

# *In Situ* and Remote Sensing of Turbulence

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# Turbulence Measurements

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- Turbulence is a stochastic process, and hence must be studied via the *statistics* of the process.
- Homogeneity, isotropy, eddy dissipation rate: these are all *defined* via the statistics.

# Important Statistical Characterizations

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## ■ Homogeneity:

“The probabilities for the components of the vector field at a number of points are invariant under rigid translations of the points and the vectors along which the components are computed.”

## ■ Isotropy:

“The probabilities for the components of the vector field at a number of points are invariant under isometries of the points and the vectors along which the components are computed.”

# Important Statistical Characterizations, cont'd

- An *isometry* is a length-preserving operator, e.g., translations, rotations, and/or reflections. That is, an orthogonal transformation,  $\mathbf{O}$ , followed by a translation,  $\mathbf{T}$ .
- Let the components at the points be:

$$\omega_1 = \mathbf{u}(\mathbf{x}_1) \square \mathbf{a}_1, \quad \omega_2 = \mathbf{u}(\mathbf{x}_2) \square \mathbf{a}_2 \cdots \omega_n = \mathbf{u}(\mathbf{x}_n) \square \mathbf{a}_n$$

$$\omega'_1 = \mathbf{u}((\mathbf{T} \circ \mathbf{O})\mathbf{x}_1) \square \mathbf{O}\mathbf{a}_1, \quad \omega'_2 = \mathbf{u}((\mathbf{T} \circ \mathbf{O})\mathbf{x}_2) \square \mathbf{O}\mathbf{a}_2 \cdots \omega'_n = \mathbf{u}((\mathbf{T} \circ \mathbf{O})\mathbf{x}_n) \square \mathbf{O}\mathbf{a}_n$$

- Then, isotropy is given by, (homogeneity is when  $\mathbf{O} = \mathbf{I}$ ):

$$P(c_1 \leq \omega_1 \leq d_1, c_2 \leq \omega_2 \leq d_2, \cdots c_n \leq \omega_n \leq d_n)$$

$$= P(c_1 \leq \omega'_1 \leq d_1, c_2 \leq \omega'_2 \leq d_2, \cdots c_n \leq \omega'_n \leq d_n)$$

# Importance of eddy dissipation rate as an intensity parameter

- Under appropriate simplifications (homogeneous, solenoidal, constant density flow; no transport, diffusion, and convection of energy), the turbulent kinetic energy equation can be written as:

$$\frac{de}{dt} = -\varepsilon$$

- Furthermore, the Kolmogorov energy spectrum (isotropic turbulence in the inertial subrange) is:

$$E(k) = A \varepsilon^{2/3} k^{-5/3}$$

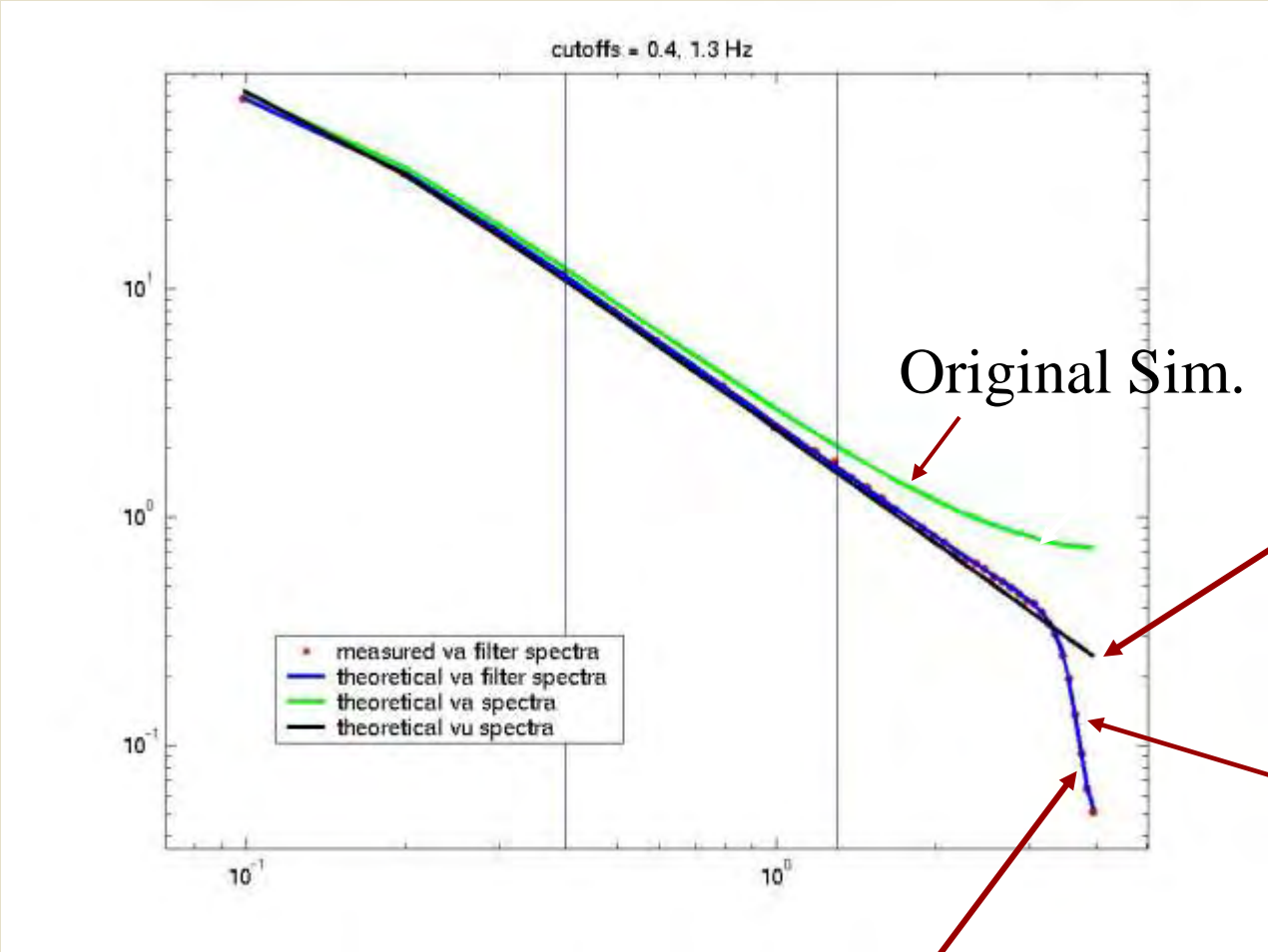
# Simulation Examples – A Cautionary Tale

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- Wind field is generated so that it has *correct spatial statistics*
  - This means that it will be aliased.
  - An optional simulation incorporating an anti-aliasing filter is also used.
- The spectral model is a von Karman one.  
Inputs:  $\sigma = 5.0ms^{-1}$ ,  $L_u = 500m$ ,  $\varepsilon^{2/3} = 0.545m^{4/3}s^{-2}$
- $\varepsilon^{2/3}$  estimates from maximum likelihood method.

# Spectra from model and simulated data



Original Sim.

Theoretical Von Karman Spectrum (black)

Calculated Averaged Spectra (red dots)

Theoretical Sim. w/ anti-alias filter (blue)

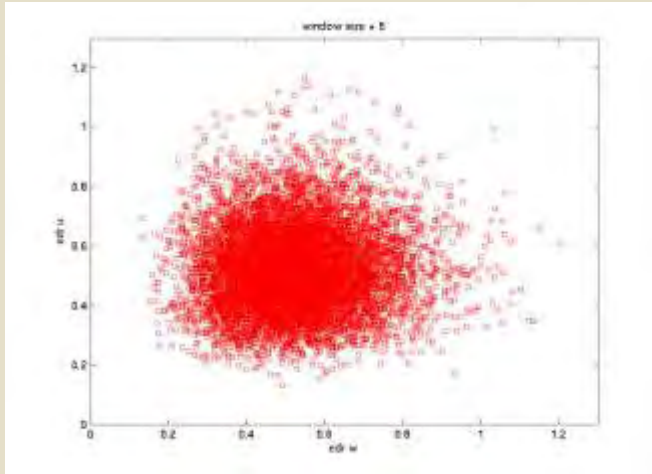
# Isotropy. Simulated data with $\varepsilon_u^{2/3}$ versus $\varepsilon_w^{2/3}$ along a line.

Theory says,  $\langle \varepsilon_u^{2/3} \rangle = \langle \varepsilon_w^{2/3} \rangle$

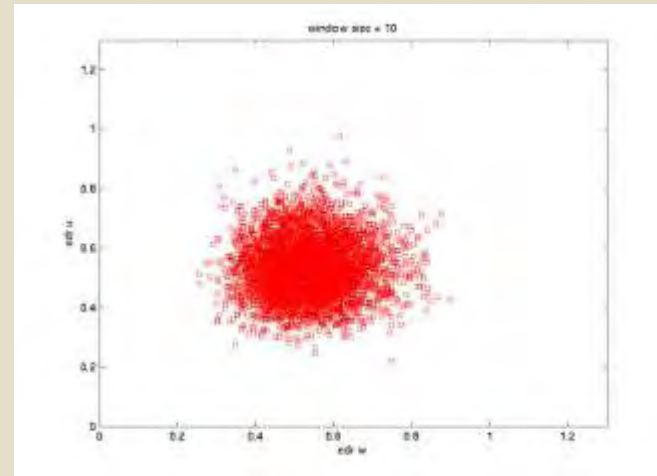
5 sec. windows

10 sec. windows

$\varepsilon_u^{2/3}$

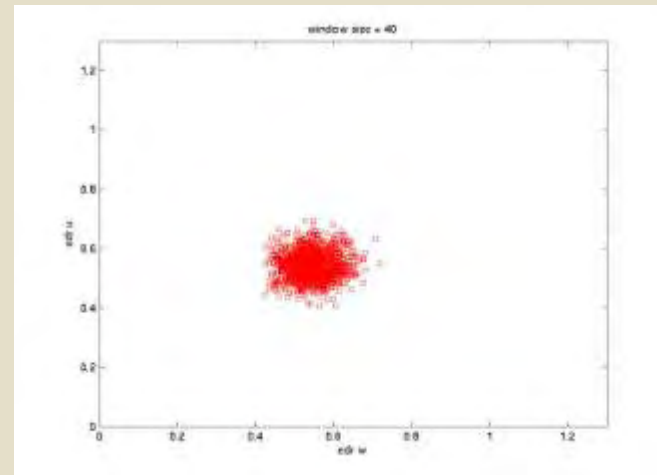
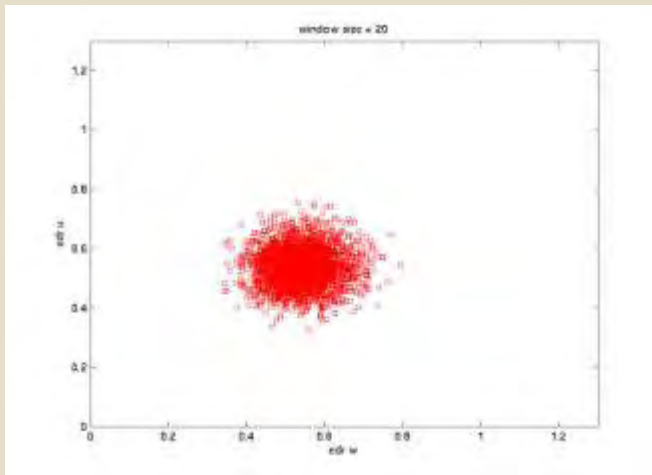


$\varepsilon_w^{2/3}$



20 sec. windows

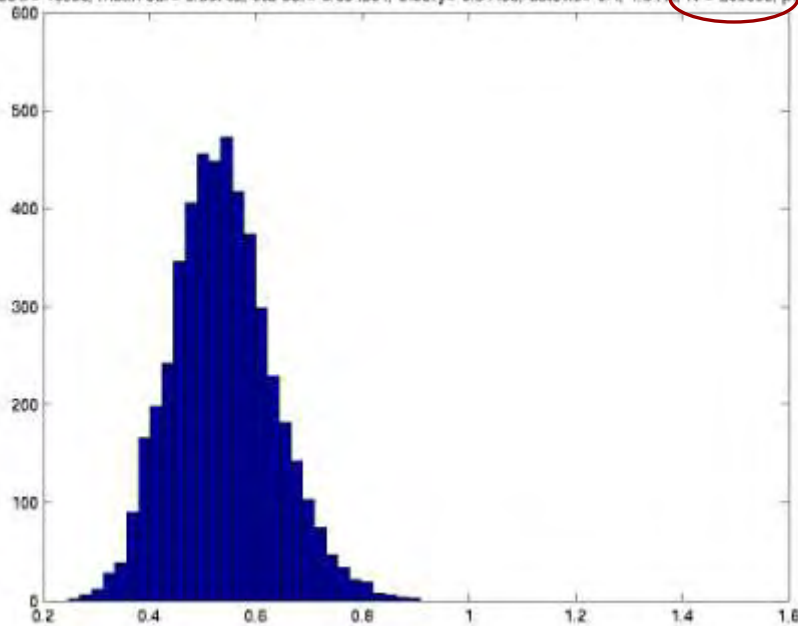
40 sec. windows





# Homogeneity and sample size statistics: Histogram of $\varepsilon_u^{2/3}$ over 10 and 40 second intervals.

window = 10sec, mean edr = 0.53732, std edr = 0.094981, theory = 0.54495, cutoffs = 0.4, 1.3 Hz, N = 200000, percent = 0

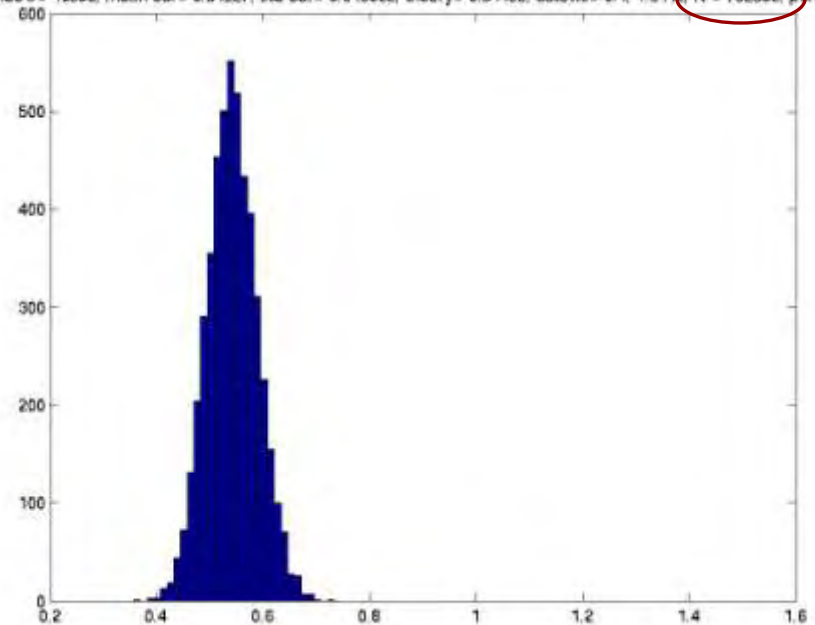


**Number of samples**

**10 Second Windows**



window = 40sec, mean edr = 0.54227, std edr = 0.046602, theory = 0.54495, cutoffs = 0.4, 1.3 Hz, N = 792500, percent = 0



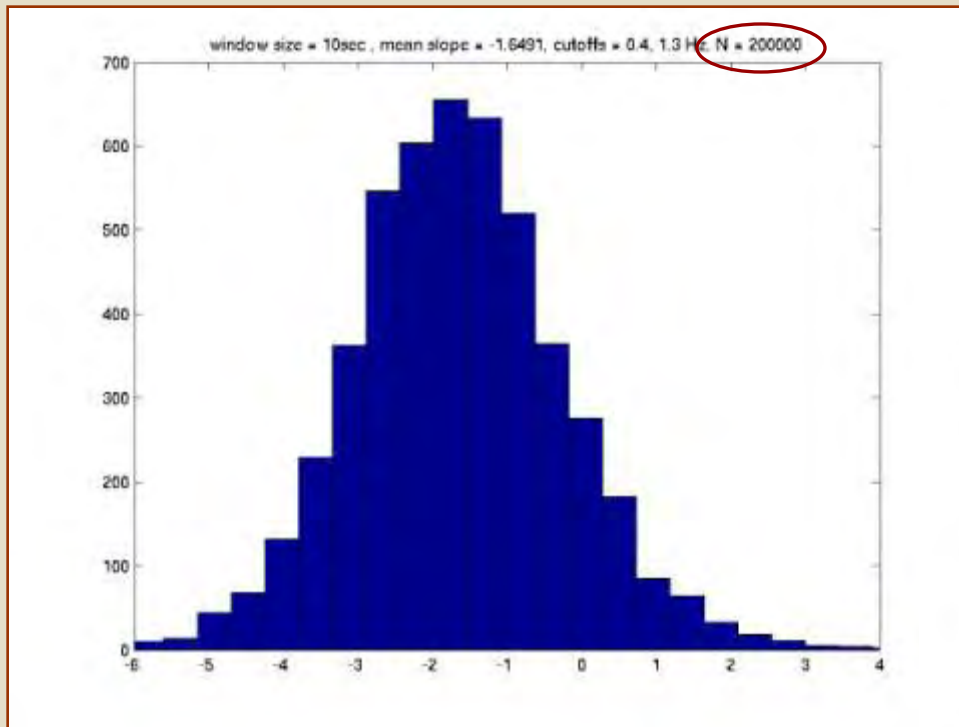
**40 Second Windows**



**Kolomogorov energy spectrum:**  $E(k) = A\varepsilon^{2/3}k^{-5/3}$

Hence, the slope in log-log should be  $-5/3$  – *on average*.

## 10 Second Windows



## 40 Second Windows



# Spectral Averaging

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## Procedure:

- Take as long of a segment of the time series where the data is reasonably stationary.
- Divide it into equal non-overlapping, or half-overlapping, sub-segments.
- Apply a window function – especially important for shorter and/or overlapping windows.
- Compute the spectrum for each segment.
- Average the spectra frequency-by-frequency.

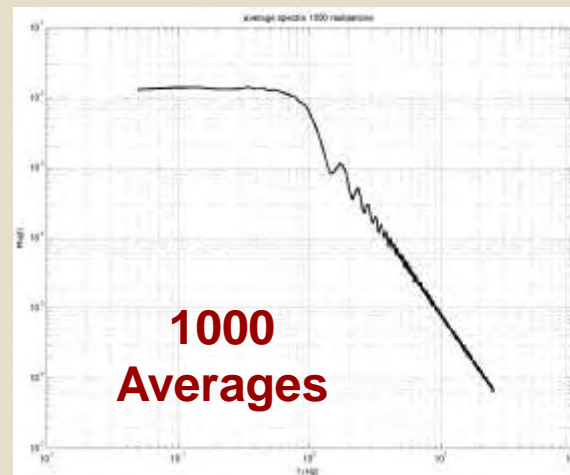
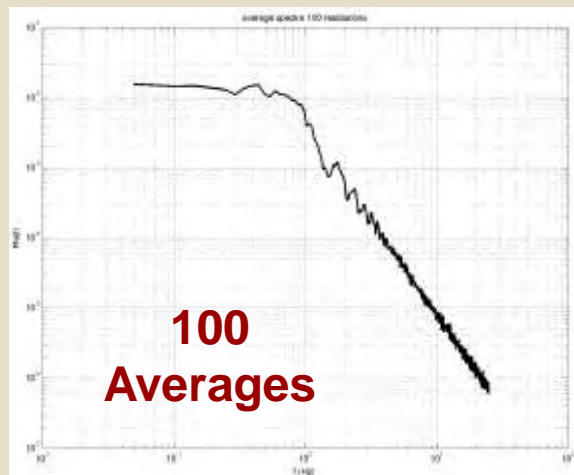
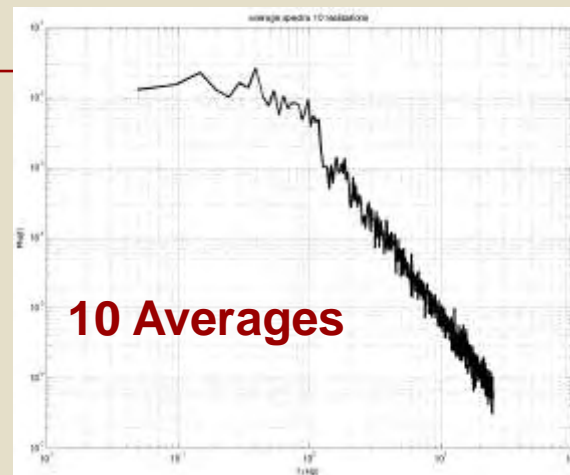
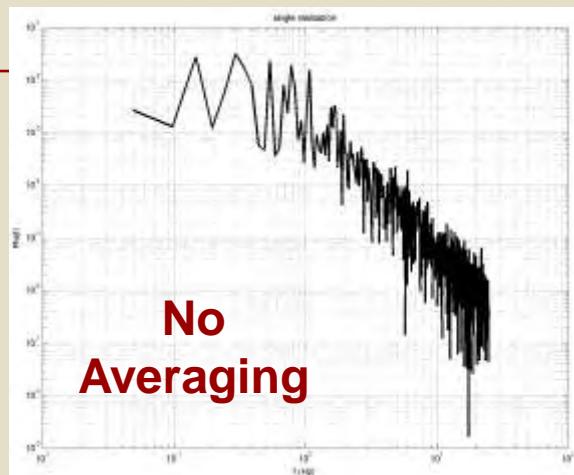
# Spectral Averaging

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## Benefits:

- Reduces the random effects of turbulence.
- It does *not* reduce the noise level but improves signal detectability, which leads to better parameter estimation.

# Another simulation example – Illustrating the beneficial effects of spectral averaging



# Parameter Estimation

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- This is a very deep and important subject, however, consider a simple one-parameter estimation approach: the *maximum likelihood estimator (ML)*.
- Assume that the turbulence can be approximated as a product-model:

$$y(t) = a(t)x(t)$$

- Where  $a(t)$  is deterministic amplitude function, and  $x(t)$  is a realization from a zero-mean, unit-amplitude, stationary random process.

# Parameter Estimation, cont'd.

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- If  $a(t)$  is constant or varies at a much more lower frequency than  $x(t)$  does, then the auto-spectrum of  $y(t)$  can be written as:

$$\Phi_{yy}(f) = a^2(t)\Phi_{xx}(f)$$

- Assume that  $a(t)$  is a constant over the sampling interval, and it is the parameter of interest.

# Parameter Estimation, cont'd.

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- Then, the maximum likelihood estimator for  $a$  is given by:

$$\hat{a} = \frac{1}{N} \sum_{i=1}^N \frac{S_{yy}(f_i)}{\Phi_{xx}(f_i)}$$

- Where,  $S$  is the measured spectrum, and the sum is taken over a frequency range wherein it is assumed that the model spectrum is valid.



# Parameter Estimation, cont'd.

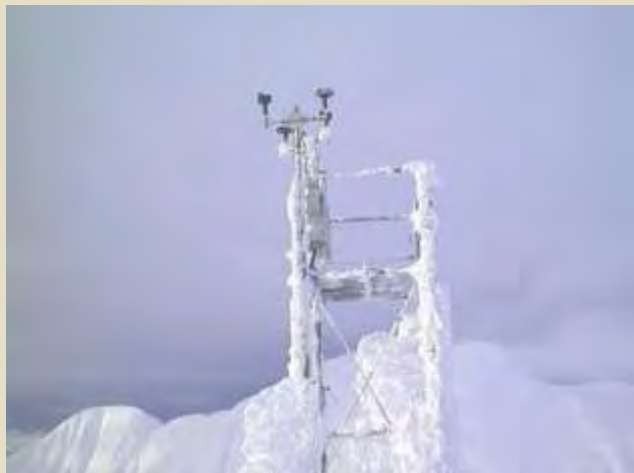
- For example, take a time series of a component of the wind taken from an anemometer.
- Assume that the turbulence at higher frequencies is given by the Kolmogorov form:  $\Phi_{xx}(f) = C\varepsilon^{2/3} f^{-5/3}$
- The ML estimate of  $\varepsilon^{2/3}$  is given by:

$$\hat{\varepsilon}^{2/3} = \frac{1}{N} \sum_{i=1}^N \frac{S_{yy}(f_i)}{C f_i^{-5/3}}$$

# Quality Control of Measurement Data

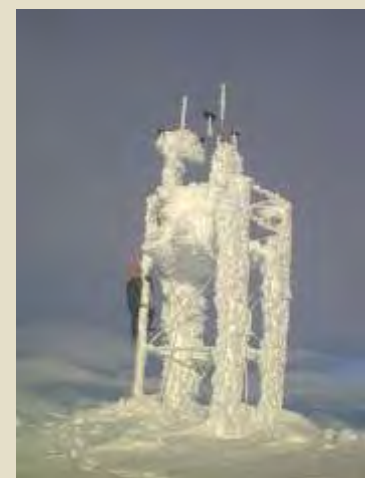
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# Quality Control of Measurement Data



**“Hmmm... These winds  
look a bit screwy.”**

**Rime ice accumulation  
on Mt. top anemometers  
in Juneau AK**



# Quality Control of Measurement Data

- Examples of quality control problems.
- Time series and “lag plots.”
- Our eyes do a good job of finding the good data amongst the outliers. Except for the third one from the top – the “good” data is hidden in the noise.



# The Top 10 List for “Success with Sensors”

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- Garbage in – garbage out. The more critical the application, the more important the QC.
- Whatever can go wrong with a sensor will go wrong.
- Your list of potential sensors failure modes is not long enough.

# “Success with Sensors” Cont’d

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- Do not believe the manufacturer:
  - They’re trying to sell devices at the minimum cost to them.
  - They usually use very simplistic QC methods.
  - Processing at the rawest data level is best – may require modification to the device to make that available externally.
  - Don’t trust their specifications.

# “Success with Sensors” Cont’d

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- Statistical methods are fine - as presented in the literature – the real world is another matter altogether:
  - A theorem does not make for a good algorithm.
  - Assumptions made in proving theorems are often violated with real data.
  - Data with lots of outliers will kill their methods.
  - Empirical methods may not make for a good journal article – but they work.
  - No one publishes their failures.

# “Success with Sensors” Cont’d

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- *A priori* assumptions are dangerous – model-free methods are more robust.
- Simulation is a great tool to start with, but you need to learn from real data.
- Typically, no one’s solved your problem – but you can learn from other’s mistakes and successes. *Use the literature – talk to colleagues.*



# “Success with Sensors” Cont’d

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- There’s no simplistic or black box solution – you have to understand the problem inside-out.
- **There’s no substitute for hard work:** *“Do you want to explain why the wind farm had to be taken down because a bird defecated on a wind sensor?”*

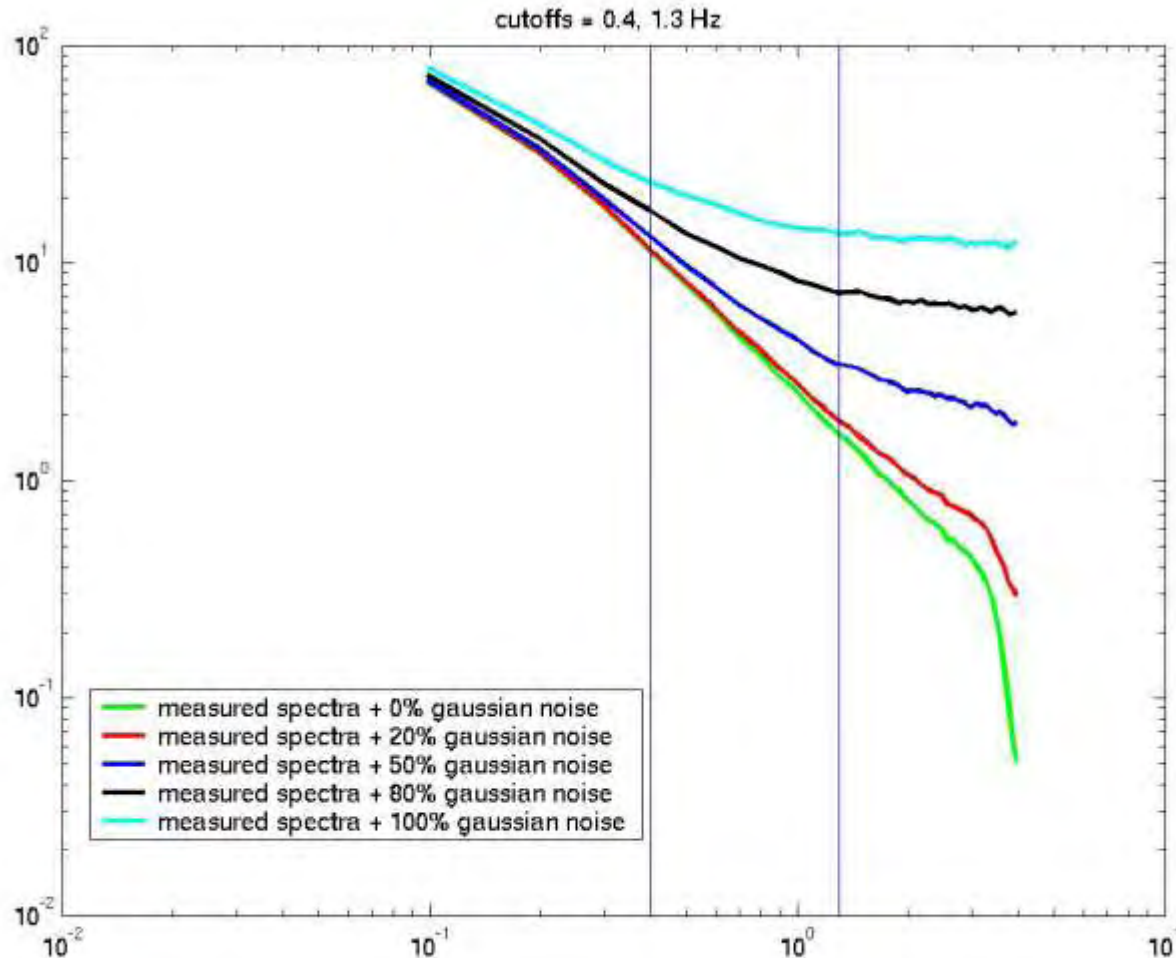
# Issues Regarding Low SNR Measurements

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- Low SNR means different things for different devices
  - Radar: low reflectivity (hydrometeors)
  - Lidar: low backscatter (aerosol density)
  - Wind profilers: low index of refraction (temperature and/or humidity fluctuations)
  - Anemometers: low wind speed

**In each case, it is important to recognize when the device is giving meaningful information – *and when it isn't.***

# Simulated Spectra with added Gaussian noise



**Averaged  
Spectra with  
Anti-  
Aliasing  
Filter.**

**EDR estimation algorithms (e.g., ML) can be modified to accommodate additive noise – but this only works to a point.**

# Some time series QC algorithms that are useful

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## ■ Running Median Filter

- Simple to implement.
- Fairly efficient for reasonable window size.
- Works well for isolated outliers.
- Has “saturation” problem when  $> 1/2$  of the samples in the window are bad.
- Inherent  $1/2$  window length lag.

# Some time series QC algorithms that are useful, Cont'd

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## ■ Auto-regression methods

- Simple to implement.
- Very efficient.
- Works well for isolated outliers.
- Has problems with abrupt changes in data.
- Can “lock onto” bad signal if it is relatively smooth.

# Some time series QC algorithms that are useful, Cont'd

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## ■ Wavelets

- Fairly efficient.
- Works well for isolated outliers and step changes.
- Can also be used to filter low or high frequencies.
- Works well for non-stationary signals.
- Some “art” required in implementation.

## Some time series QC algorithms that are useful, Cont'd

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- **NCAR Improved Moment Algorithm (NIMA)**
  - Developed for wind profiler application – Doppler second moments can be used to calculate turbulence.
  - Fuzzy logic image processing algorithm for finding atmospheric part of Doppler spectra in the presence of contaminants.
  - Not very efficient.
  - Very specific for the application.

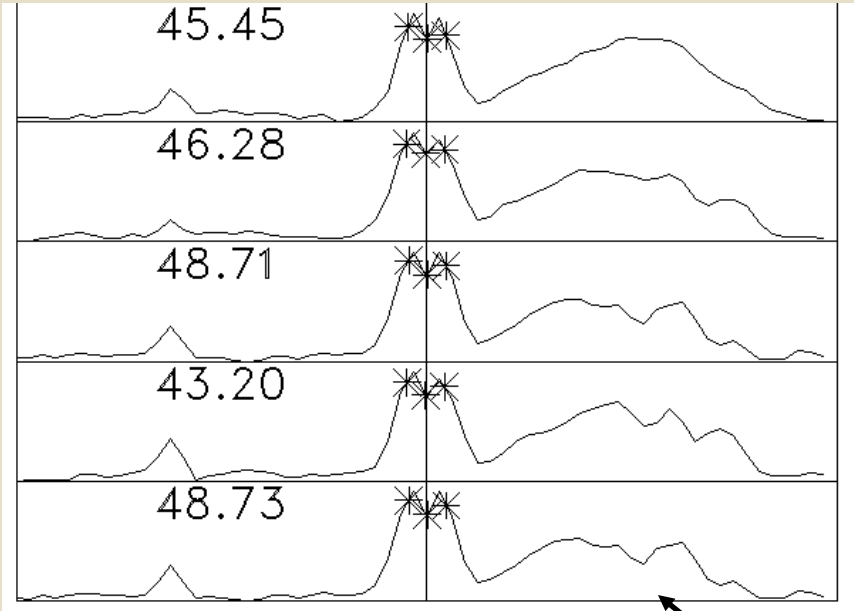
# The NIMA Method

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- ■ **NIMA tries to imitate human experts using**
  - Mathematical analysis
  - Fuzzy logic synthesis
  - Global image processing
- **Problem is broken down into smaller sub-problems**
  - Doppler peak detection
  - Clutter feature detection
  - RFI feature detection
  - Atmospheric feature detection
  - Continuity assurance
  - **Confidence estimation**



# Doppler spectra as a function of range and velocity

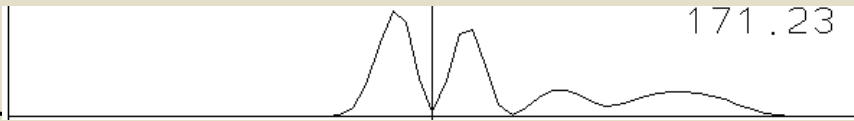
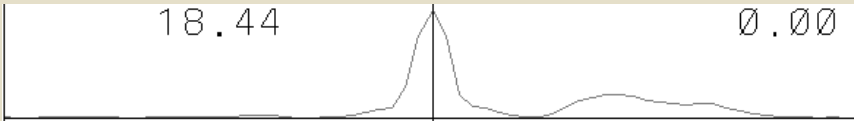
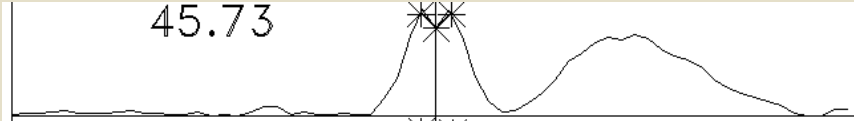


RFI contamination

Ground clutter contamination

Wind

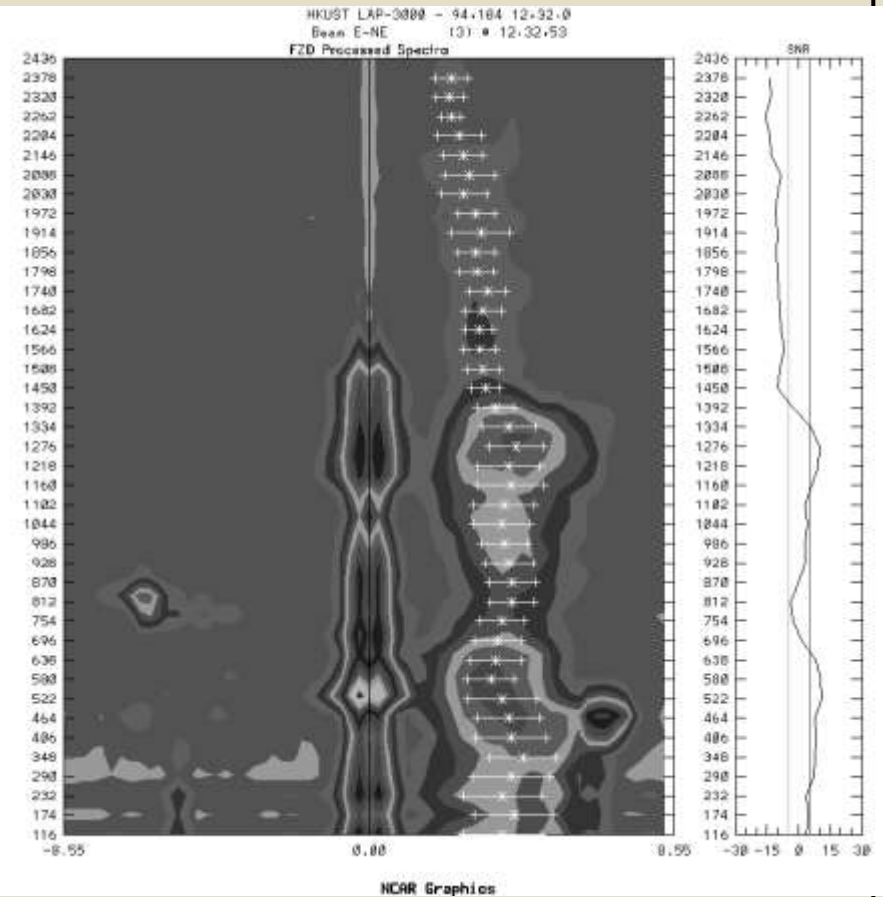
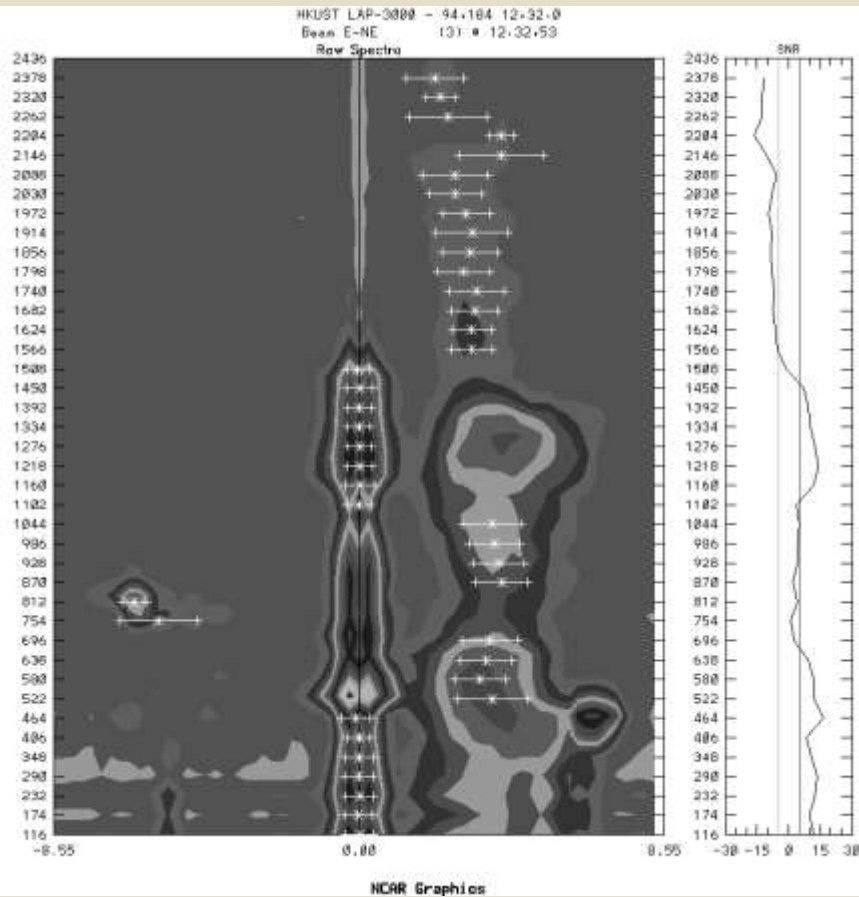
# Doppler spectra at one range



Slope information

Curvature information

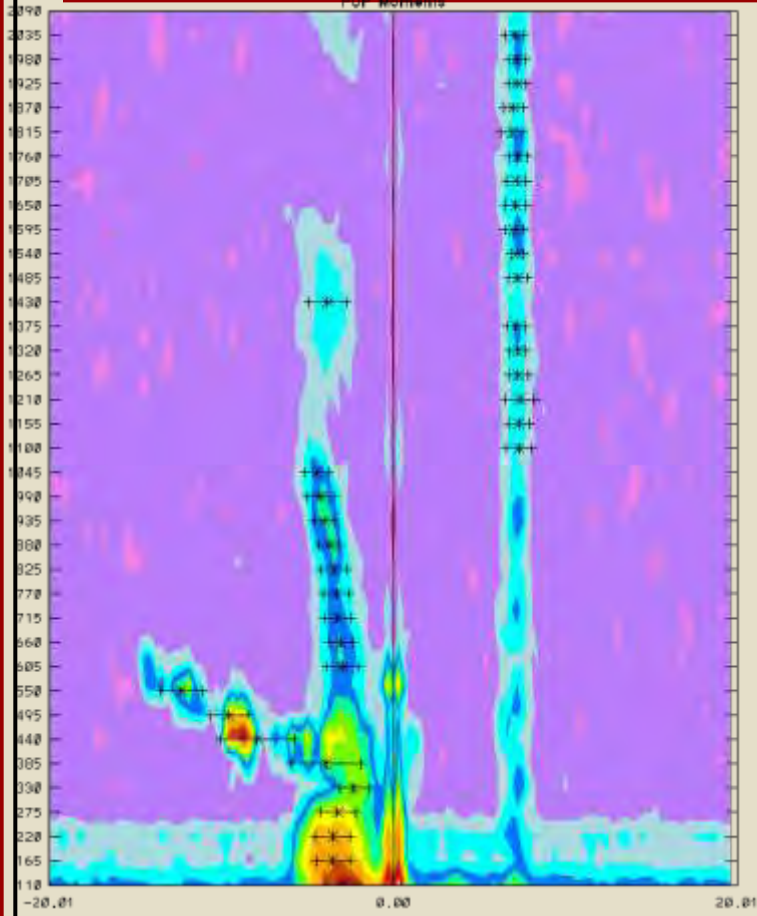
# Before and after NIMA processing



# Before and after NIMA processing

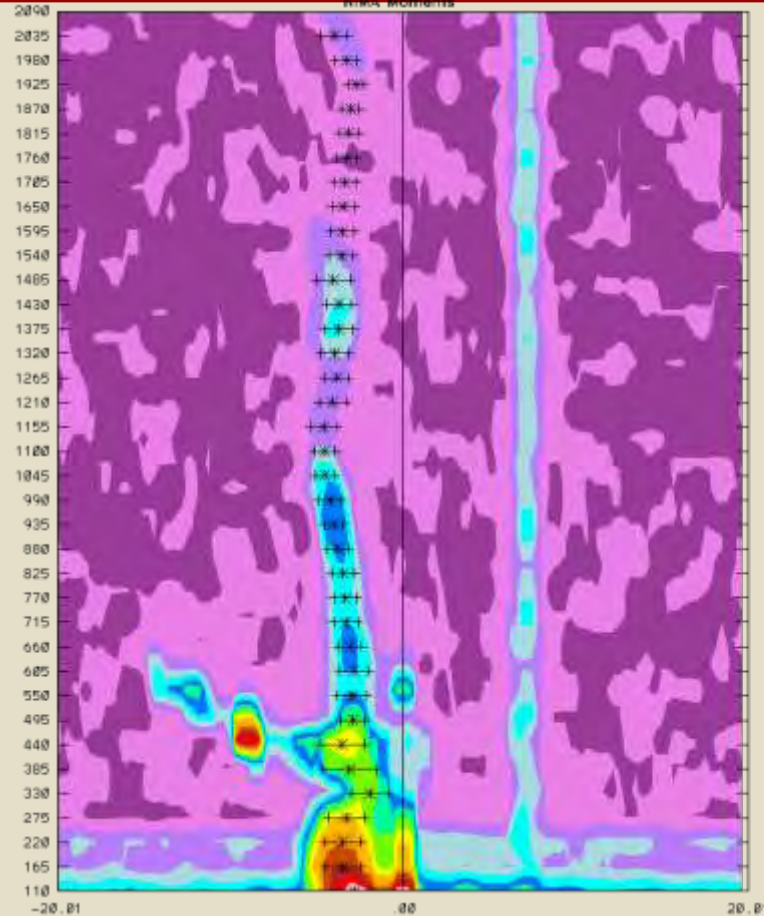
Lenon Creek LAP-3000 - Start @ 98.043 18.55  
Beam West (3) @ 18.55.03

POP Moments



Lenon Creek LAP-3000 - Start @ 98.043 18.55  
Beam West @ 18.55.03

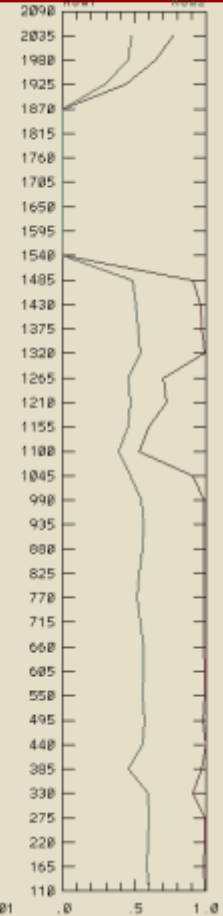
NIMA Moments



## Confidence

Confidence

West East



# Remote Sensing of Turbulence

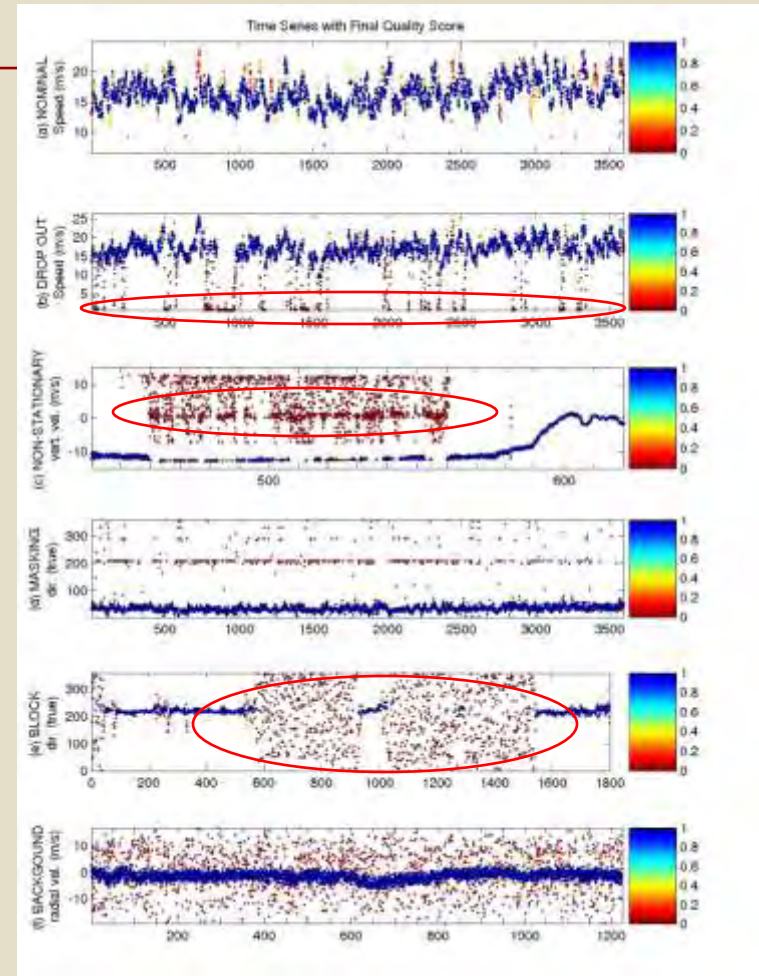
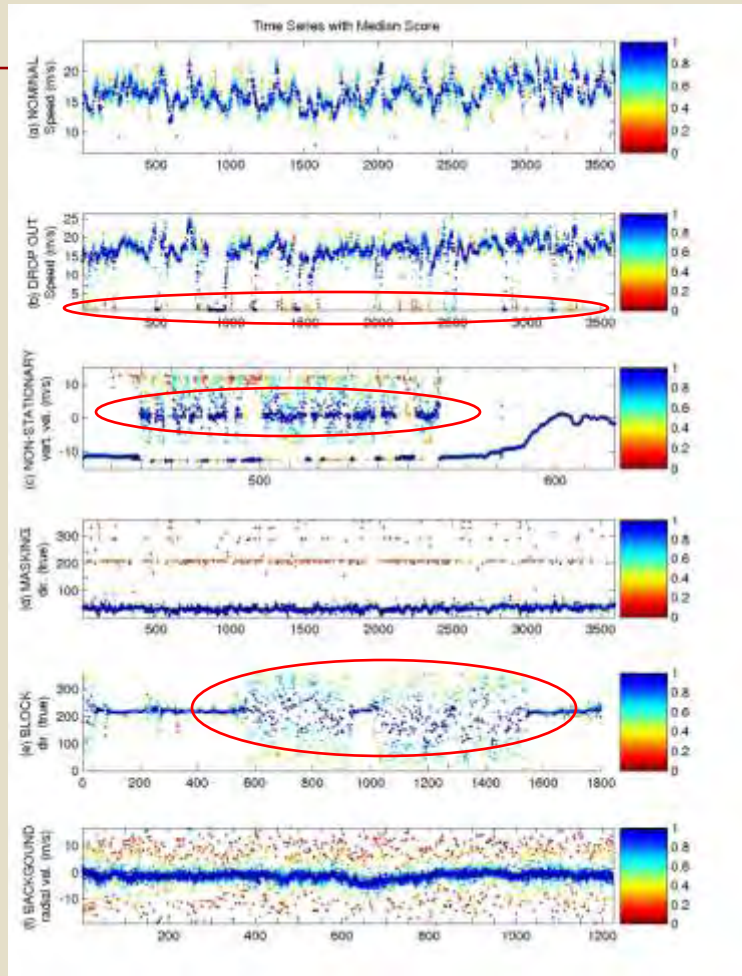
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## Some time series QC algorithms that are useful, Cont'd

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- **Intelligent Outlier Detection Algorithm (IODA)**
  - Very powerful times series QC and data classification algorithm.
  - Handles larger numbers of outliers.
  - Identifies some failure modes.
  - Requires correlation structure in data.
  - Not very efficient.

# Running Median (left) vs. IODA (right)



# Remote Sensing of Turbulence: A Primer

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- Remote sensing of turbulence is performed via active or passive sensors, and hence the measurements are typically highly affected by the sensor.
- The sensor acts as a “filter” on the atmospheric turbulence, and hence we are solving an inverse problem:

*Given sensor measurements, and a model of the sensor, what was the turbulence that produced the measurements?*

# Remote Sensing of Turbulence: A Primer

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- Measuring turbulence with *active sensors* can be broken-down into two types of problems:
  - **Backscatter**. Transmitter and receiver are typically at the same place.
  - **Propagation**. Transmitter and receiver are on opposite sides of the sample volume. (E&M, optical, acoustic)



# Remote Sensing of Turbulence: A Primer

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- Typical backscatter devices include: radars, lidars, and sodars.
  - **Radars** reflect microwaves (mm to cm) off of hydrometeors (rain, snow, ice, etc.) and index of refraction fluctuations (typically, temperature or humidity variations).
  - **Lidars** reflect photons off of aerosols (micron-sized particulates)
  - **Sodars** reflect acoustic waves off of index of refraction fluctuations (typically, temperature variations).

# Remote Sensing of Turbulence: A Primer

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- Typical **propagation** measurements of turbulence are obtained by mm-cm radio waves, optical and acoustic waves which are diffracted by temperature and/or humidity fluctuations in the index of refraction.

# Remote Sensing of Turbulence: A Primer

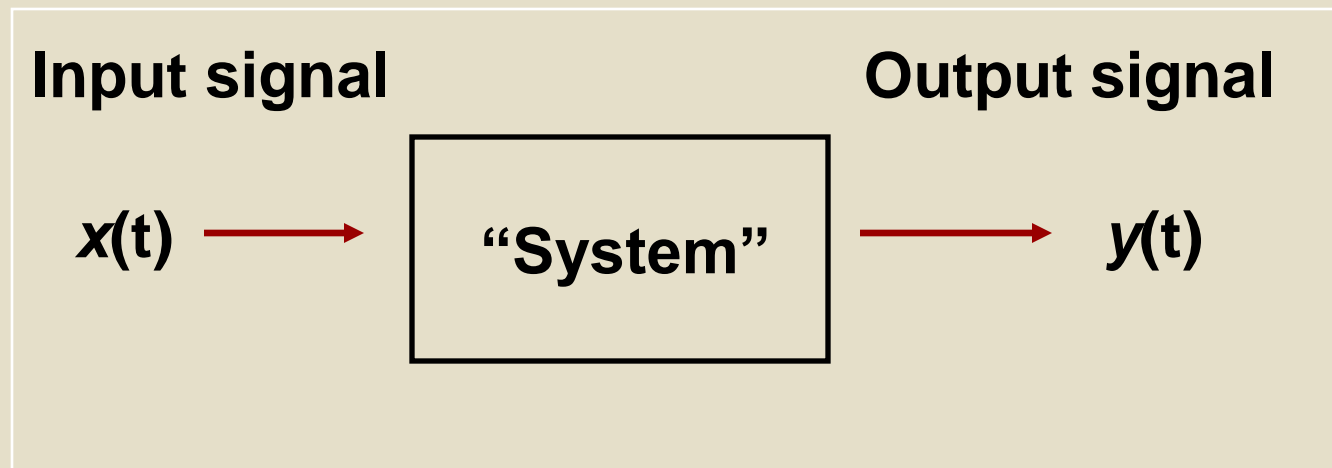
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- **Passive** sensors used to obtain measurements of turbulence are typically:
  - Infrared devices which measure the re-radiated field from atmospheric atoms and molecules. The source typically being solar radiation re-radiated from the Earth's surface.
  - Optical sensors which measure the light coming from stars.

# Remote Sensing of Turbulence: A Primer

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- Returning to the “sensor as turbulence filter” problem... —
- Consider a single input – single output relation:



# Remote Sensing of Turbulence: A Primer

- For a linear system, the mathematical relationship is a convolution integral:

$$y(t) = \int_0^{\infty} h(\tau) x(t - \tau) d\tau$$

- Where,  $h(t)$ , is the *unit impulse response function* for the linear system.
- Taking the Fourier transform of both sides gives:  
$$Y(f) = H(f)X(f)$$

# Remote Sensing of Turbulence: A Primer

- Considering  $x(t)$  as being a stochastic and stationary signal, the input-output correlation and spectral relationships are:

$$R_{yy}(\tau) = \int_0^{\infty} \int_0^{\infty} h(\alpha)h(\beta)R_{xx}(\tau + \beta - \alpha)d\alpha d\beta$$

$$\Phi_{yy}(f) = |H(f)|^2 \Phi_{xx}(f)$$

Note the simplicity of the spectral relation.

# Remote Sensing of Turbulence: A Primer

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- Where for a stationary signal  $s(t)$ ,

$$R_{ss}(\tau) = \langle s(t)s(t + \tau) \rangle$$
$$\Phi_{ss}(f) = \int_0^{\infty} R_{ss}(\tau) e^{-2\pi f\tau} d\tau$$
$$= \langle Y(f)Y^*(f) \rangle$$

# Remote Sensing of Turbulence: A Primer

- The same relationships work for the components of a homogeneous spatial field,  $\mathbf{x}(\mathbf{r})$ .

$$y_i(\mathbf{r}) = \int_0^\infty \int_0^\infty \int_0^\infty h(\boldsymbol{\rho}) x_i(\mathbf{r} - \boldsymbol{\rho}) d\boldsymbol{\rho}$$

$$Y_i(\mathbf{k}) = |H(\mathbf{k})|^2 X_i(\mathbf{k})$$



# Remote Sensing of Turbulence: A Primer

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- Where the correlation and spectral relationships for a homogeneous field are:

$$R_{s_i s_j}(\boldsymbol{\rho}) = \left\langle s_i(\mathbf{r}) s_j(\mathbf{r} + \boldsymbol{\rho}) \right\rangle$$

$$\Phi_{s_i s_j}(\mathbf{k}) = \int_0^{\infty} R_{s_i s_j}(\boldsymbol{\rho}) e^{-2\pi \mathbf{k} \cdot \boldsymbol{\rho}} d\boldsymbol{\rho}$$

# Remote Sensing of Turbulence: A Primer

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- Consider a specific example: Doppler (backscatter) radar measurements of turbulence.
- Two approaches:
  - Spatial spectrum of the radial velocities.
  - Doppler second moment.

# Remote Sensing of Turbulence: A Primer

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- For both approaches, we need the following relationship:  $\bar{v}(\mathbf{r}_0) \equiv M_1(\mathbf{r}_0) = \int u S_n(u, \mathbf{r}_0) du$
- Where  $S_n(u, \mathbf{r}_0)$  is the normalized Doppler spectrum for the pulse volume centered at  $\mathbf{r}_0$ ,  $M_1(\mathbf{r}_0)$  is the first moment of the Doppler spectrum, and  $u$  is the radial component of the wind field.

# Remote Sensing of Turbulence: A Primer

- Assuming uniform reflectivity (i.e., uniform hydrometeor size and distribution), it can be shown that the first moment can be written as a spatial integral:

$$S_n(u, \mathbf{r}_0) = \int \delta(u - u(\mathbf{r})) I_n(\mathbf{r} - \mathbf{r}_0) d\mathbf{r}$$

- So that, 
$$v(\mathbf{r}_0) = \int u(\mathbf{r}) I_n(\mathbf{r} - \mathbf{r}_0) d\mathbf{r}$$

i.e., a spatial convolution.  $I_n(\mathbf{r} - \mathbf{r}_0)$  is the pulse-volume illumination function – the response function for this problem.

# Remote Sensing of Turbulence: A Primer

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- From the convolution integral, the correlation and spectral relations can be obtained:

$$R_{vv}(\mathbf{r}_0, \boldsymbol{\rho}) = \int R_{uu}(\boldsymbol{\rho}) I_n(\mathbf{r} - \mathbf{r}_0) I_n(\mathbf{r} + \boldsymbol{\rho} - \mathbf{r}_0) d\mathbf{r}$$

$$\Phi_{vv}(\mathbf{r}_0, \mathbf{k}) = \Phi_{uu}(\mathbf{r}_0, \mathbf{k}) I_n^2(\mathbf{r}_0, \mathbf{k})$$

# Remote Sensing of Turbulence: A Primer

- Assume that the radar pulse volume is narrow enough so that the radial velocities can be approximated via a Cartesian component.
- Assume that the velocity spectrum is isotropic, so that the coordinate system can be rotated such that the  $x$ -axis points along the radar beam.
- Assume that the turbulence spectrum can be modeled as  $\Phi_{u_x u_x}(\mathbf{r}_0, \mathbf{k}) = \varepsilon^{2/3} F_{u_x u_x}(\mathbf{r}_0, \mathbf{k})$  then,

$$\varepsilon^{2/3} = \frac{\Phi_{vv}(\mathbf{r}_0, \mathbf{k})}{F_{u_x u_x}(\mathbf{r}_0, \mathbf{k}) I_n^2(\mathbf{r}_0, \mathbf{k})}$$

# Remote Sensing of Turbulence: A Primer

- Next, consider the Doppler 2<sup>nd</sup> moment method (think of  $I_n$  as a PDF):

$$\begin{aligned}\langle M_2(\mathbf{r}_0) \rangle &= \int \langle [u(\mathbf{r}) - M_1(\mathbf{r}_0)]^2 \rangle I_n(\mathbf{r} - \mathbf{r}_0) d\mathbf{r} \\ &= \int \langle u^2(\mathbf{r}) \rangle I_n(\mathbf{r} - \mathbf{r}_0) d\mathbf{r} - \langle M_1^2(\mathbf{r}_0) \rangle \\ &= \langle u^2(\mathbf{r}) \rangle - \int R_{uu}(\mathbf{r}, \boldsymbol{\rho}) I_n(\mathbf{r} - \mathbf{r}_0) I_n(\mathbf{r} + \boldsymbol{\rho} - \mathbf{r}_0) d\mathbf{r} d\boldsymbol{\rho}\end{aligned}$$

- Where in the last step it has been assumed that the turbulence is isotropic.

# Remote Sensing of Turbulence: A Primer

- Next, take the inverse Fourier transform of the right-hand side.

$$\langle u^2(\mathbf{r}) \rangle = \int \Phi_{uu}(\mathbf{r}_0, \mathbf{k}) d\mathbf{k}$$

$$\langle M_1^2(\mathbf{r}_0) \rangle = \int \Phi_{uu}(\mathbf{r}_0, \mathbf{k}) I_n^2(\mathbf{r}_0, \mathbf{k}) d\mathbf{k}$$

- Using the same assumptions as above:

$$\varepsilon^{2/3} = \frac{\langle M_2(\mathbf{r}_0) \rangle}{\int F_{u_x u_x}(\mathbf{r}_0, \mathbf{k}) [1 - I_n^2(\mathbf{r}_0, \mathbf{k})] d\mathbf{k}}$$



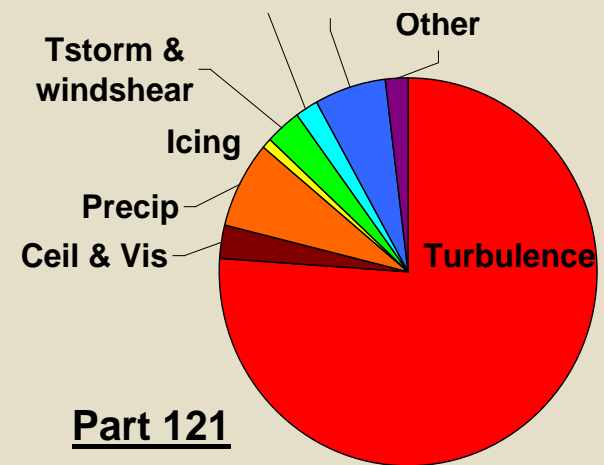
# **Examples of Measurements and Remote Sensing of Turbulence in Support of the Aviation Community.**

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# The Motivation

- Turbulence is the main cause of in-flight injuries – for both passengers and flight attendants.
- After a severe encounter, the airline has to perform a structural check on the aircraft.
- Pilots will try to re-route around an area if there have been reports of moderate or greater turbulence.

**Bottom-line: Turbulence is a safety problem as well as having a large financial impact on the airlines.**



**Weather-related accidents for large transport aircraft**

# DC-8 Cargo Aircraft Damaged Due to Extreme Turbulence



**Missing  
Something?**



# The Turbulence Problem for Aviation (Grossly Oversimplified)

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“Turbulent eddies larger than 100 meters and smaller than 3000 meters (approximately) produce aircraft motions which can be difficult -- or impossible -- to control.

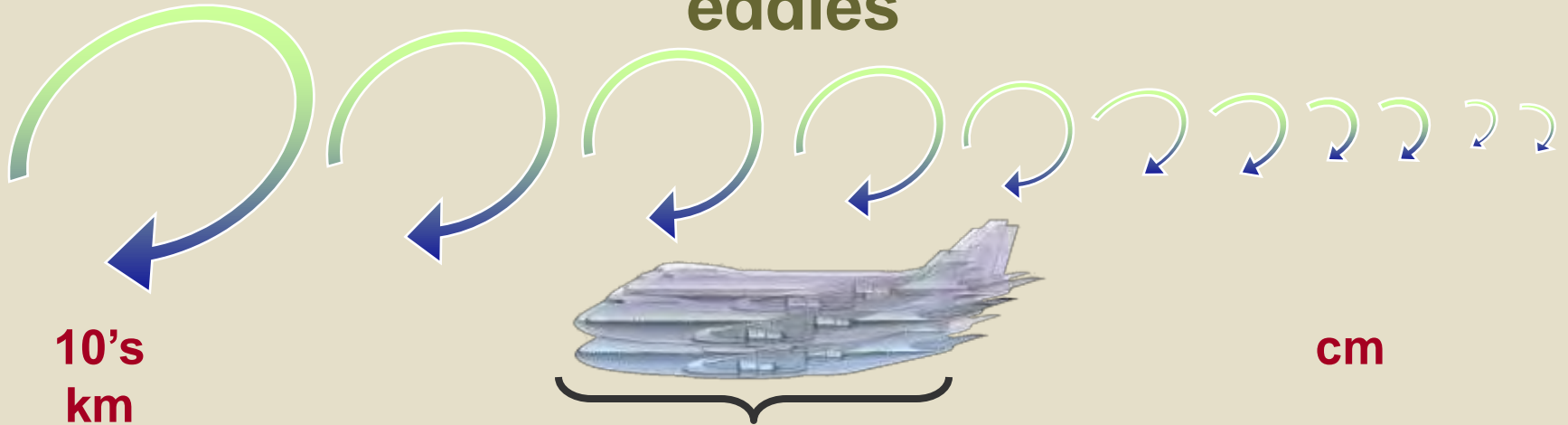
With small-amplitude eddies, these induced motions may be simply uncomfortable to passengers. Large amplitude eddies, on the other hand, can result in passenger injuries or even structural damage to the aircraft.”

# Turbulence Scales of Motion

Large eddies

“Turbulent”  
eddies

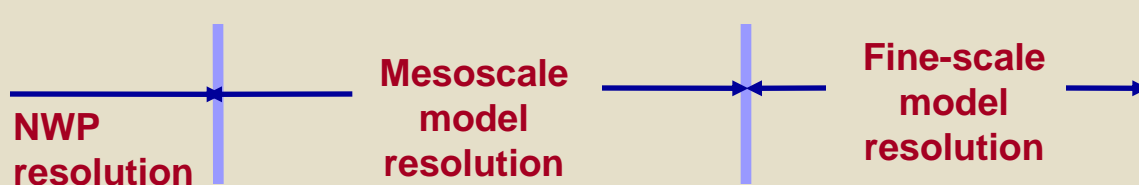
Small eddies



10's  
km

cm

Aircraft responds to scales  
from approx. 100 m – 3 km



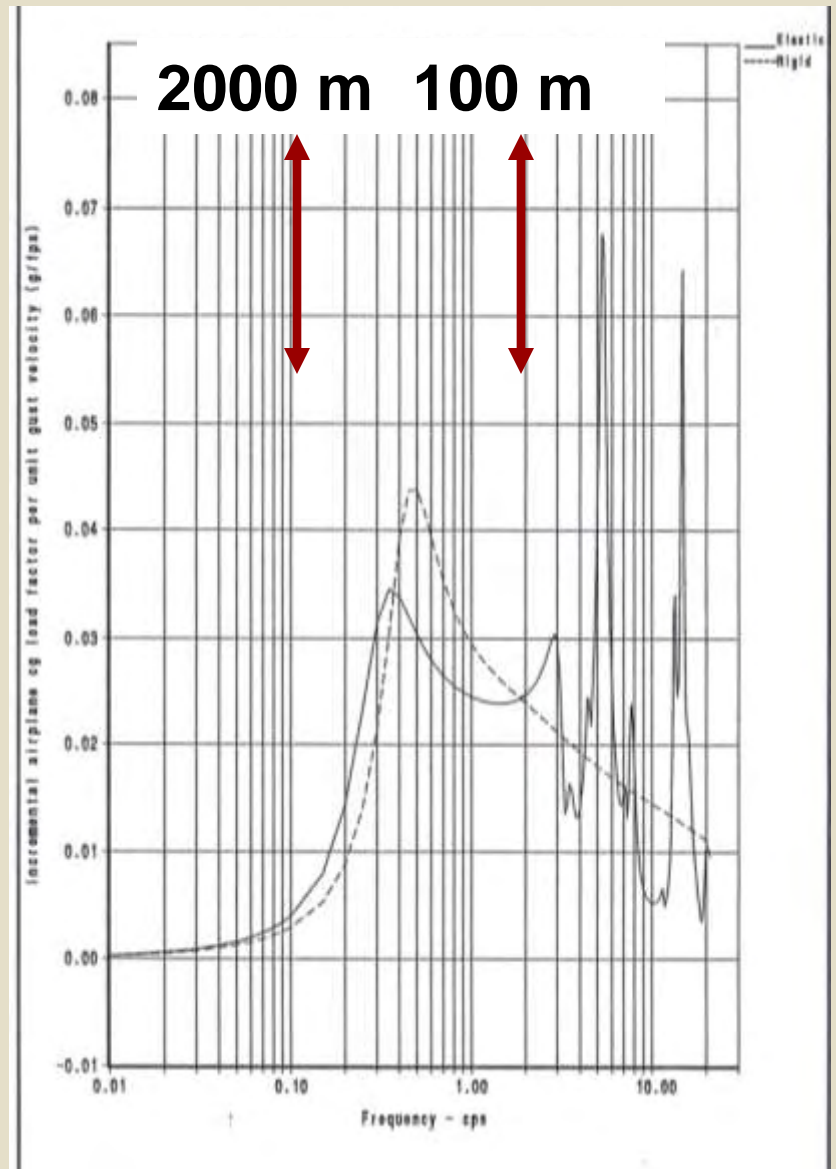
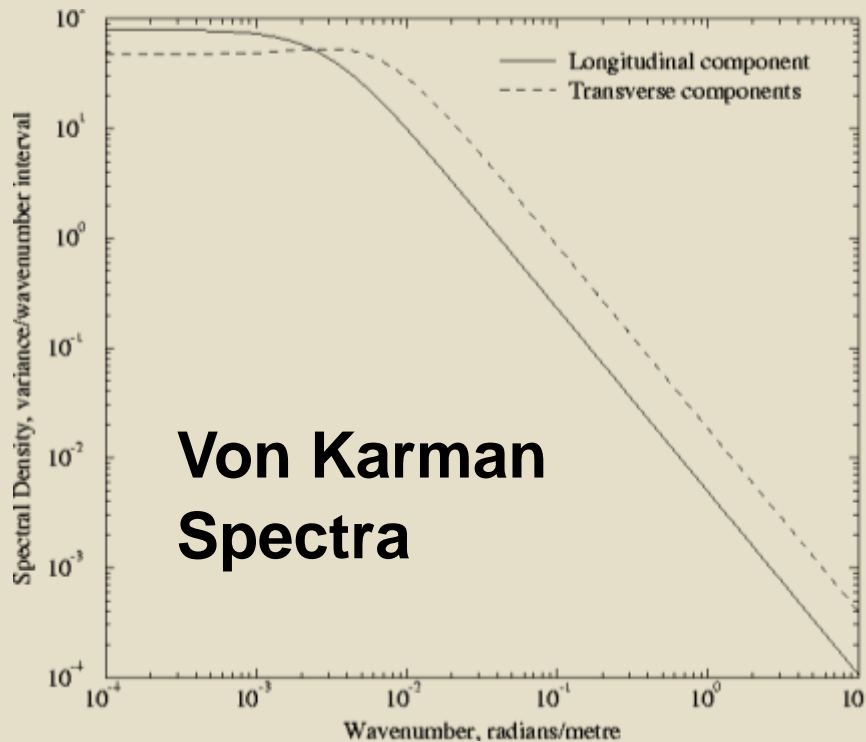
NWP  
resolution

Mesoscale  
model  
resolution

Fine-scale  
model  
resolution

# Response Function for Transport Aircraft

The product of the response function and the wind spectrum gives the output spectrum.



# The Need For Turbulence Measurements

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## ■ Tactical:

- Real time alerts of eminent encounter (< 1 min.)
  - Turn seat belt sign on.
  - Get passengers seated and in seatbelts.
  - Get service carts stowed and flight attendants seated.
- Real time alerts/nowcast of impending encounter (< 15 min.)
  - All of the above.
  - Change altitude.
  - Change flight path.

# The Need For Turbulence Measurements

---

## ■ Strategic:

- Nowcast/Forecast of potential encounter (en route)
  - Increase pilot awareness.
  - Discussions with airline Dispatch personnel.
  - Discussions with en route air traffic personnel.
  - Consider altitude/course change.
- Forecast of potential encounter (pre-flight)
  - Pre-flight awareness for pilot/Dispatch.
  - Consider re-routing flight path.

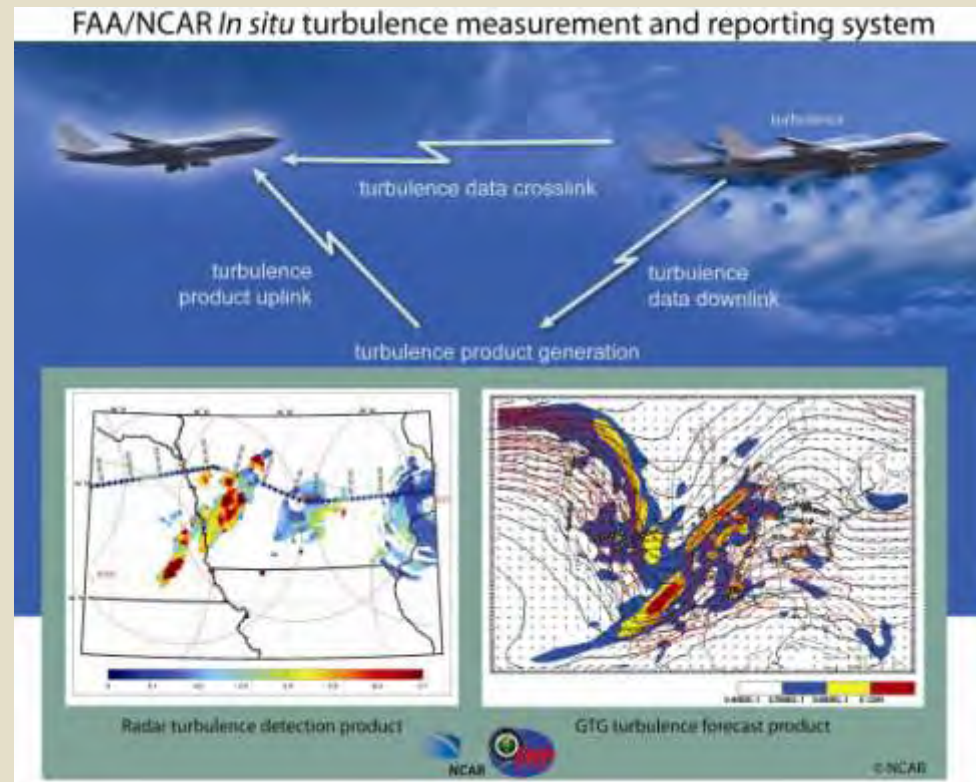


# *In situ* Turbulence Measurement and Reporting System

**Goal:** To augment/replace subjective PIREPs with objective and precise turbulence measurements.

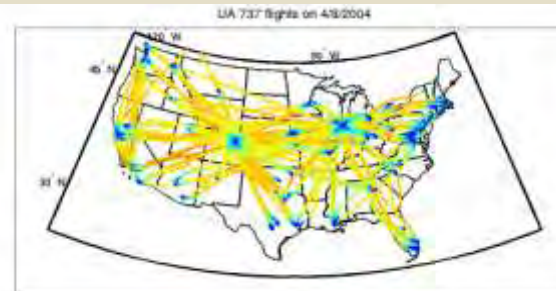
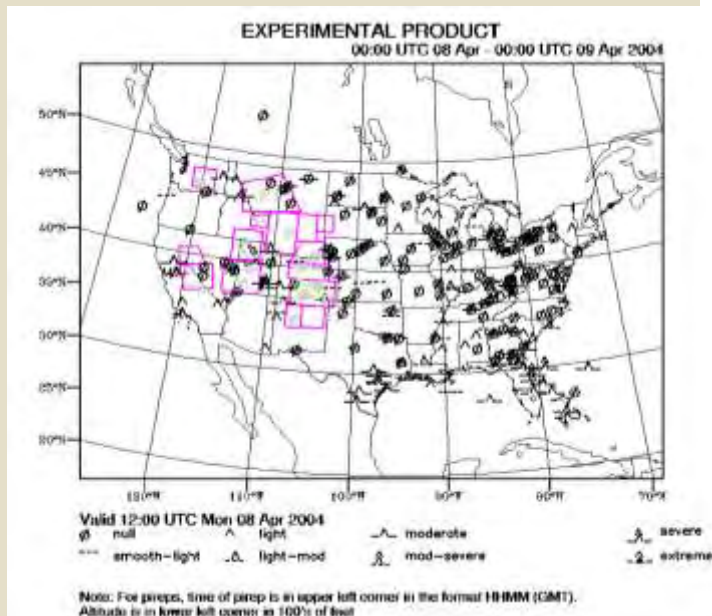
**Features:**

- Atmospheric turbulence metric: eddy dissipation rate (EDR).
- EDR can be scaled into aircraft turbulence response metric (RMS-g).
- **Adopted as ICAO Standard**

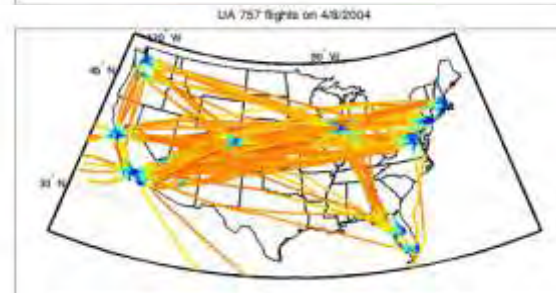


# Increase in Spatial/temporal Coverage: UAL EDR Reports Compared to pireps

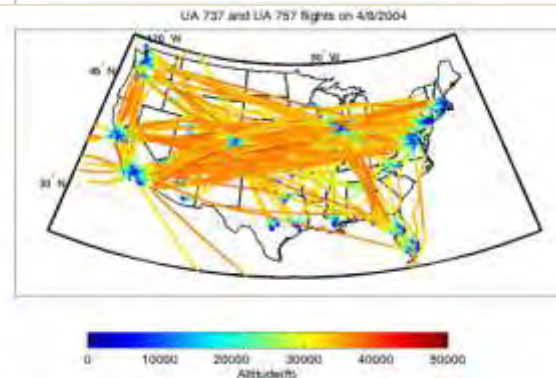
1.3 million EDR reports/month from ~~100 or so aircraft~~ - compared to 55k pireps from all aircraft.



737



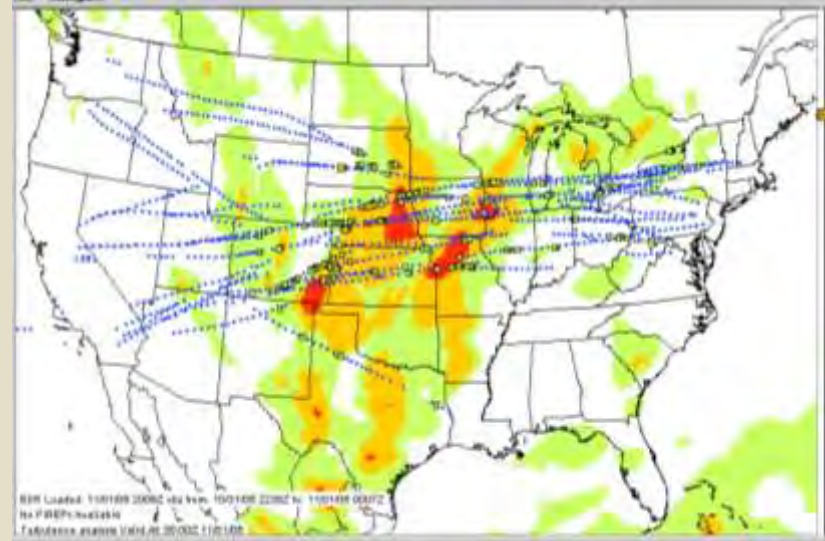
757



737  
+  
757

# *In situ* measurement and reporting system

- Implemented on ~ 200 UAL aircraft since 2000
- Implementation status
  - SWA: Fleet implementation on ~ 280 737-700s in CY08
  - DAL: Fleet implementation on ~ 120 737-800s in CY08
  - NWA: Discussions ongoing for implementation on ~ 140 Airbus 319/320s and 56 787s
  - UAL: 757 ACMS replacement
  - AAL: Discussions ongoing



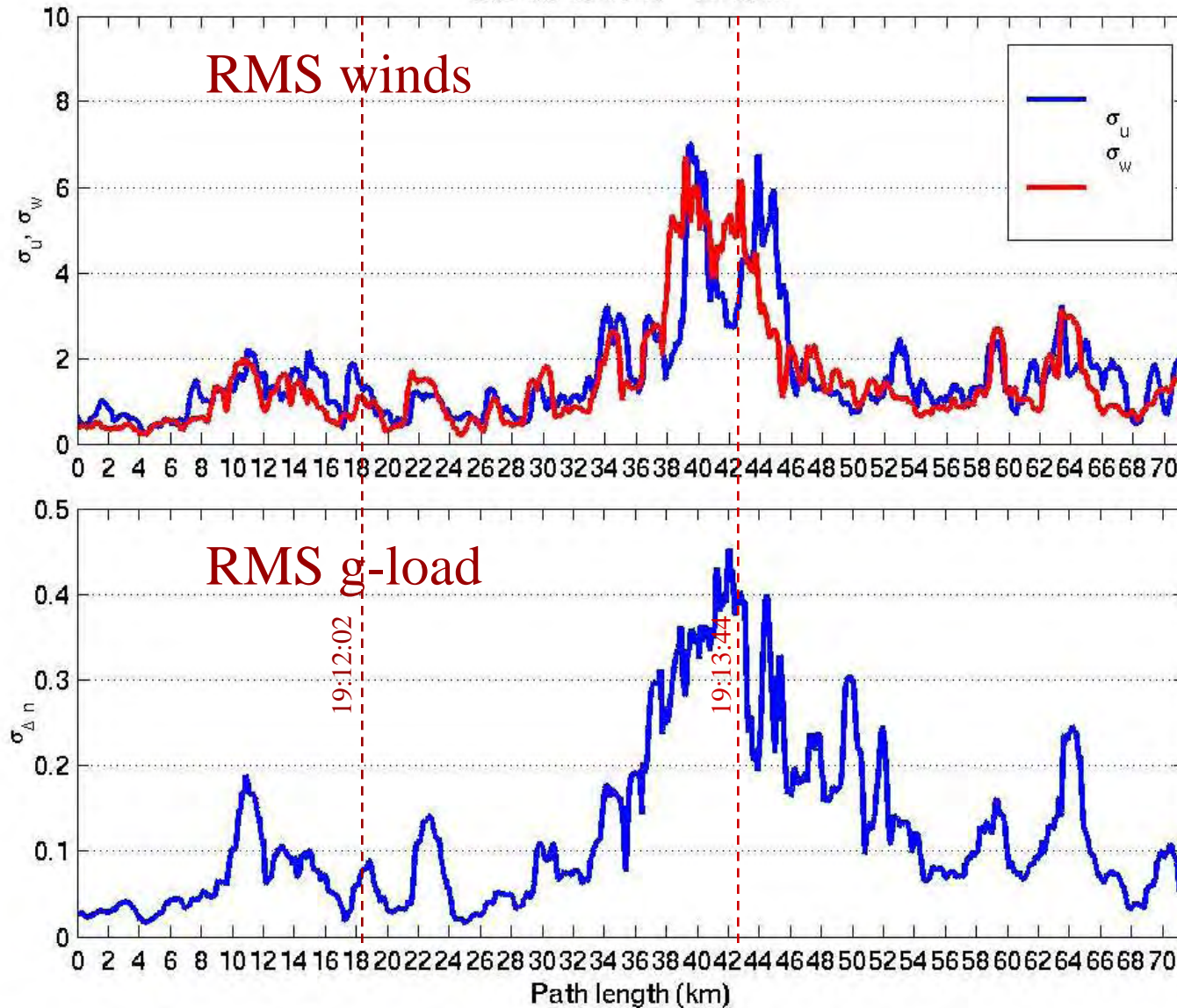
**Website: UAL 757 edr flight tracks overlaid on GTG forecasting product**

# **Radar Measurements of Turbulence**

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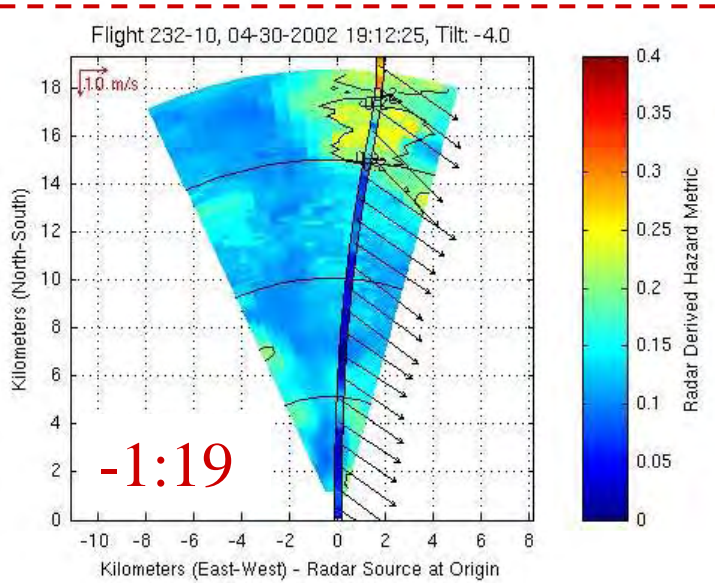
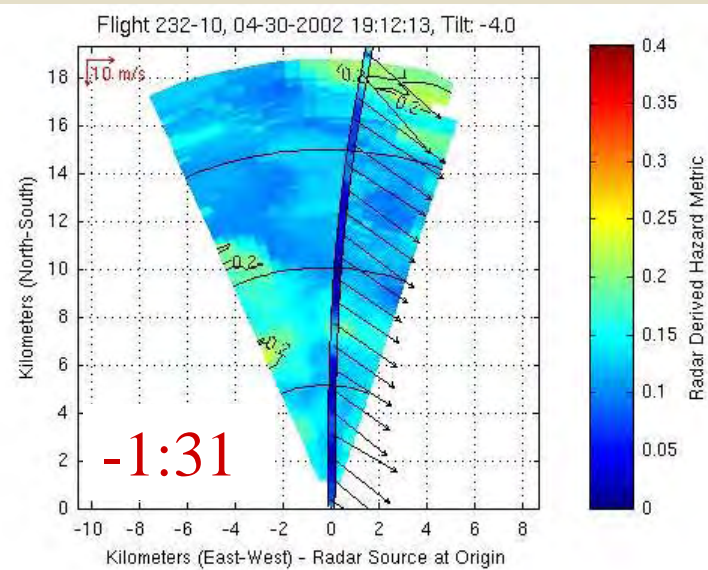
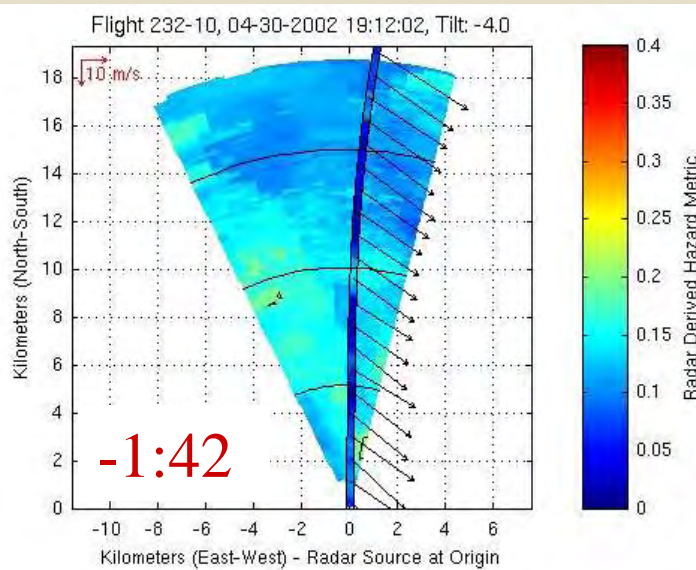
# NASA Airborne Radar Detection of Turbulence Program

232-10: 19:11:10 - 19:16:14



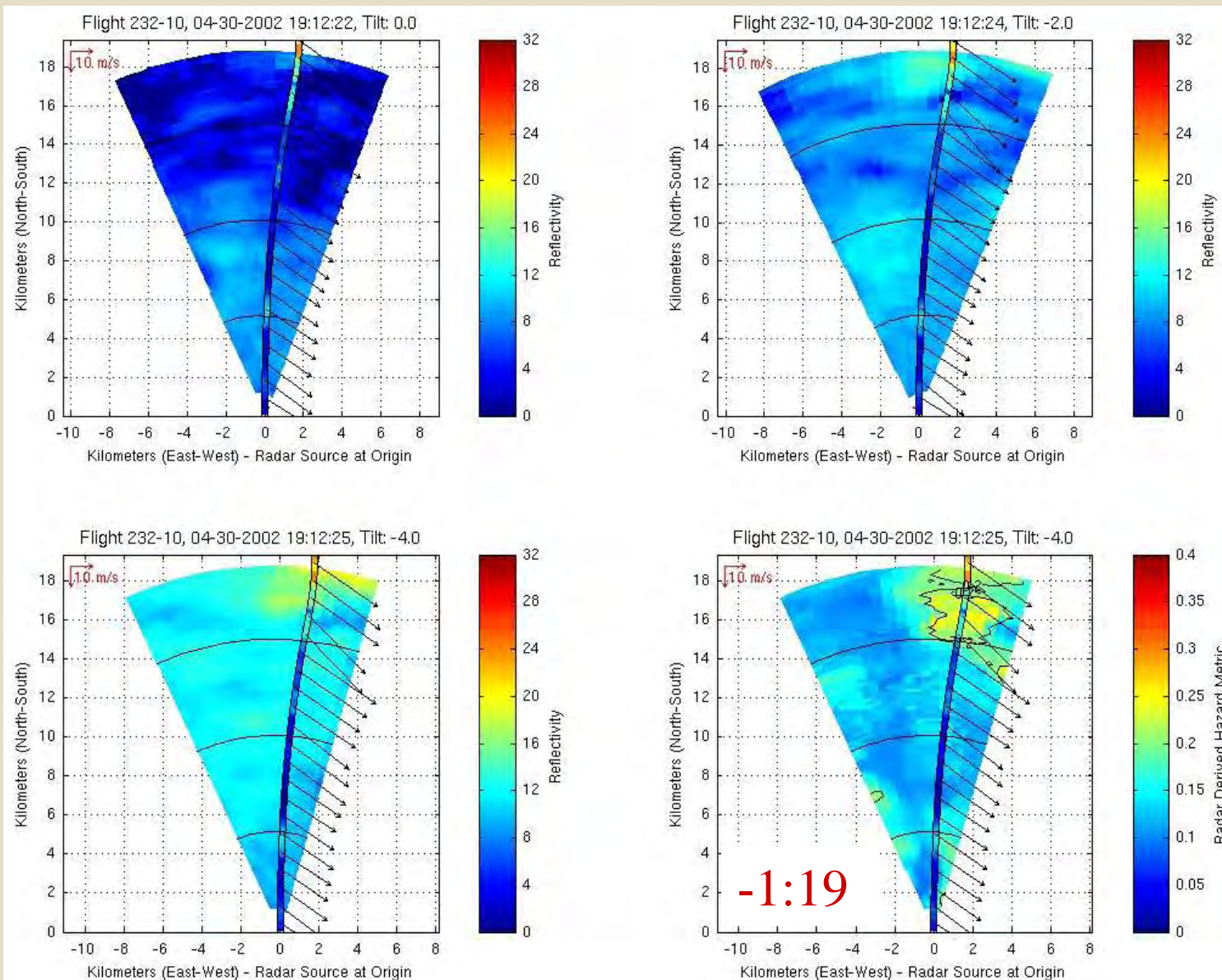
From  
NASA  
B-757  
Aircraft

# Event 232-10 (19:12:02, 19:12:13, 19:12:25)

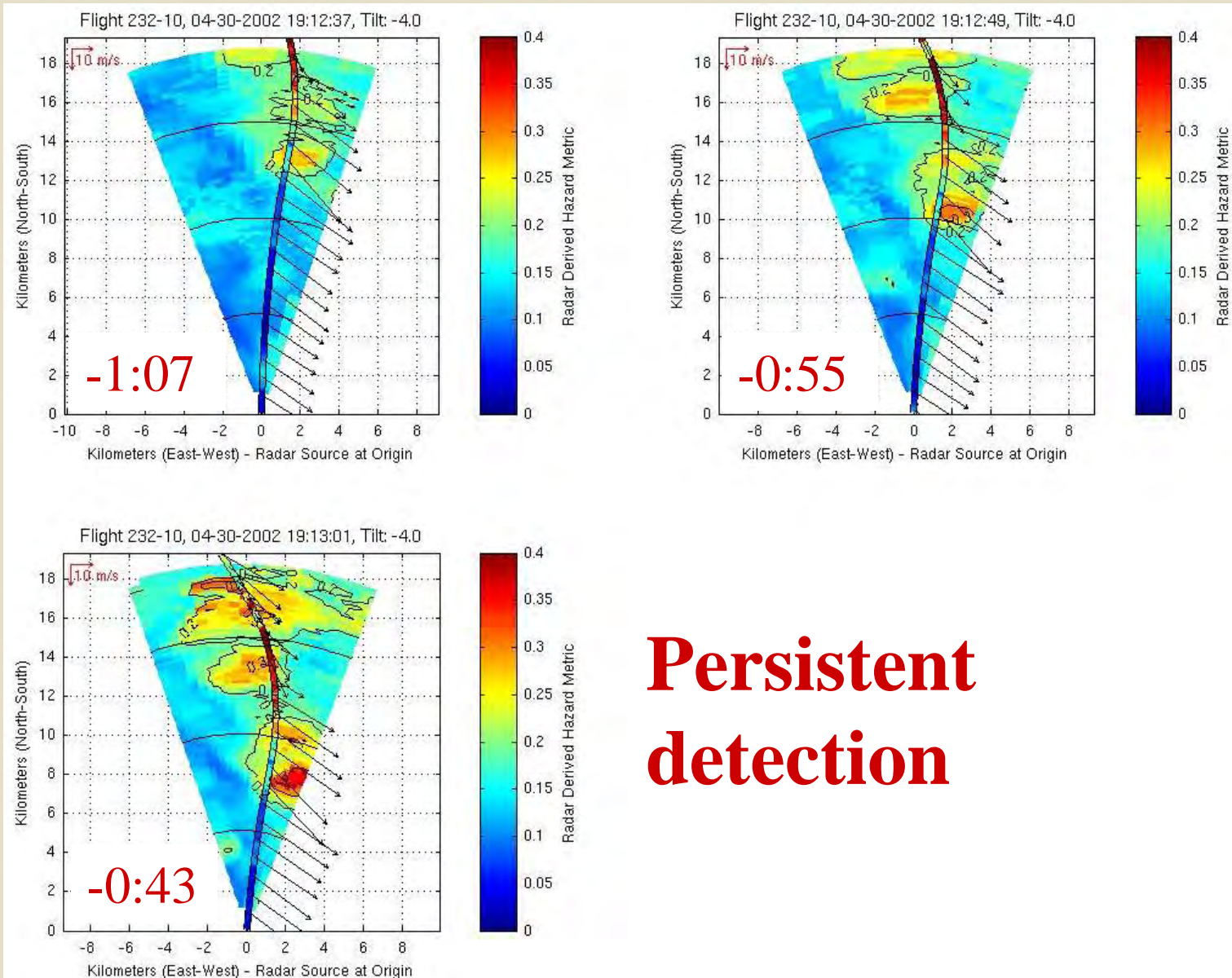


**Hazard detected**  
**1:19, 18 km to**  
**encounter**

# Event 232-10 (reflectivities at 19:12:25)

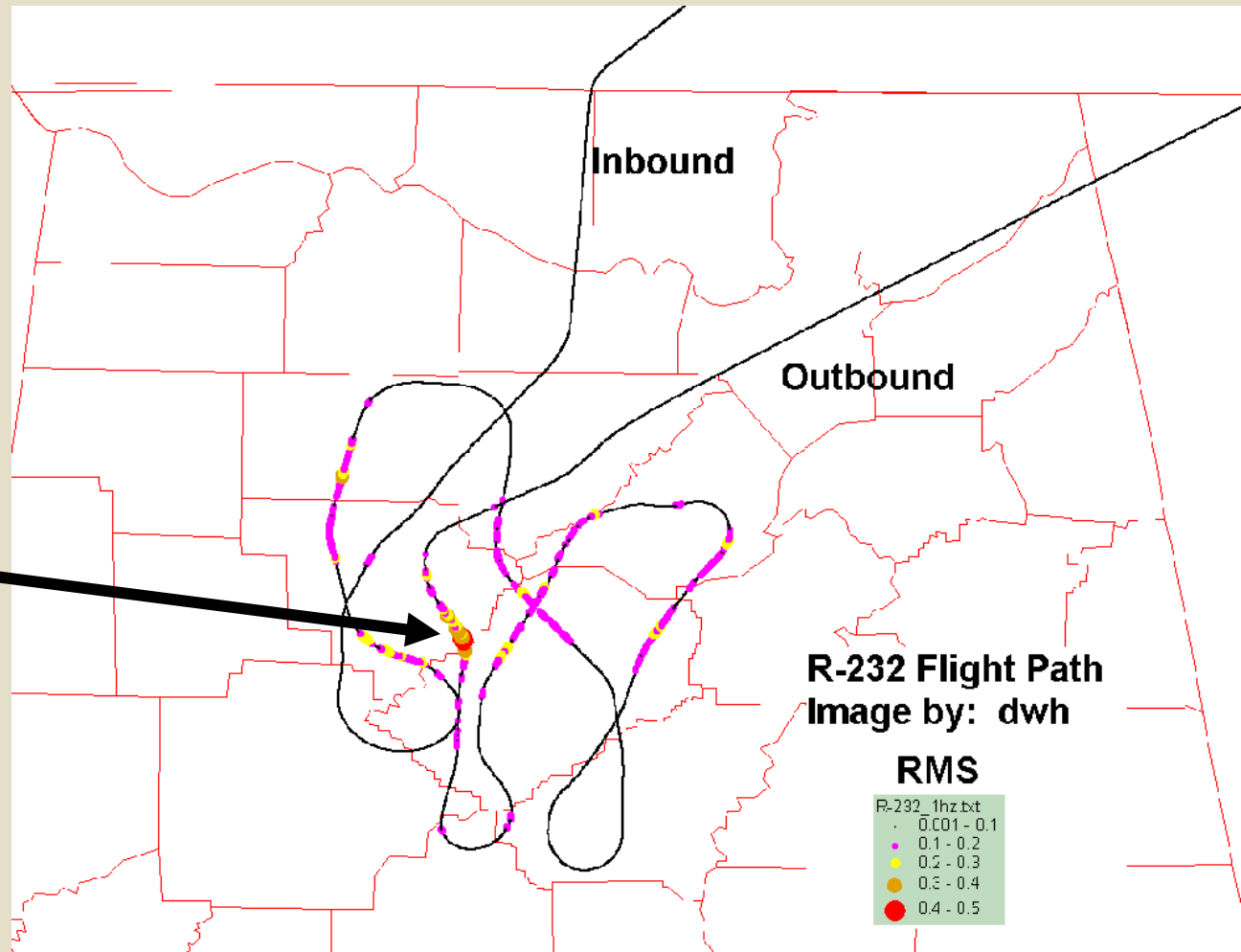


# Event 232-10 (19:12:37, 19:12:49, 19:13:01)





# Flight Track for NASA flight R232



**Event  
R232-10**

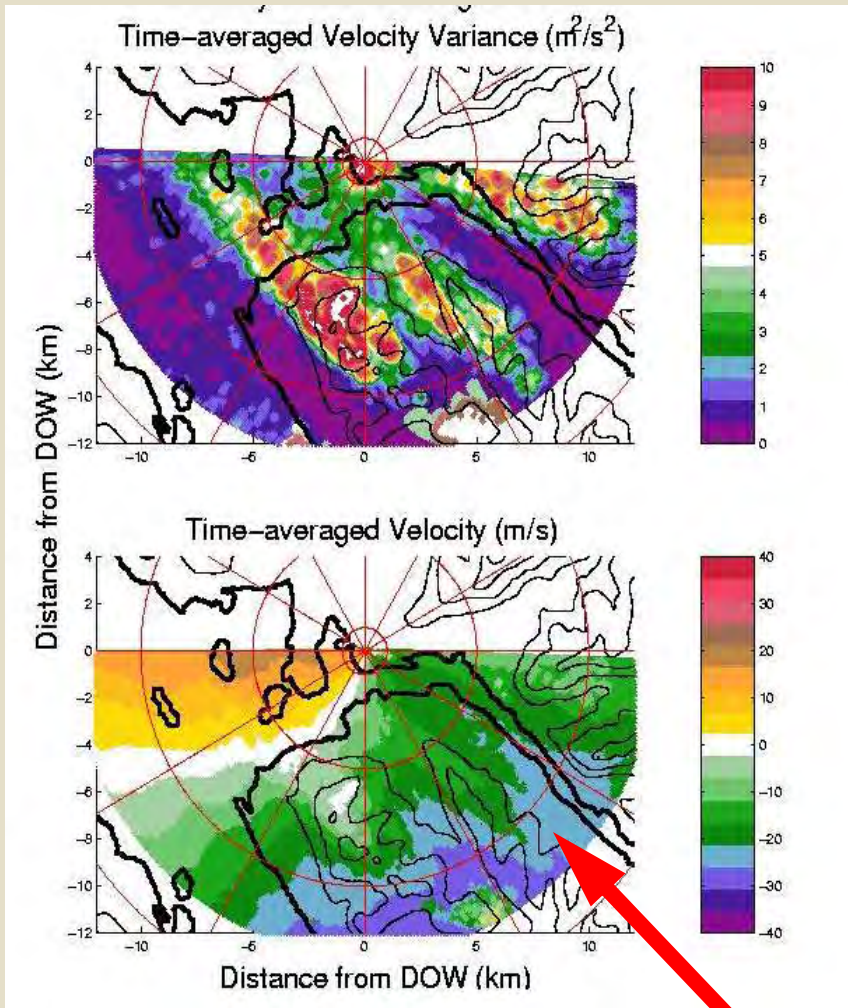
**North-Central  
Alabama**

# FY00 Juneau Field Project Equipment

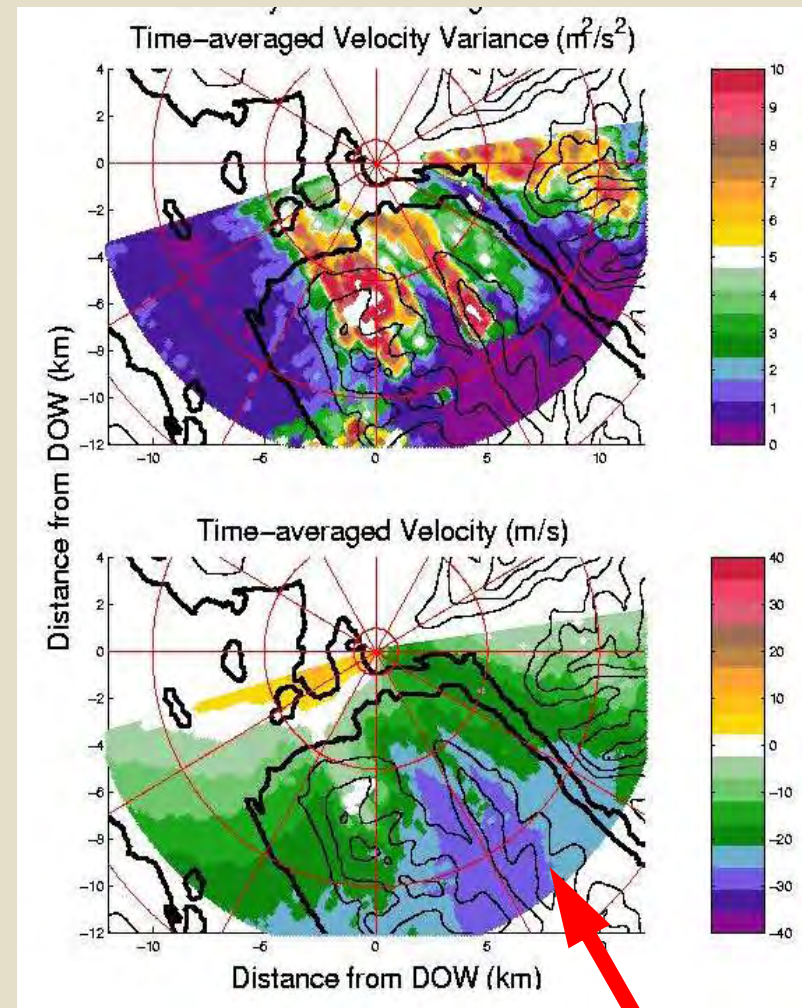
## Doppler on Wheels (DOW)



# DOW Radar Data Showing Terrain-Induced Turbulence for Two Different Wind Directions



**Wind Direction**



**Wind Direction**

# Detection of Turbulence Using an Airborne Forward-Looking IR Sensor

## ■ Possible Approach:

- Derive equation that relates the statistics of the atmospheric turbulence (e.g., temperature field) to those of the sensor measurables.
- Consider the irradiance ( $H$ ) at a given frequency, measured at the aircraft ( $x=0$ ):

$$H(0) = \int_L^0 W_B(T(x)) f(x) dx$$

Where  $W_B$  is the Planck function,  $L$  is the path length over which the measurement is made, and  $f$  is a combined function of the non-turbulent atmosphere and the response characteristics of the sensor.

# Detection of Turbulence Using an Airborne Forward-Looking IR Sensor

- Next, consider the same measurement when the aircraft has moved a distance  $\rho$  :

$$H(\rho) = \int_{L+\rho}^{\rho} W_B(T(x))f(x)dx$$

- Assuming that the Planck function is linear in the temperature, the correlation function of the irradiances can be computed:

$$\langle H(0)H(\rho) \rangle = K \int_L^0 \int_{L+\rho}^{\rho} \langle T(x)T(x') \rangle f(x)f(x')dx dx'$$

# Detection of Turbulence Using an Airborne Forward-Looking IR Sensor

- Assuming that standard turbulence theory applies,  $\langle T(x)T(x') \rangle = C_T^2 g(x' - x)$ , where  $C_T^2$  is the intensity parameter of the turbulent temperature field. In principle the turbulence intensity parameter is given by:

$$C_T^2 = \frac{\langle H(0)H(\rho) \rangle}{K \int_L^0 \int_{L+\rho}^\rho g(x'-x) f(x) f(x') dx dx'}$$

# Detection of Turbulence Using an Airborne Forward-Looking IR Sensor

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## ■ Issues:

- Aircraft respond to vertical wind motions, not temperature fluctuations - the relationship between the two is not well-understood.
- To what spatial scales are these IR devices sensitive?

# **Turbulence Detection via Airborne GPS Receivers**

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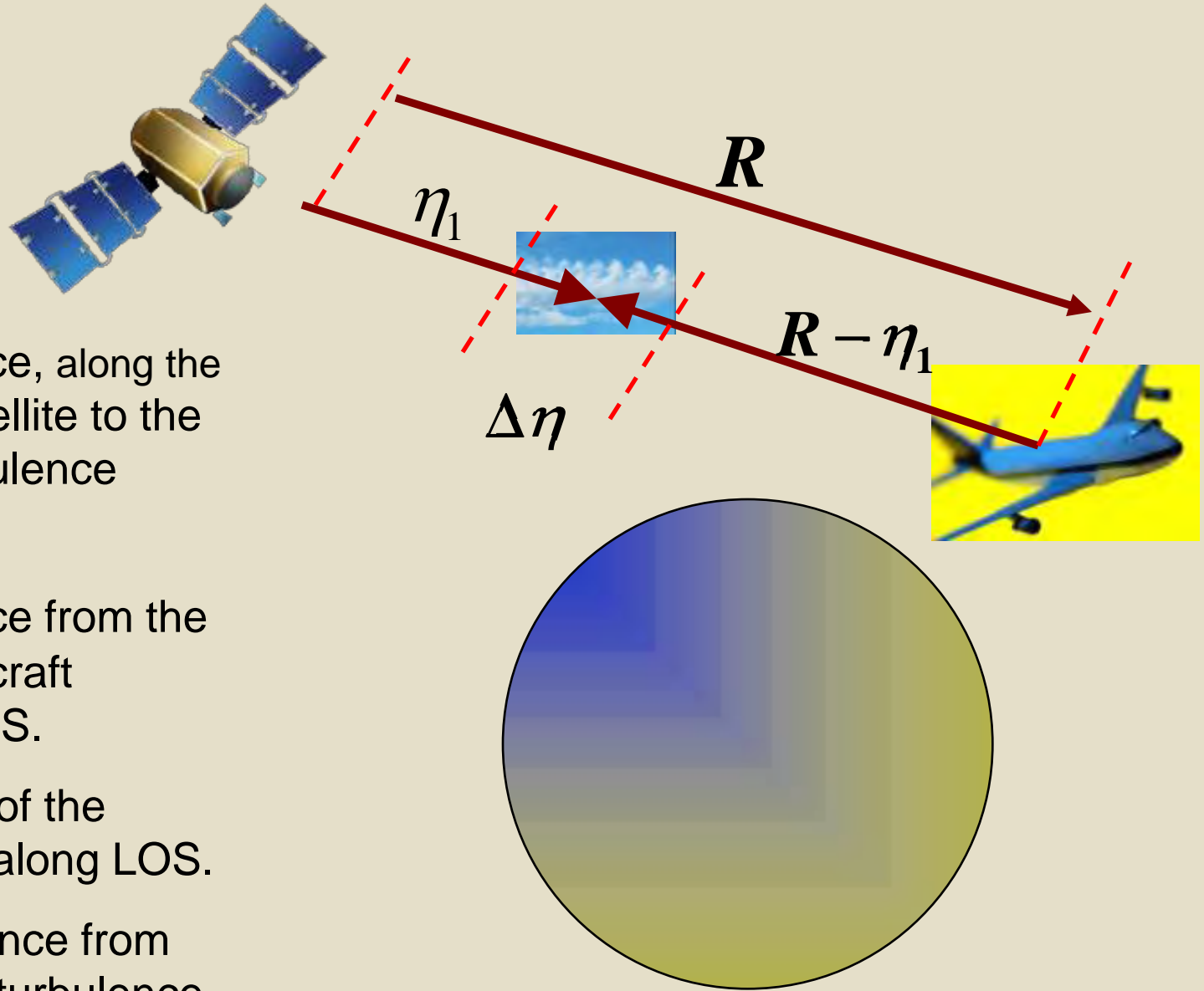


# Turbulence Detection via Airborne GPS Receivers: The Concept

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- Airborne receivers would be a platform of opportunity to collect occultations in the cruise regime of commercial aviation, e.g., 20-40 kft. AGL.
- The turbulence measurements from these occultations would probably not be used as stand-alone information, but integrated into operational nowcast/forecast products.

# Geometry of the Problem



$\eta_1$  is the distance, along the LOS, from the satellite to the center of the turbulence patch.

$R$  is the distance from the satellite to the aircraft receiver along LOS.

$\Delta\eta$  is the width of the turbulence patch along LOS.

$R - \eta_1$  is the distance from the aircraft to the turbulence patch.

# Theory (Condensed Version)

- Use standard weak scattering wave propagation theory - modified to accommodate moving transmitter and receiver, and localized turbulence patch.
- Log-amplitude frequency spectrum for a turbulence patch in the neutral atmosphere at  $\eta_1$  and width  $\Delta\eta$  is functionally given by:

$$\Phi_{\chi}(f) = C \int_{\eta_1 - \Delta\eta/2}^{\eta_1 + \Delta\eta/2} F \left[ C_n^2(\eta), V_{eff}(\eta), L_0, R, k, f; \eta \right] d\eta$$

- Where  $k$  is the transmitter wavenumber,  $L_0$  is the turbulence length scale,  $C_n^2(\eta)$  is the turbulence intensity, and  $V_{eff}(\eta)$  is an effective velocity.

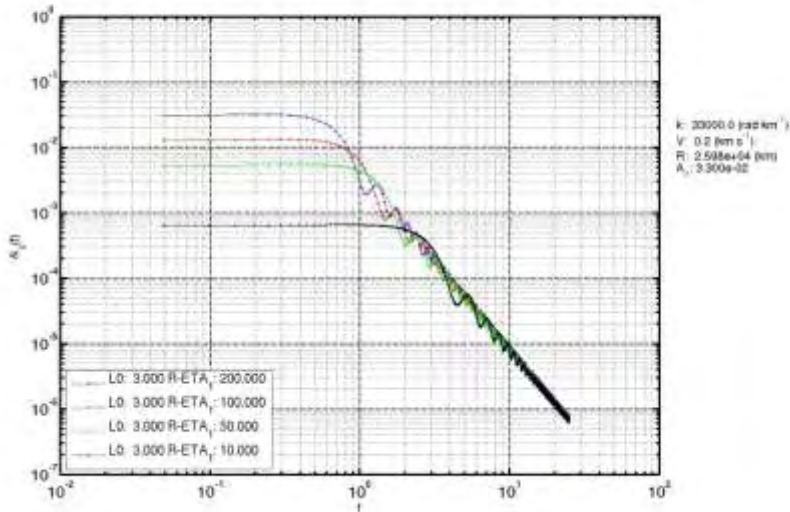
# Analysis

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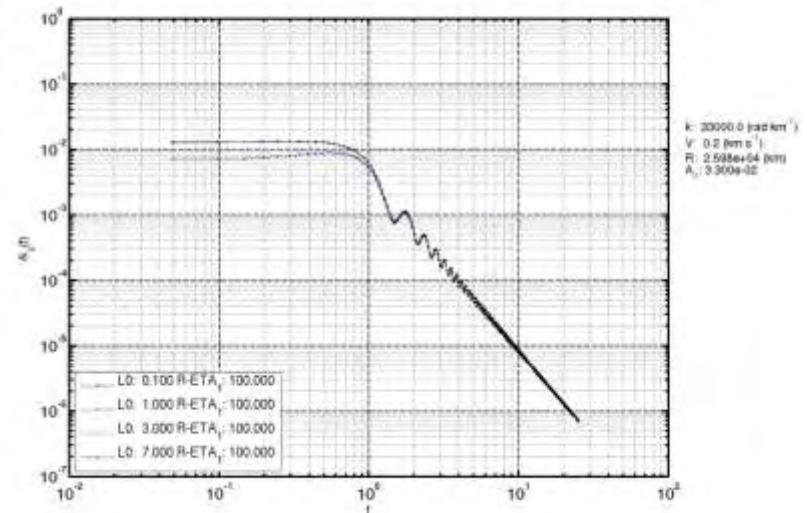
- The objective is to determine where the turbulence is along the LOS,  $\eta_1$  , and what the intensity  $C_n^2(\eta_1)$  , is at that location.
- Note that the comparable expression for turbulence in the ionosphere is given by a different constant and a change in the wavenumber functionality  $k^2$  to  $k^{-2}$  .

**In the following, a mid-point approximation to the integral was used.**

# Example Spectra (constant intensity)



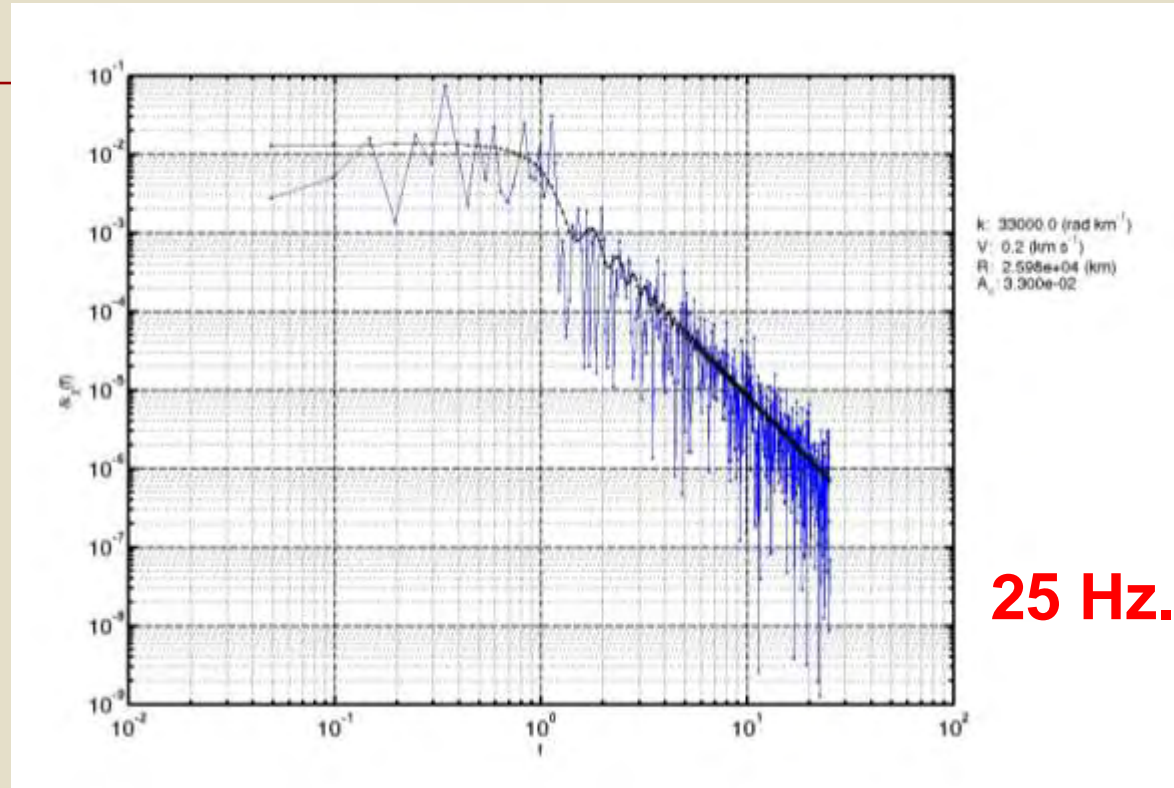
Holding  $L_0 = 3 \text{ km}$  fixed and varying  $R - \eta_1$ .



Holding  $R - \eta_1 = 100 \text{ km}$  fixed and varying  $L_0$ .

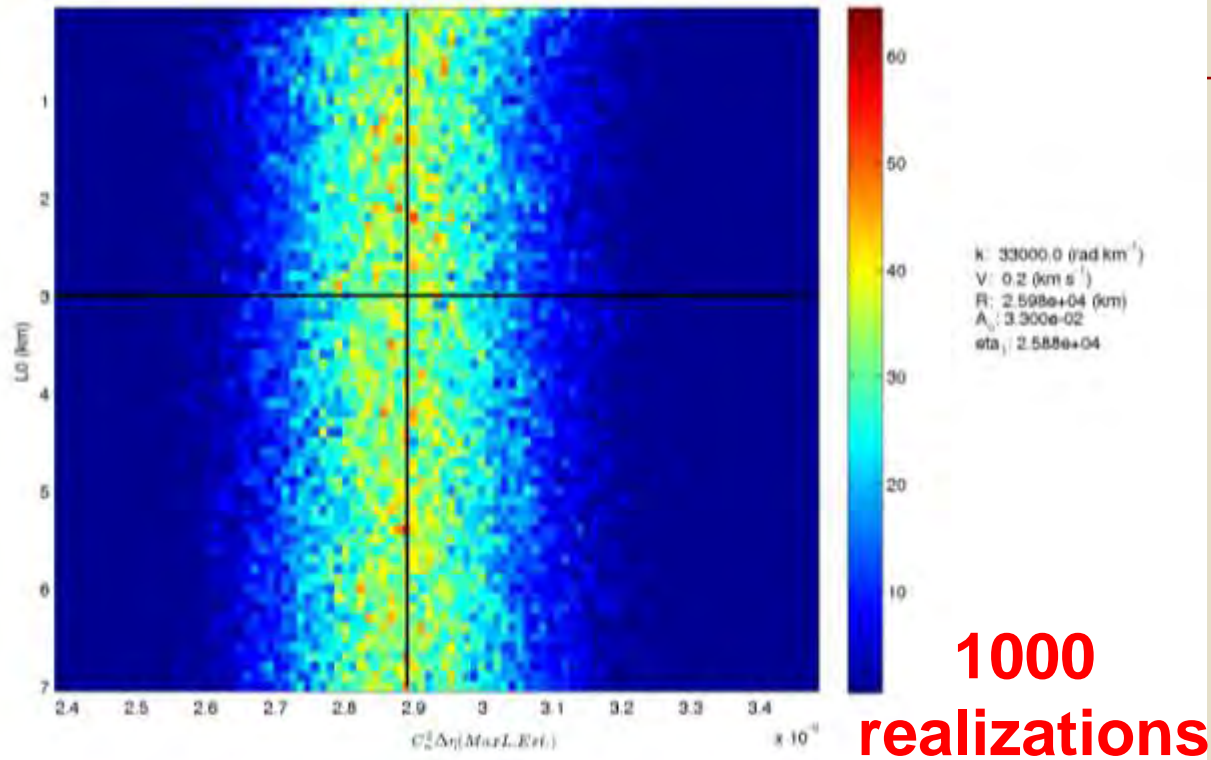
**Note strong functionality on  $R - \eta_1$  and weak on  $L_0$ .**

# Spectral Simulation



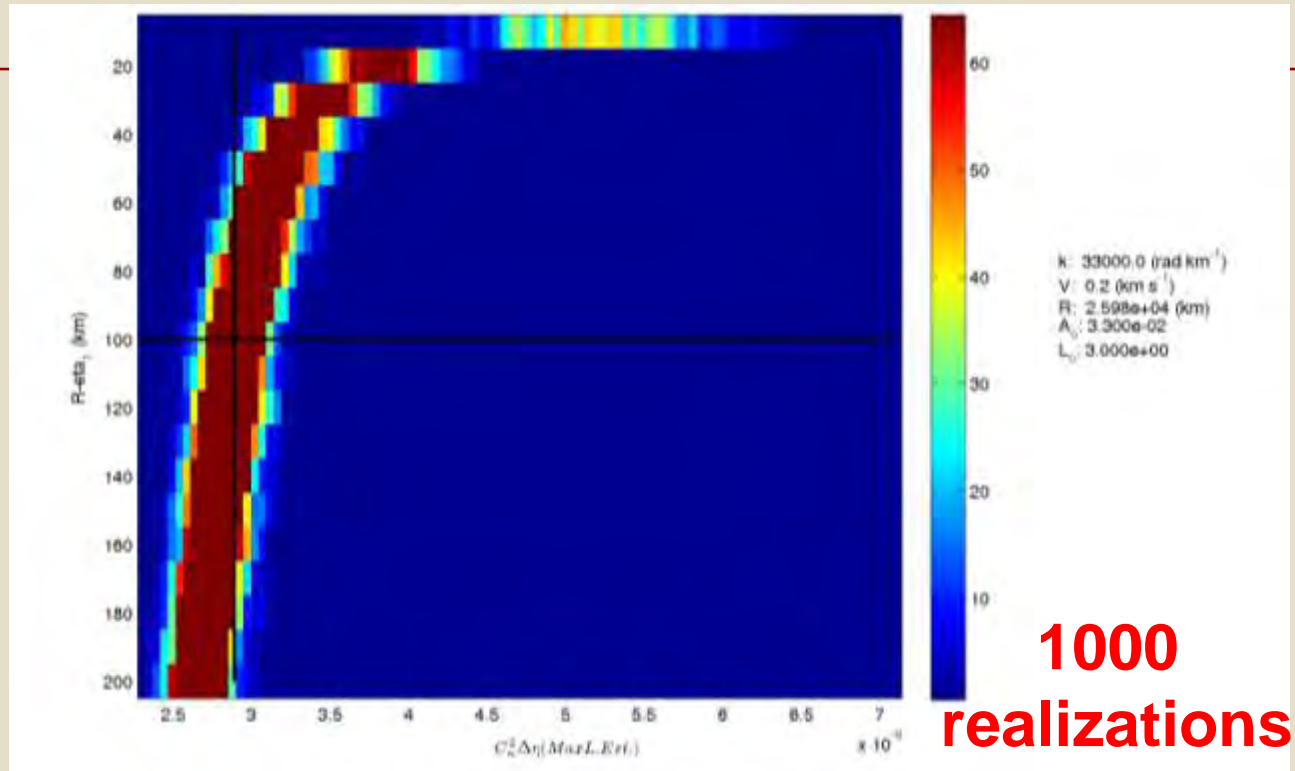
**Log-Amplitude Spectrum: theory and simulated.**

# Parameter Estimation



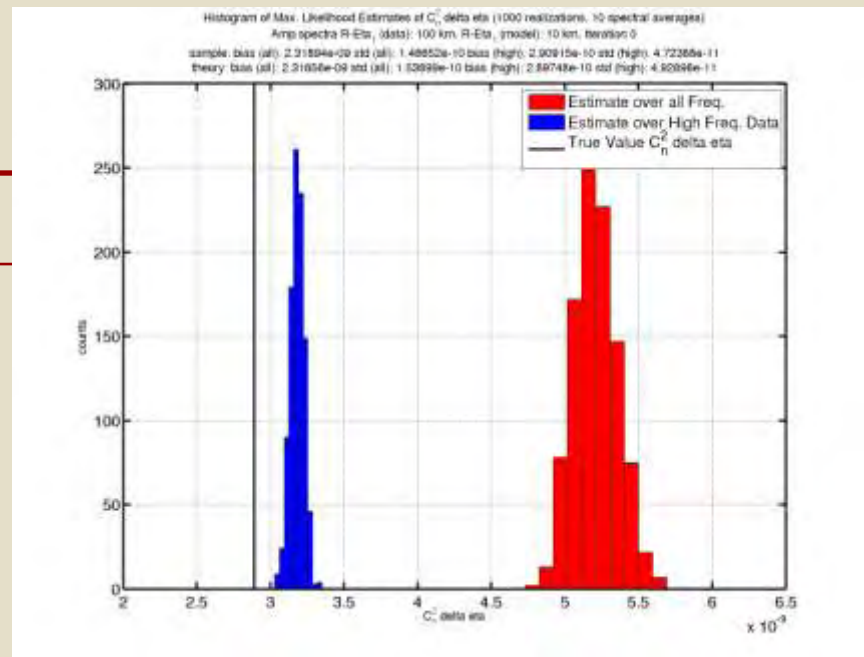
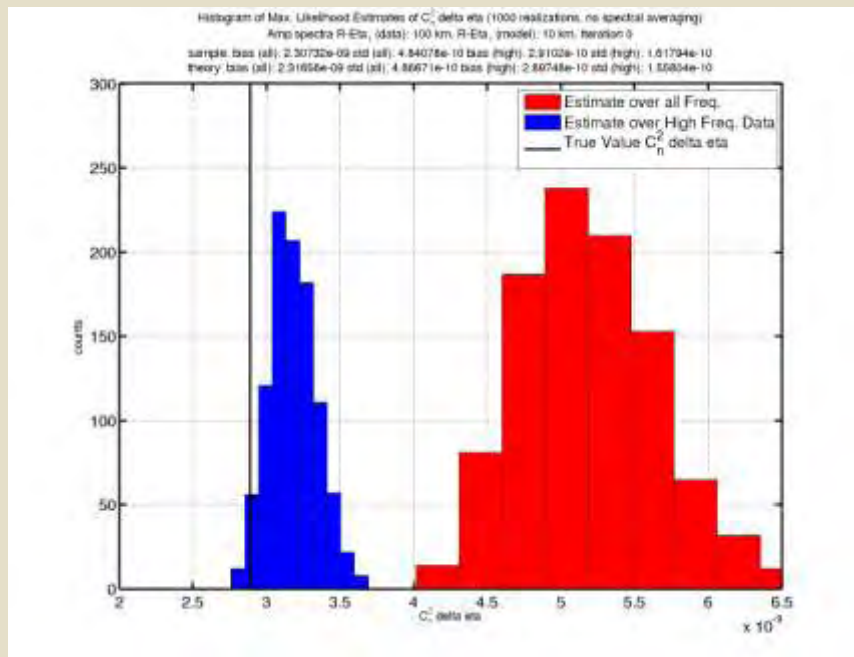
Maximum Likelihood (ML) estimation of intensity with fixed  $R - \eta_1 = 100 \text{ km}$  while varying  $L_0$ . Horizontal and vertical lines are simulated (i.e., “true”) values.

# Parameter Estimation (cont'd)



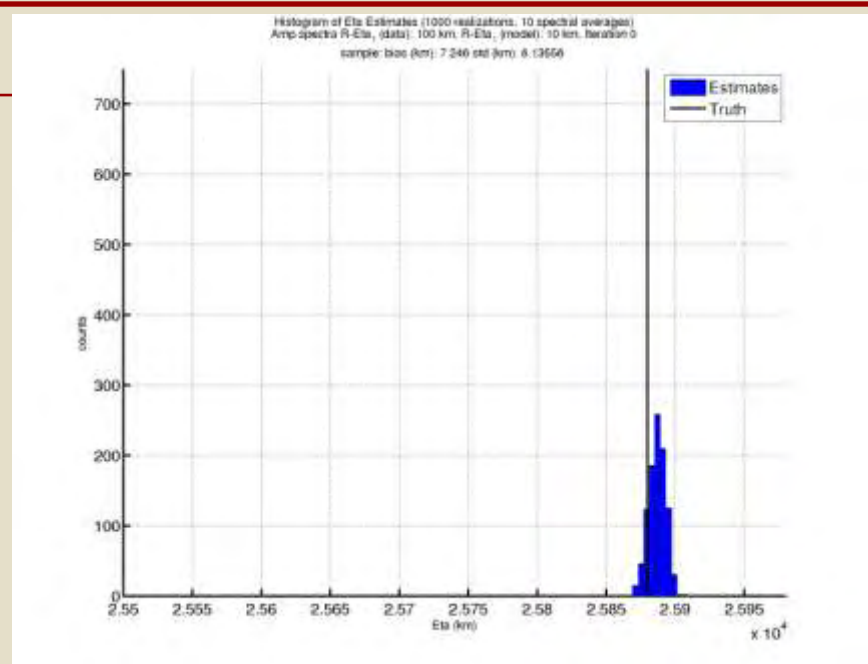
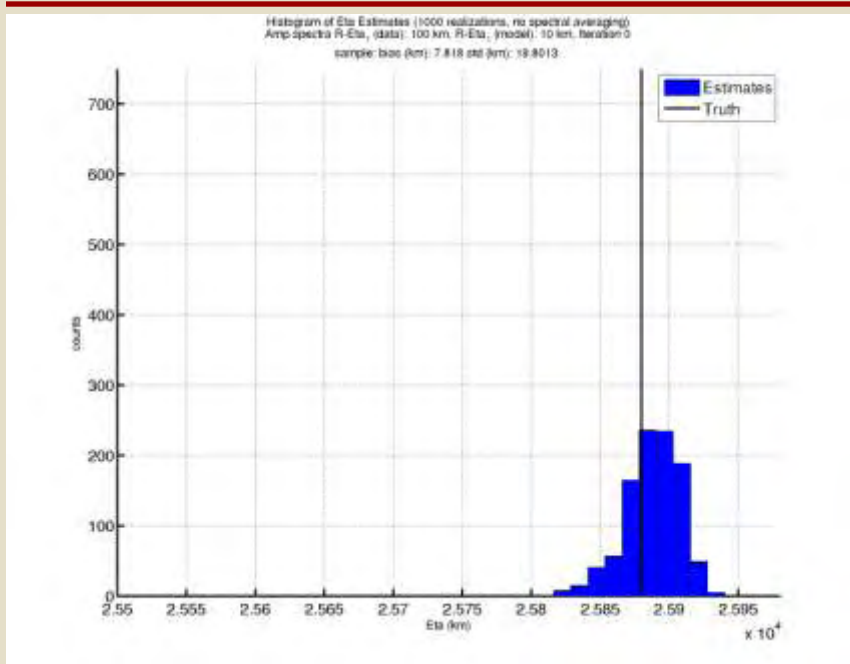
ML estimation of intensity with fixed  $L_0 = 3 \text{ km}$  while varying  $R - \eta_1$ .





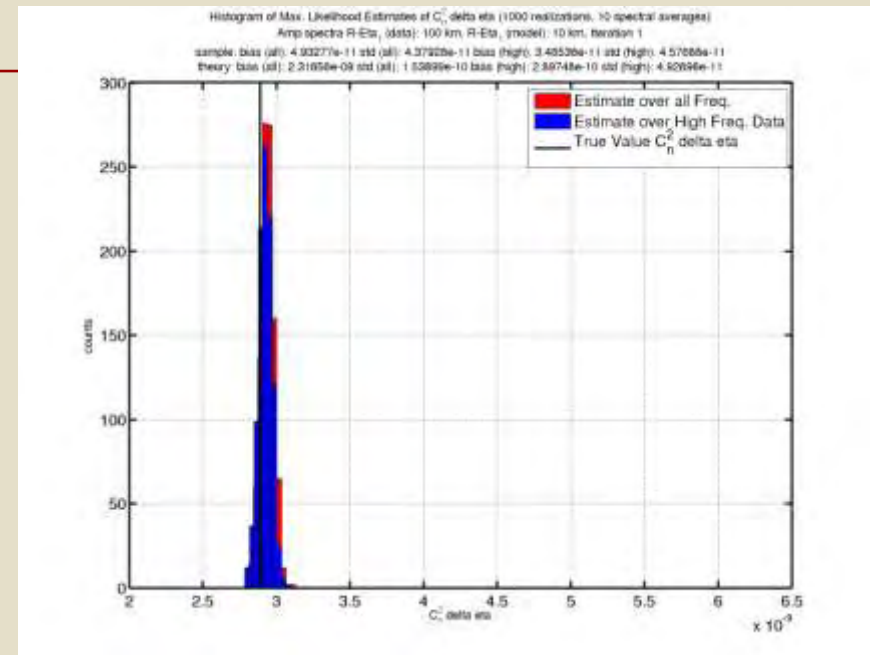
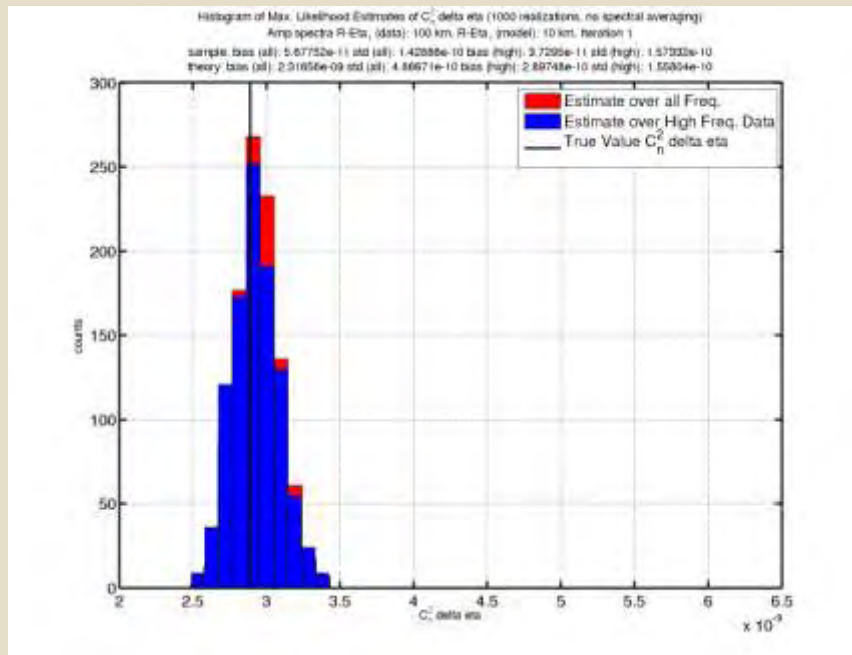
- ML estimation of intensity for un-averaged spectra (left) and 10 spectral averages (right).
- Solid vertical line is true value.
- **Blue** values are from using the high-frequency portion of the spectrum, **red** values use all the spectral points.
- ~~Simulated  $R - \eta_1$  value is 100 km, “guess” is 10 km - i.e., underestimate.~~

# Iterative Parameter Estimation



**Estimation of  $\eta_1$  by minimization method. Un-averaged spectra (left) and 10 spectral averages (right).**

# Iterative Parameter Estimation



**ML estimation of intensity after using  $\eta_1$  estimates. Un-averaged spectra (left) and 10 spectral averages (right).**

# Summary

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- Turbulence measurements are critical in providing accurate and operationally useful tactical and strategic information to users.
- *In situ* turbulence measurements are now available and used operationally – more to come.
- A number of proof-of-concept sensor demonstrations have occurred, with positive results.
- Other technologies in development/evaluation:
  - Airborne lidar
  - Satellite
  - GPS/Iridium
  - Airborne IR