

# USING HATS DATABASES TO EVALUATE SUBFILTER-SCALE RATE EQUATIONS FOR LES

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# LARGE-EDDY SIMULATIONS AND OBSERVATIONS

## Why LES of the PBL:

- Outdoor 4-D measurements are challenging
- Unsteady nature of the atmosphere and ocean
- Systematic investigation of the parameter space
- *Advances in parallel computing*

## Validating/Improving LES with observations:

- Test the output
- Test the input subgrid-scale parameterizations



Gulf of Tehuantepec U ~[20-25] m/s

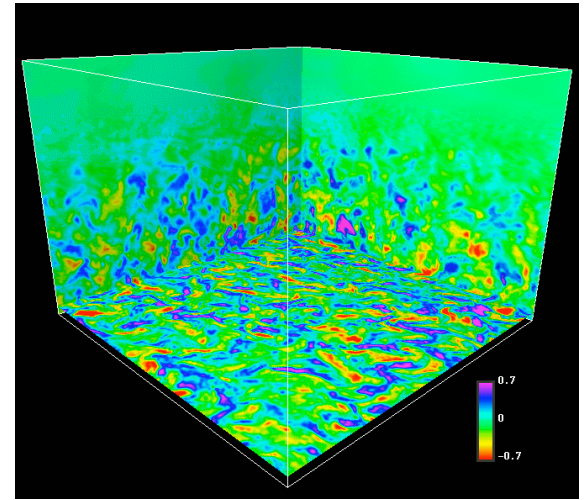
## LES APPLICATIONS AND THE PBL

- Turbulence dynamics, stratification, entrainment
- Surface-atmosphere interactions
- Dispersion, chemistry
- Clouds
- ...

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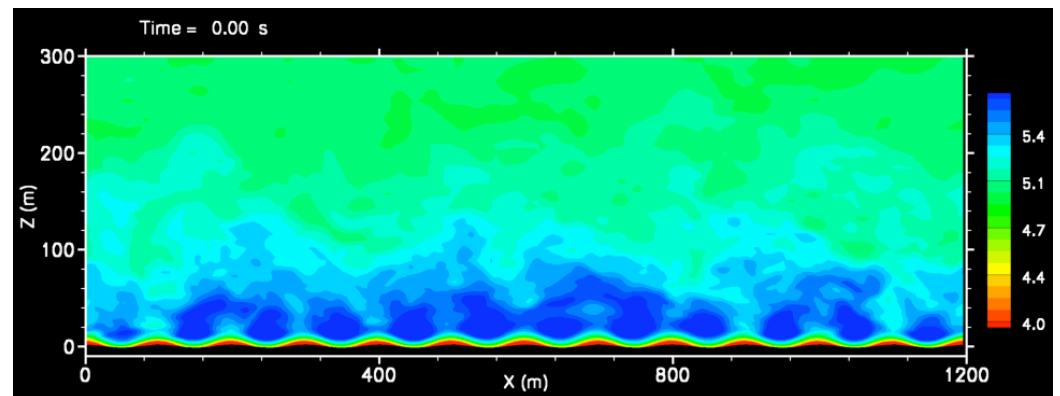
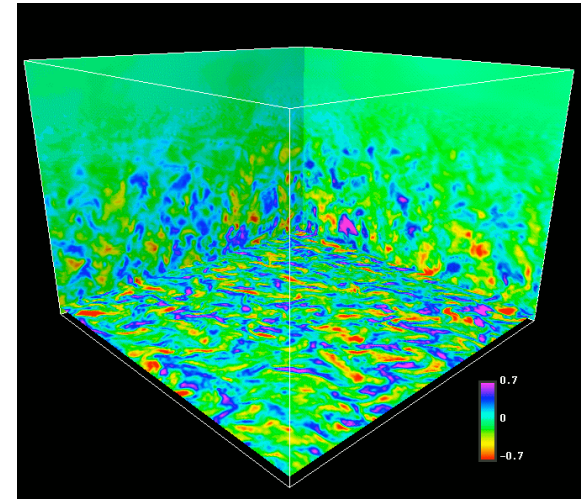
Stable boundary layers  $z_i/L \sim 1.2$



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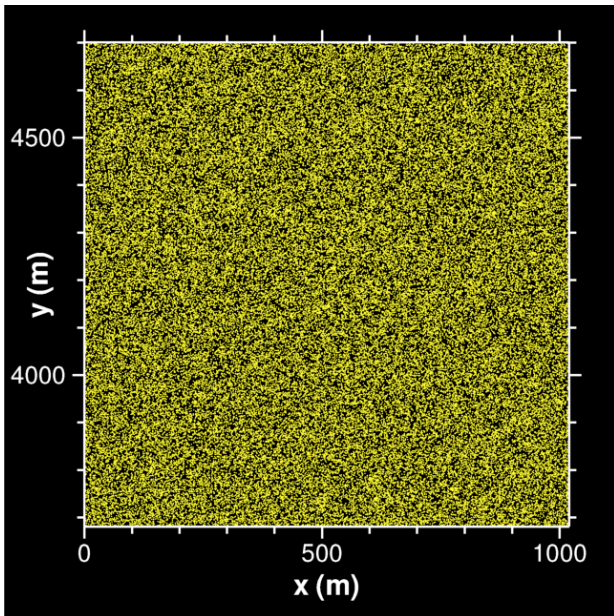
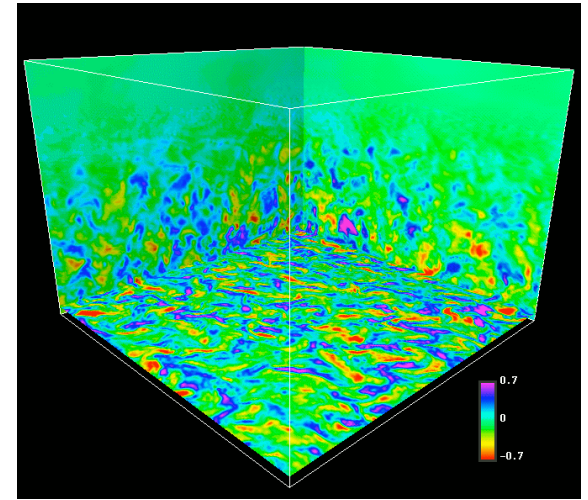


Non-equilibrium winds and waves

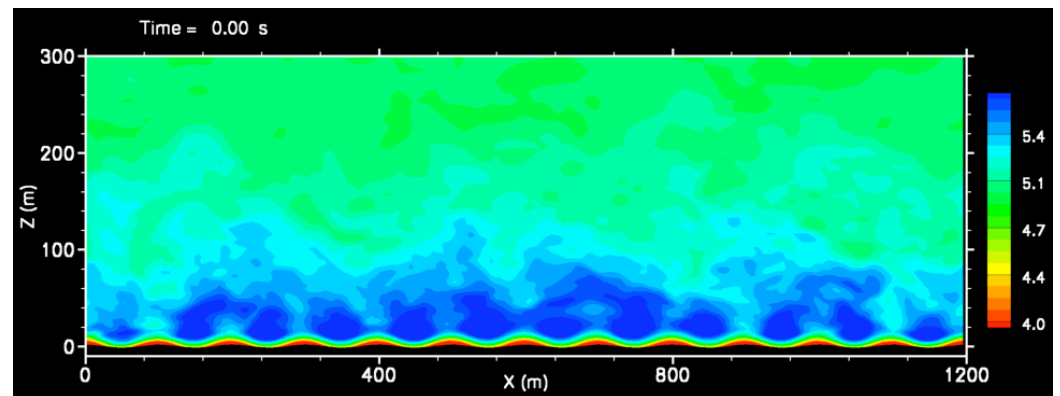
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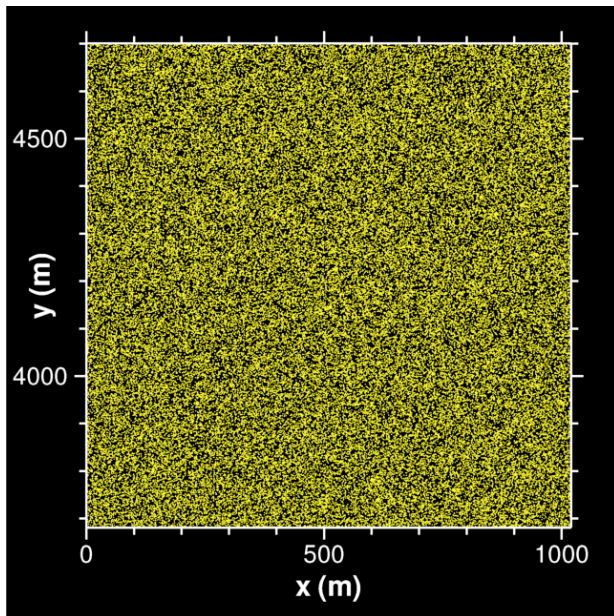
Plumes and dust-devils in a convective PBL  $1024^3$  simulation on 4096 cpus



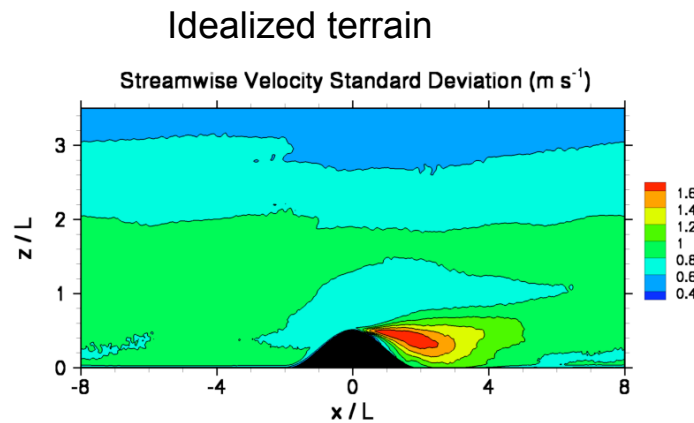
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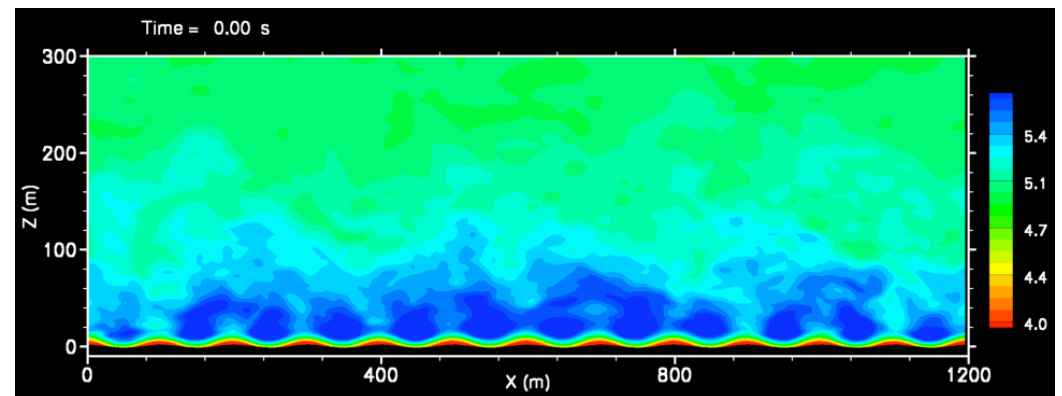
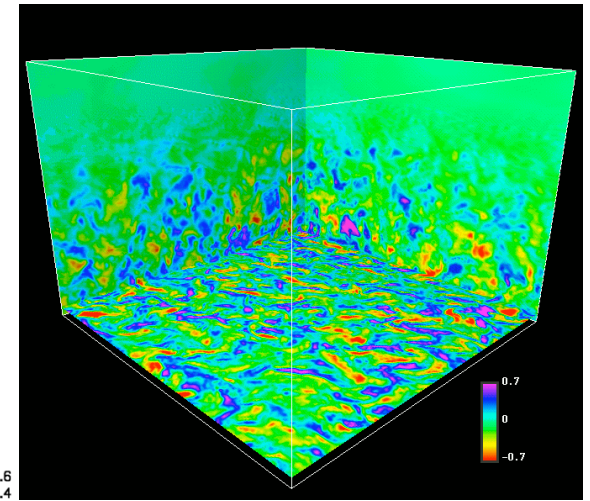
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Plumes and dust-devils in a convective PBL  $1024^3$  simulation on 4096 cpus



Stable boundary layers  $z_i/L \sim 1.2$



Non-equilibrium winds and waves

# LES EQUATIONS FOR DRY ATMOSPHERIC PBL

Momentum

$$\frac{D\bar{\mathbf{u}}}{Dt} = -\mathbf{f} \times \bar{\mathbf{u}} - \nabla\pi + \hat{\mathbf{z}}g\frac{\bar{\theta}}{\theta_*} - \nabla \cdot \mathbf{T}$$

Scalar

$$\frac{D\bar{b}}{Dt} = -\nabla \cdot \mathbf{B}$$

TKE

$$\frac{De}{Dt} = -\mathbf{T} : \mathbf{S} + \mathbf{B} \cdot \hat{\mathbf{z}} - \mathcal{E} + \nabla \cdot (2\nu_t \nabla e)$$

Subgrid-scale momentum and scalar fluxes

$$\begin{aligned}\mathbf{T} &= \overline{u_i u_j} - \overline{u_i} \overline{u_j} \\ \mathbf{B} &= \overline{u_i b} - \overline{u_i} \overline{b}\end{aligned}$$

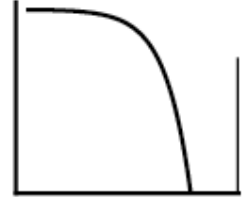
Random variables, require a parameterization



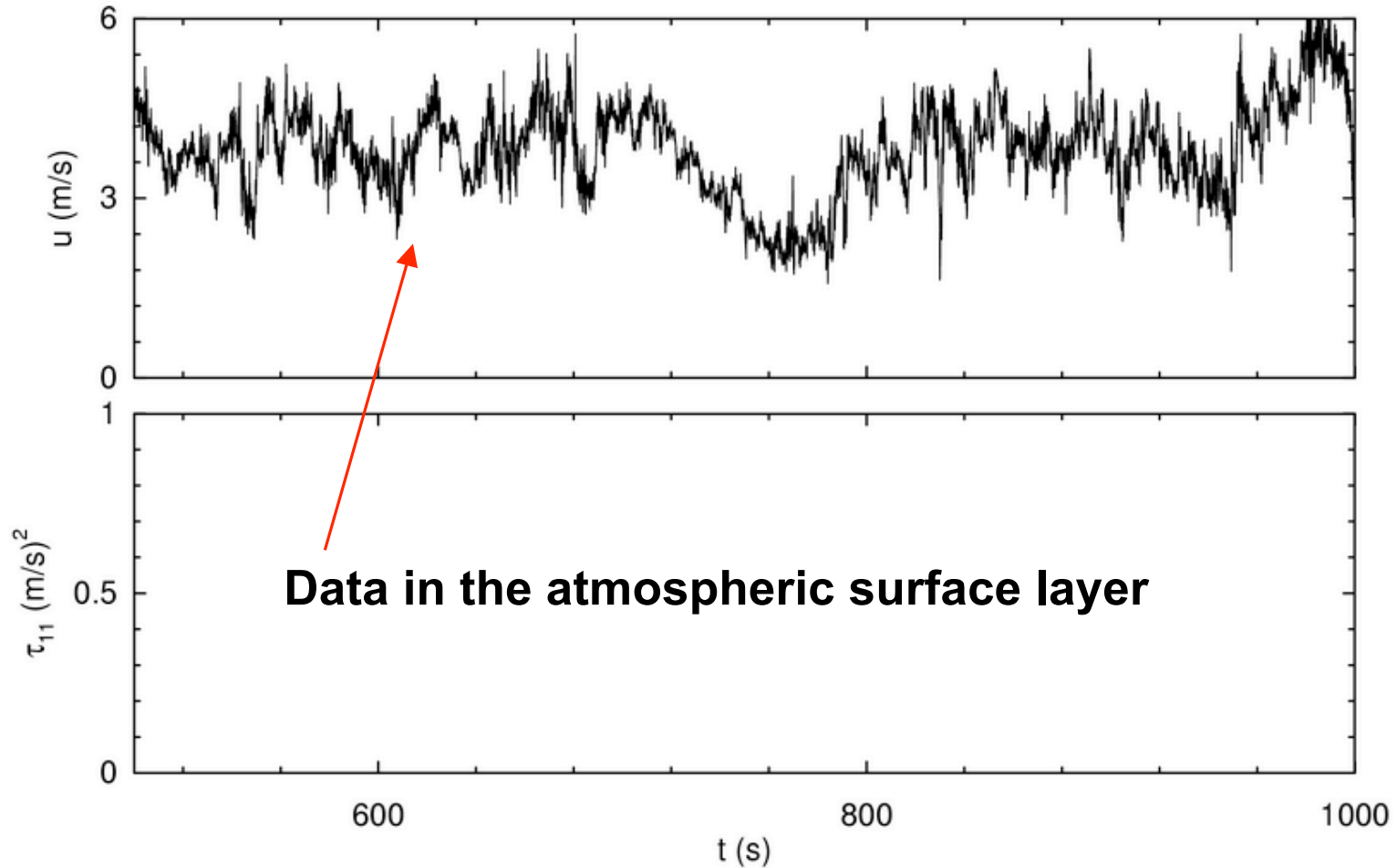
***SIMPLE (CHEAP) FILTERING  
EXAMPLE ...***

$$\tau_{11} = \overline{u_1 u_1} - \bar{u}_1 \bar{u}_1$$

# MOVING BETWEEN DNS $\iff$ LES $\iff$ RANS



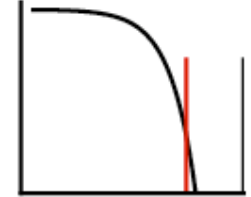
What happens to  $\bar{u}_i$  and  $\mathcal{T}_{ij}$  as we vary the filter cutoff  $k_c$  ?



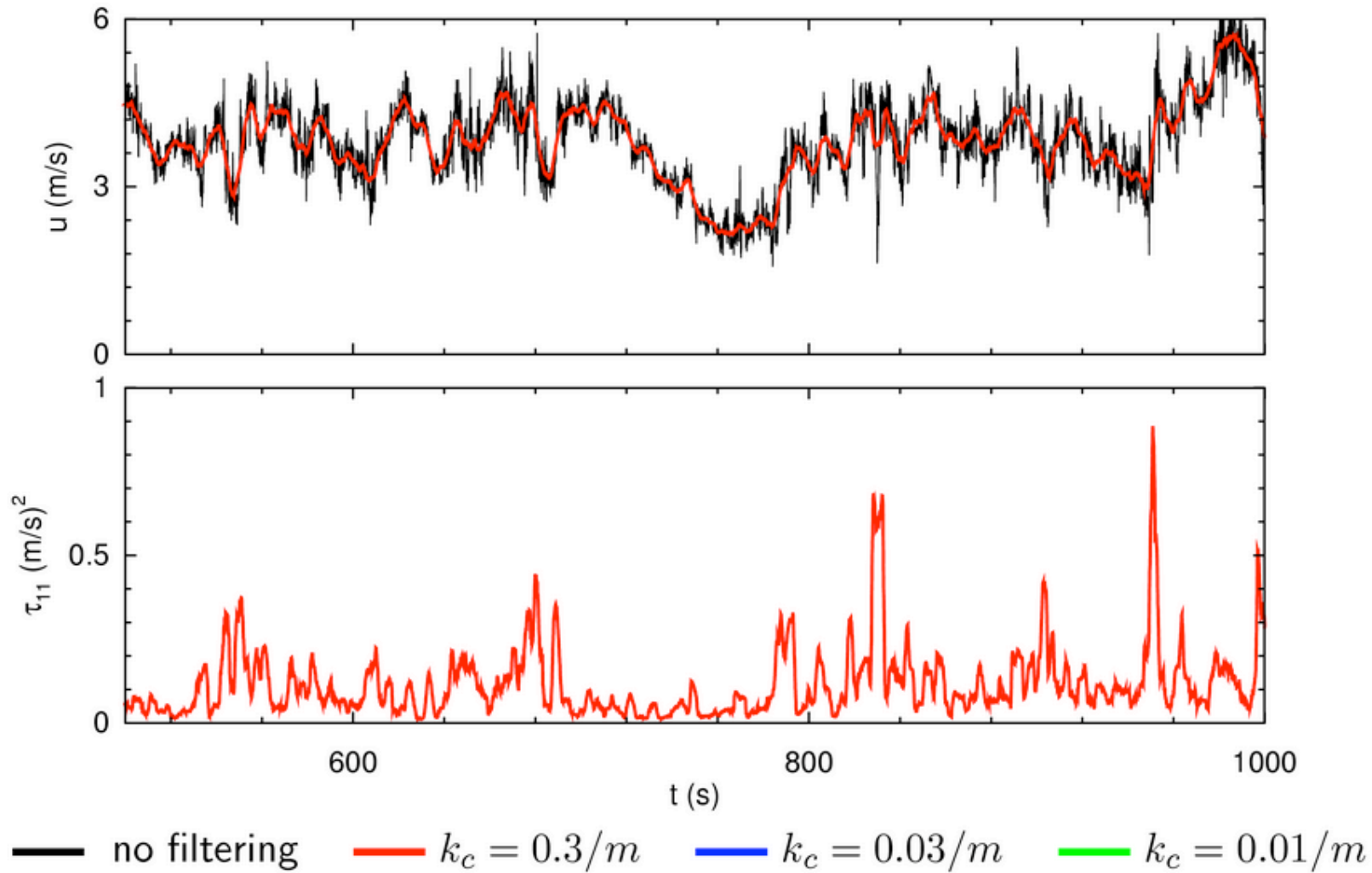
**Data in the atmospheric surface layer**

— no filtering    —  $k_c = 0.3/m$     —  $k_c = 0.03/m$     —  $k_c = 0.01/m$

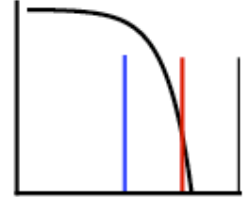
# MOVING BETWEEN DNS $\iff$ LES $\iff$ RANS



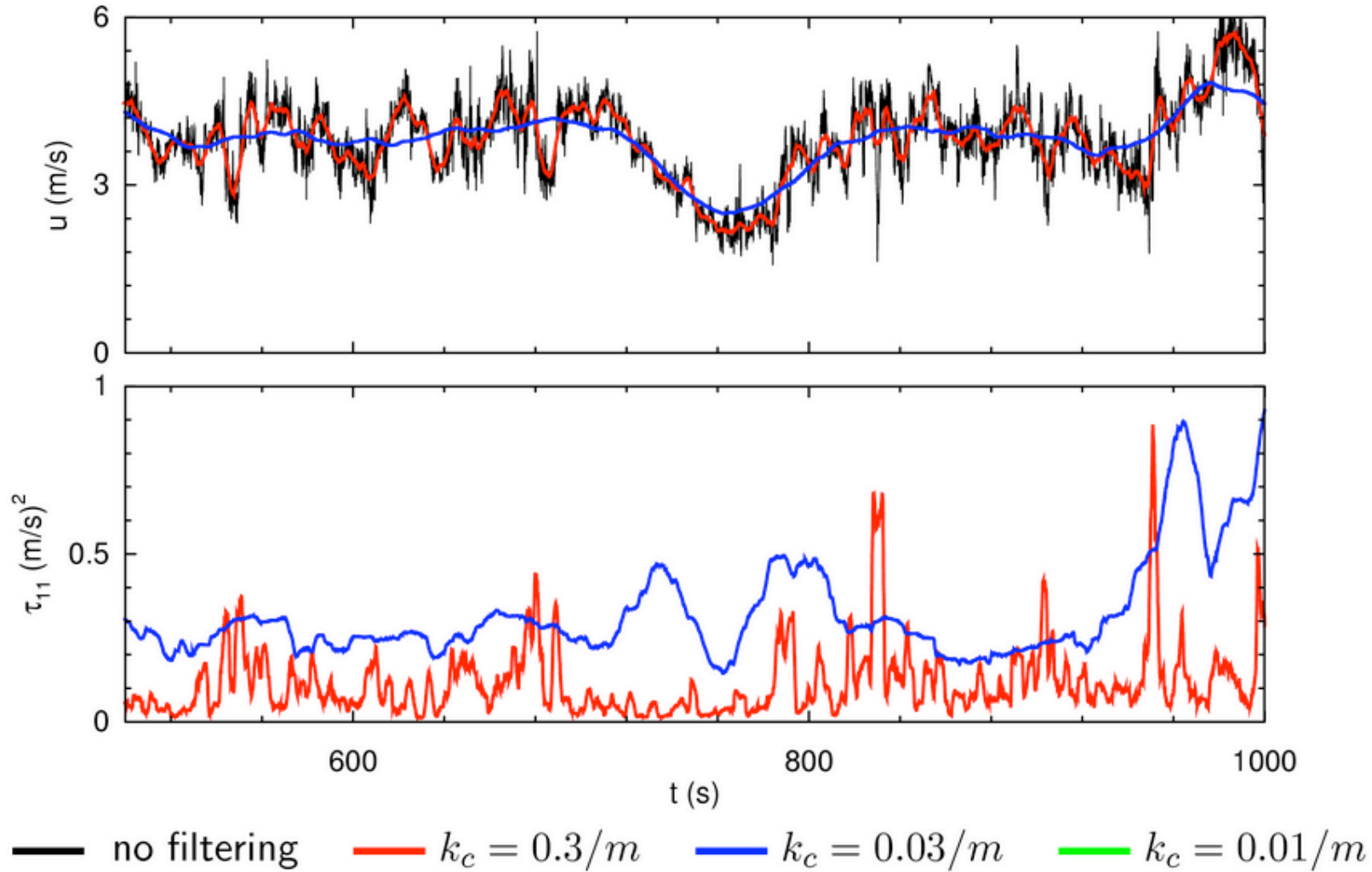
What happens to  $\bar{u}_i$  and  $\mathcal{T}_{ij}$  as we vary the filter cutoff  $k_c$  ?



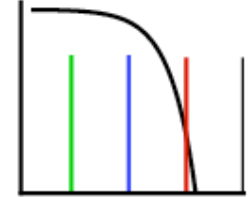
# MOVING BETWEEN DNS $\iff$ LES $\iff$ RANS



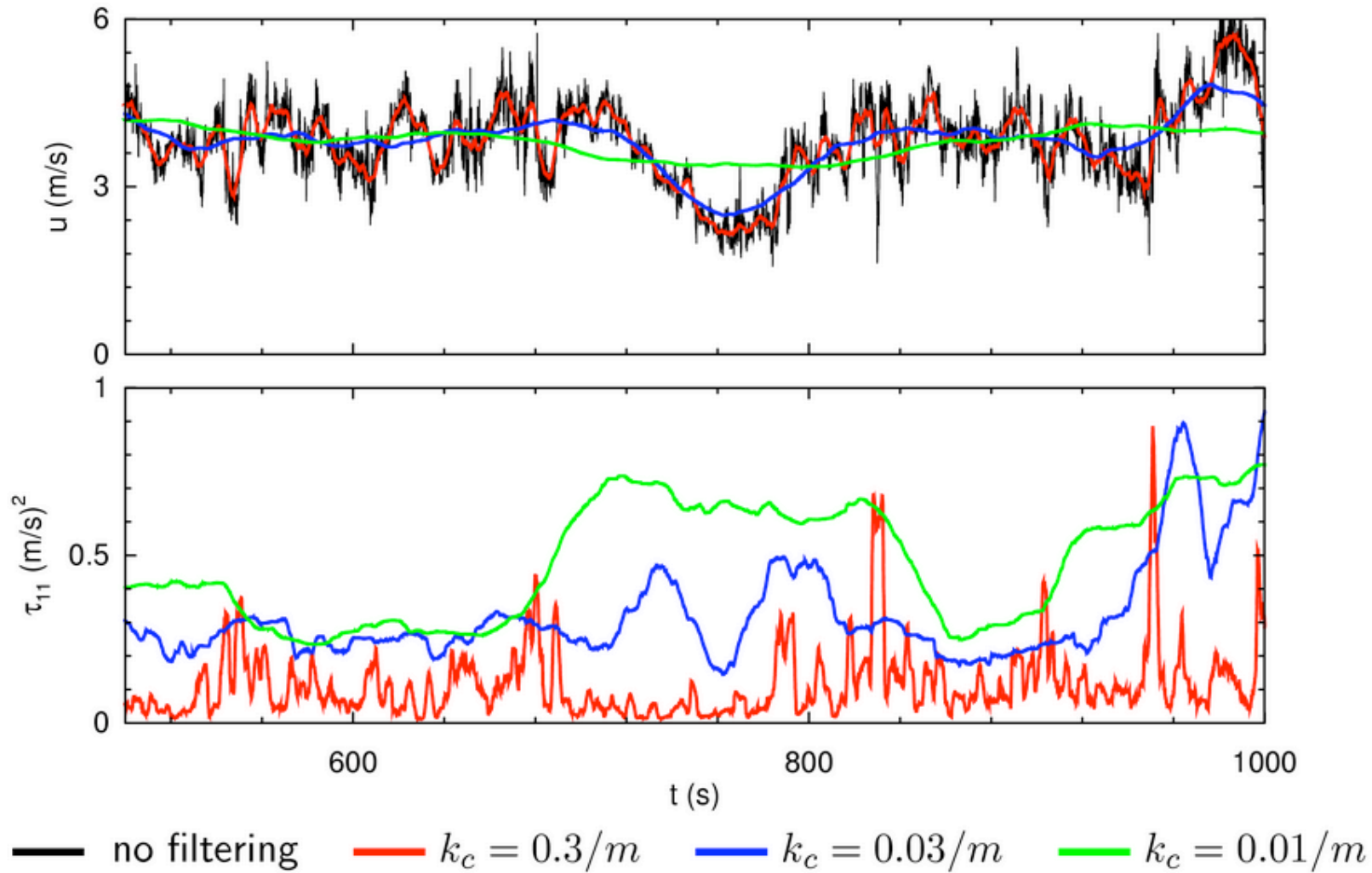
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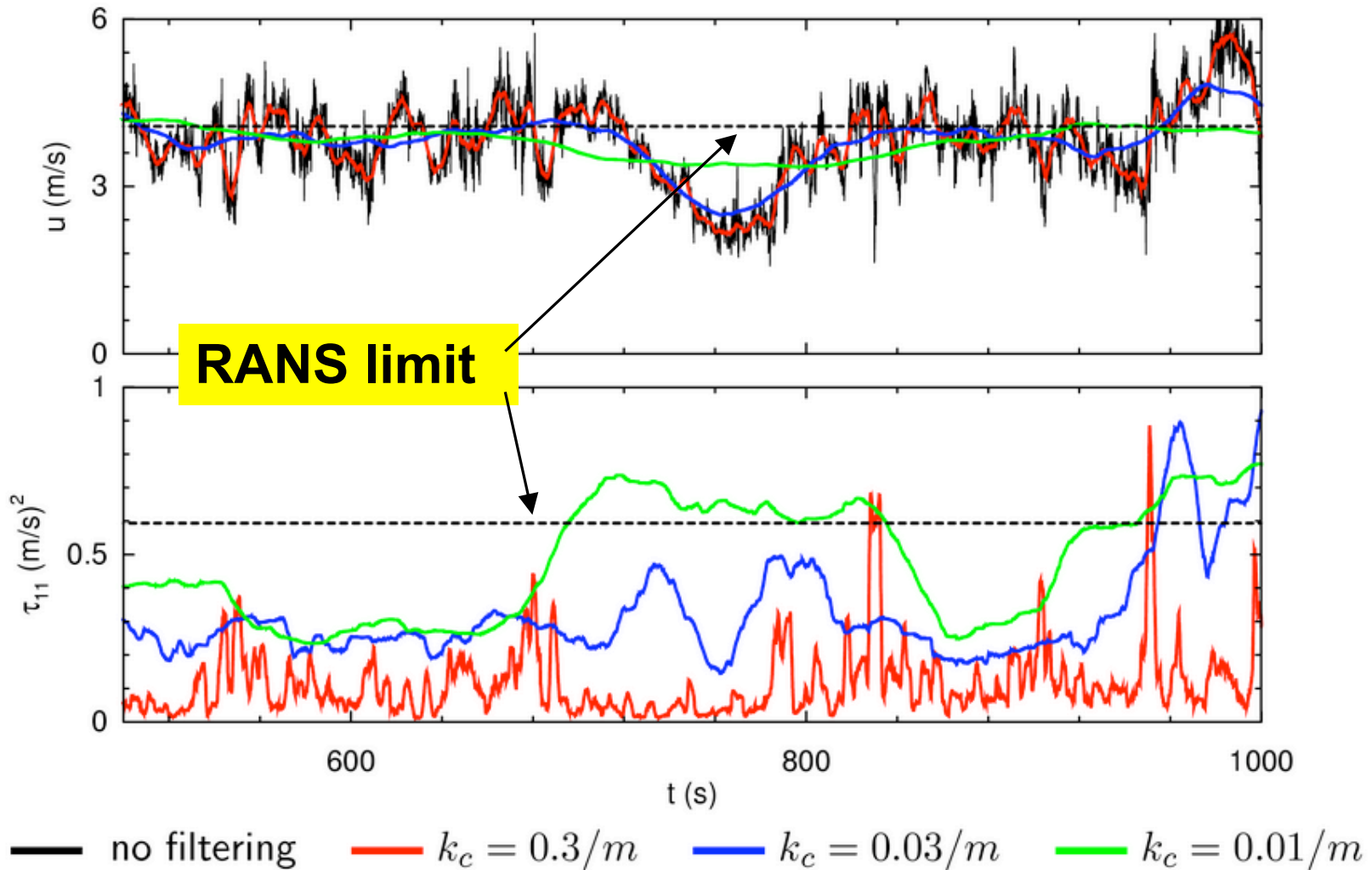


What happens to  $\bar{u}_i$  and  $\mathcal{T}_{ij}$  as we vary the filter cutoff  $k_c$  ?



# MOVING BETWEEN DNS $\iff$ LES $\iff$ RANS

What happens to  $\bar{u}_i$  and  $\mathcal{T}_{ij}$  as we vary the filter cutoff  $k_c$  ?



# HIGH REYNOLDS NUMBER OBSERVATIONS AND LES

- **SINGLE-POINT MEASUREMENTS**

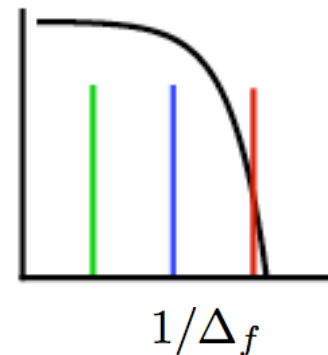
- Cannot be used directly to improve LES

- **MULTI-POINT MEASUREMENTS**

- Span a range of filter widths, *e.g.*,  $\mathcal{O}(\text{m})$  to  $\mathcal{O}(100\text{m})$

- Ideally 3-D, time varying “volume” of turbulence and scalars in canonical flows with shear, stratification, near boundaries, ...

- Horizontal Array Turbulence Study field campaigns, HATS (2000), OHATS (2004), CHATS (2007), AHATS (2008)

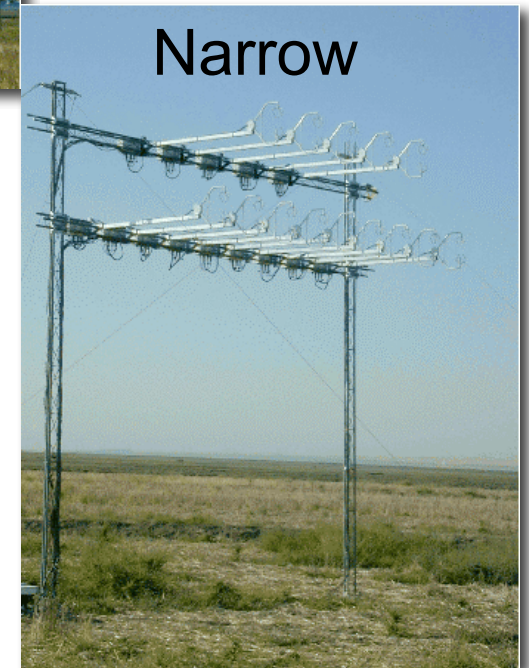
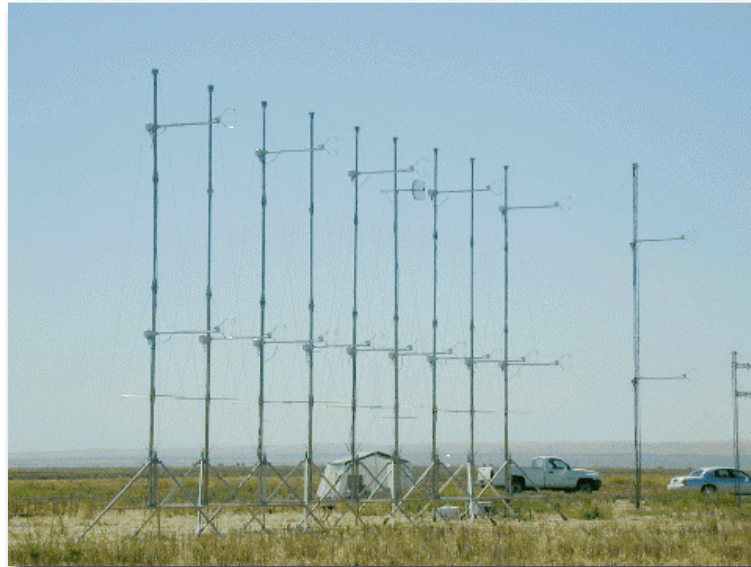


# HATS CONFIGURATIONS

$\sim 36$  cases

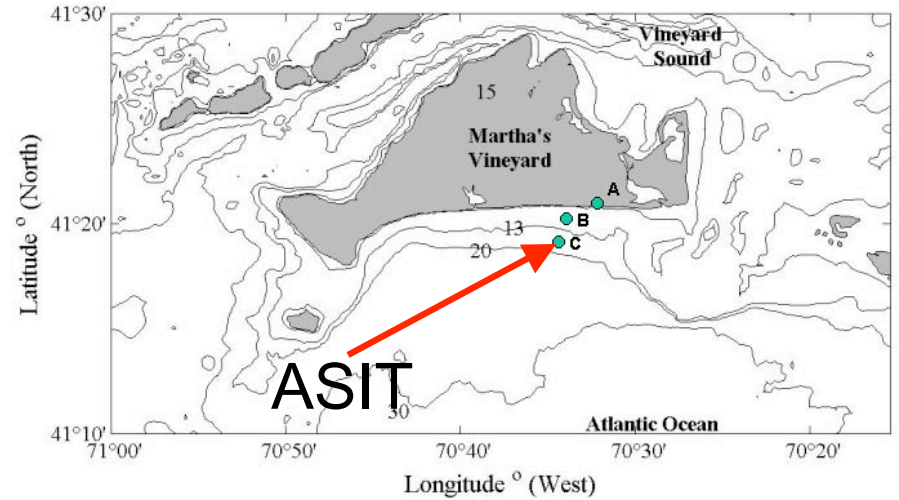
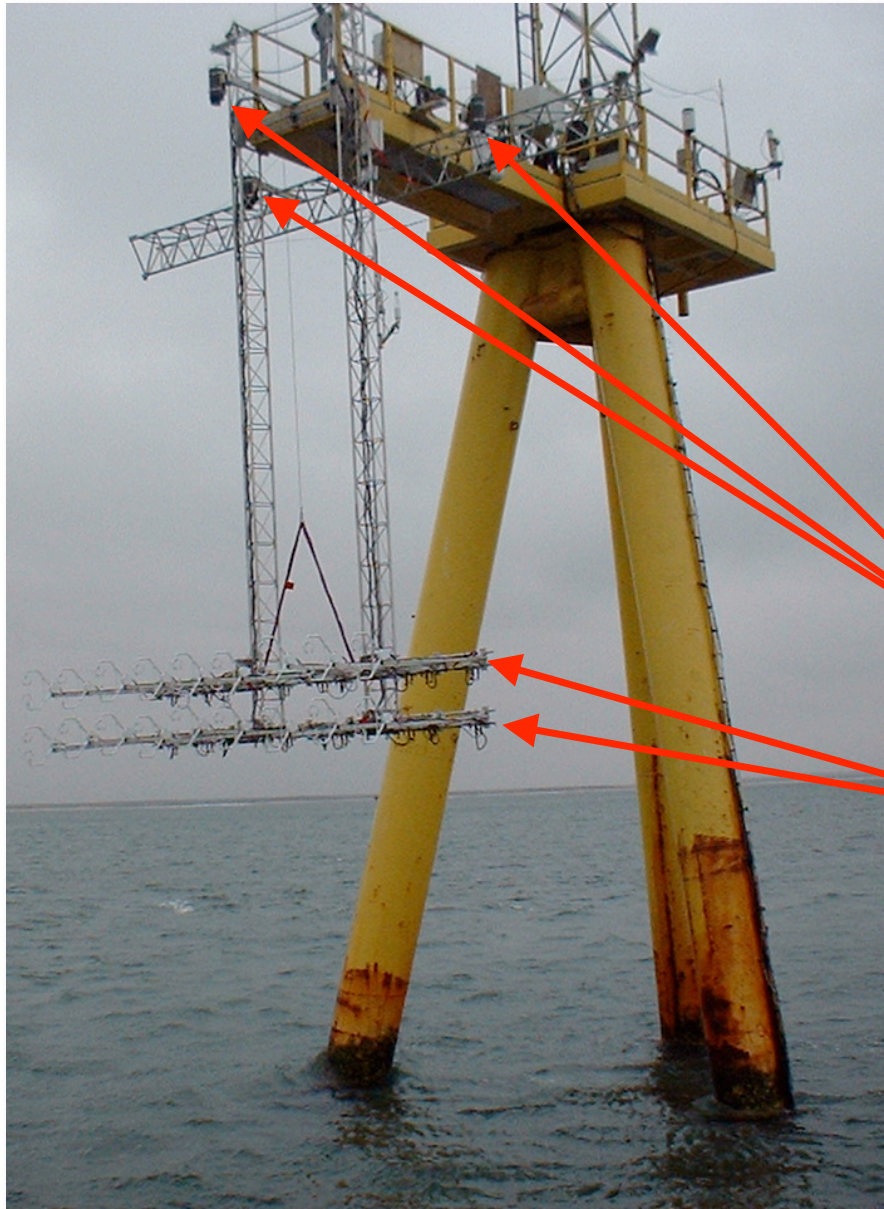
$$-1.2 < z/L < 1.6$$

$$0.15 < \Lambda_w/\Delta_f < 15$$





# OHATS FIELD CAMPAIGN



Laser altimeters

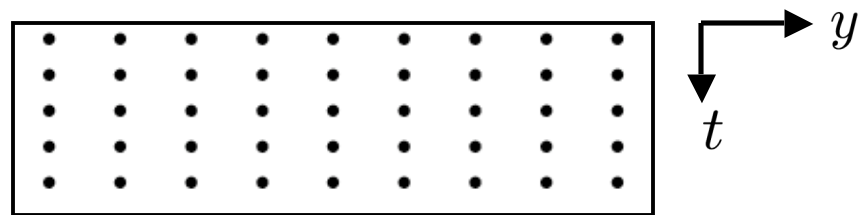
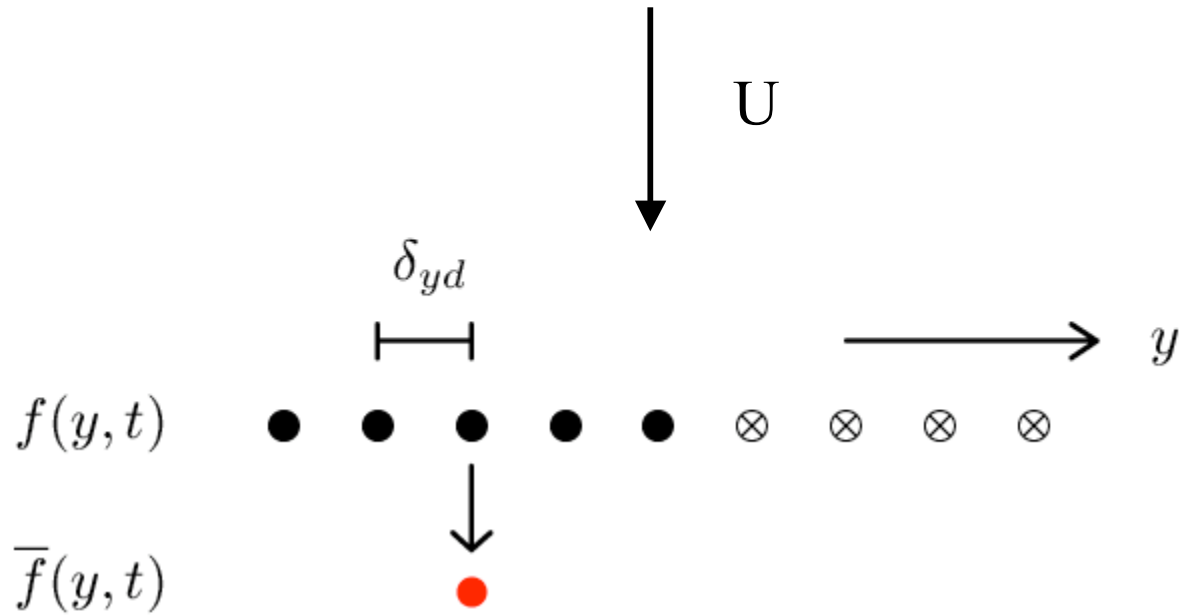
18 CSATS

275 hours ``12 days of data''  
analyzed

# CANOPY HORIZONTAL ARRAY TURBULENCE STUDY

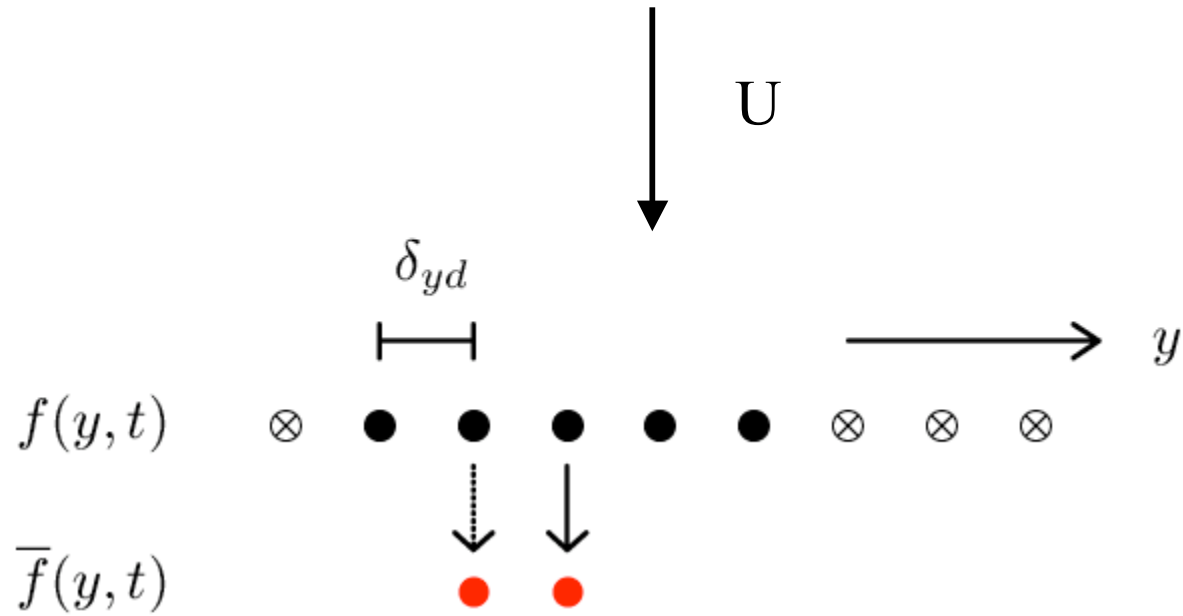


# AN EXAMPLE OF LATERAL (Y) FILTERING

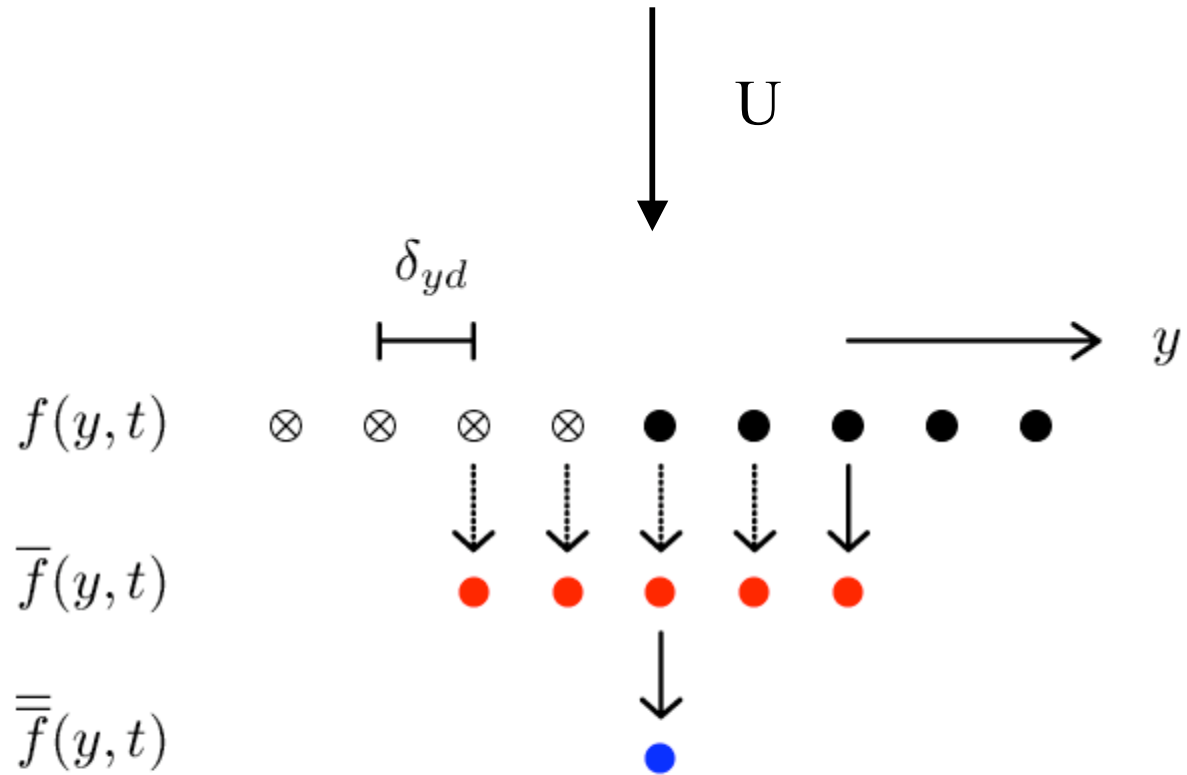


“2D plane of turbulence”

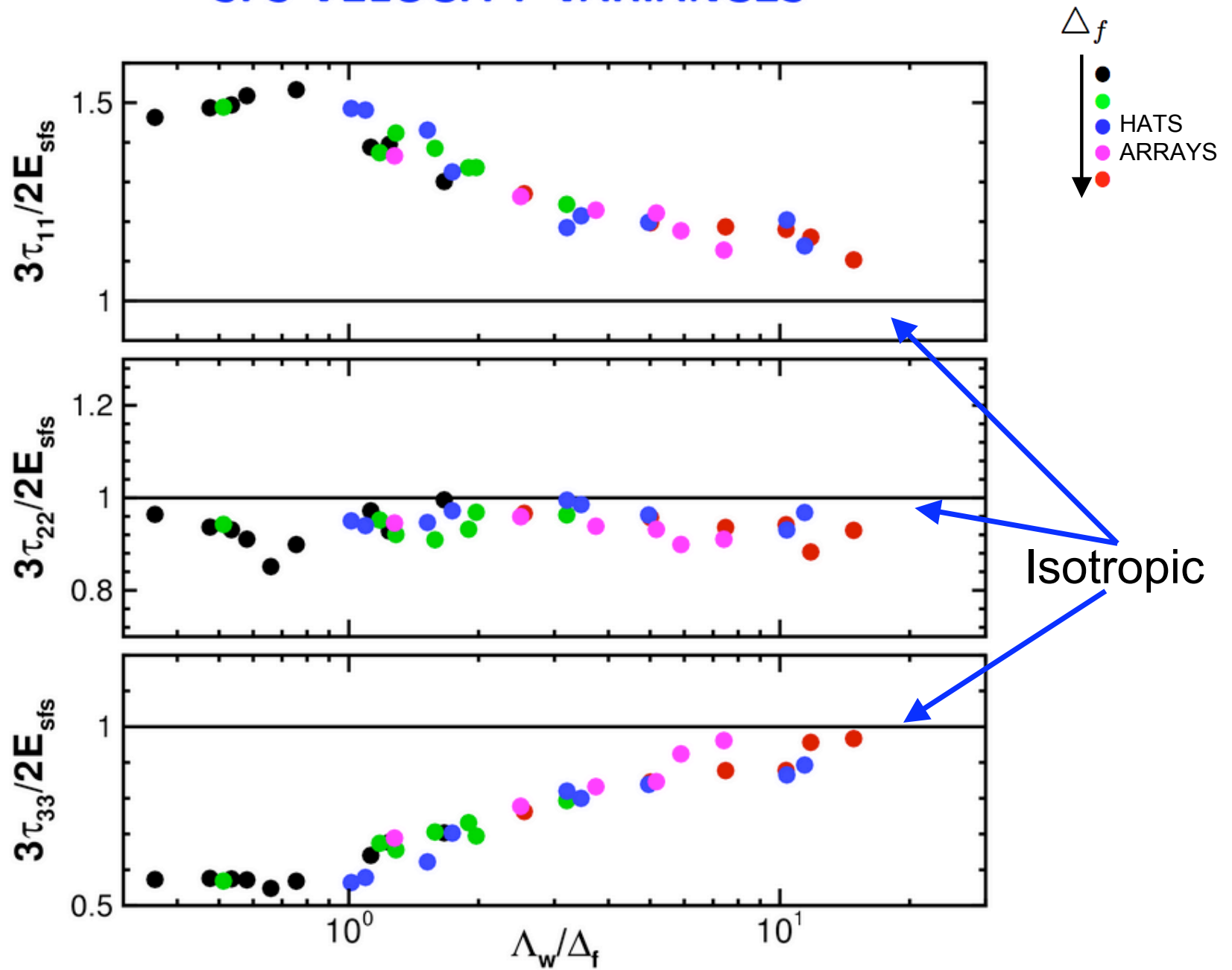
# AN EXAMPLE OF LATERAL (Y) FILTERING



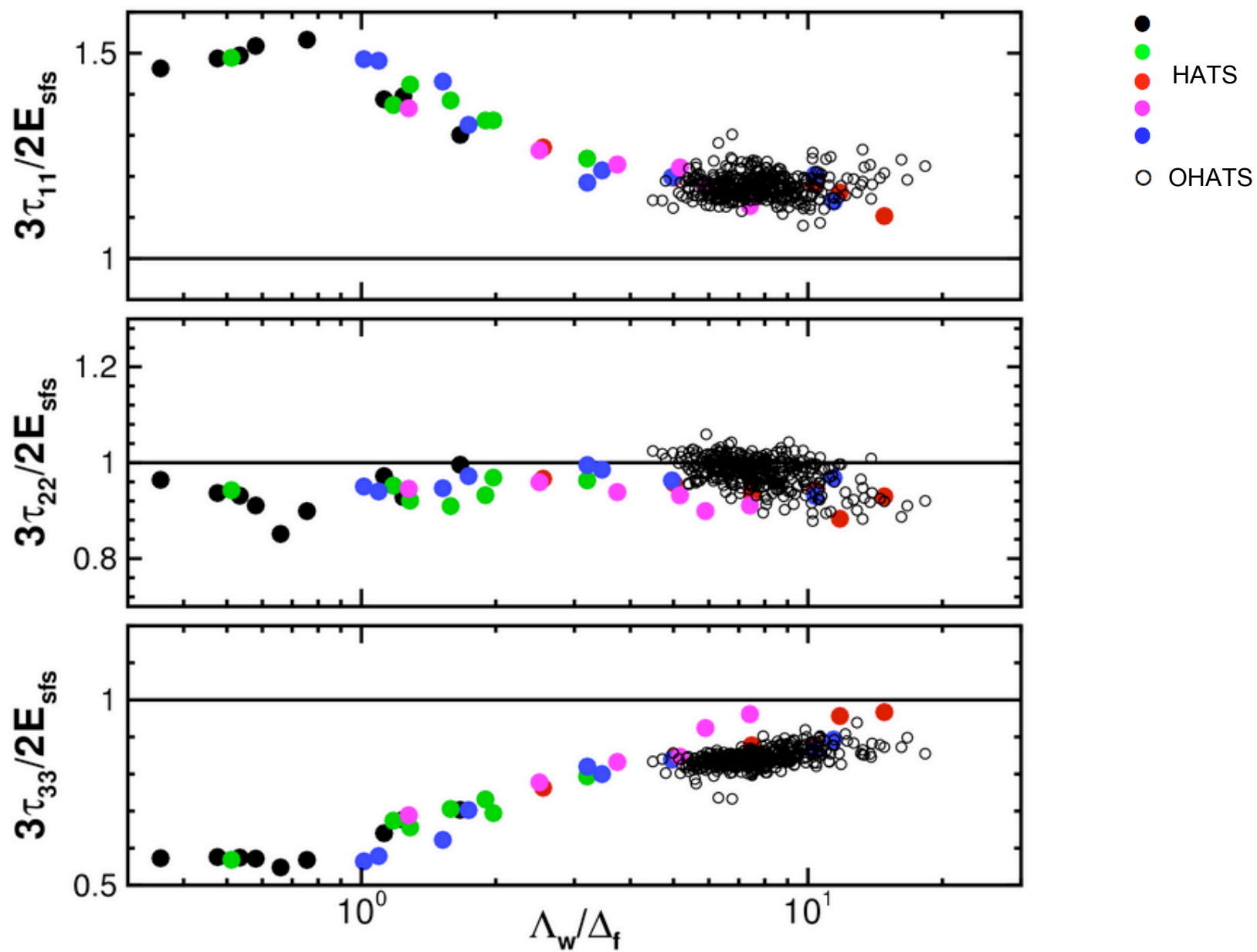
# AN EXAMPLE OF LATERAL (Y) FILTERING



# SFS VELOCITY VARIANCES



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## RATE EQUATIONS FOR SUBGRID DEVIATORIC STRESS

- What are the parent equations for the Smagorinsky model?



# RATE EQUATIONS FOR SUBGRID DEVIATORIC STRESS

- **What are the parent equations for the Smagorinsky model?**

- Lilly (1967), Deardorff (1973), Wyngaard (2004), Hatlee & Wyngaard (2007)

$$\begin{aligned}
 \frac{D\tau_{ij}}{Dt} = & \frac{2}{3}e \left( \frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right) \leftarrow \text{Isotropic production} \\
 & - \left[ \tau_{ik} \frac{\partial \bar{u}_j}{\partial x_k} + \tau_{jk} \frac{\partial \bar{u}_i}{\partial x_k} - \frac{1}{3} \delta_{ij} \tau_{kl} \left( \frac{\partial \bar{u}_k}{\partial x_l} + \frac{\partial \bar{u}_l}{\partial x_k} \right) \right] \\
 & - \frac{1}{\rho} \left[ \overline{p \left( \frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)} - \bar{p} \left( \frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right) \right] \\
 & + \text{transport} + \text{buoyancy production}
 \end{aligned}$$

Pressure destruction

Anisotropic deviatoric production

# RATE EQUATIONS FOR SUBGRID DEVIATORIC STRESS

- **What are the parent equations for the Smagorinsky model?**

- Lilly (1967), Deardorff (1973), Wyngaard (2004), Hatlee & Wyngaard (2007)

$$\begin{aligned}
 \frac{D\tau_{ij}^0}{Dt} &= \frac{2}{3}e \left( \frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right) \\
 &\quad - \left[ \tau_{ik} \frac{\partial \bar{u}_j}{\partial x_k} + \tau_{jk} \frac{\partial \bar{u}_i}{\partial x_k} - \frac{1}{3} \delta_{ij} \tau_{kl} \left( \frac{\partial \bar{u}_k}{\partial x_l} + \frac{\partial \bar{u}_l}{\partial x_k} \right) \right] \\
 &\quad - \frac{1}{\rho} \left[ \overline{p \left( \frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)} - \bar{p} \left( \frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right) \right] \\
 &\quad + \text{transport}^0 + \text{buoyancy production}^0
 \end{aligned}$$

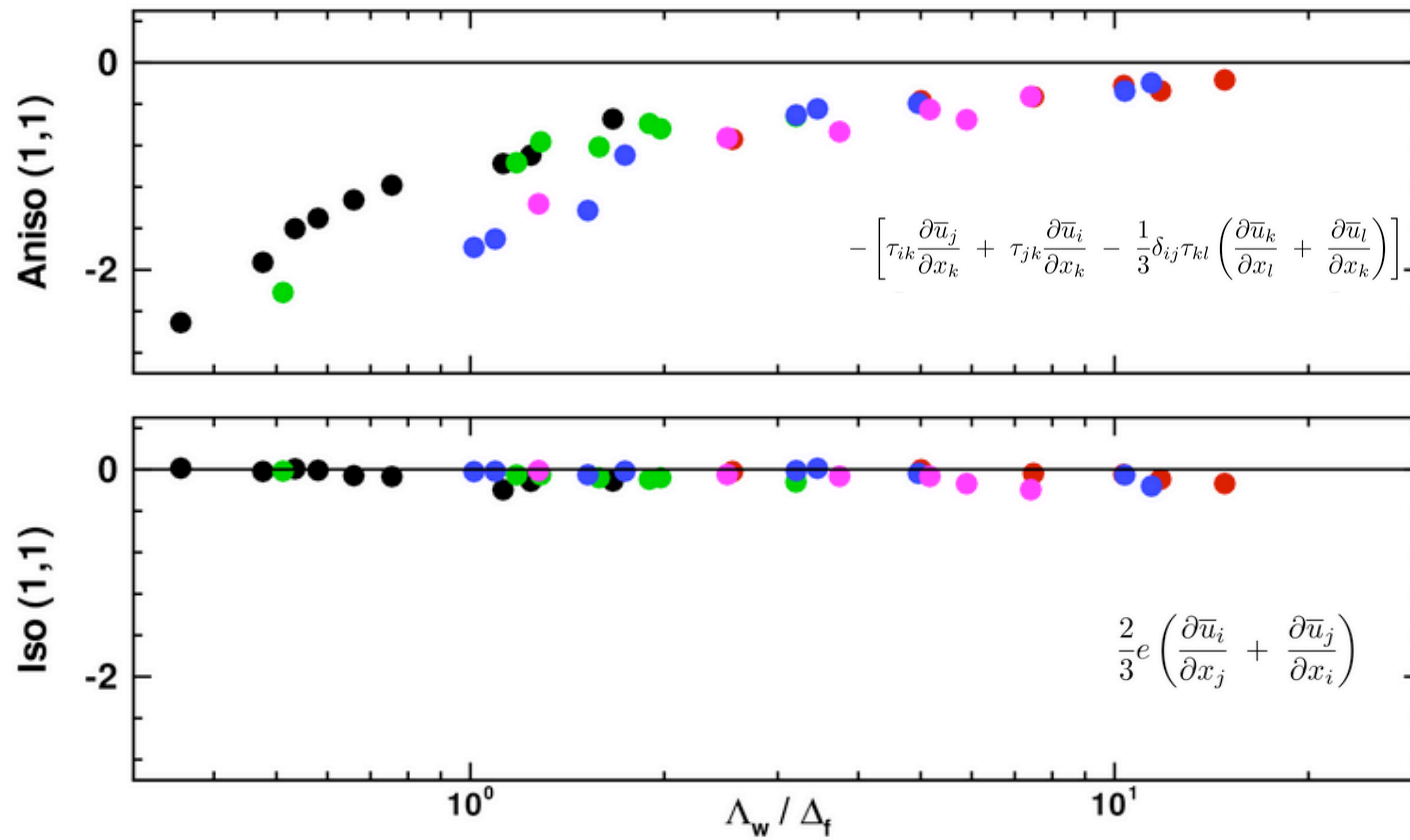
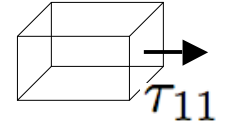
Rotta model

$$\frac{\tau_{ij}}{T} = \frac{2}{3}e \left( \frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right)$$

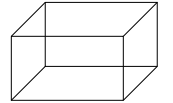
$$T = c \frac{\Delta_f}{\sqrt{e}}$$

Time scale

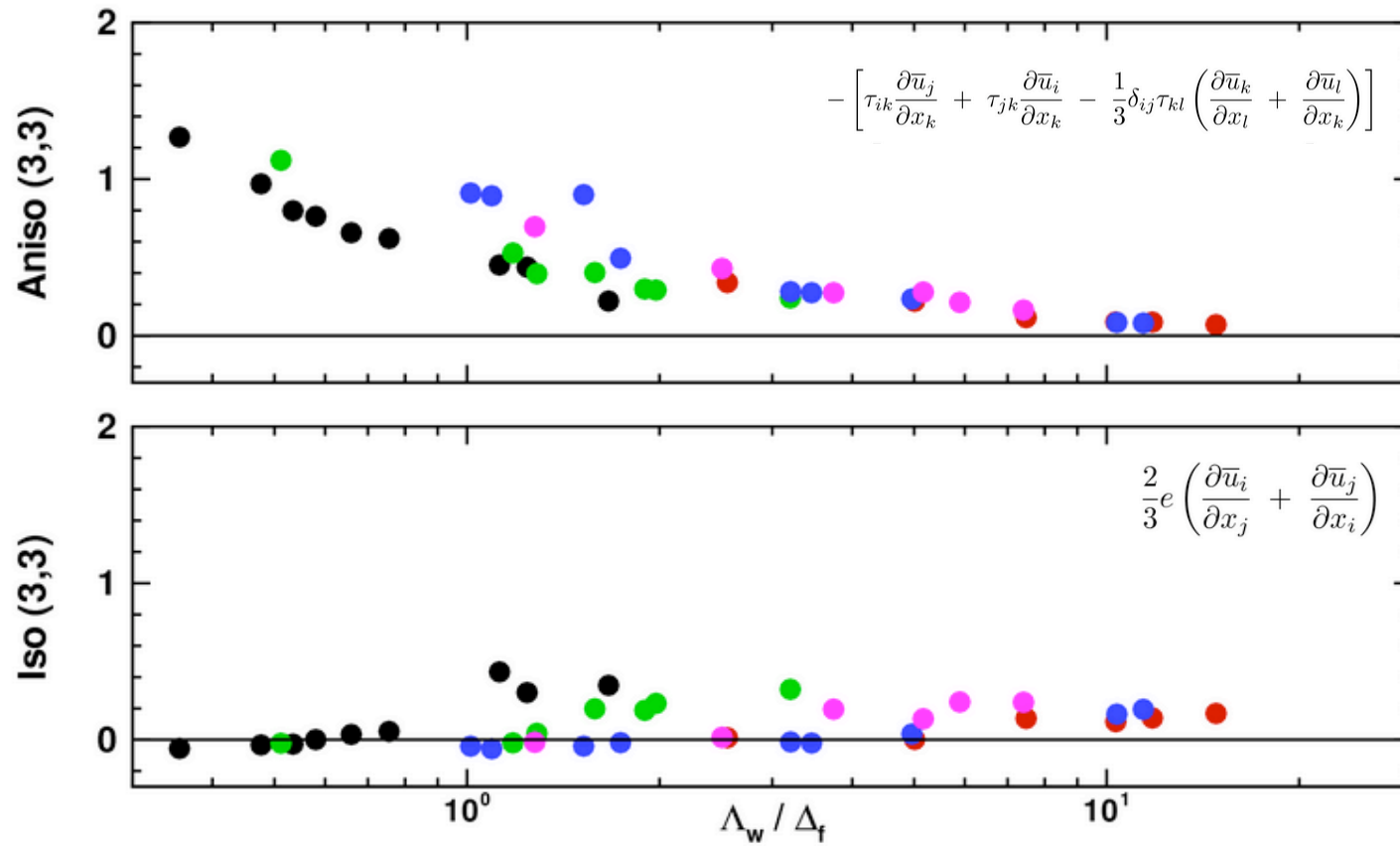
# PRODUCTION OF SUBFILTER SCALE FLUX $\tau_{11}$



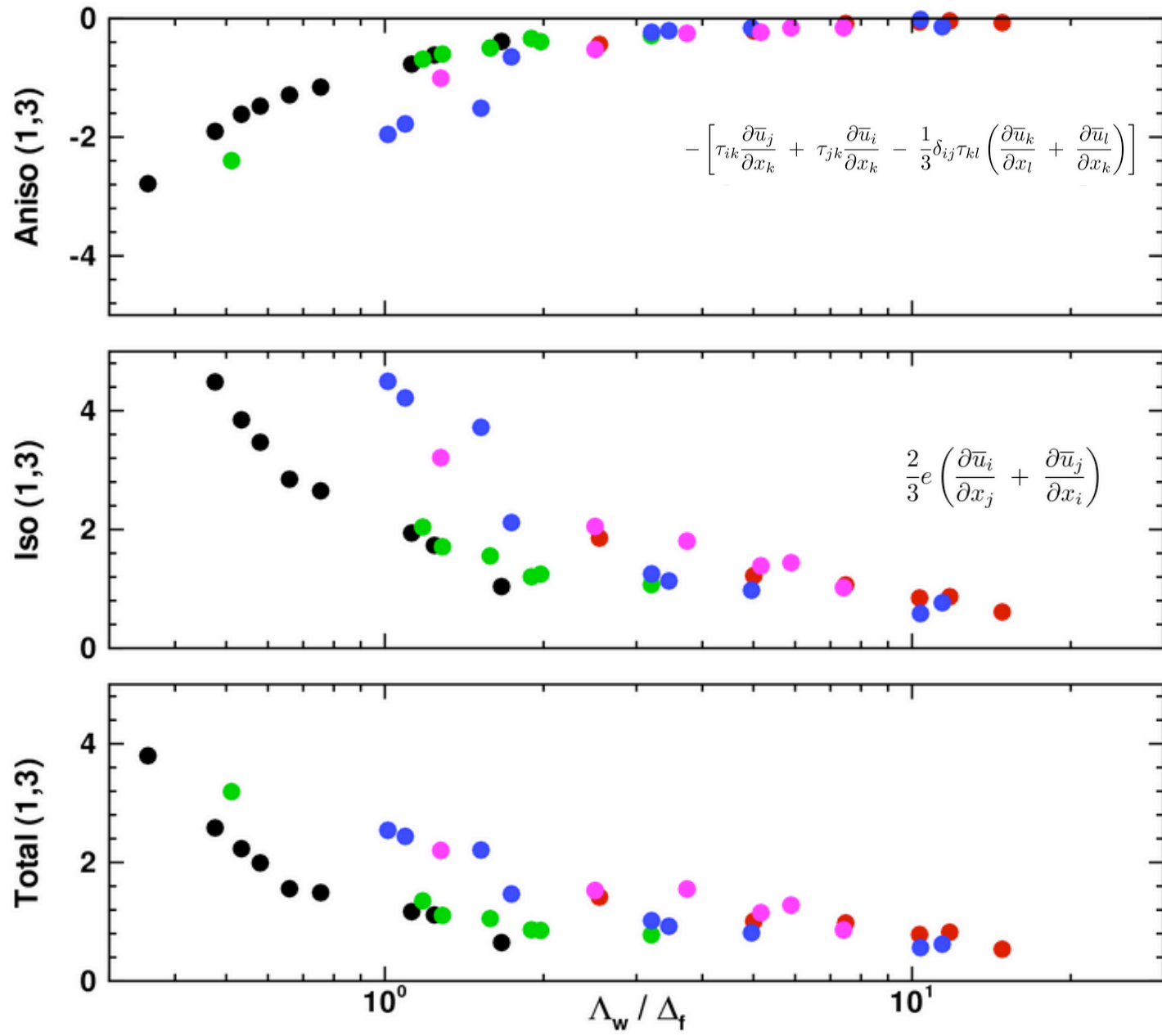
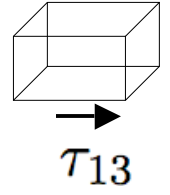
# PRODUCTION OF SUBFILTER SCALE FLUX $\tau_{33}$



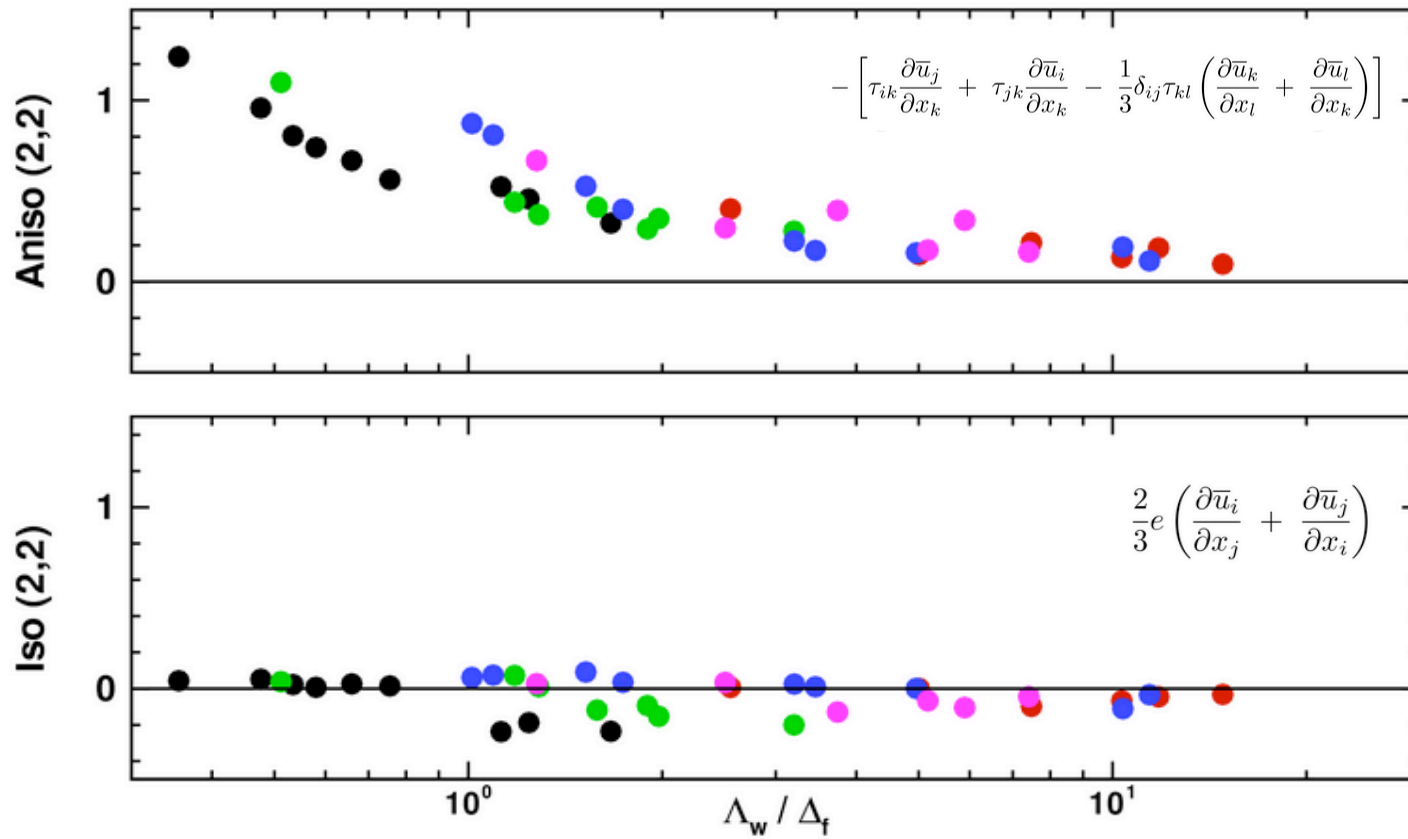
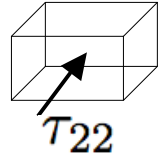
↑  $\tau_{33}$



# PRODUCTION OF SUBFILTER SCALE FLUX $\tau_{13}$



# PRODUCTION OF SUBFILTER SCALE FLUX $\tau_{22}$



# VARIATION OF DEVIATORIC STRESS IN LIMIT $\Lambda_w/\Delta_f \rightarrow 0$

$\langle \tau_{11} \rangle = T \left( -2\langle \tau_{13} \rangle \frac{\partial U}{\partial z} + \frac{2}{3}\epsilon \right)$	$\langle \tau_{11} \rangle = 0$
$\langle \tau_{22} \rangle = T \left( \frac{2}{3}\epsilon \right)$	$\langle \tau_{22} \rangle = 0$
$\langle \tau_{33} \rangle = T \left( \frac{2}{3}\epsilon \right)$	$\langle \tau_{33} \rangle = 0$
$\langle \tau_{13} \rangle = T \left( \frac{2}{3}e \frac{\partial U}{\partial z} - \langle \tau_{33} \rangle \frac{\partial U}{\partial z} \right)$	$\langle \tau_{13} \rangle = T \left( \frac{2}{3}e \frac{\partial U}{\partial z} \right)$

Steady-state rate equations

Smagorinsky model

***WHAT ABOUT SCALARS?***



# RATE EQUATIONS FOR SUBGRID SCALAR FLUX

- What are the parent equations for subgrid-scale scalar flux?

$$f_i = \overline{u_i c} - \bar{u}_i \bar{c}$$

$$\frac{Df_i}{Dt} = -\frac{2}{3}e \frac{\partial \bar{c}}{\partial x_i} \quad \leftarrow \text{Isotropic production}$$

$$-f_j \frac{\partial \bar{u}_i}{\partial x_j} + \tau_{ij} \frac{\partial \bar{c}}{\partial x_j}$$

$$+ \frac{1}{\rho} \left( \overline{p \frac{\partial c}{\partial x_i}} - \bar{p} \frac{\partial \bar{c}}{\partial x_i} \right)$$

+ transport + buoyancy

Pressure destruction

Anisotropic production

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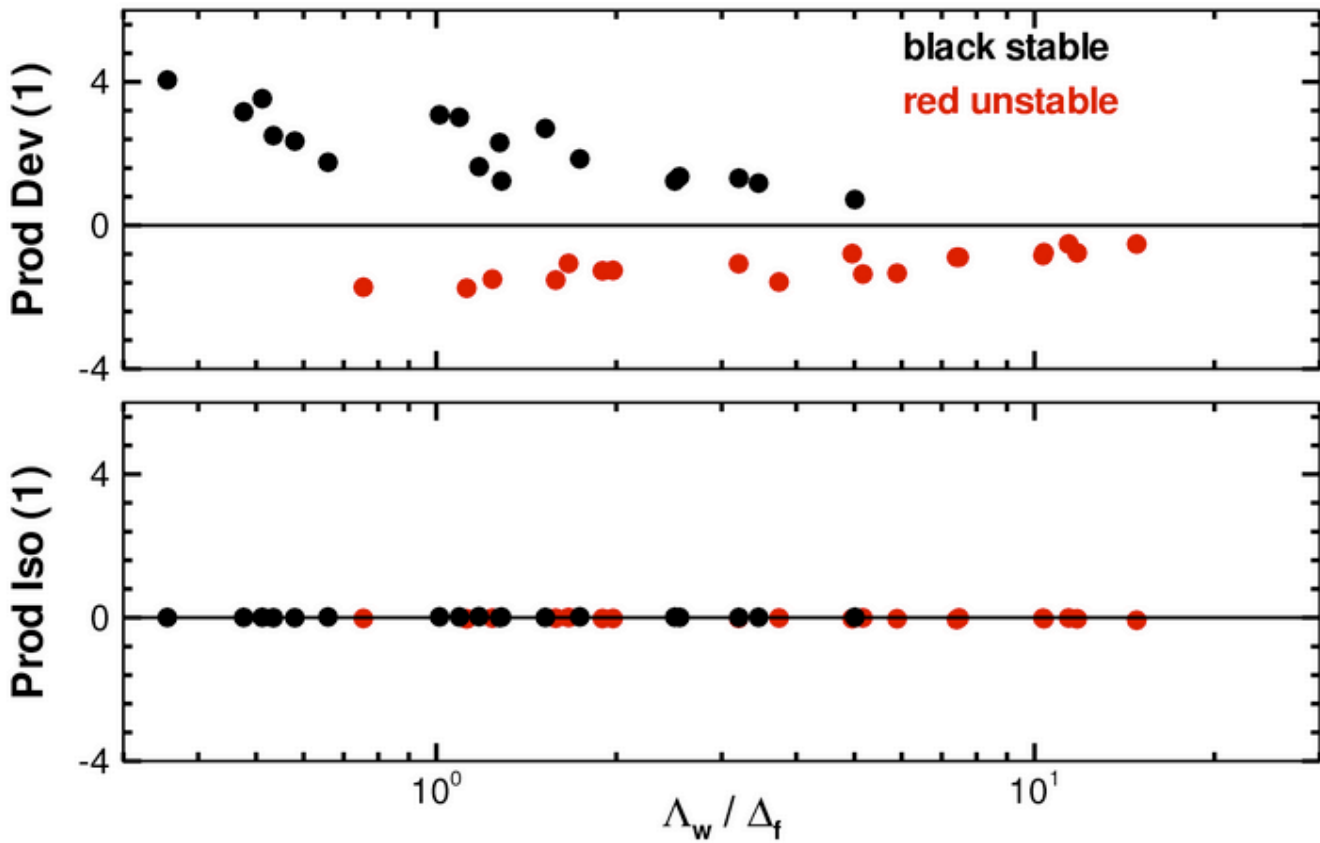
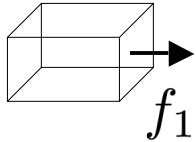
$$f_i = \overline{u_i c} - \bar{u}_i \bar{c}$$

$$\frac{Df_i}{Dt} = -\frac{2}{3}e \frac{\partial \bar{c}}{\partial x_i} - f_j \frac{\partial \bar{u}_i}{\partial x_j} + \tau_{ij} \frac{\partial \bar{c}}{\partial x_j} + \frac{1}{\rho} \left( \overline{p \frac{\partial c}{\partial x_i}} - \bar{p} \frac{\partial \bar{c}}{\partial x_i} \right) + \text{transport} + \text{buoyancy}$$

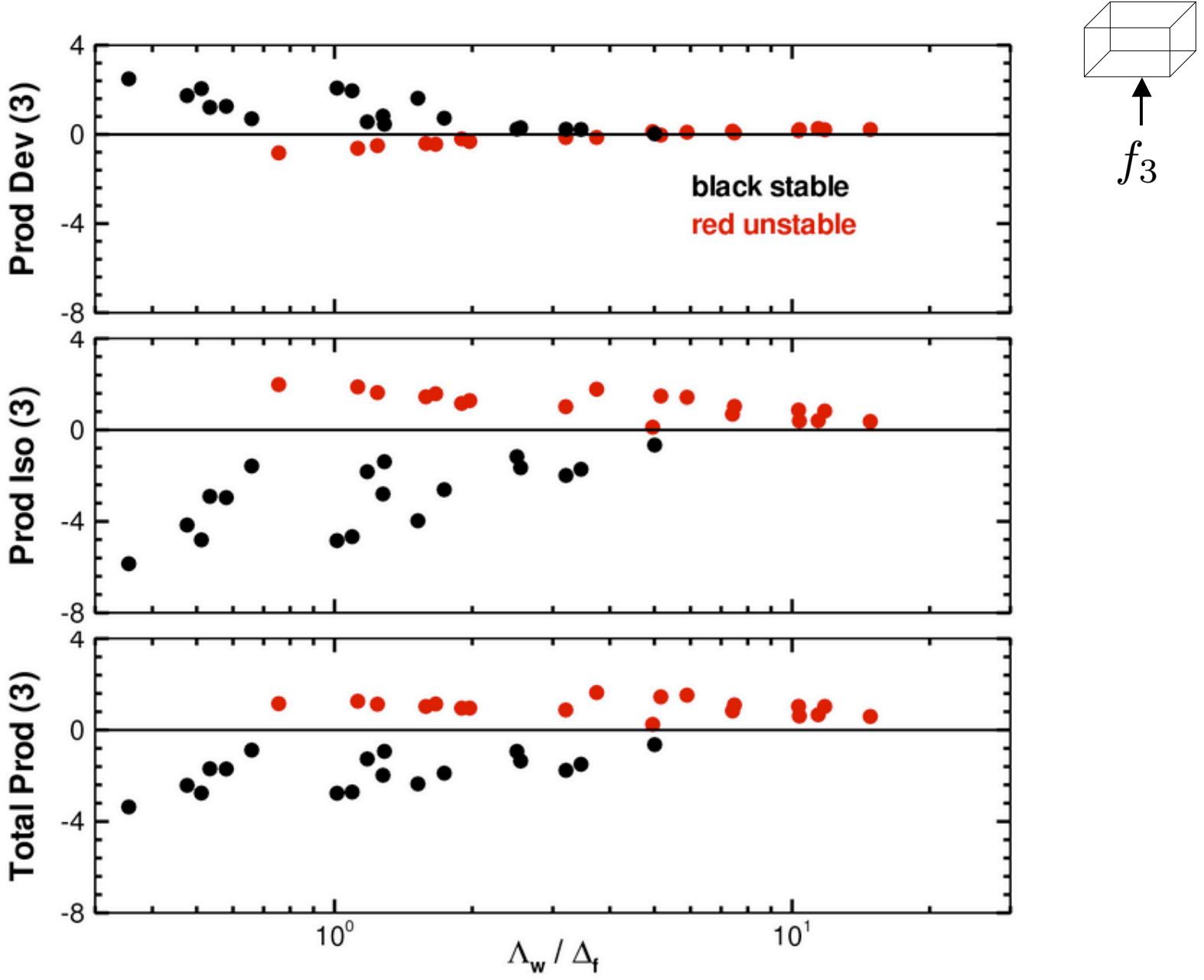
Eddy viscosity model

$$f_i = -\nu_h \frac{\partial \bar{c}}{\partial x_i} \quad \nu_h = \frac{2c_h \Delta_f \sqrt{e}}{3}$$

# PRODUCTION OF SUBFILTER SCALE SCALAR FLUX $f_1$



# PRODUCTION OF SUBFILTER SCALE SCALAR FLUX $f_3$



# SUBGRID-SCALE SCALAR FLUX

## Comments:

- Net horizontal scalar flux  $f_1 = \langle \overline{uc} - \bar{u}\bar{c} \rangle \neq 0$  even horizontally homogeneous PBLs, *i.e.*,  $\frac{\partial}{\partial x} \langle C \rangle = 0$
- Tilting of vertical flux by vertical shear is important  
$$f_1 \sim -f_3 \frac{\partial \bar{u}}{\partial z} T$$
- No eddy viscosity model, including the “dynamic approach”, can capture anisotropic production

# SUBFILTER-SCALE PRESSURE DESTRUCTION

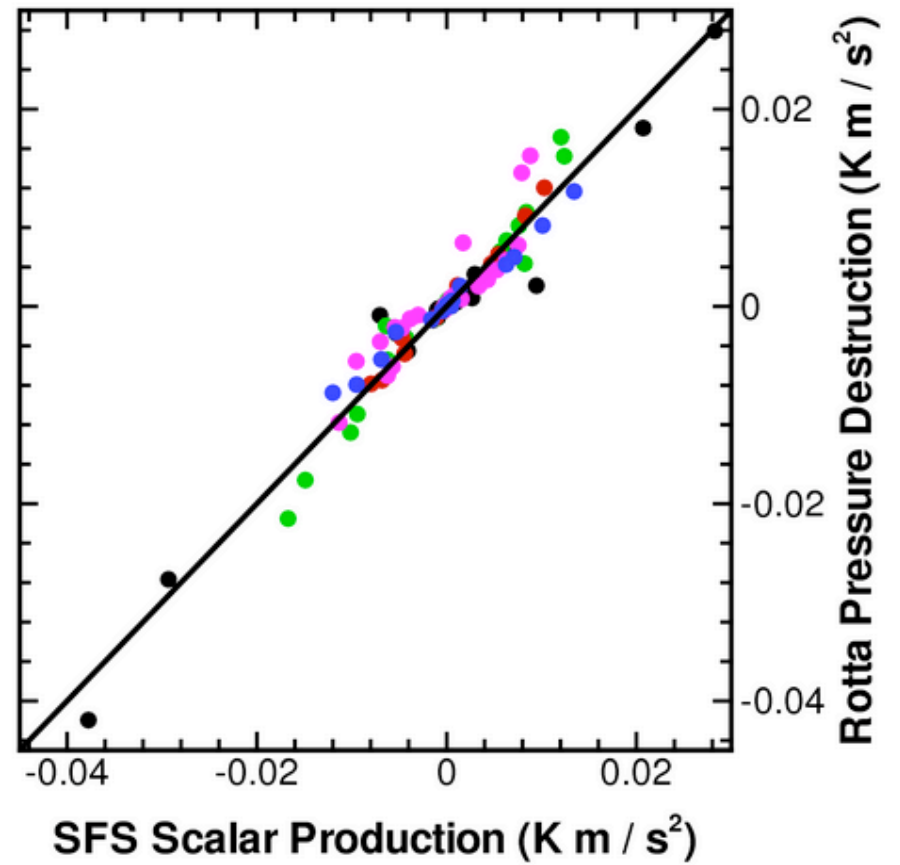
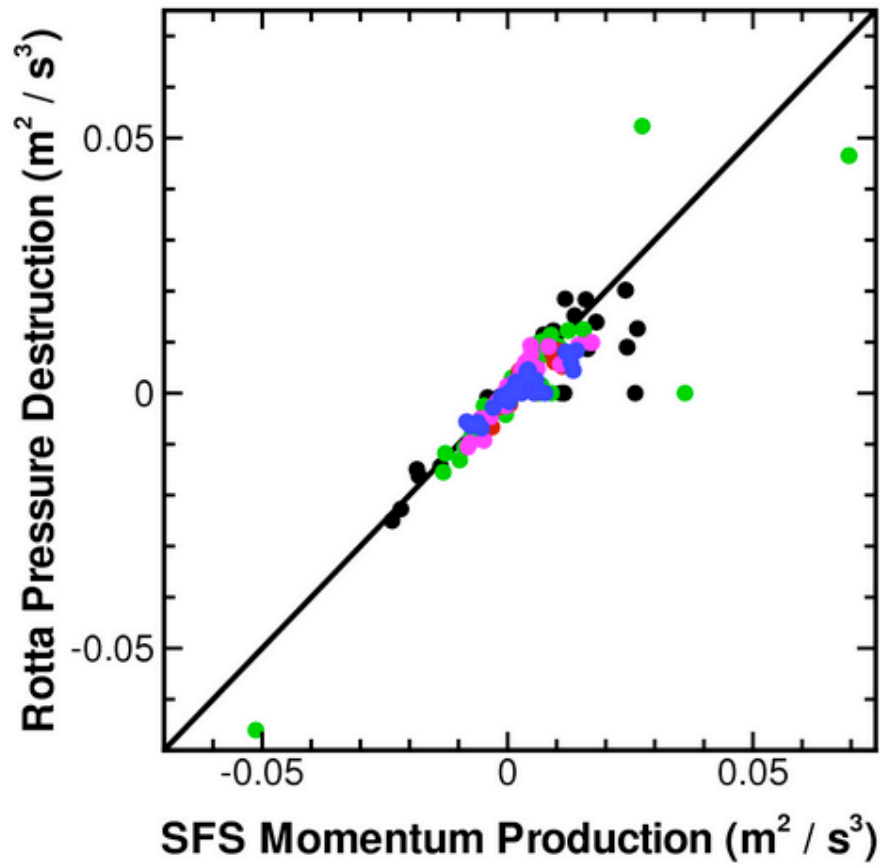
$$\begin{aligned}
 -\frac{1}{\rho} \left[ \overline{p \left( \frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)} - \bar{p} \left( \frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right) \right] &= -\frac{\tau_{ij} \sqrt{e}}{C_m \Delta_f} && \text{Momentum} \\
 +\frac{1}{\rho} \left( \overline{p \frac{\partial c}{\partial x_i}} - \bar{p} \frac{\partial \bar{c}}{\partial x_i} \right) &= -\frac{f_i \sqrt{e}}{C_s \Delta_f} && \text{Scalar}
 \end{aligned}$$



CHATS PRESSURE SENSOR (Steven Oncley)

AHATS (2008) "HORIZONTAL ARRAY" OF PRESSURE SENSORS

# VALIDATION OF ROTTA MODEL FOR MOMENTUM AND SCALARS



*Production  $\approx$  Destruction*

## SUMMARY

- LES is being applied to a richer set of boundary layer flows because of advances in parallel computing
- Subgrid-scale parameterizations in LES need to be validated/improved for geophysical applications
- Multi-point measurements from the HATS field campaigns compliment our ability to compute
  - Evaluation of subgrid scale models with high  $Re$  data
  - Rate equations provide insight into SGS dynamics
  - Importance of anisotropic production for stress and scalar especially for  $\Lambda_w/\Delta_f \sim \mathcal{O}(1)$  or less
  - Data highlights the shortcomings of an eddy viscosity approach