

Theoretical Studies of Strongly Nonlinear Langmuir Circulation

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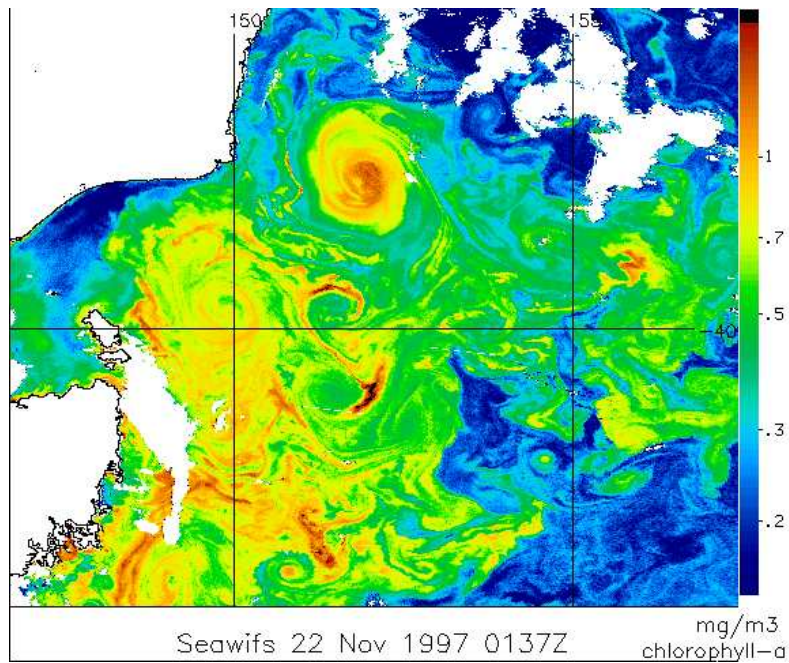
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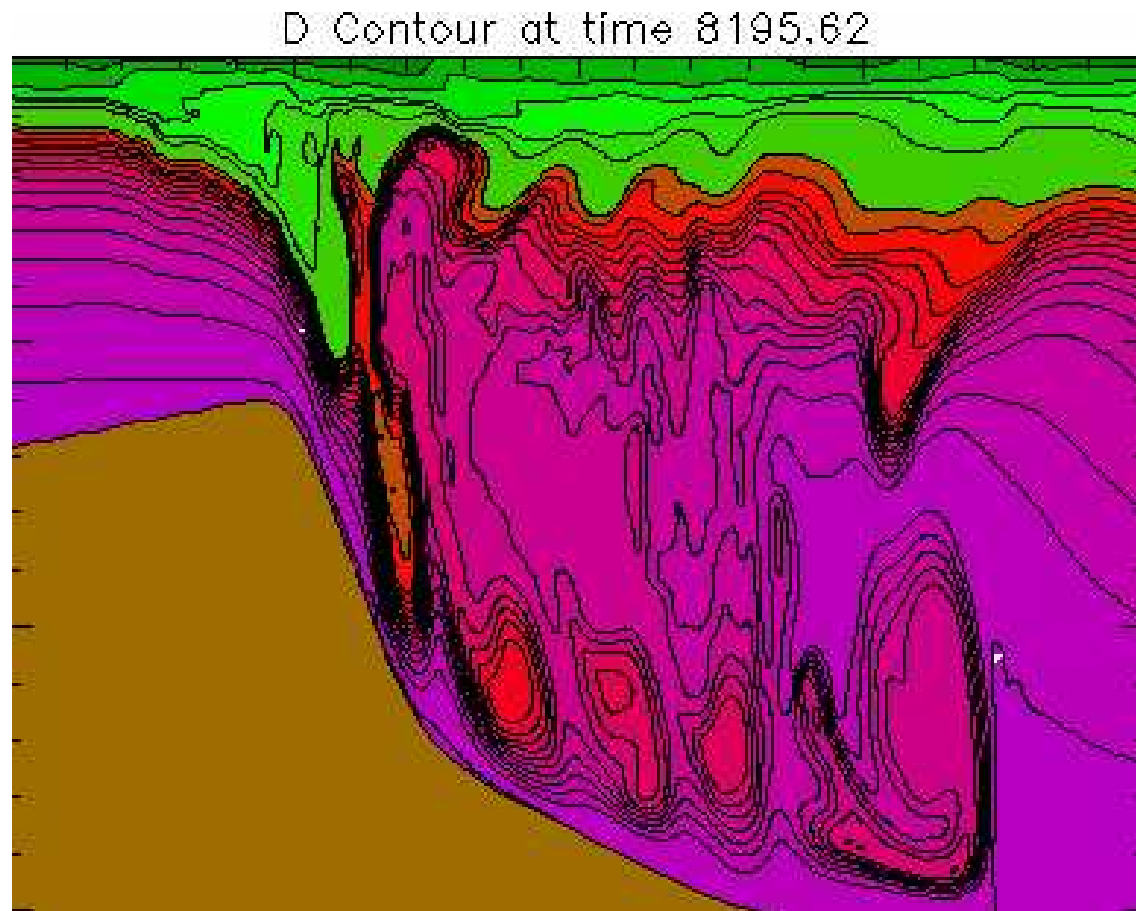
Ocean Surface Mixed Layer

- Crucial small-scale process: sub-mesoscale dynamics of ocean surface boundary (or “mixed”) layer – the O(100)-m deep ocean surface layer which connects atmosphere to the deep ocean.
- According to the NSF report *Ocean Sciences at the New Millennium* (2001), **“Forecasting the evolution, in space and time, of the mixed layer is perhaps the biggest individual challenge in ocean prediction.”**
- Key challenges include: (1) estimating mixing and transport across the mixed layer; (2) understanding interaction between the mixed layer and stratified ocean interior; (3) connecting mixed-layer/mesoscale dynamics.

Mesoscale Eddies & Spiral Vortices (NASA)

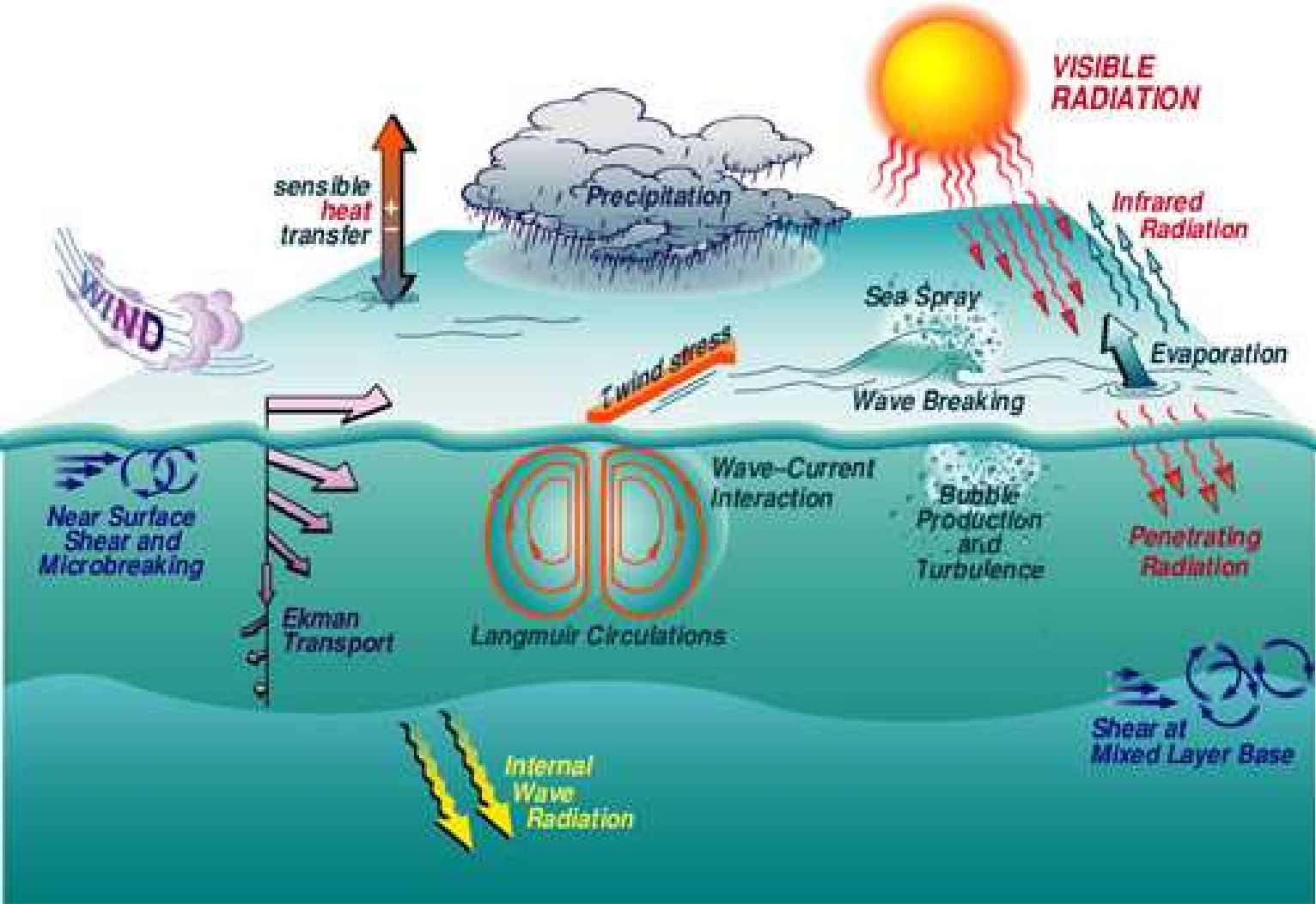


Submesoscale Phenomena – Vertical Transport/Mixing



- Image from numerical simulation by K. Lamb (U. Waterloo).

Schematic of Mixed Layer Processes (B. Weller, WHOI)



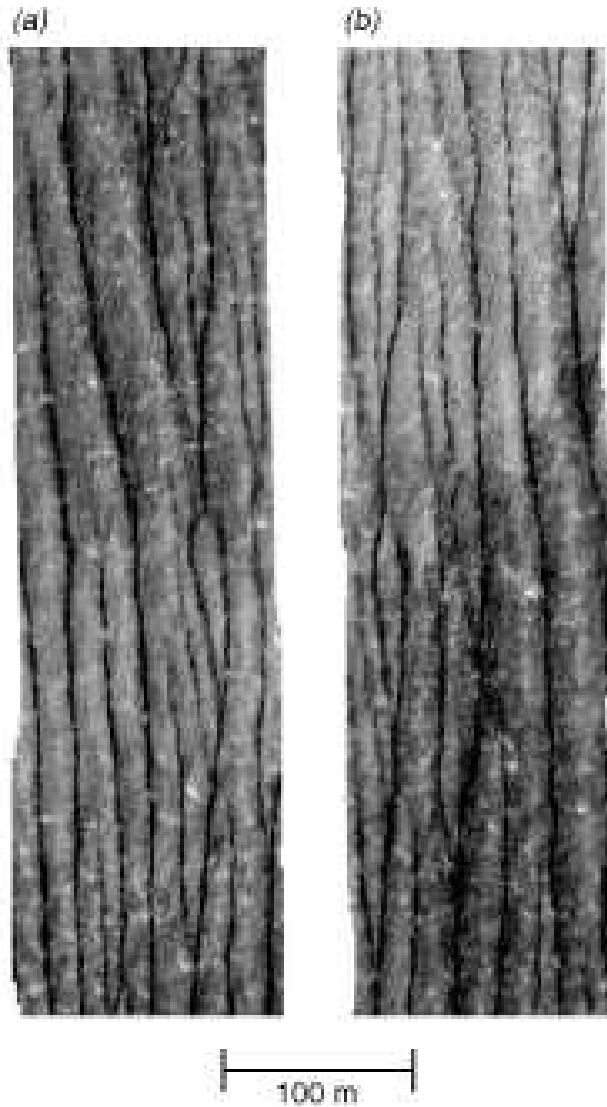
Langmuir Circulation (LC)

- Wind- and surface-wave driven convective motion (array of wind-aligned counter-rotating vortices) commonly occurring in natural water bodies.
- Range from $O(10\text{ cm})$ to $O(500\text{ m})$, with $O(\text{cm/s})$ velocities.
- Crucial for establishment and maintenance of oceanic mixed layer.
- Moderates air–sea exchanges of heat, mass and momentum, ultimately influencing weather, climate, ocean biology, pollutant (oil-spill) dispersal.

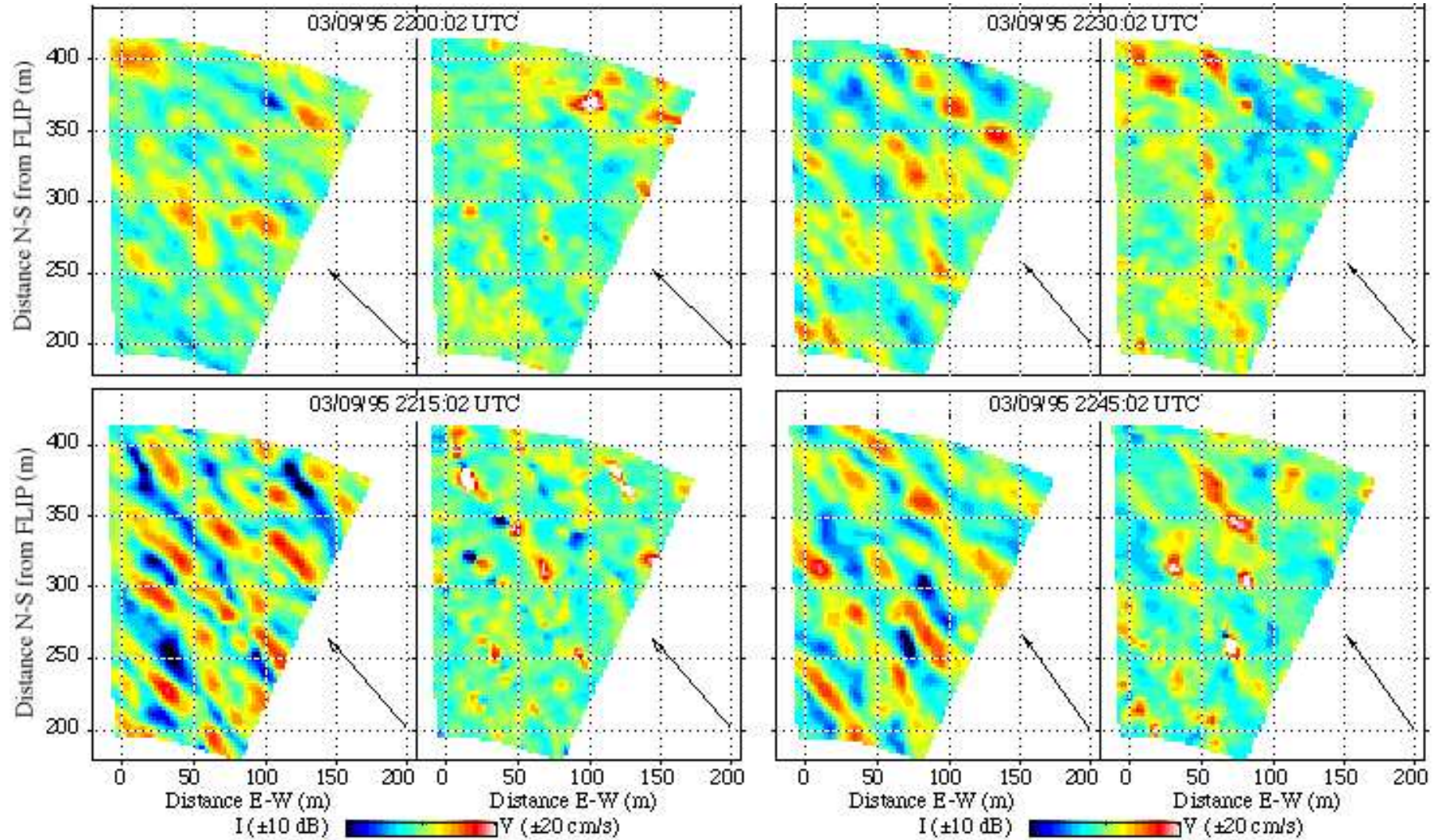
Photo of LC Windrows (A. Szeri)



IR Images of LC Windrows (G. Marmorino, NRL)



Sonar Images of LC Windrows (J. Smith, SIO)



LC Banding – Implications for Oil Spills



Craik-Leibovich (CL) Theory

Assumption

Dominant upper ocean motion due to **irrotational surface waves**: $O(\text{m/s})$

Wind-driven shear flow, LC, other **rotational motions** weaker: $O(\text{cm/s})$

Multiple-Time Scale Analysis $t = \varepsilon^2 t_f$

▷ Expansion: $\mathbf{v}(x, y, z, t, t_f) = \varepsilon \mathbf{u}_w(x, y, z, t_f) + \varepsilon^2 \mathbf{u}(x, y, z, t, t_f)$

▷ Substitute into Navier-Stokes equations (NSE).

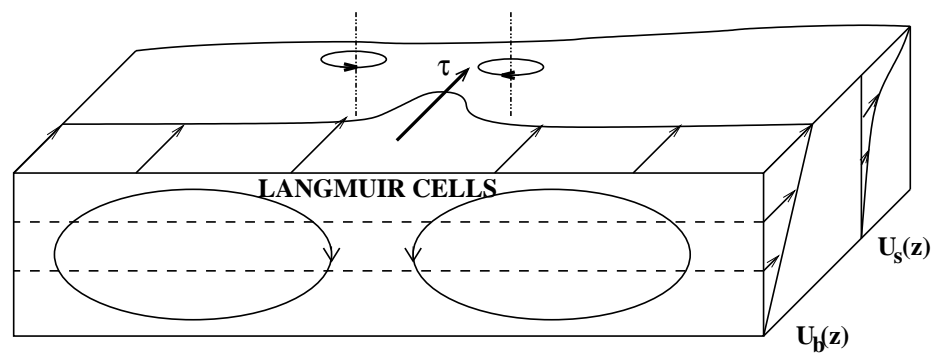
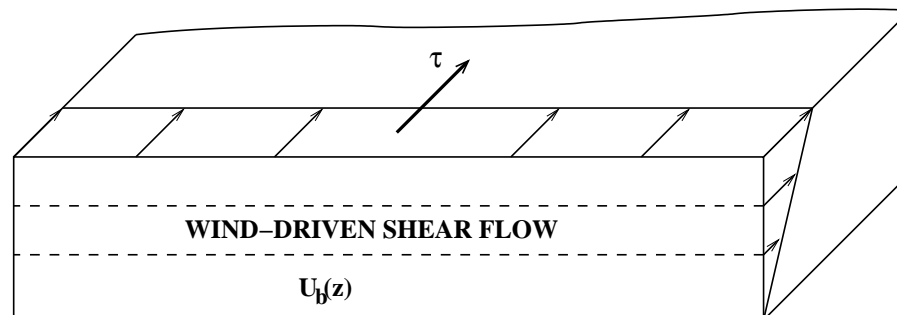
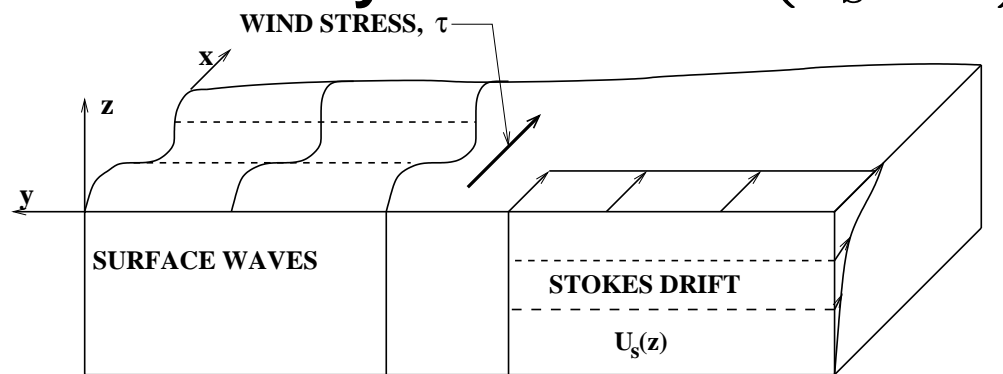
▷ Average over t_f , fast time scale associated with surface-wave period.

▷ Obtain equations formally identical to NSE... except for **vortex force**:

$$\mathbf{u}_s \times \boldsymbol{\omega}(x, y, z, t)$$

CL eqns show that LC arises as an **instability** of a wind-driven shear flow on which surface waves propagate.

LC Instability Mechanism ($u_s \times \omega$)



Robustness of CL Theory

Following the original derivation using formal multiple scale analysis, the CL eqns have been re-derived using more sophisticated mathematical approaches:

1. Generalized Lagrangian Mean (GLM) theory – an exact theory of nonlinear wave/mean-flow interactions.
2. A time-averaged version of the Kelvin circulation theorem.
3. Lagrangian averaging of Hamilton's principle.

Open Issues

1. $O(1)$ shear – feedback of LC on waves, rotational waves.
2. Finite-amplitude waves, wave-breaking.
3. LC wavelength selection mechanism: connection to surface waves?
4. Quantitative comparisons with field or laboratory data.

Related Work – Theory and Simulation

Quasi-Laminar Simulations of 2D Craik–Leibovich (CL) Equations

- Li & Garrett (*J. Mar. Res.* 1993, *JPO* 1995,1997)
- Gnanadesikan & Weller (*JPO* 1995)

Weakly Nonlinear or Small Wavenumber 2D and 3D Investigations

- 2D: Leibovich, Lele & Moroz (*JFM* 1989)
- 3D: Bhaskaran & Leibovich (*Phys. Fluids*, 2002)
- 3D: Cox & Leibovich (*Phys. Fluids*, 1997)

Simulations of Full 3D Craik–Leibovich (CL) Equations

- DNS: Tandon & Leibovich (*JGR*, 1995)
- LES: Skillingstad & Denbo (*JGR*, 1995), McWilliams *et al.* (*JFM*, 1997), Tejada-Martinez & Grosch (*JFM*, 2007)

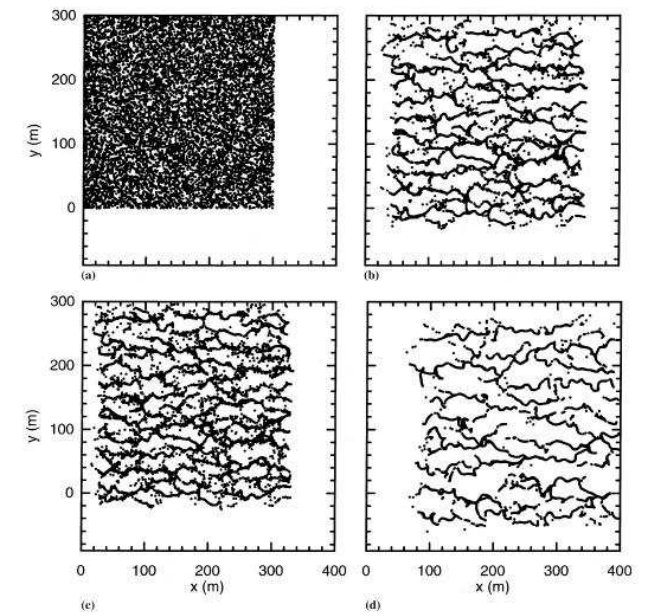
Anisotropic LC Dynamics



–A. Szeri (1996)



–G. Marmorino



–McWilliams *et al.* (1997)

Goals and Motivation

Objective Obtain reduced PDE model capable of describing coarse-grained, strongly anisotropic but otherwise turbulent LC dynamics.

Motivation

- Secondary stability analysis by Tandon & Leibovich (*JPO*, 1995)
- Reduced PDEs for rapidly-rotating thermal convection by Julien, Knobloch & Werne (*Theoret. Comput. Fluid Dyn.*, 1998), Sprague *et al.* (*JFM*, 2006)

Purpose

- Reveal dominant 3D physics.
- More amenable to (e.g. upper-bound) analysis.
- Less expensive numerical simulations for multi-scale process studies.
- Incorporation into formal multiscale numerical scheme.

Isotropically Scaled CL Equations

- Consider full 3D, isotropically-scaled CL equations, where two parameters $R_* \equiv u_* H / \nu_e$, $La_t = \sqrt{u_* / u_{s0}}$ replace single parameter $La \equiv La_t R_*^{-3/2}$:

$$\frac{D\mathbf{u}}{Dt} = -\nabla p + \frac{1}{La_t^2} (\mathbf{U}_s \times \boldsymbol{\omega}) + \frac{1}{R_*} \nabla^2 \mathbf{u}$$

- Two turbulence regimes:

Shear flow turbulence regime: $La_t \gg 1$ with $R_* \gg 1$.

Langmuir turbulence regime: $La_t = O(0.1)$ with $La \ll 1$.

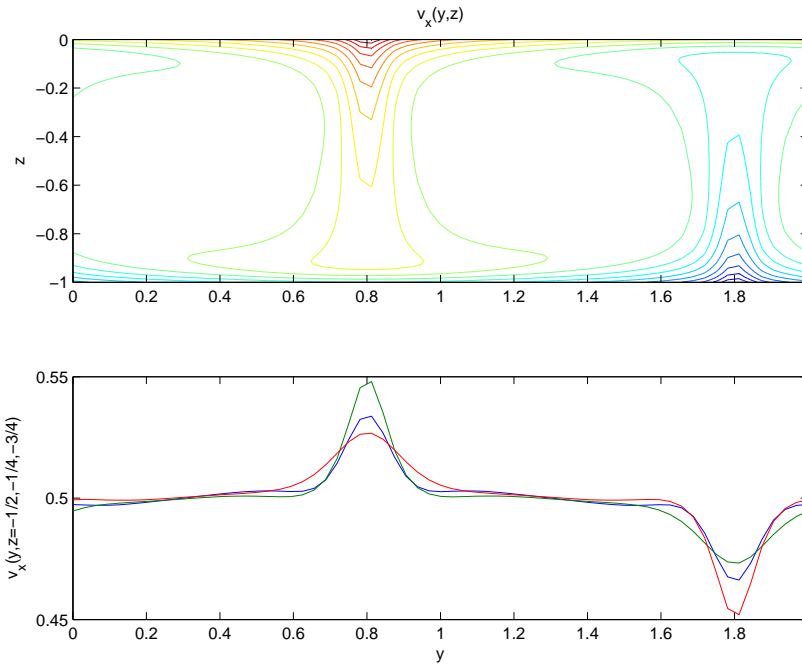
- Motivates consideration of formal limit $La_t \rightarrow 0$ with R_* or La fixed:

2D dynamics: $\Omega \neq 0$, $\partial_x \Omega = 0$, u -fluctuations $\ll (v, w)$ -fluctuations.

Constraint on Downwind Velocity Fluctuations

Consider 2D ($\partial/\partial x = 0$) CL downwind vorticity (Ω) equation as $La_t \rightarrow 0...$

$$\frac{\partial \Omega}{\partial t} + v \frac{\partial \Omega}{\partial y} + w \frac{\partial \Omega}{\partial z} = -\frac{1}{La_t^2} U'_s(z) \frac{\partial u}{\partial y} + \frac{1}{R_*} \nabla_{\perp}^2 \Omega$$
$$\Rightarrow \frac{\partial u}{\partial y} \approx 0$$



Transient Growth of Cellular Velocity Fluctuations

As $La_t \rightarrow 0$, there is a “rapid-distortion” (strong Stokes “shear”) transient that drives strong cellular (v,w) velocity fluctuations.

$$\partial_\tau u = 0 \Rightarrow u(y, z, \tau) = U_0(y, z)$$

$$\partial_\tau v = -\partial_y \mathcal{P} + U_s \partial_y u$$

$$\partial_\tau w = -\partial_z \mathcal{P} + U_s \partial_z u \Rightarrow \partial_\tau \Omega = -U'_s(z) \partial_y u$$

$$\Omega(y, z, \tau) = - [U'_s(z) \partial_y U_0(y, z)] \tau + \Omega_0(y, z)$$

- (v,w) grow linearly with (fast) time τ , while u remains constant.

Anisotropic Velocity Scalings

- Employ anisotropic velocity scales to capture nonlinear, spatially anisotropic reduced dynamics:

$$L_x = H, \quad (L_y, L_z) = H, \quad \mathcal{T} = H/\mathcal{V}$$
$$\mathcal{U} = u_* R_*, \quad (\mathcal{V}, \mathcal{W}) = \sqrt{\mathcal{U} u_{s0}}, \quad \mathcal{P} = \rho \mathcal{V}^2$$

- In essence, perturbing off of strictly 2D $[\partial(\cdot)/\partial x = 0]$ problem.

- Identify $\mathcal{U}/\mathcal{W} = La_t^{4/3} La^{-1/3} \equiv \epsilon La^{-1/3} \ll 1$

(cf. Tejada-Martinez & Grosch 2007).

Rescaled CL Equations in Strong CL Vortex-Force Limit

$$\begin{aligned}
 \partial_t u + \epsilon L a^{-1/3} u \partial_x u + (\mathbf{v}_\perp \cdot \nabla_\perp) u &= -\epsilon^{-1} L a^{1/3} \partial_x P + L a \left[\partial_x^2 + \nabla_\perp^2 \right] u \\
 \partial_t \mathbf{v}_\perp + \epsilon L a^{-1/3} u \partial_x \mathbf{v}_\perp + (\mathbf{v}_\perp \cdot \nabla_\perp) \mathbf{v}_\perp &= -\nabla_\perp P + L a \left[\partial_x^2 + \nabla_\perp^2 \right] \mathbf{v}_\perp \\
 &\quad + U_s \left(\nabla_\perp u - \epsilon^{-1} L a^{1/3} \partial_x \mathbf{v}_\perp \right) \\
 \epsilon L a^{-1/3} \partial_x u + \nabla_\perp \cdot \mathbf{v}_\perp &= 0
 \end{aligned}$$

- Wind stress BC: $\partial_z u = 1$ along $z = 0, -1$.
- x -invariance at leading-order: $\partial_x P = \partial_x v = \partial_x w = 0$ and $\nabla_\perp \cdot \mathbf{v}_\perp = 0$.

Multiple Scale Expansion

1. Limit process: $\epsilon \equiv La_t^{4/3} \rightarrow 0$ with La fixed.
2. Introduce slow x scale: $X \equiv \epsilon La^{-1/3}x$ so that $\partial_x \rightarrow \partial_x + \epsilon La^{-1/3}\partial_X$.
3. Expand fields:

$$\begin{aligned}u(x, y, z, t) &= u_0(x, X, y, z, t) + \epsilon u_1(x, X, y, z, t) + \dots \\ \mathbf{v}_\perp(x, y, z, t) &= \mathbf{v}_{0\perp}(X, y, z, t) + \epsilon \mathbf{v}_{1\perp}(x, X, y, z, t) + \dots \\ P(x, y, z, t) &= P_0(X, y, z, t) + \epsilon P_1(x, X, y, z, t) + \dots\end{aligned}$$

4. Substitute into PDEs, collect terms of like order and **average** over fast x .
5. Obtain closed set of equations for $\bar{u}_0 \equiv U(X, y, z, t)$, $\mathbf{v}_{0\perp} \equiv \mathbf{V}_\perp(X, y, z, t)$ and $P_0 \equiv \Pi(X, y, z, t)$.

Example: Averaging Downwind Momentum Equation

- At $O(1)$, the x momentum equation becomes:

$$\partial_T u_0 + (\mathbf{v}_{0\perp} \cdot \nabla_{\perp}) u_0 = -La^{1/3} \partial_x P_1 - \partial_X P_0 + La \left[\partial_x^2 + \nabla_{\perp}^2 \right] u_0,$$

- Decompose all (fast) x -varying fields into fast- x average plus fluctuation, e.g.,

$$u_0(x, X, y, z, T) \equiv \bar{u}_0(X, y, z, T) + u'_0(x, X, y, z, T),$$

- Averaging in x , using the x -invariance of P_0 and $\mathbf{v}_{0\perp}$, yields an equation for $\bar{u}_0(X, y, z, T)$:

$$\partial_T \bar{u}_0 + (\mathbf{v}_{0\perp} \cdot \nabla_{\perp}) \bar{u}_0 = -\partial_X P_0 + La \nabla_{\perp}^2 \bar{u}_0.$$

Reduced PDEs

- Define:

$$D_t^\perp(\cdot) \equiv \partial_t(\cdot) + (\mathbf{V}_\perp \cdot \nabla_\perp)(\cdot) \equiv \partial_t(\cdot) + J[(\cdot), \psi],$$

where $J[(\cdot), \psi] = \partial_z \psi \partial_y(\cdot) - \partial_y \psi \partial_z(\cdot)$.

- Reduced dynamics governed by:

$$\begin{aligned} D_t^\perp U &= -\partial_X \Pi + La \nabla_\perp^2 U \\ D_t^\perp \Omega + U_s(z) \partial_X \Omega &= U_s'(z) (\partial_X V - \partial_y U) + La \nabla_\perp^2 \Omega \\ \nabla_\perp^2 \Pi &= 2J[\partial_y \psi, \partial_z \psi] + \nabla_\perp \cdot (U_s(z) \nabla_\perp U) + U_s'(z) \partial_X (\partial_y \psi) \\ \nabla_\perp^2 \psi &= -\Omega, \quad \mathbf{V}_\perp \equiv \nabla_\perp \times \psi \hat{i} \end{aligned}$$

- Fast x averaged BCs along $z = 0, -1$: $\partial_z U = 1$, $\Omega = 0$, $\psi = 0$.
- Advection by U and stretching of Ω are subdominant processes.

Linear Stability Analysis of Reduced PDEs

- Linearize about wind-driven base state shear flow $U_B(z) = z+1$.
- For simplicity, set $U_s(z) = z+1$.
- Decompose all fields into a base state plus perturbation, e.g.
 $U(X, y, z, T) = U_B(z) + u(X, y, z, T)$

- Perturbations satisfy

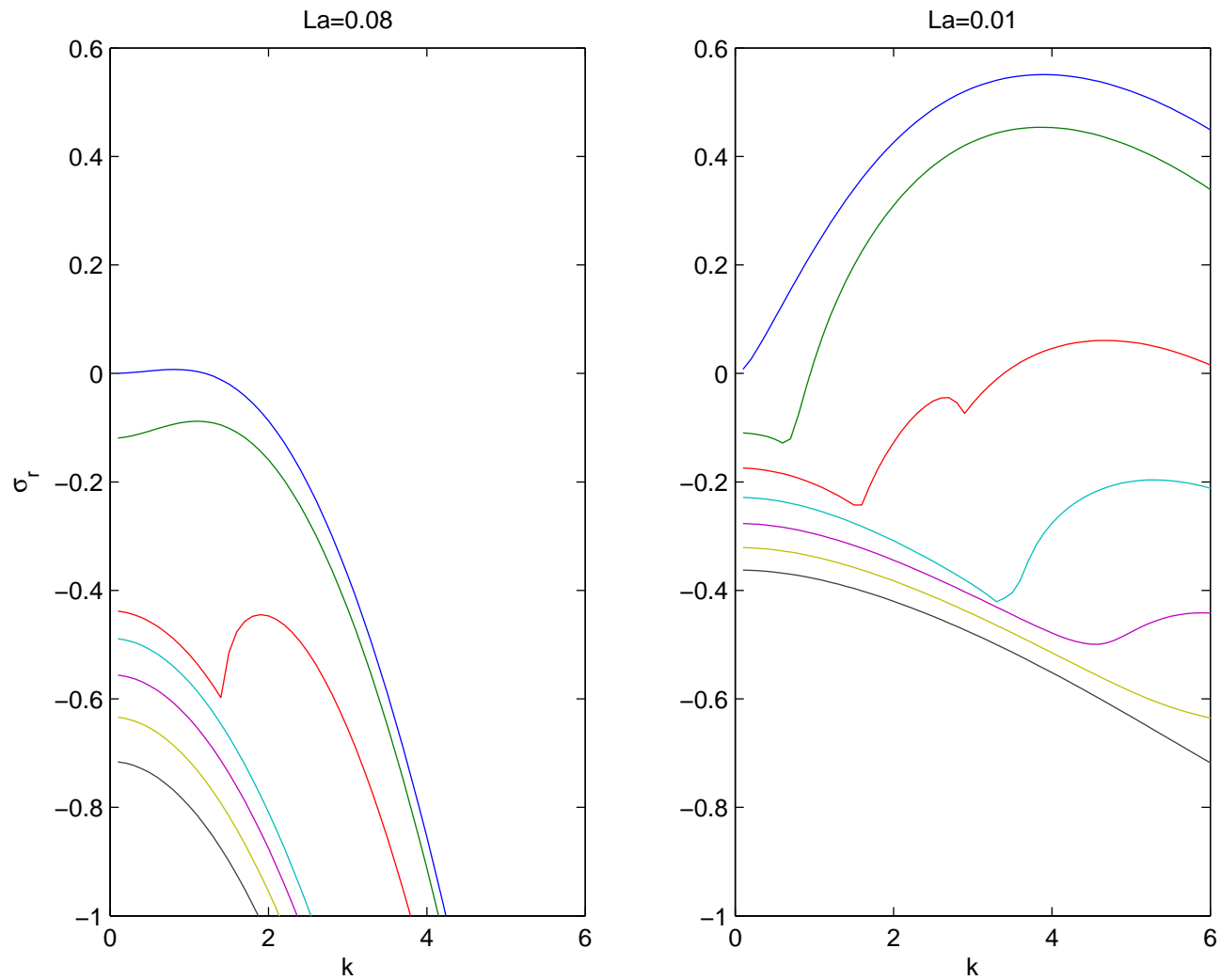
$$\begin{aligned}\partial_T u - \partial_y \phi &= -\partial_x p + La \nabla_{\perp}^2 u, \\ \partial_T \omega + (z+1) \partial_x \omega &= \partial_x (\partial_z \phi) - \partial_y u + La \nabla_{\perp}^2 \omega, \\ \nabla_{\perp}^2 \phi &= -\omega, \\ \nabla_{\perp}^2 p &= \nabla_{\perp} \cdot [(z+1) \nabla_{\perp} u] + \partial_x (\partial_y \phi),\end{aligned}$$

subject to $\partial_z u = \omega = \phi = \partial_z p = 0$ along $z = 0, -1$.

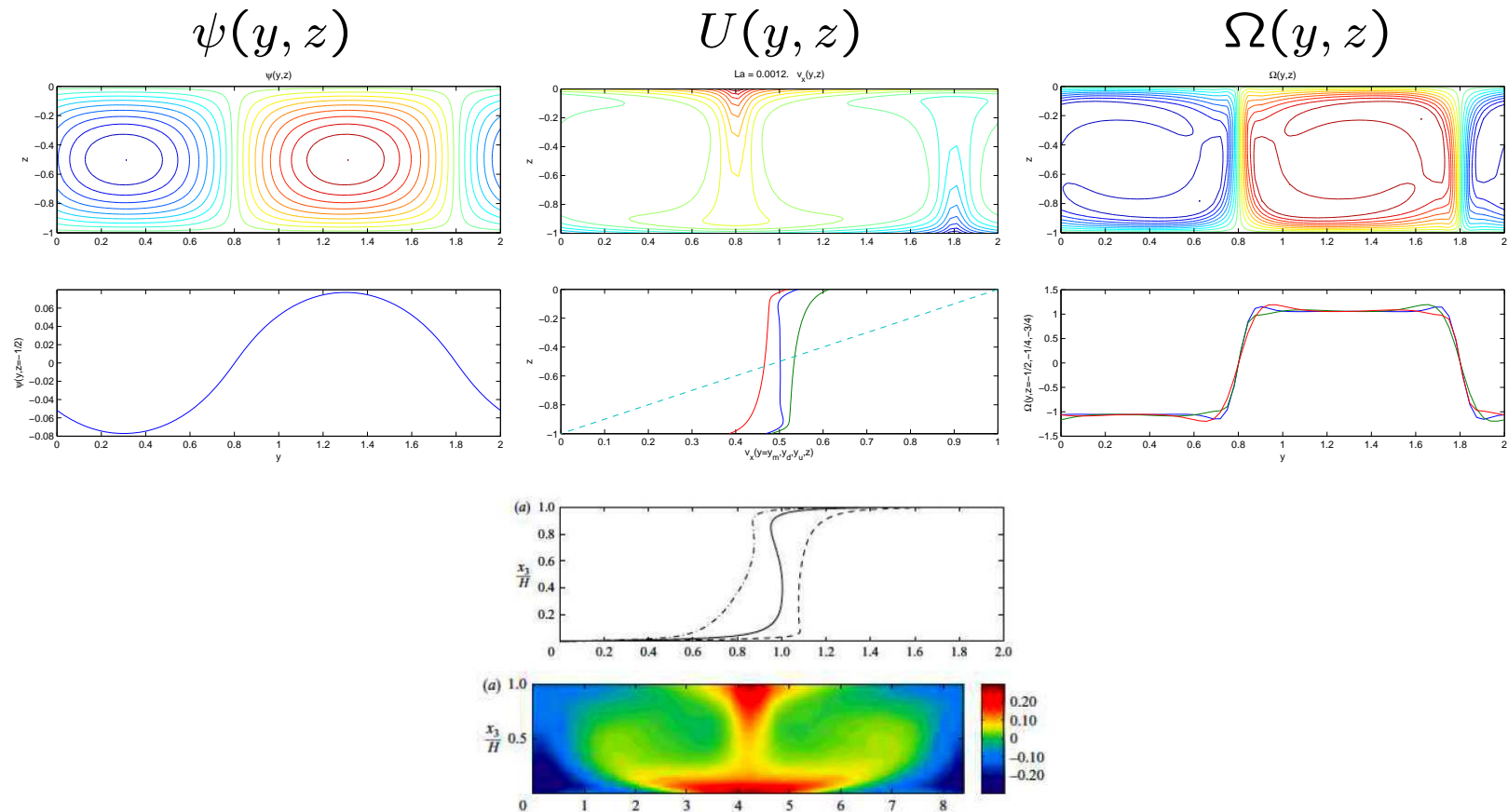
- Normal-mode *ansatz*, e.g.

$$u(X, y, z, T) = \hat{u}(z) e^{i(ky+lX)} e^{\sigma t} + \text{c.c.}$$

Reduced PDEs: Linear Stability Results



Strongly Nonlinear, Strictly 2D Convective States



- Steady-state $U(y, z)$ profiles show excellent qualitative agreement with $x-t$ averaged LES profiles of Tejada-Martinez & Grosch (2007).

Matched Asymptotic Analysis

$$\frac{\partial \psi}{\partial z} \frac{\partial U}{\partial y} - \frac{\partial \psi}{\partial y} \frac{\partial U}{\partial z} = La \left(\frac{\partial^2 U}{\partial y^2} + \frac{\partial^2 U}{\partial z^2} \right)$$

$$\frac{\partial \psi}{\partial z} \frac{\partial \Omega}{\partial y} - \frac{\partial \psi}{\partial y} \frac{\partial \Omega}{\partial z} = -U'_s(z) \frac{\partial U}{\partial y} + La \left(\frac{\partial^2 \Omega}{\partial y^2} + \frac{\partial^2 \Omega}{\partial z^2} \right)$$

$$\frac{\partial^2 \psi}{\partial y^2} + \frac{\partial^2 \psi}{\partial z^2} = -\Omega$$

where: $U(y, z) = U_B(z) + u(y, z) = [z + 1] + u(y, z)$

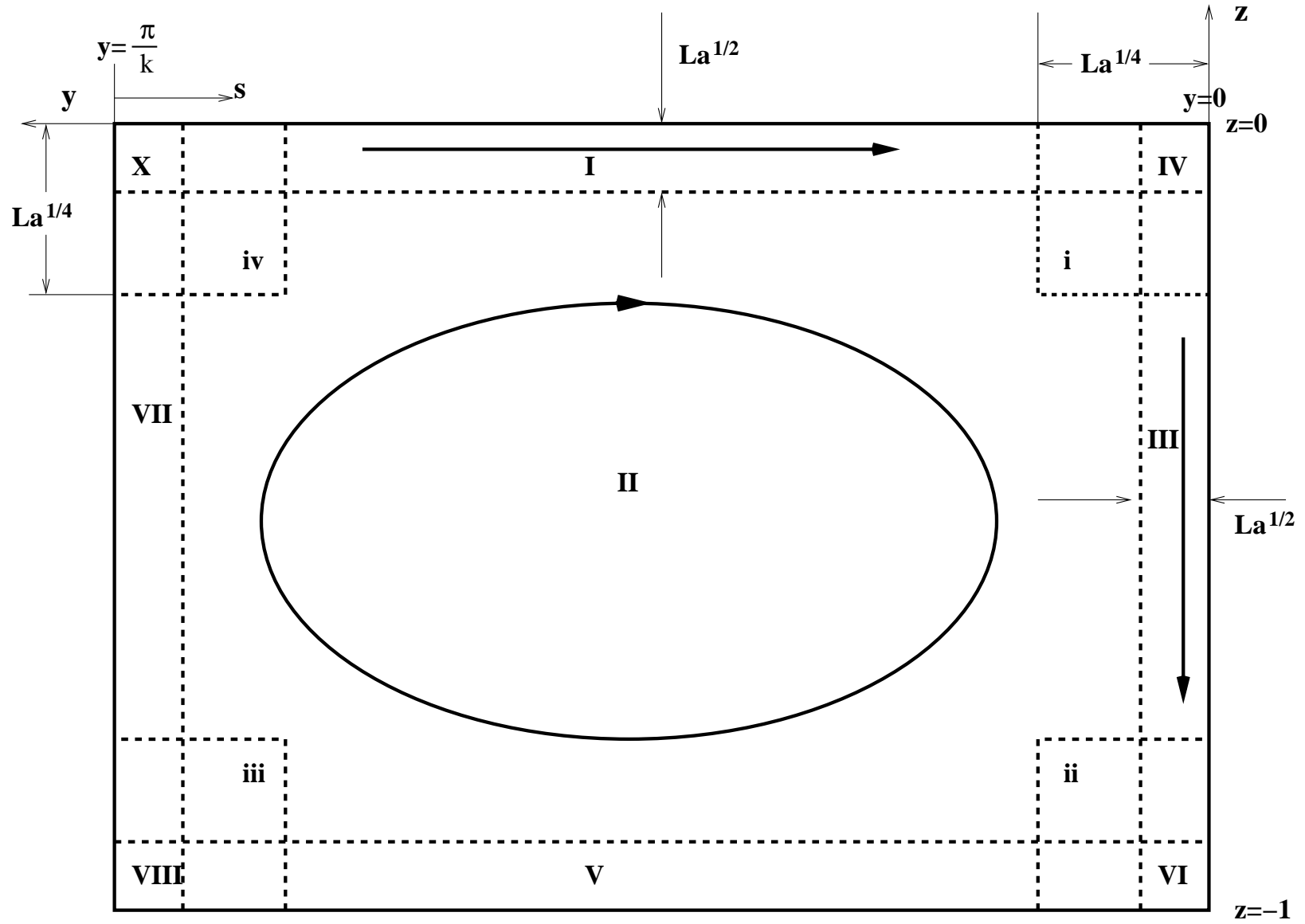
$$\underline{z = 0, -1} : \frac{\partial U}{\partial z} = 1, \psi = 0, \Omega = 0$$

$$\underline{y = 0, \pi/k} : \frac{\partial U}{\partial y} = 0, \psi = 0, \Omega = 0$$

Non-dimensional parameter:

$$La \equiv \frac{(\nu_e/H)^{3/2}}{u_* u_{s0}^{1/2}} = La_t R_*^{-3/2}$$

Hypothesized Small- La Asymptotic Structure



Region II: Inviscid Vortex Core

$$U(y, z) \sim \bar{U}; \quad \Omega(y, z) \sim \bar{\Omega}; \quad \psi(y, z) \sim \psi_{II}(y, z)$$

Equations/BCs

$$\frac{\partial^2 \psi_{II}}{\partial y^2} + \frac{\partial^2 \psi_{II}}{\partial z^2} = -\bar{\Omega}$$

$$\psi_{II}(y, z) = 0 \text{ along } z = 0, -1 \text{ and } y = 0, \pi/k$$

Matching with I, III

$$\psi(y, z) \sim v_{II}(y, 0) z \equiv V(y) z \text{ as } z \rightarrow 0^-$$

$$\psi(y, z) \sim -w_{II}(0, z) y \equiv -W(z) y \text{ as } y \rightarrow 0^+$$

$$V(y) = \sum_{n=1(\text{odd})}^{\infty} \left(-\frac{4\bar{\Omega}}{\pi k n^2} \right) \tanh\left(\frac{nk}{2}\right) \sin(nky)$$

$$W(z) = \sum_{n=1(\text{odd})}^{\infty} \left(-\frac{4\bar{\Omega}}{\pi k n^2} \right) \left[1 - \frac{\cosh\left(nk\left(z + \frac{1}{2}\right)\right)}{\cosh\left(\frac{nk}{2}\right)} \right]$$

Analysis in Regions I and III

Region I. Near-Surface Boundary Layer: $y = O(1)$, $z \equiv La^{1/2}Z$

$$U(y, z) - \bar{U} \sim La^{1/2}u_I(y, Z); \quad \Omega(y, z) \sim \Omega_I(y, Z); \quad \psi(y, z) \sim La^{1/2}\psi_I(y, Z)$$

- BL equations linearize since $\psi_I(y, Z)$ known.
- BL equations completely de-couple, since $\partial U / \partial y$ weak, $O(La^{1/2})$.

Region III. Downwelling Jet/Plume: $y \equiv La^{1/2}Y$, $z = O(1)$

$$U(y, z) - \bar{U} \sim La^{1/2}u_{III}(Y, z); \quad \Omega(y, z) \sim \Omega_{III}(Y, z); \quad \psi(y, z) \sim La^{1/2}\psi_{III}(y, Z)$$

- BL equations linearize since $\psi_I(y, Z)$ known.
- $\Omega(Y, z)$ coupled to $U(Y, z)$, since $\partial U / \partial y = O(1)$ (but converse is not true).

...Need $\bar{\Omega}$ to proceed...

Determination of Core Vorticity

1. Integrated Energy Balance.

$$La \int \int \Omega^2 dA = - \int \int \psi \frac{\partial U}{\partial y} dA \quad (\text{Exact Relation})$$

Asymptotic Approximation:

$$\begin{aligned} \bar{\Omega}^2 &\sim -\frac{k}{\pi La} \int \int \psi \frac{\partial U}{\partial y} dA \\ &\sim -\frac{2k}{\pi} \int_{-1}^0 \int_0^{\infty} \psi_{III} \frac{\partial U_{III}}{\partial Y} dY dz \\ &= -\frac{2k}{\pi} \int_{-1}^0 \int_0^{\infty} \psi_{III} \frac{\partial U_{III}}{\partial \psi_{III}} d\psi_{III} dz \\ &= -\frac{2k}{\pi} \int_{-1}^0 \left[\psi_{III} U_{III} \Big|_0^{\infty} - \int_0^{\infty} U_{III} d\psi_{III} \right] dz \end{aligned}$$

Determination of Core Vorticity (Cont'd)

2. Downwind Momentum Flux Conservation.

Exact Relation:

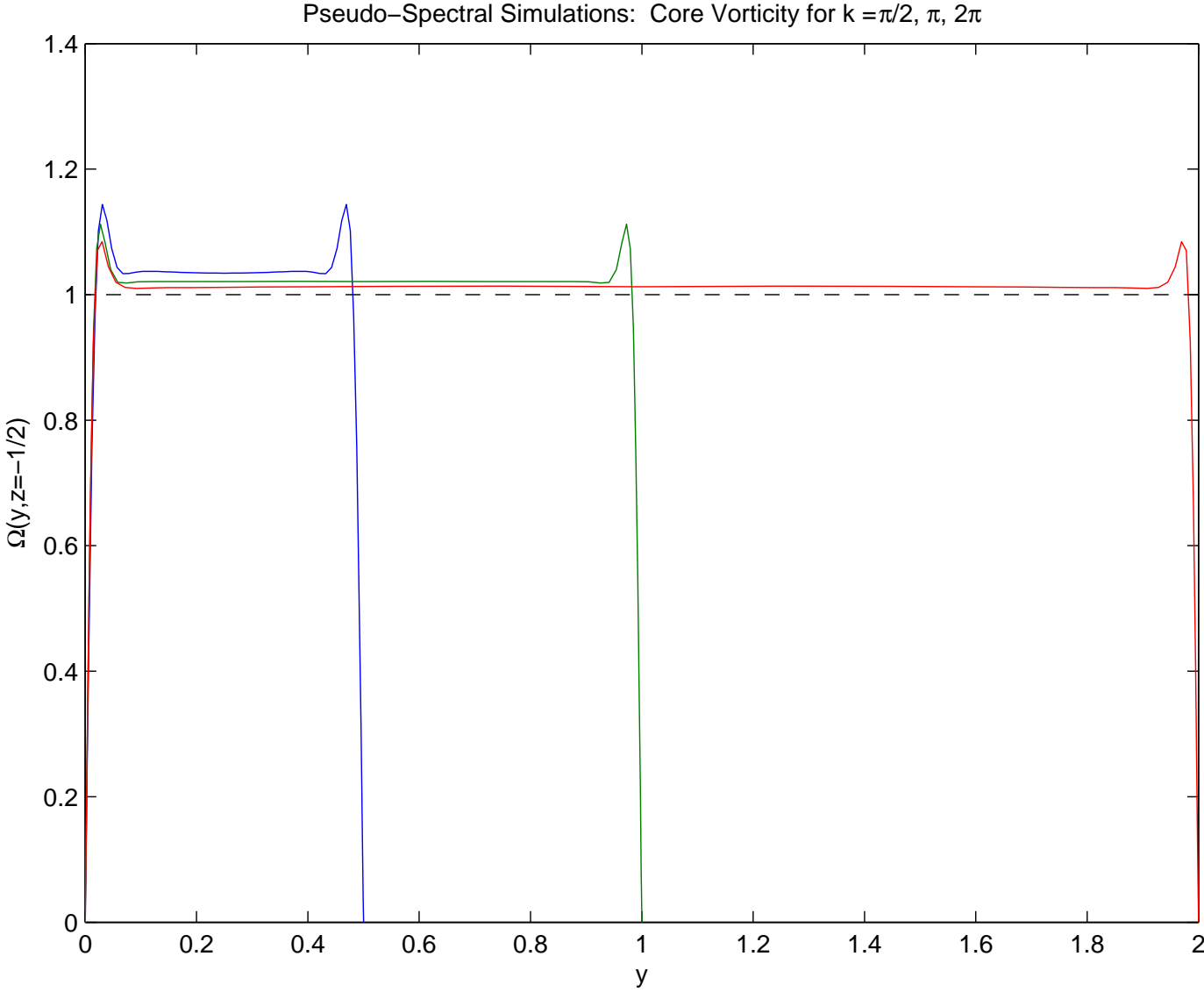
$$La \frac{\pi}{k} = \int_0^{\pi/k} \left[La \frac{\partial U}{\partial z} - w(U - \bar{U}) \right] dy$$

Asymptotic Approximation (True for *any* $\bar{\Omega}$):

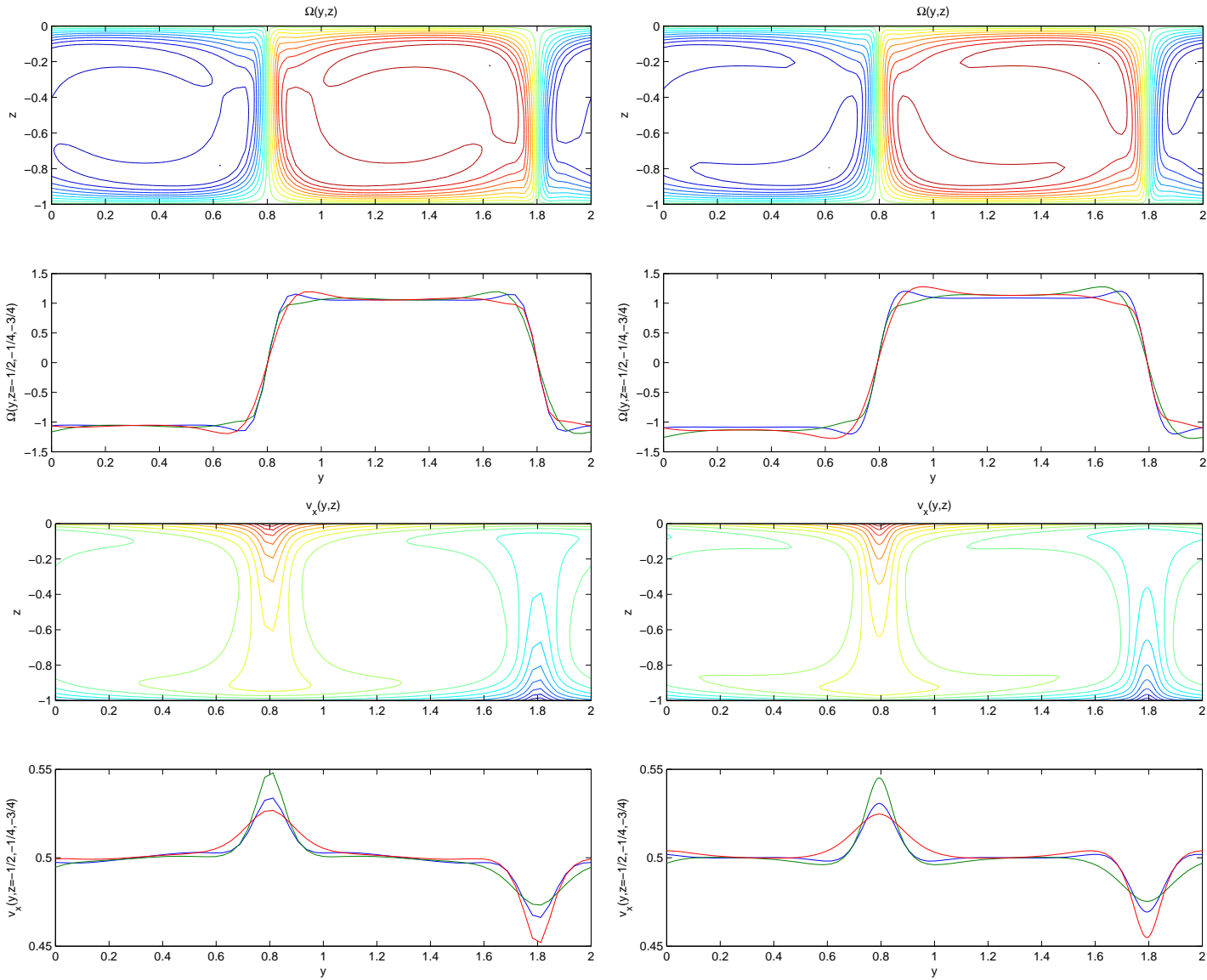
$$\frac{\pi}{2k} \sim -W(z) \int_0^\infty u_{III}(Y, z) dY = \int_0^\infty U_{III}(\psi_{III}, r) d\psi_{III}$$

$\Rightarrow \bar{\Omega} \sim 1$ (or -1), *independently of k .*

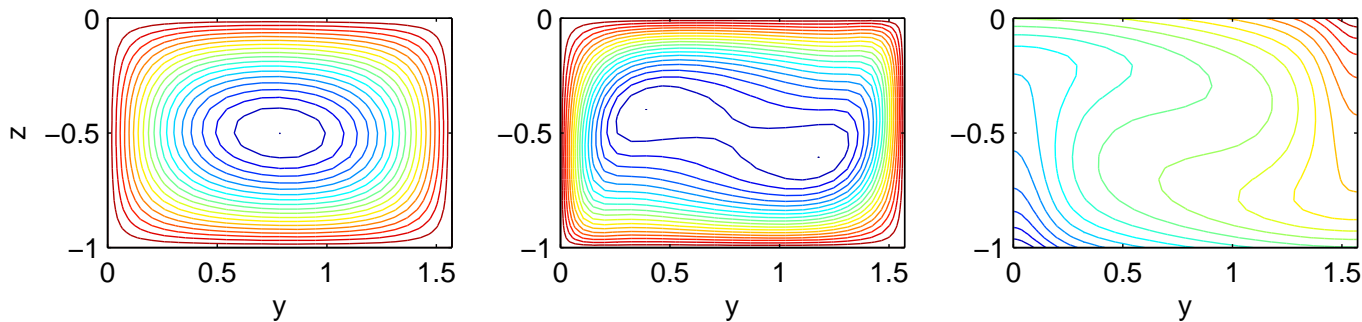
Numerical Simulations: Core Vorticity vs. k



Numerical Fields vs. Semi-Analytical Fields



Secondary Stability Analysis of Reduced PDEs



- Linearize about **fully nonlinear 2D** cellular flow, e.g.
 $U(X, y, z, T) = U_{2D}(y, z) + u(X, y, z, T)$
- Since coefficients in disturbance equations vary periodically in y , employ **Floquet theory**, e.g.

$$u(X, y, z, T) = e^{i\gamma y} \left[\sum_{n=-\infty}^{\infty} \hat{u}_n(z) e^{i(nky)} \right] e^{ilX} e^{\sigma t} + \text{c.c.}$$

- σ, l are the temporal growth rate, downwind wavenumber.
- k is the fundamental wavenumber of underlying 2D convective base flow.
- $i\gamma$ is the Floquet exponent, with the real parameter γ providing the freedom to modify the crosswind wavenumber.

Secondary Stability Equations

Small-amplitude 3D disturbances to fully nonlinear 2D roll solutions satisfy:

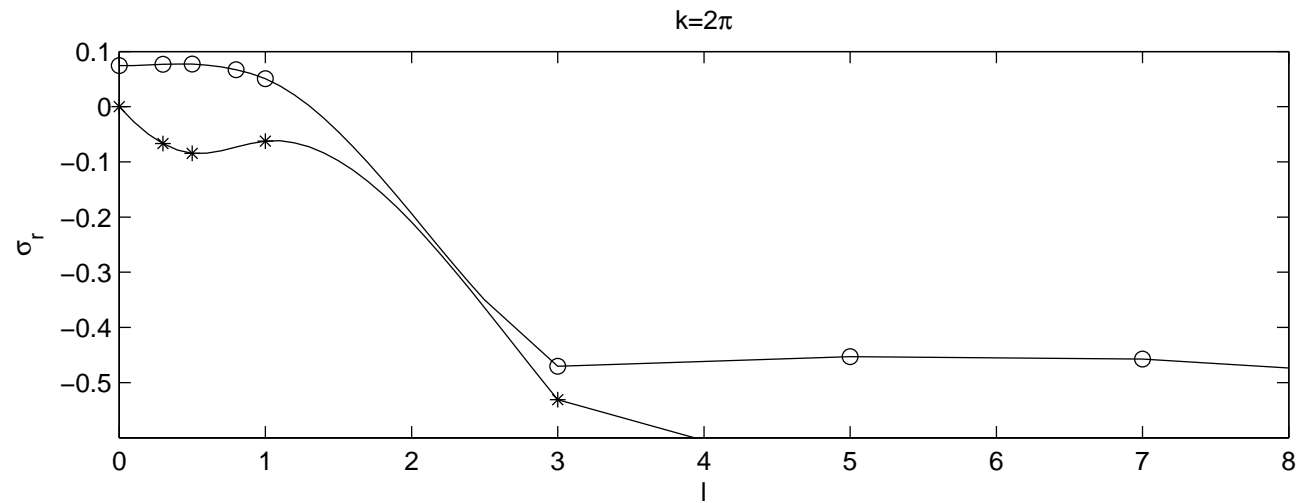
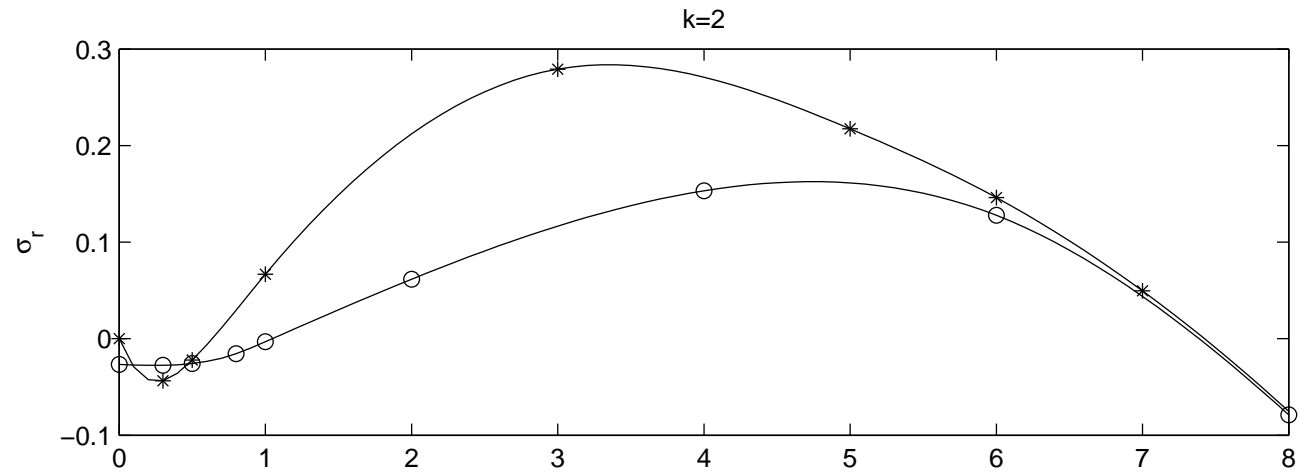
$$\partial_T u + \partial_z \psi_{2D} \partial_y u - \partial_y \psi_{2D} \partial_z u + \partial_y U_{2D} \partial_z \phi - \partial_z U_{2D} \partial_y \phi = -\partial_X p + La \nabla_{\perp}^2 u$$

$$\begin{aligned} \partial_T (\nabla_{\perp}^2 \phi) + U_s(z) \partial_X (\nabla_{\perp}^2 \phi) + \partial_z \psi_{2D} \partial_y (\nabla_{\perp}^2 \phi) \\ - \partial_y \psi_{2D} \partial_z (\nabla_{\perp}^2 \phi) + \partial_y \Omega_{2D} \partial_z \phi - \partial_z \Omega_{2D} \partial_y \phi = U'_s(z) [\partial_y u - \partial_X (\partial_z \phi)] \\ + La \nabla_{\perp}^4 \phi \end{aligned}$$

$$\begin{aligned} \nabla_{\perp}^2 p = -4 \partial_y (\partial_z \psi_{2D}) \partial_y (\partial_z \phi) + 2 \partial_y^2 \psi_{2D} \partial_z^2 \phi + 2 \partial_z^2 \psi_{2D} \partial_y^2 \phi + U_s(z) \nabla_{\perp}^2 u \\ + U'_s(z) [\partial_z u + \partial_X (\partial_y \phi)] \end{aligned}$$

with $\partial_z u = \phi = \partial_z^2 \phi = \partial_z p = 0$ along $z = 0, -1$.

Reduced PDEs: Secondary Stability Results



Summary

- Derived reduced PDEs for anisotropic turbulent Langmuir circulation in strong vortex-force limit.
- Reduced PDEs capture dominant linear and secondary instabilities.
- Reduced PDEs offer several analytical and computational advantages:
 1. Filter rapid-distortion (i.e. fast) transients \Rightarrow larger Δt (even in 2D).
 2. Filter fine x -scale variability \Rightarrow larger Δx and Δt .
 3. Vortex stretching is sub-dominant process \Rightarrow facilitates analysis (e.g. homogenization theory, upper-bound analysis).
 4. *Limiting process effectively suppresses “classical” shear-flow instability mechanisms.*

Ongoing Investigations and Future Work

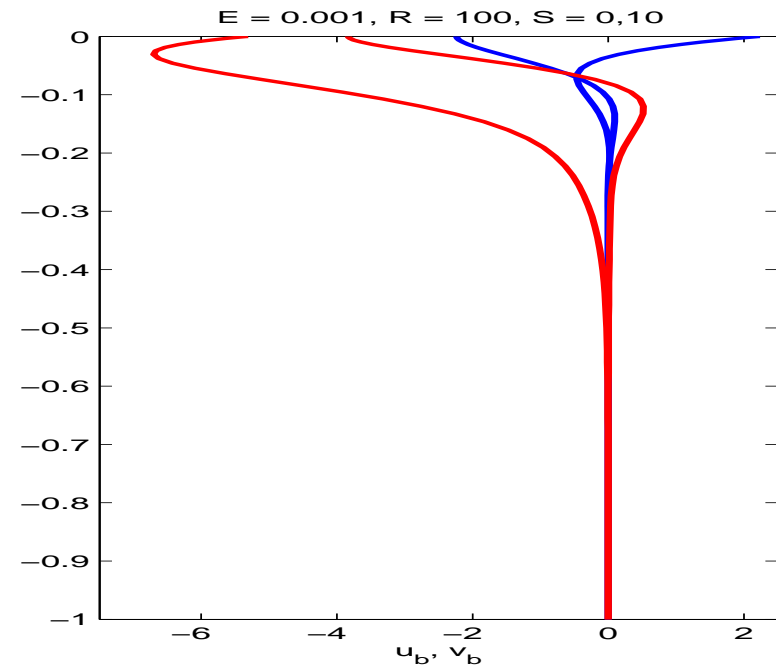
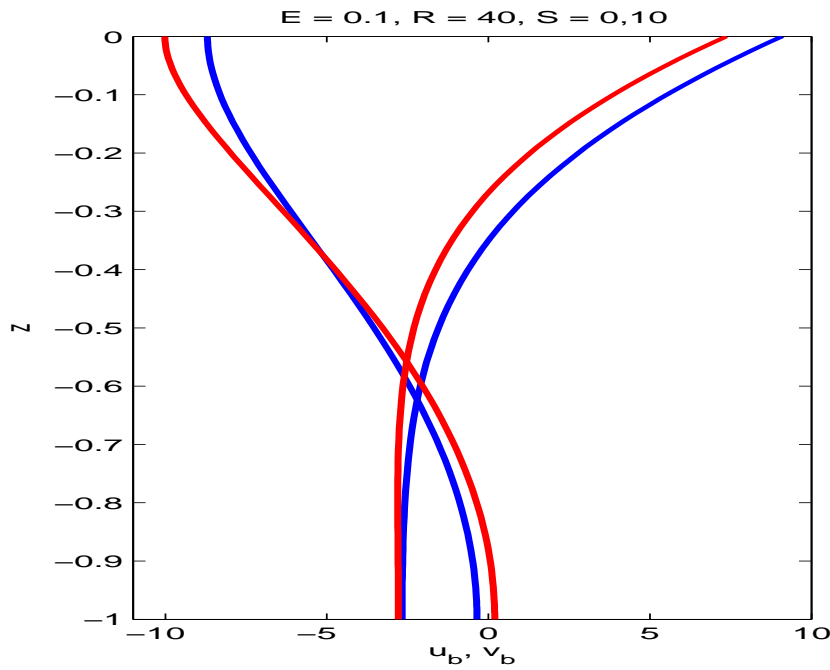
1. Influence of density stratification – coupling with internal gravity waves.
2. Influence of background rotation (Coriolis accelerations) – both kinematic and dynamic effects.
3. Coupling b/w submesoscale convective turbulence and submesoscale and mesoscale eddies using homogenization theory.
4. Application to other boundary-layer instability phenomena (e.g. Ekman rolls in the atmospheric boundary layer, classical shear flow rolls/streaks).

Langmuir Circulation–Internal Wave Interactions

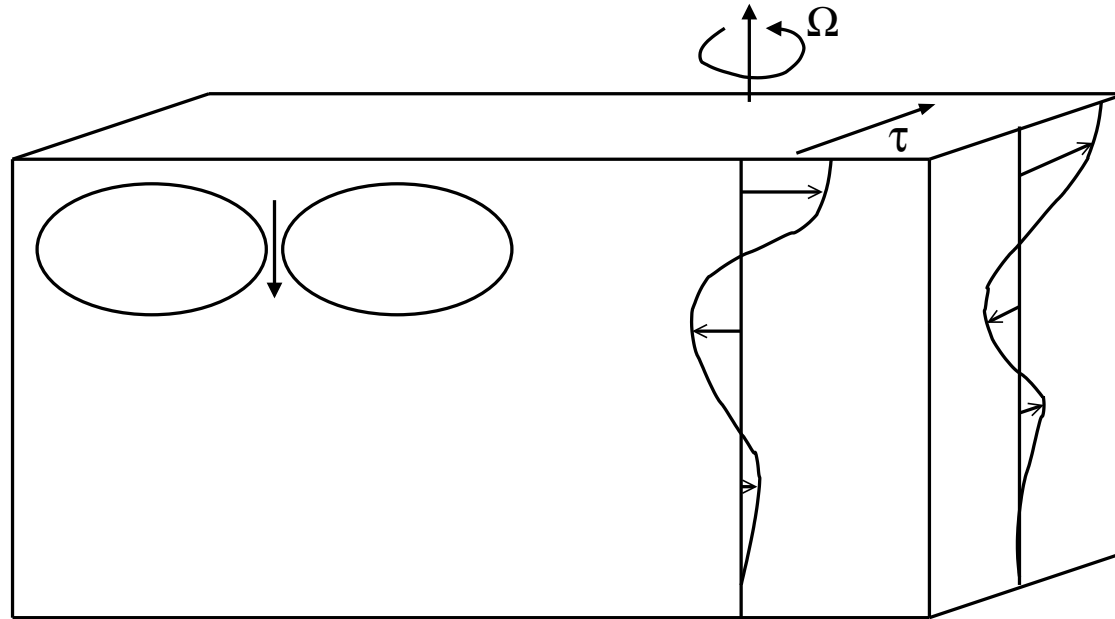
Rotating CL Equations

$$\partial_t \mathbf{u} + \mathbf{u} \cdot \nabla \mathbf{u} + \frac{1}{Ro} \left[\mathbf{f} \times \left(\mathbf{u} + \frac{1}{La_t^2} \mathbf{U}_s \right) \right] = -\nabla P + \frac{1}{La_t^2} (\mathbf{U}_s \times \boldsymbol{\omega}) + \frac{1}{R_*} \nabla^2 \mathbf{u}$$

Inclusion of Stokes-drift-induced Coriolis force implies classical Ekman spiral is modified.



Ekman Instability Modes

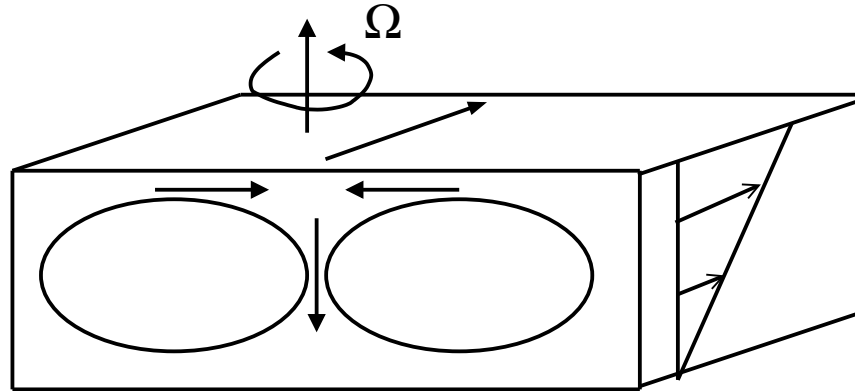


1. **Mode I – Inviscid** instability arising from **inflectional shear** in component of Ekman spiral **perpendicular** to roll axis:

$$\hat{R}_I = O(100), \hat{k}_I = O(0.5).$$

2. **Mode II – “Viscous”** instability arising from shear in component of Ekman spiral **parallel** to roll axis: $\hat{R}_C = 11.8, \hat{k}_C = 0.32$.

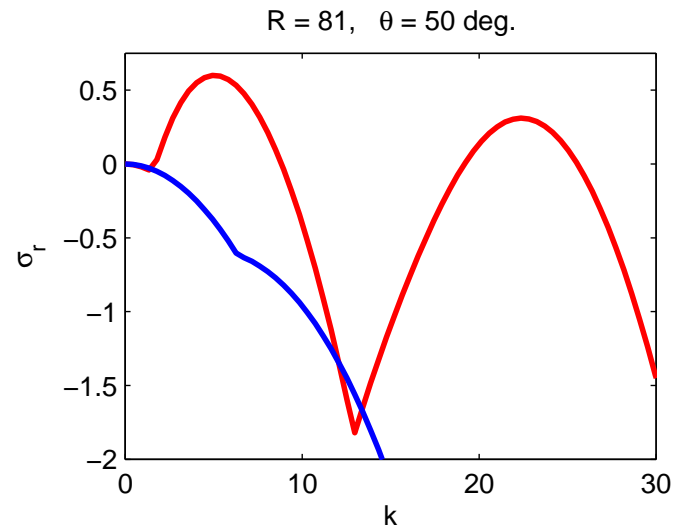
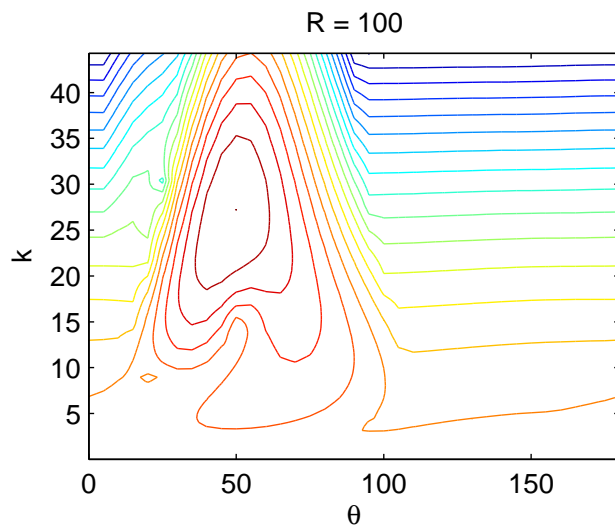
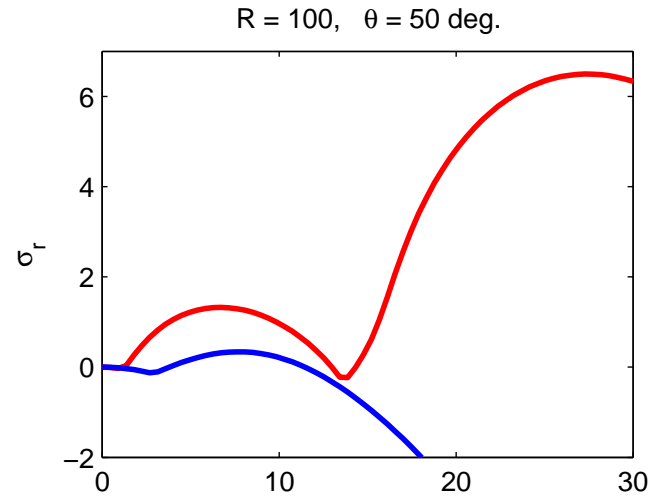
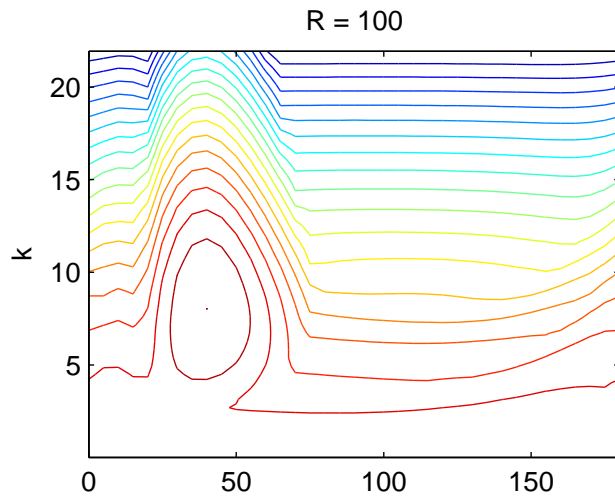
Vortex-Force Driven Instabilities: LC and Mode II



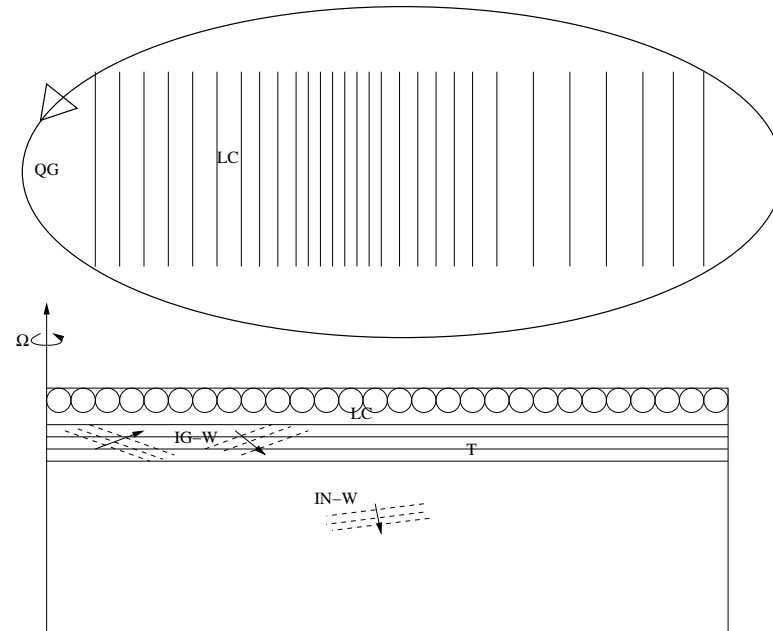
Common features:

1. A component of shear parallel to the roll axes.
2. A vortex force that couples along-roll velocity perturbations to the cross-roll cellular flow.

Influence of Stokes Drift ($E = 0.001$, $R = 100$)



Multiscaling via Homogenization Theories



- Develop homogenization theory for asymptotic strongly-nonlinear 2D LC solutions, exploiting **linearization** of CL equations using known $\psi(y, z)$.
- Employ novel Lagrangian homogenization formalism (T. Hou) that does **not** require strict scale separation and that can treat a nonlinear, dynamic (turbulent) “cell problem.”