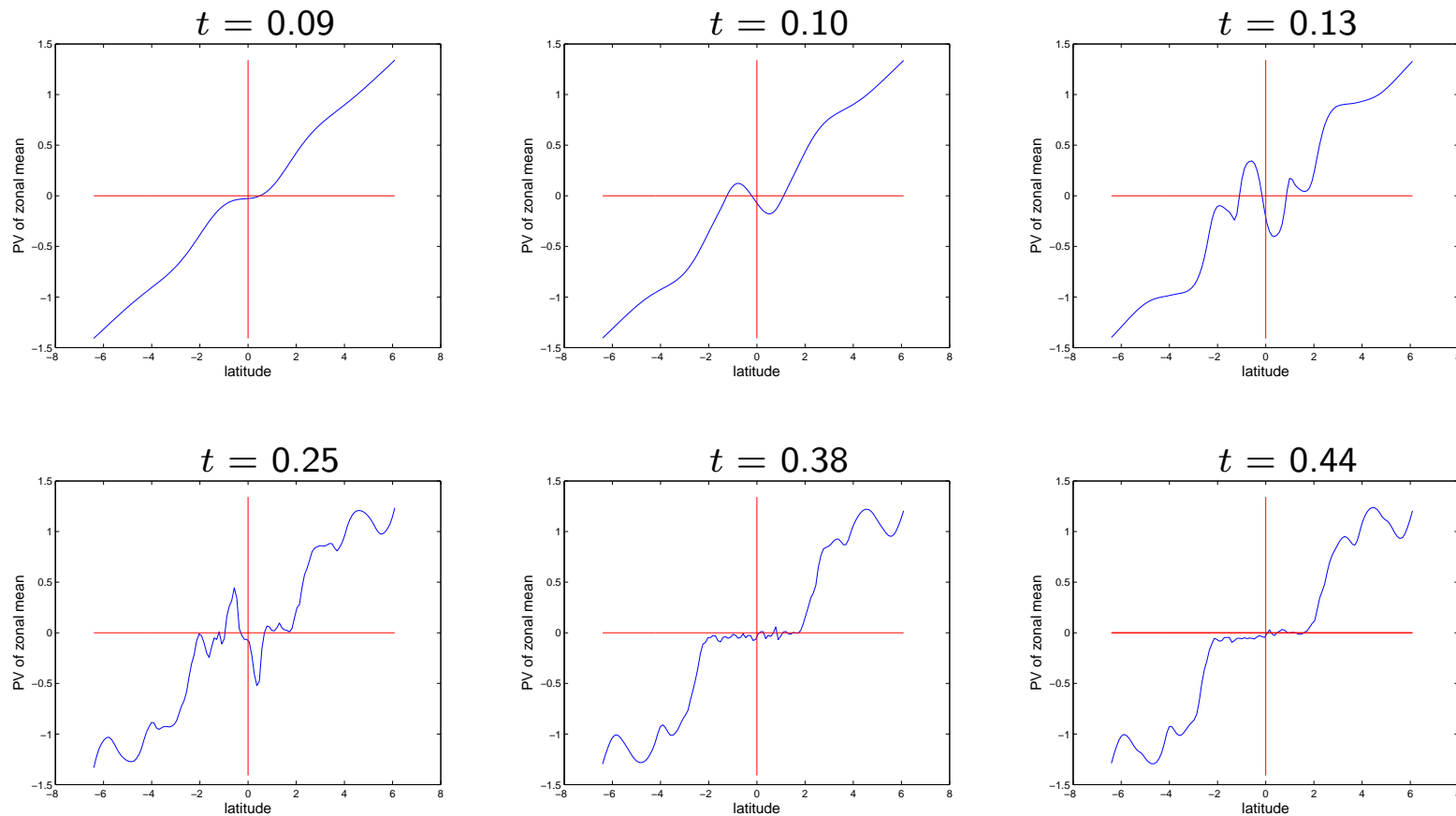
“Curvature”
inertial instability



Jet amplitude and inertial instability thresholds of linear modes:



Time evolution of zonal mean PV:



Note equatorially symmetric **inertial instability** ($t > 0.10$) and subsequent adjustment to uniform zero PV over interval wider than initially unstable interval (implies a widening and/or weakening of westward jet).

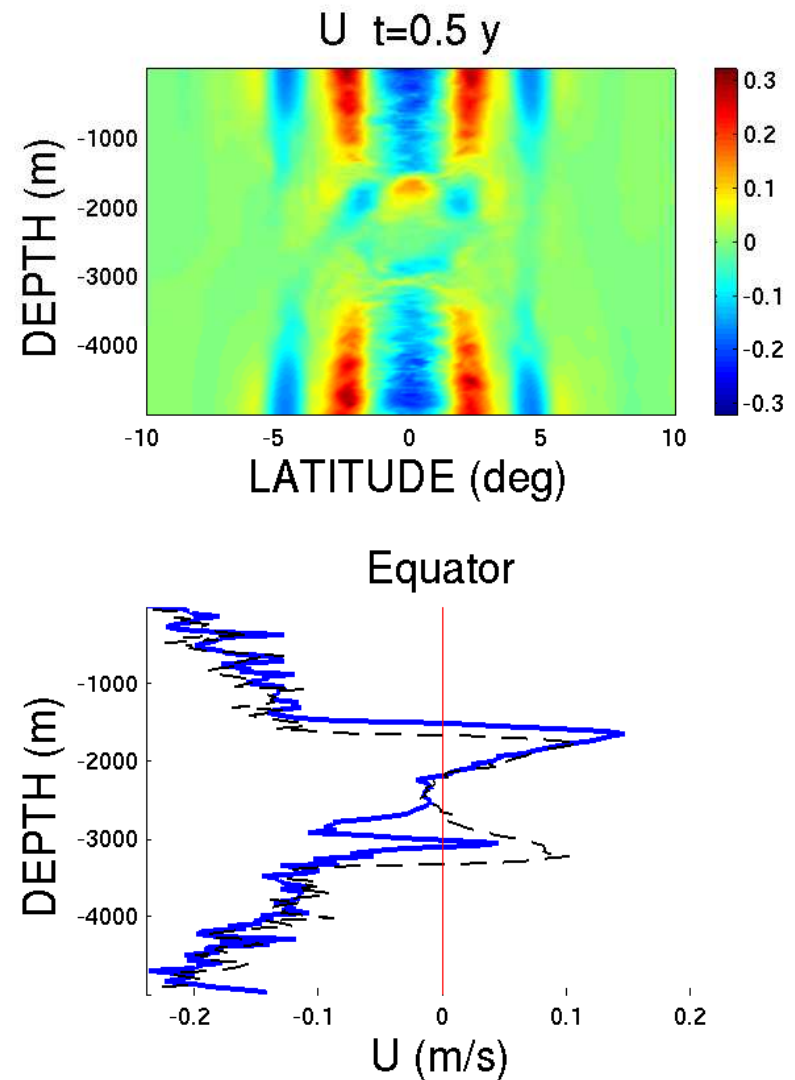
4. Super-rotation and the non-traditional Coriolis force

Amplitude 0.36 cm, $k = -6.3$ (3.3° wavelength)

- Clue may be effect on super-rotating jet of non-traditional Coriolis force terms (much greater than $\mathcal{O}(\Omega/N)$).
- What is reason for asymmetric vertical redistribution of angular momentum

$$M_\gamma \equiv U - \frac{1}{2}\beta y^2 + \gamma z ?$$

- Explanation in terms of lateral mixing of PV extending beyond unstable zone in both depth and latitude?



Summary

- **Zonal jets** with realistic spatial scales can be generated by the **destabilization of mixed Rossby-gravity waves** of frequencies consistent with forcing on equatorial ocean.
- Considerations of inertial stability and planetary potential vorticity mixing explain sign and position of equatorial jet.
- Shape of equatorial westward jet well described by truncated Floquet solution to linearized barotropic vorticity equation.
- Equatorial jet saturates at **lower** amplitude for **more unstable** initial waves – consistent with saturation due to **inertial instability**.
- **Super-rotating equatorial jets** form for very short waves, phenomenon particularly sensitive to traditional Coriolis force approximation.